



Interpretation of NO_3 – N_2O_5 observation via steady state in high-aerosol air mass: the impact of equilibrium coefficient in ambient conditions

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Abstract. Steady-state approximation for interpreting NO_3 and N_2O_5 has large uncertainty under complicated ambient conditions and could even produce incorrect results unconsciously. To provide an assessment and solution to the dilemma, we formulate datasets based on in situ observations to reassess the applicability of the method. In most of steady-state cases, we find a prominent discrepancy between K_{eq} (equilibrium coefficient for reversible reactions of NO_3 and N_2O_5) and correspondingly simulated $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$, especially under high-aerosol conditions in winter. This gap reveals that the accuracy of K_{eq} has a critical impact on the steady-state analysis in polluted regions. In addition, the accuracy of γ (N_2O_5) derived by steady-state fit depends closely on the reactivity of NO_3 (k_{NO_3}) and N_2O_5 ($k_{\text{N}_2\text{O}_5}$). Based on a complete set of simulations, air mass of k_{NO_3} less than 0.01 s^{-1} with high aerosol and temperature higher than 10°C is suggested to be the best suited for steady-state analysis of NO_3 – N_2O_5 chemistry. Instead of confirming the validity of steady state by numerical modeling for every case, this work directly provides appropriate concentration ranges for accurate steady-state approximation, with implications for choosing suited methods to interpret nighttime chemistry in high-aerosol air mass.

1 Introduction

The nitrate radical (NO_3), an extremely reactive species prone to buildup at night, is an ideal candidate for steady-state analysis in combination with dinitrogen pentoxide (N_2O_5) due to fast equilibrium reactions between them (Reaction R1).



Under the steady-state condition, the lifetime of NO_3 (denoted as $\tau_{\text{ss}}(\text{NO}_3)$) can be calculated as the ratio of NO_3 concentration over the production rate ($k_{\text{NO}_2+\text{O}_3}[\text{NO}_2][\text{O}_3]$) or over the removal rate of both NO_3 and N_2O_5 , as indicated in Eq. (1). A similar representation of N_2O_5 steady-state lifetime is also shown in Eq. (2). The loss frequencies of various sink pathways of NO_3 and N_2O_5 are integrated as total first order in the following equations, represented by the k_{NO_3} and $k_{\text{N}_2\text{O}_5}$ terms. Briefly, k_{NO_3} is contributed by the reaction of the NO_3 radical with NO and hydrocarbons and uptake on particles at night, ranging from hundredths of recip-

rocal seconds to several reciprocal seconds depending on the air mass. Due to its large rate constant with NO, the concentration usually dominates the lifetime of the NO₃ radical in urban areas with fresh NO emission. Otherwise, the reactions with hydrocarbons, especially unsaturated hydrocarbons, are preferential for NO₃ in rural areas. The K_{eq} denotes the equilibrium coefficient for Reactions (R1a) and (R1b), derived by Eq. (3).

$$\begin{aligned}\tau_{\text{ss}}(\text{NO}_3) &\equiv \frac{[\text{NO}_3]}{k_{\text{NO}_2+\text{O}_3}[\text{NO}_2][\text{O}_3]} \\ &\approx (k_{\text{NO}_3} + K_{\text{eq}}[\text{NO}_2]k_{\text{N}_2\text{O}_5})^{-1}\end{aligned}\quad (1)$$

$$\begin{aligned}\tau_{\text{ss}}(\text{N}_2\text{O}_5) &\equiv \frac{[\text{N}_2\text{O}_5]}{k_{\text{NO}_2+\text{O}_3}[\text{NO}_2][\text{O}_3]} \\ &\approx \left(k_{\text{N}_2\text{O}_5} + \frac{k_{\text{NO}_3}}{K_{\text{eq}}[\text{NO}_2]}\right)^{-1}\end{aligned}\quad (2)$$

$$K_{\text{eq}} = \frac{k_{\text{R1a}}}{k_{\text{R1b}}} = \frac{[\text{N}_2\text{O}_5]}{[\text{NO}_2][\text{NO}_3]}\quad (3)$$

Numerous works have taken advantage of the steady-state calculation to quantify the total first-order loss rate for NO₃ or N₂O₅ such that they drew conclusions about the oxidation capacity and reactive nitrogen budgets contributed by this chemical system (Allan et al., 1999, 2000; Carslaw et al., 1997; Platt et al., 1984; Vrekoussis et al., 2007; Wang et al., 2013). Since the steady-state approximation was used to interpret atmospheric observation of NO₃–N₂O₅ (Brown et al., 2003; Platt et al., 1981), this method was also widely implemented to quantify the N₂O₅ uptake coefficient ($\gamma(\text{N}_2\text{O}_5)$) (Brown et al., 2009, 2003; Li et al., 2020; McDuffie et al., 2019; Phillips et al., 2016; Wang et al., 2017a, c, 2020a).

