

Geological Control upon Groundwater Flow and Major Ion Chemistry with Influence on Basin Management in a Coastal Aquifer, South Australia

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Received August 7, 2013; revised September 11, 2013; accepted October 6, 2013

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ABSTRACT

Estimation of natural recharge and potential for seawater intrusion are critical considerations for management of coastal freshwater aquifers. We show hydrochemical signatures of groundwater to identify the influence of geological control on chemical processes in a coastal groundwater system. We used dominant hydrochemical facies, salinity and magnesium ions to determine two main groundwater flow paths with different origins and ages. Mixing of groundwater with different origins and ages results in unreliable recharge estimates using chlorofluorocarbon (CFC) and chloride mass balance (CMB) methods, thus limiting available methods for recharge assessment. Interpretation of hydrochemical data suggests that calcium carbonate dissolution, ion exchange processes and mixing with sea aerosol in coastal zones are the main influencing factors on groundwater chemistry. Restricted groundwater flows due to occurrence of a basement high at the southern side of the basin boundary influence the distance to the toe of the saline wedge. Thus, knowledge of geological control over groundwater systems forms an important part of characterising basins and contributes toward effective management of groundwater resources.

Keywords: Hydrochemistry; Groundwater Mixing; Groundwater Management; Recharge

1. Introduction

Spatial distribution of groundwater chemistry is primarily influenced by mineralogical composition of sediments in contact with groundwater [1]. The nature and distribution of aquifers and aquitards in a geologic system are controlled by the lithology, stratigraphy and structure of geologic formations. The main geological controls are the stratigraphy and structural features such as cleavages, fractures, faults, and solution features which are the geometrical properties of the geologic systems. The influence of geological control changes natural water chemistry and therefore knowledge of main hydrochemical processes that control groundwater chemistry is a prerequisite for effective basin management [2]. Several studies examined the influence of geological controls: the presence of faults [3,4]; geology and structure of bedrock [5,6]; fracture control [7] which influences groundwater flow and water chemistry; and hydrogeological study of

groundwater flow diversion caused by anticlinal structures [8].

Among geological controls, the presence of faults and their influence on flow paths and water chemistry are widely reported and these are known to be effective flow barriers [3]. Green *et al.* [4] reported flows across regional scale faults and found that groundwater flow either enhanced or inhibited, depended on orientation of fractures with respect to the cross fault hydraulic gradient. Bense *et al.* [9] and Meeder [10] related the impact of faults on near surface processes to surface features such as vegetation patterns and drainage networks; and thus, faults can have great influence on transport processes in the sub-surface and are important in water management.

Eaton *et al.* [7] noted that there are few studies on the hydrogeology of sedimentary rock aquitards, although they are important controls in regional ground water flow systems. Cherry *et al.* [11] highlighted that determination of aquitard integrity is an important aspect of groundwater management because it influences water quality. Thus,

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the presence or absence of an aquitard, which is not necessarily a confining unit, can have a significant impact on basin water quality. We examined effects of basement rock outcrop and discontinuity of the clay aquitard on groundwater salinity, water chemistry, influence on recharge estimation methods and extent of the saline wedge in a coastal aquifer in South Australia.

2. Study Area

The Uley South basin is located in the Southern Eyre Peninsula in South Australia (Figure 1). Average annual rainfall and pan evaporation is 550 mm and 1550 mm, respectively. The basin is used for town water supply with current allocation of about 6.8×10^6 m³ per year. The geology and hydrogeology of the basin is characterised based on: Segnit [12], Johns [13], Morton and Steel [14], Wilson [15], and Shepherd [16]. These works include development of a geological model from borehole information, geological surface mapping, delineation of saturated limestone boundaries and pumping tests with attention focussed on Uley Wanilla, Uley East and Uley South groundwater lenses.

The aquifer system in Uley South is comprised of Quaternary Bridgewater Formation limestone (QL aquifer), beneath this is the Upper Tertiary Clay (Uley Formation) unit with thickness of 5 - 25 m, which forms an aquitard between the Tertiary Sand (TS) of Wanilla Formation and the Quaternary aquifer systems [14]. Recent studies [17,18] show that there is a hydraulic connection between the Uley South unconfined QL aquifer and the Uley Wanilla and Big Swamp areas through the TS aquifer (Figures 1 and 2). The basin is topographically closed and the low-lying central portion contains numerous sinkholes. Runoff is highly ephemeral, occurs only after moderate to high intensity rainfall and persists only tens to hundreds of meters before entering a sinkhole [17, 19,20]. Surface drainage is poorly developed as is typical of semi-arid karstic terrains that support only thin soils combined to produce scant vegetation. The groundwater flow direction in Uley South is from north-east to south-west.