However, with the influence induced by complicated atmospheric conditions and emission, steady state in ambient air mass will not always be the case (as illustrated in Sect. S1 and Fig. S1 in the Supplement). These situations are prevalent in nocturnal boundary layer (Phillips et al., 2016; Stutz et al., 2004; Wang et al., 2017a, c) and therefore increase the difficulty of applying steady state directly to NO₃–N₂O₅ observation data, whereas few studies have systematically characterized the error source and application conditions of this method (Brown et al., 2009).

Due to a faster approach to equilibrium than steady state, the application of K_{eq} in the calculation of steady-state equations seems to be reasonable (Brown et al., 2003). For example, the ambient NO₃ concentration was usually calculated based on ambient N₂O₅ concentration with $K_{\text{eq}} \times [\text{NO}_2]$ when determining their budgets or characterizing the lifetime or sink attribution of these two reactive nitrogen compounds (Brown et al., 2011; Osthoff et al., 2006; Wang et

al., 2018, 2017c, d; Yan et al., 2019). In addition, the mathematical conversion between NO₃ and N₂O₅ concentration via K_{eq} coefficient can simplify the calculation in the iterative box model, which derives $\gamma(\text{N}_2\text{O}_5)$ by iterating its value in the model until the predicted N₂O₅ concentration matches the observation (Wagner et al., 2013; Wang et al., 2020b). However, considerable uncertainty could be associated with the quantification of K_{eq} and its different parameterizations (Cantrell et al., 1988; Pritchard, 1994). The impact of K_{eq} value on steady-state fit or concentration conversion has not been explored to date in the analysis of NO₃–N₂O₅ steady state.

In this study, we formulate a half artificial dataset with expected properties based on field campaigns. Specifically, most of the species contained in the dataset are observed values while only NO₃ and N₂O₅ were calculated by the steady-state model (illustrated in the Sect. 2.2). With the dataset, we illustrate the reasons for deviation of parameterized K_{eq} from $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ in ambient conditions, the possible uncertainties of linear fit based on steady-state Eqs. (4) and (5) (the related variables are explained in Sect. 2.1) that resulted from different K_{eq} values, and the influence of relevant atmospheric variables on $\gamma(\text{N}_2\text{O}_5)$ derivation via the steady-state method. Furthermore, a series of ambient condition tests specify the exact ranges suited for steady-state analysis according to not only the validity of steady state but also K_{eq} values, which optimize the validity check by numerical modeling in previous research (Brown et al., 2009, 2003) and develop complete standard for data filtering.

2 Methods

2.1 $\gamma(\text{N}_2\text{O}_5)$ derivation by steady-state approximation

The framework of steady-state approximation for the NO₃–N₂O₅ system is basically built on its chemical production and removal pathways, in case of extremely weak physical processes (e.g., transport, dilution, and deposition) relative to its chemical processes. With simultaneous measurements of NO₃, N₂O₅, and relevant precursor concentrations, the steady-state lifetime $\tau_{\text{ss}}(\text{NO}_3)$ and $\tau_{\text{ss}}(\text{N}_2\text{O}_5)$ can be quantified for a targeted period as shown in Eqs. (1) and (2). By substituting the $k_{\text{N}_2\text{O}_5}$ with $0.25 \times c \times S_a \times \gamma(\text{N}_2\text{O}_5)$, the $\gamma(\text{N}_2\text{O}_5)$ and the reactivity of NO₃ (k_{NO_3} , including the reactions of NO₃ with NO and hydrocarbons) can therefore be determined by Eqs. (4) and (5).

$$\tau_{\text{ss}}^{-1}(\text{NO}_3) \approx k_{\text{NO}_3} + 0.25c S_a K_{\text{eq}}[\text{NO}_2]\gamma(\text{N}_2\text{O}_5)\quad (4)$$

$$\begin{aligned}(0.25c S_a \tau_{\text{ss}}(\text{N}_2\text{O}_5))^{-1} \\ \approx \gamma(\text{N}_2\text{O}_5) + k_{\text{NO}_3} (0.25c S_a K_{\text{eq}}[\text{NO}_2])^{-1}\end{aligned}\quad (5)$$

Here c represents the mean molecular velocity of N₂O₅, S_a represents the aerosol surface area, and K_{eq} is calculated from the rate constant of reversible Reactions (R1a)

(k_{R1a}) and (R1b) (k_{R1b}), which is a temperature-dependent parameter. It should be noted that the photolysis of NO_3 is not considered in the $k\text{NO}_3$ due to weak radiation at night, and the homogeneous hydrolysis was also ignored due to its small contribution in comparison to the heterogeneous pathway. A similar presumption was also implemented in previous studies (Brown et al., 2009; Mentel et al., 1996; Wahner et al., 1998). In the form of these two equations, the potential covariance between S_a and NO_2 concentration can be avoided to decrease the uncertainty (Brown et al., 2009). By being fit to these two equations, γ (N_2O_5) can be directly derived from slope of the plot of $\tau_{\text{ss}}^{-1}(\text{NO}_3)$ against $0.25c S_a K_{\text{eq}}[\text{NO}_2]$ or from the intercept of the plot of $(0.25c S_a \tau_{\text{ss}}(\text{N}_2\text{O}_5))^{-1}$ against $(0.25c S_a K_{\text{eq}}[\text{NO}_2])^{-1}$ respectively. In the following analysis, the linear fit based on Eq. (5) is preferred in steady-state approximation.

2.2 Steady-state model and half-artificial datasets

The steady-state model is reformed from a zero-dimensional box model to produce NO_3 and N_2O_5 , which are in steady state as far as possible. It is constrained by measurements of NO , NO_2 , O_3 , CO , CH_4 , VOCs , HCHO , S_a , relative humidity (RH), temperature (T), and pressure, coupled with the Regional Atmospheric Chemistry Mechanism, version 2 (RACM2). Each data point is treated as an independent air mass, aging 10 h and keeping input constraint unchanged. As NO_3 – N_2O_5 chemistry, the interest of this work, usually shows marked impacts during the night, only the time periods with negligible photolysis frequency are under consideration. In the standard simulation (herein referred to as Mod0), the uptake coefficient of N_2O_5 is set to 0.02, as a reasonable value of literature (Brown et al., 2006; Chen et al., 2020; McDuffie et al., 2018; Morgan et al., 2015; Phillips et al., 2016; Wagner et al., 2013; Wang et al., 2017c; Yu et al., 2020).

Two half-artificial datasets are derived from PKU2017 and TZ2018 field campaigns (see Sect. S2) based on the steady-state model for analysis in the following sections. The simulated NO_3 and N_2O_5 and other observed values used for the constraints of the steady-state model jointly formulate these half-artificial datasets. Specifically, the NO_3 and N_2O_5 concentrations in this dataset are the output of the steady-state model simulation and guaranteed to be in steady state with respect to other observed precursors. To verify the steady state of NO_3 and N_2O_5 for each data point, we filtered the dataset according to the deviation between the steady-state lifetime of N_2O_5 ($\tau_{\text{ss}}(\text{N}_2\text{O}_5) = \frac{[\text{N}_2\text{O}_5]}{k_{R1}[\text{NO}_2][\text{O}_3]}$) and calculated lifetime of N_2O_5 ($\tau_{\text{calc}}(\text{N}_2\text{O}_5) = (k_{\text{N}_2\text{O}_5} + \frac{k_{\text{NO}_3}}{K_{\text{eq}}[\text{NO}_2]})^{-1}$). If the deviation exceeds 10 % for a data point, it will be excluded from the following analysis. We presume that if any data point output from the model is still out of steady state in terms of NO_3 and N_2O_5 , the sink rate constant of air mass represented by this data point should be too weak for steady-state analysis within a reasonable timescale. In addition, the

data higher than 5 ppbv NO are filtered out in the following calculation, since the resulting large variation in $k\text{NO}_3$ can bias the linear fit even though NO_3 and N_2O_5 approach steady state rapidly under high NO (discussed in Sect. 3.2). The fraction of excluded data is less than 8 %, and they are expected to have little influence on our results. The calculated nighttime loss fraction accounted for by NO_3 and N_2O_5 show large discrepancy (see Sect. S3 and Fig. S2) between these two half-artificial datasets, which provide us a good opportunity to investigate the factors impacting steady-state approximation across different conditions.