An important surface feature is the Big Swamp (Figure 1) with an underlying Tertiary Clay with the primary mechanism of discharge through evaporation. In wetter

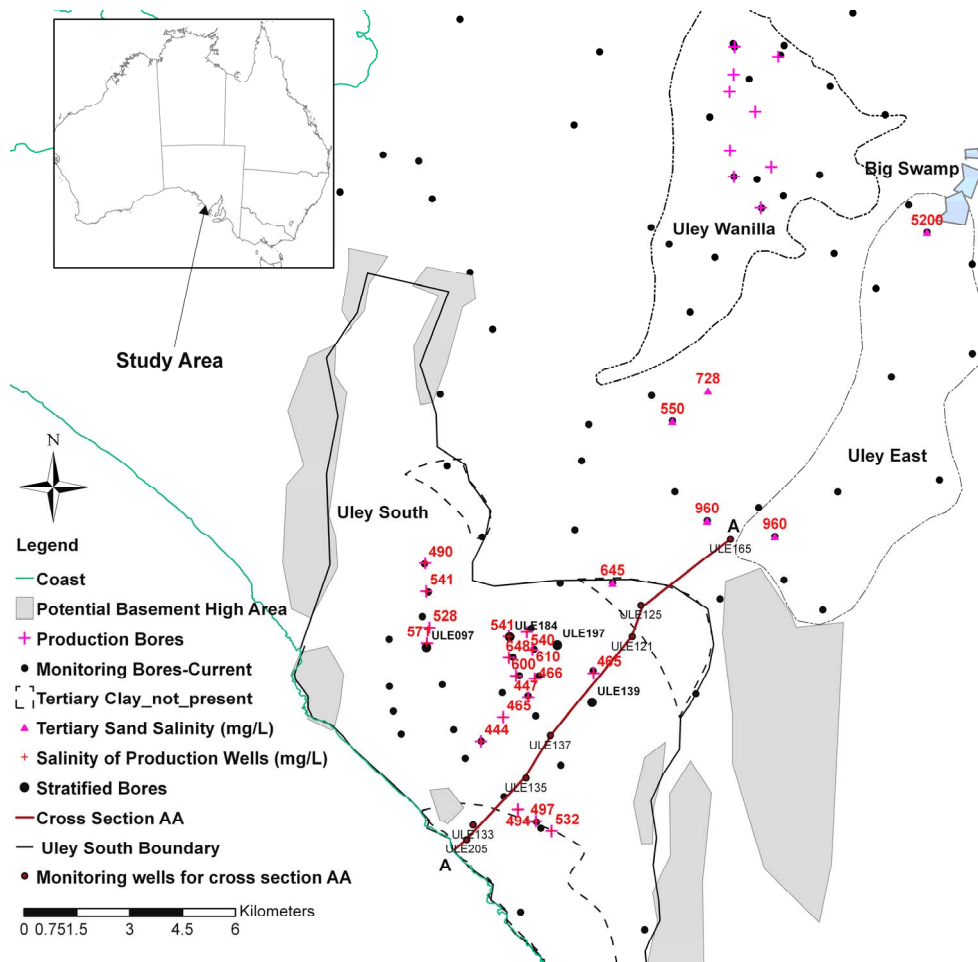


Figure 1. Location of the study area.