Rather than using observation data directly, a half-artificial dataset can provide a larger amount of valid data for steady-state analysis with a known γ (N_2O_5) value. In addition, this method avoids the impacts from steady-state deviation, which helps to analyze the factors influencing γ (N_2O_5) quantification via steady-state approximation backwards from a known steady-state condition.

3 Results and discussion

3.1 Varying equilibrium coefficient under steady state

The rates of NO_3 – N_2O_5 reversible reactions are expected to be equal for the steady-state case, so that the equilibrium coefficient K_{eq} can be determined from either the rate constant ratio of Reactions (R1a) and (R1b) or the ratio of $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$. Although this approach is reasonable under ideal conditions, the exact same rates between reversible reactions and the following calculation based on K_{eq} scaling are not so appropriate for ambient atmosphere where the removal pathway for NO_3 – N_2O_5 is not negligible, especially under the high-aerosol-loading condition. NO_3 and N_2O_5 achieve steady state after 1.5 h evolution, when concentration and rates remain constant (Fig. 1). In this simulation, the starting mixing ratios of NO_2 and O_3 are 10 and 23 ppbv, respectively, which is the average level for the nighttime conditions in PKU2017. The concentrations of these two precursors are held constant in the simulation to better illustrate the influence of removal rates. This result will stay almost the same no matter whether these starting values are initialized to be constant or allowed to vary. Under steady state, the net equilibrium reaction rate in Fig. 1b and c stays negative and positive for NO_3 and N_2O_5 , respectively. In addition, the absolute values and difference of the forward and backward reaction rates remain unchanged after achieving steady state. This result is similar to a previous numerical calculation study (Brown et al., 2003), while the deviation between reversible reaction rates becomes larger in our case.

In this case, the original equilibrium is imperfectly realized (a perfect realization of the original equilibrium condition is that K_{eq} and the ratio of $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ are equivalent as in Eq. 6), leading to errors on the projection of NO_3 and N_2O_5 concentration via $K_{\text{eq}} \times [\text{NO}_2]$. In fact, we note that a new equilibrium between NO_3 and

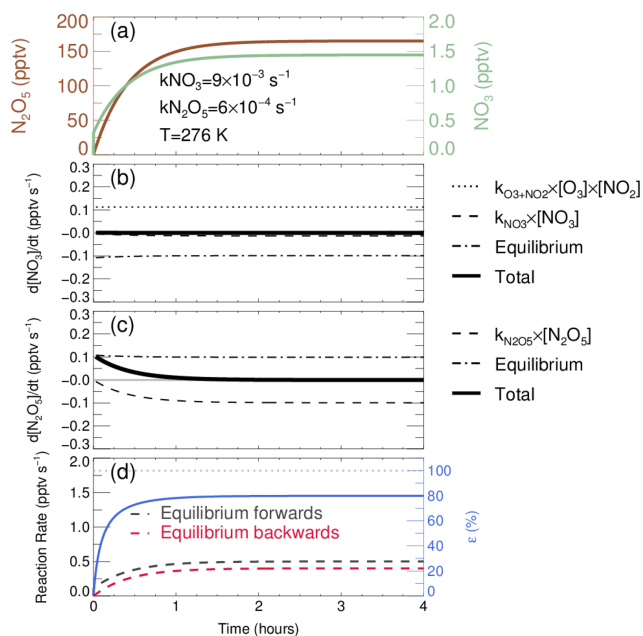


Figure 1. Evolution of the NO₃-N₂O₅ system simulated by the steady-state model for an average case. **(a)** Temporal profiles of N₂O₅ and NO₃; the constraint of simulation is displayed as the text. **(b)** Evolution of $d[\text{NO}_3]/dt$ calculated from the source of $k_{\text{O}_3+\text{NO}_2} \times [\text{O}_3] \times [\text{NO}_2]$, sink of $k_{\text{NO}_3} \times [\text{NO}_3]$, and equilibrium terms, detailed in the text. **(c)** Evolution of $d[\text{N}_2\text{O}_5]/dt$ calculated from equilibrium terms, sink of $k_{\text{N}_2\text{O}_5} \times [\text{N}_2\text{O}_5]$. **(d)** Forward (N₂O₅ formation) and backward (N₂O₅ decomposition) equilibrium rates are represented as black and red dashed lines, and the equilibrium completeness ε is calculated by the ratio of backward rate over forward rate, shown as blue full line.