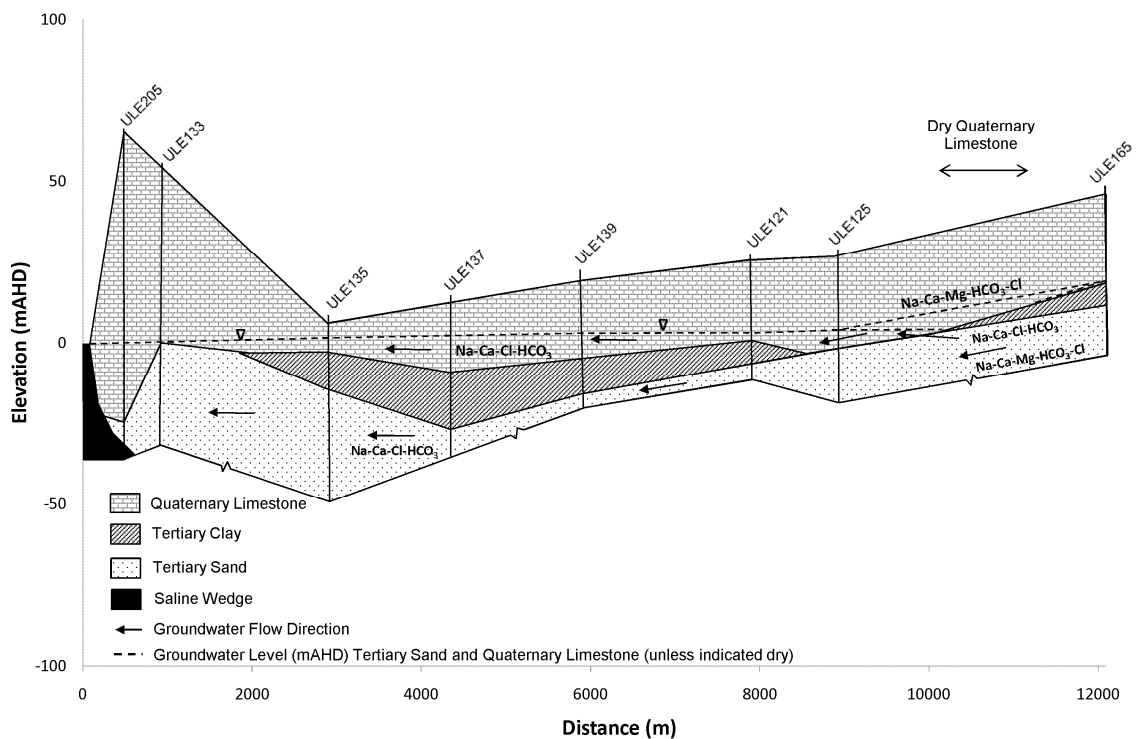


Figure 2. Cross section AA from the coast in the Uley South basin to Uley East basin's western edge.

years, Big Swamp overflows and infiltrates into the Quaternary Limestone unconfined aquifer and leaks into the TS aquifer raising salinity up to 5200 mg/L near Big Swamp.

3. Results and Discussion

The influence of geological control is divided into three sections: influence on QL aquifer salinity, hydrochemistry of QL groundwater, and the effect of extension of the saline wedge in TS aquifer. These are important factors in managing the Uley South basin.

3.1. Influence on Salinity

In Uley South, the occurrence of basement high at the basin boundary and discontinuity of aquitards and solution features constitute the most important controls upon salinity and water chemistry. Salinity measurements taken from the town water supply (TWS) wells indicate higher salinity in the north-central portion than elsewhere in the basin (Figure 1). Salinity distribution in the basin is influenced by two main geological controls: the occurrence of basement high in the southern half of the basin, and the area of Tertiary Clay absence in the central part (Figures 1 and 2).

Harrington *et al.* [17] and Somaratne [18] show that the high-salinity plume of groundwater in the TS aquifer originates from the Big Swamp area as a result of downward leakage of Big Swamp's surface water through the

clay aquitard. As the plume moves down gradient along its main flow path, the western edge of the Uley East basin, it is diluted from 5200 mg/L salinity at Big Swamp to 960 mg/L at the southern boundary of the Uley East basin. The area where aquitard is absent in Uley South, the TS aquifer receives direct recharge, resulting in further dilution of salinity to 645 mg/L (Figure 1). As a result, pronounced salinity stratification is found in QL aquifer monitoring well ULE197, down-gradient to the Tertiary Clay absence area (Figure 3(a)). In the ULE197 well, salinity of the QL base is identical to that in TS aquifer monitoring well ULE109, located at the same site. Harrington *et al.* [17] indicated that residence time in Big Swamp and Uley Wanilla groundwater age is >45 years. Mixing of old waters originating from Uley Wanilla and Big Swamp with younger water (<20 years) in Uley South has important ramifications on applicable recharge estimation methods. When mixing occurs of different groundwater systems with different ages and different chlorides concentrations, this results in unreliable recharge estimates using chlorofluorocarbon [21,22] and chloride mass balance methods [23,24], thus limiting the methods available for recharge assessment.

In contrast, the southern half of the QL aquifer has no link to the Big Swamp area through either QL or TS aquifers due to occurrence of the basement high at the Uley South boundary (Figure 1). Therefore, QL water in the southern part of Uley South is primarily derived from direct rainfall recharge within Uley South. ULE139 is a

QL aquifer monitoring well, which shows no evidence of salinity stratification (Figure 3(b)).

3.2. Influence on Hydrochemical Signatures

The combination of aquifer heterogeneity, low hydraulic gradient, recharge through solution features in addition to diffuse recharge, mixing with different water types and modification along flow paths results in Uley South being hydrochemically complex. Hydrochemistry was assessed principally by major ion chemistry within the Uley Wanilla, Uley East and Uley South lenses, and how these are related to evolution of hydrochemical processes in the Uley South basin.

Groundwater can be classified into “water types” on the basis of the absolute and relative abundance of major ions [25]. However, due to the predominance of Na and Cl ions in both QL and TS aquifers, the major ion differences present are inconclusive in distinguishing between QL and TS waters. For all QL and TS groundwaters, Ordens *et al.* [20] has shown that Cl/Br in both rainfall and groundwater samples are close to the Cl/Br ratio of seawater indicating groundwater chloride are marine derived implying that both QL and TS waters may have similar chemical histories.

Discrimination based on relative major ion abundance is more successful where recharging freshwater is of continental (Ca-HCO₃ type) rather than marine (Na-Cl type) origin as in Uley South. QL and TS groundwater are analysed for pH, electrical conductivity, anions (HCO₃⁻, SO₄²⁻, Cl⁻) and cations (Ca²⁺, Mg²⁺, Na⁺, K⁺). Although the majority of water has no dominant ion composition, Uley Wanilla water contains distinctively high Mg ions (Figure 4) as observed by Harrington *et al.* [17]. High chloride in TS waters is derived from the Big Swamp area and the Na-Cl type water is diluted towards Na-Cl-HCO₃ waters at Uley South. This represents either waters that have chemically evolved along flow paths or that result from mixing between Na-HCO₃ waters and Na-Cl waters.

Classification of groundwater into water types on the basis of major ion composition and other selected characteristics provides a framework for interpretation of groundwater pathways [26]. Eighty-five groundwater samples were used for QL aquifer. This includes the data of Evans [19] and town water supply wells’ water quality data. For the TS aquifer, twenty-seven water samples were taken from Evans [19].