N₂O₅ is developed with constant but unequal rates. Under this new equilibrium condition, the ratio of (R1b) reaction rate (the red dashed line in Fig. 1d) over (R1a) reaction rate (the black dashed line in Fig. 1d) can be regarded as the degree of approaching original equilibrium (the blue line in Fig. 1d). In addition, this value is also the ratio of $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ against original K_{eq} ; therefore we defined this ratio as a correction factor ε , implemented to calculate accurate $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ with significant N₂O₅ removal pathways. The value of K_{eq} after scaled by ε can be used for converting the concentration of NO₃ and N₂O₅ via Eq. (6):

$$\varepsilon \times K_{\text{eq}} = \varepsilon \times \frac{k_{\text{R1a}}}{k_{\text{R1b}}} = \frac{[\text{N}_2\text{O}_5]}{[\text{NO}_2][\text{NO}_3]}. \quad (6)$$

Sensitivity tests are conducted to demonstrate the dependence of ε on relevant variables based on the steady-state model. The averages ambient conditions observed at the wintertime Peking University (PKU) site and summertime Tai Zhou (TZ) site are taken as two basic constraints for sensitivity tests (Table S2 in the Supplement). By separately altering variables, such as NO₂, O₃, kN₂O₅, kNO₃, and T , the sensi-

tivity of the ε value can be obtained as shown in Figs. 2 and S4. The ε value depends primarily on kN₂O₅ and T in both scenarios, where ε increases with T (approaching 1 under relatively high T) and decreases with kN₂O₅. In comparison, the ε value behaves insensitively as well as NO₂ and O₃ concentration, at least within the range of reasonable ambient conditions. High kN₂O₅ results from high-aerosol events, usually occurring in winter accompanied by low temperature and high relative humidity in some populated areas (Baasandorj et al., 2017; Huang et al., 2014; Wang et al., 2017b, 2014), further decreasing the accuracy of original K_{eq} values. It can be inferred that in order to accurately interpret the relationship of NO₃ and N₂O₅, calculation relying on the equilibrium equation and steady-state approximation should consider the dependence of ε on ambient conditions.

Even if the K_{eq} value serves as a good representation of the ratio of $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ or ε can be readily quantified in the field, the discrepancy among different databases in calculating K_{eq} still increases the uncertainties of NO₃-N₂O₅ calculation through steady-state approximation or equilibrium, which has not been carefully considered. Here, we apply a set of uniform formulas to describe k_{R1a} and k_{R1b} (see Sect. S4) from preferred values of several popular atmospheric chemistry mechanisms (Mozart, CB05, Sapr07, RACM2 and kinetic databases JPL2015 as well as IUPAC2017) and finally calculating K_{eq} . As is shown in Figs. S5 and S6, K_{eq} variations derived from these six different databases reflect considerable discrepancy from each other, especially in colder conditions. Because parameterized K_{eq} values are only dependent on ambient temperature, they continuously increase with time due to the decrease in temperature. In addition to the discrepancy between different K_{eq} parameterizations, the ε value varies dissimilarly with each K_{eq} , ranging from 70 % to 90 %. All these results demonstrate that, in most cases, K_{eq} values simply derived from an existing database would fail to reproduce an accurate relationship between NO₃ and N₂O₅.

To further elucidate the impact of K_{eq} on deriving γ (N₂O₅) via steady-state approximation (hereafter defined as γ_{ss} (N₂O₅)), Fig. S6 shows the steady-state fit based on all six database-derived K_{eq} values and in the same time periods as Fig. S5 through Eqs. (4) and (5), respectively (both equations can derive a pair of γ_{ss} (N₂O₅) and kNO₃). K_{eq} (corrected with ε) is calculated with NO₃ and N₂O₅ concentration simulated based on RACM2. Fits based on Eq. (4) could lead to 11 %–46 % underestimation of γ_{ss} (N₂O₅), as indicated by varying slopes in Fig. S7b and d, when using the database-derived K_{eq} . Conversely, fit by Eq. (5) (shown in Fig. S7a and c) biases the result of kNO₃ as the slopes without much influence on γ_{ss} (N₂O₅) as the intercept. Previous research ascribed inconsistent fit results between two equations to measurement uncertainty (Brown et al., 2009, 2006). However, fit with the original K_{eq} might be the primary reason for such inconsistent results and even differentiates the derived γ_{ss} (N₂O₅) and kNO₃ from true values. Therefore,