Hydrochemical examination shows that carbonate dis-

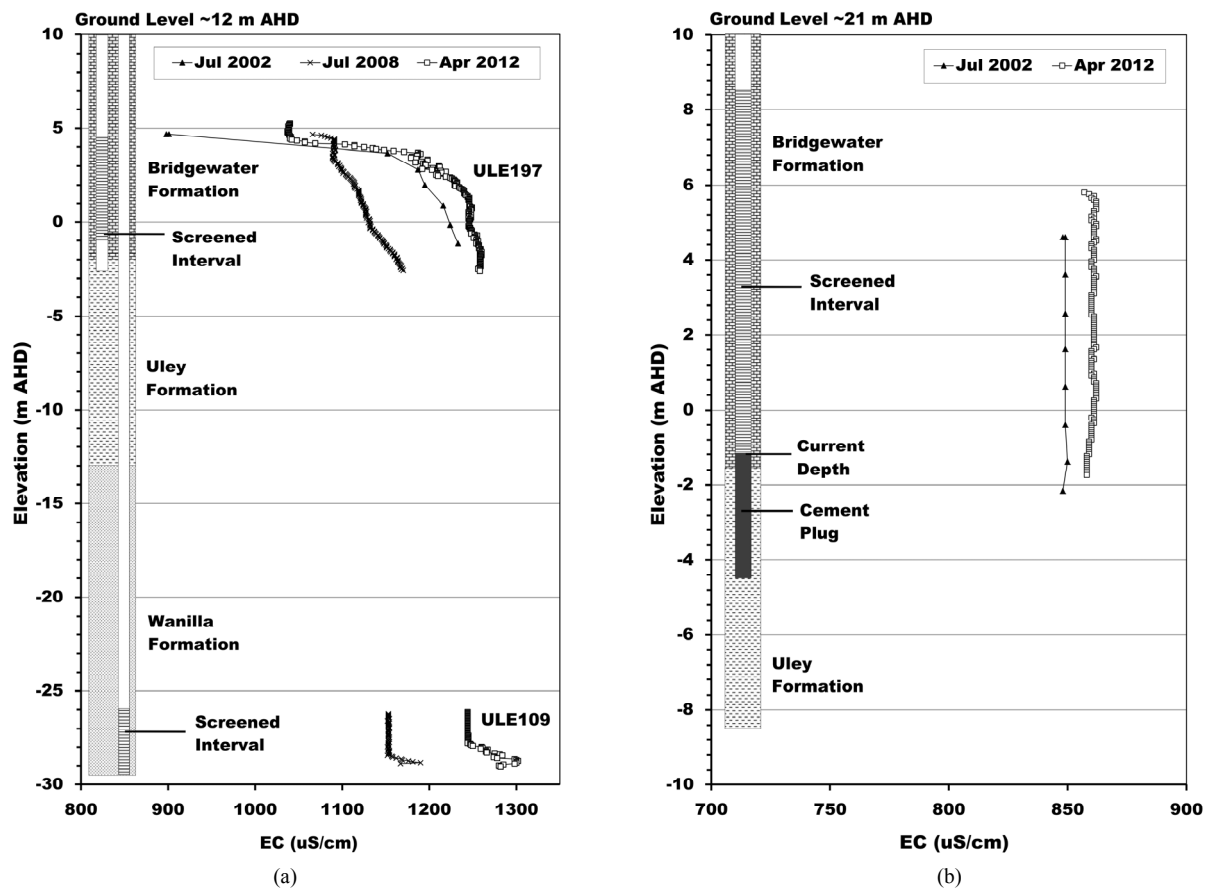


Figure 3. Salinity stratification of QL groundwater in the central part of the basin. (a) ULE109 and ULE197; (b) ULE139.

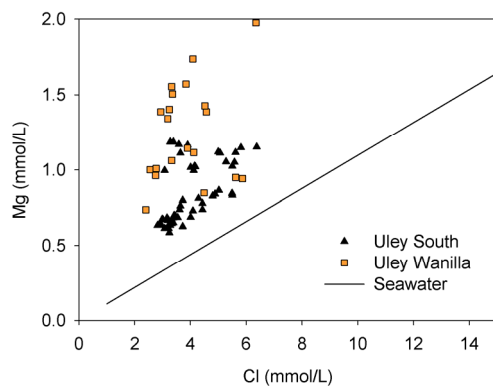


Figure 4. Cl vs Mg in Uley Wanilla and Uley South lenses.

solution and ion exchange are the two dominant chemical processes active in the QL aquifer. Water types are categorised into four types of groundwater, which represents a chemical evolution from Na-Mg-Cl-HCO₃ groundwater at Uley Wanilla lens to end member Na-Cl groundwater at Big Swamp (Figure 5).

3.2.1. Na-Cl or Na-Ca-Cl (1)

In the absence of evaporites, this type of water occurs in the Big Swamp due to enrichment of Na and Cl ions through evaporation. The end member of this water type is Na-Ca-Cl-HCO₃ or Na-Ca-HCO₃-Cl in the north-central part of Uley South basin, where salinity of town water supply wells are about 650 mg/L; about 100 mg/L higher than elsewhere in the QL aquifer.

3.2.2. Na-Mg-Ca-Cl-HCO₃/Na-Ca-Mg-HCO₃-Cl (2)

Na-Mg-Ca or Na-Ca-Mg water indicates that dissolution of dolomite (or dolomitic limestone) is taking place in the catchment area. In Quaternary Limestone, this type of water occurs in the Uley Wanilla basin (Figure 6). In the TS aquifer, this type of water occurs either within the Uley Wanilla basin or directly down-gradient in the flow direction, indicating direct leakage (interconnection between two aquifers) from QL to TS aquifer. This type of water in the Uley South QL aquifer is found only in production bores directly down gradient to Tertiary Clay absence area as a result of QL aquifer receiving TS waters. Apart from production wells directly down gradient, none of the monitoring wells nor any other town water supply wells contain this type of water.

3.2.3. Ca-Na-HCO₃-Cl/Ca-Na-Mg-HCO₃-Cl (3)

This is the typical water type found in limestone aquifers. The Ca-Na-HCO₃-Cl water type represents recently recharged water with an abundance of dissolved CO₂; thus calcite dissolution is taking place. This water type occurs in the southern part of the central basin where there is no TS water mixing with Uley South QL aquifer water. Uley South basin being a karstic limestone aquifer, sink-

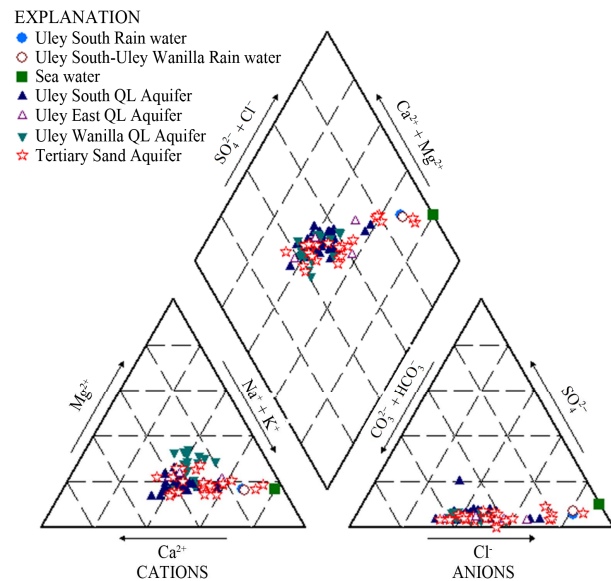
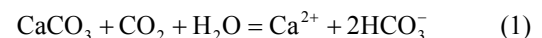


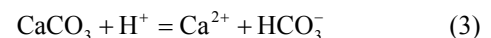
Figure 5. Piper diagram showing water types.

holes collect runoff containing organics; this provides sufficient CO₂ for the following reactions to dissolve calcite.

The main mechanism for dissolution of calcite is as follows:



The above reaction includes the following steps:



3.2.4. Na-Ca-Cl-HCO₃/Na-Ca-HCO₃-Cl (4)

Na-Ca-HCO₃-Cl water types represent Ca-HCO₃ water affected by mixing with Na-Cl waters or that has undergone cation exchange processes. The most common occurrence of such water is in the central basin (Figure 6), where high salinity waters originating from Big Swamp enter the QL aquifer. In addition another possible reason for the occurrence of this water type is cation exchanges, along with sea aerosol in the coastal zone.

3.3. Influence on Coastal Salinity

Seawater intrusion is generally identified through chemical analysis of groundwater. No single analysis definitively identifies seawater intrusion [27]; however by examining various analysis methods, it is evident when fresh groundwater mixes with seawater. According to HydroMetric WRI [27], at low chloride concentrations, it is often difficult to identify incipient seawater intrusion. This is due to the natural variation in fresh water chemistry at chloride concentrations below 1000 mg/L.

Mixing trends between groundwater and seawater are

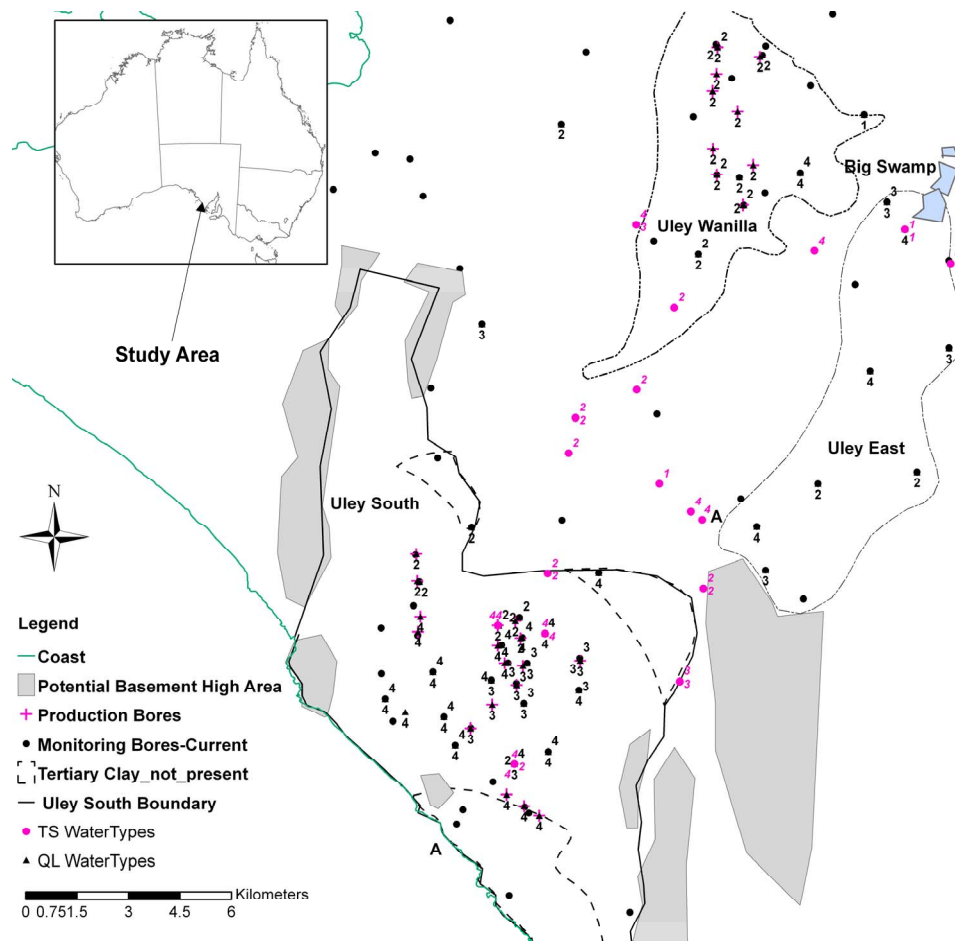


Figure 6. QL and TS aquifer water type distribution (1 = Na-Cl or Na-Ca-Cl; 2 = Na-Mg-Ca-Cl-HCO₃ or a-Ca-Mg-HCO₃-Cl or Ca-Na-Mg-HCO₃-Cl; 3 = Ca-Na-HCO₃-Cl or Ca-Na-Cl-HCO₃; 4 = Na-Ca-Cl-HCO₃ or Na-Ca-HCO₃-Cl).

more easily defined when chloride concentrations exceed 1000 mg/L, as is evident in **Figure 7** where mixing increases with depth. As shown in **Figure 7**, groundwater mixing takes place in coastal aquifer monitoring wells from Na-Ca-Cl-HCO₃ to Na-Cl water types at two different distances from the coast at the same elevation. On the northern side, TS groundwater is derived from downward leakage from the Uley Wanilla-Big Swamp system, which is approximately 22 km from the coastline. TS receives further direct recharge from areas where Tertiary Clay is absent.

Therefore, it is conceptualised that a large quantity of TS groundwater flows on the northern side with a higher hydraulic gradient (0.00458). This makes the distance to the toe of the saline wedge in TS smaller, less than 500 m on the northern side. In contrast, the TS aquifer on the southern side receives leakage from the QL aquifer within Uley South basin. The extent of the TS is limited only to the Uley South boundary, which is on average approximately 5 km from the coast, yielding an average hydraulic gradient of 0.00117. This results in comparatively small groundwater flows through TS along the

southern side, creating a longer saline wedge for pressure balance at the same elevation. Distance to the toe of the saline wedge is about 900 m from the coastline on the southern side. In Uley South coastal aquifer monitoring wells, sea water is intercepted in Tertiary Sand aquifer in SLE069 at -23 m AHD and in ULE205 at -30 m AHD (**Figure 7(b)**) showing increasing salinity with depth. No QL aquifer coastal bores intercept saline waters indicating that the saline wedge in QL aquifer is less than 500 m.

4. Conclusions

We highlight the importance of integrating knowledge of hydrochemistry of groundwater sources into estimation of recharge and the potential for saline intrusion, both of which are essential components in assisting with effective management of groundwater resources. In particular, we analysed the groundwater chemistry from the Uley South coastal aquifer in South Australia, and examined the processes that control groundwater chemistry. The major outcome of this study is identification of two distinct groundwater sources due to geological controls and

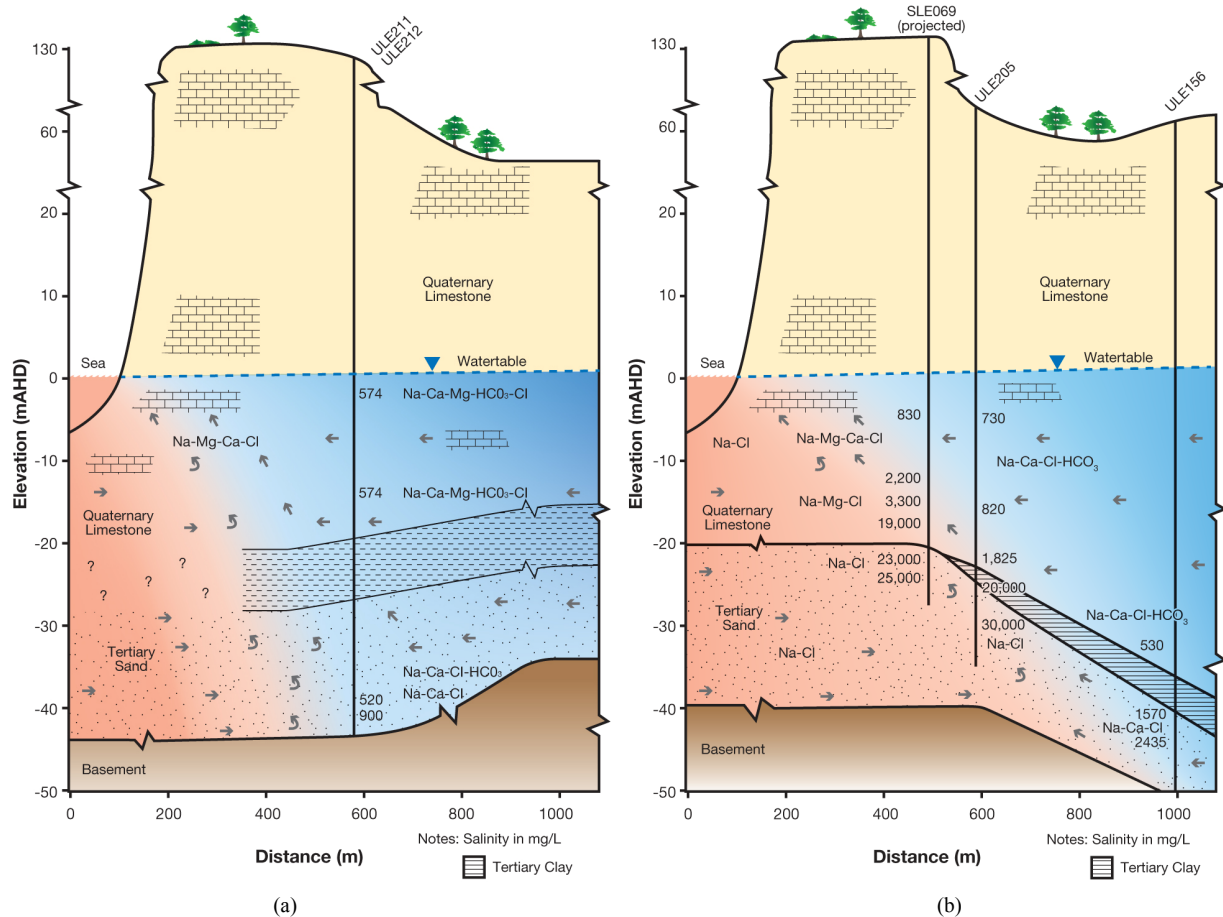


Figure 7. Uley South Saline Wedge. (a) Northern side; (b) Southern side.

their influence on groundwater chemistry. The geologic controls that led to these two distinct sources of different origin and age, were evidenced in the presence of two features, basement rock highs and absence of an aquitard.

Evaluation of hydrochemical data suggests that calcium carbonate dissolution, ion exchange processes and mixing with sea aerosols in the coastal zone are the main factors influencing groundwater chemistry. When the two groundwater sources with different ages and different chloride concentrations mix, results in unreliable recharge estimates are likely to result in using CMB and CFC methods. In addition, the subsurface basement rock high influences the currently observed seawater wedge. The toe of the saline wedge has been observed between 900 m and 500 m inland. Thus knowledge of geological control in the groundwater system, as well as the hydrogeochemical makeup of groundwater source(s), forms important parts of the characterisation of groundwater basin, ultimately contributing to effective management of the groundwater resources.

5. Acknowledgements

The editor and reviewers are thanked for evaluating this

paper.

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