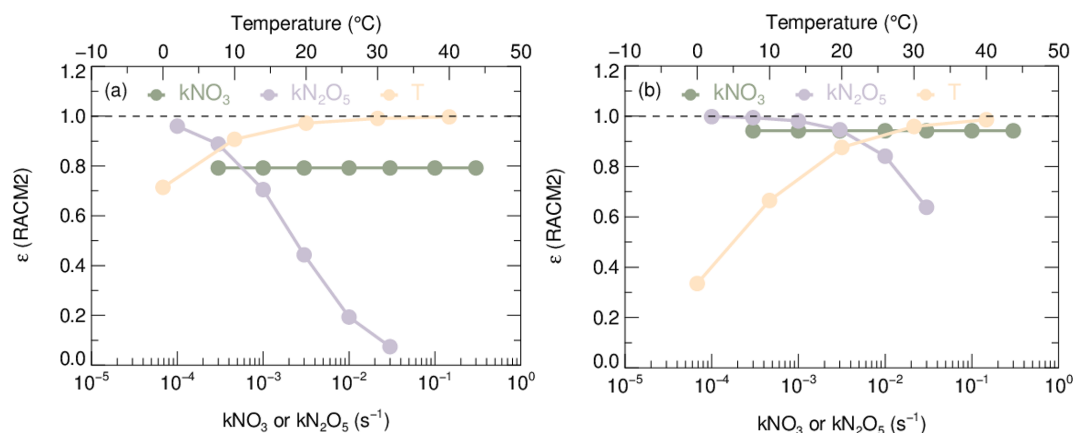


Figure 2. Sensitivity plot of $k\text{NO}_3$, $k\text{N}_2\text{O}_5$, and temperature (T) against coefficient ϵ . The trace of T is plotted against the upper horizontal axis, and the traces of the other two parameters are plotted against the lower horizontal axis. **(a)** The basic model condition is according to typical winter conditions of PKU2017. **(b)** The basic model condition is according to typical summer conditions of TZ2018. Basic model conditions including $k\text{NO}_3$, $k\text{N}_2\text{O}_5$, and temperature (T) are shown in Table S2. It should be noted that the provided ranges of each factor do not exactly equal but encompass the ambient conditions encountered during the two campaigns.

steady-state fit based on Eq. (5) might be the best choice for γ (N_2O_5) derivation via steady-state approximation. Similarly, Eq. (4) is preferred to be applied when $k\text{NO}_3$ is the final objective.

3.2 Impacts of $\text{NO}_3\text{--N}_2\text{O}_5$ reactivity on steady state

In order to further explore the impacting factors on the steady-state fit method, γ_{ss} (N_2O_5) results are derived for each 2 h time period of the PKU2017 and TZ2018 datasets based on output from the steady-state model. Since the pre-set γ (N_2O_5) in this model is 0.02, the degree of deviation from this value is supposed to reflect the accuracy of the fitted result.

It can be noticed from Eq. (5) that the variability of $k\text{NO}_3$ during the same time period leads data points to scatter on lines with different slopes, which could bias the resulted γ_{ss} (N_2O_5) from the model pre-set value. As is shown in Fig. 3, the absolute percentages of γ_{ss} (N_2O_5) deviation grow dramatically with the increase in relative standard deviation of $k\text{NO}_3$ ($k\text{NO}_3$ RSD) in both winter and summer datasets. The positive correlation even gives rise to extreme deviation in the summer dataset, with up to almost 10 times of model setting γ (N_2O_5). In fact, there remain accurate γ_{ss} (N_2O_5) values derived in each range of $k\text{NO}_3$ RSD, indicating a not strictly positive correlation between γ_{ss} (N_2O_5) deviation and $k\text{NO}_3$ RSD. This implies that large variation in $k\text{NO}_3$ only enhances the possibilities of inaccurate results from steady-state fit rather than hinder the γ_{ss} (N_2O_5) quantification all the time.

In addition to the large variation in $k\text{NO}_3$ in a short time period, the absolute level of $k\text{NO}_3$ and $k\text{N}_2\text{O}_5$ could influence the possibilities of inaccurate γ_{ss} (N_2O_5) from different aspects. Although the enhancement of $k\text{NO}_3$ and $k\text{N}_2\text{O}_5$

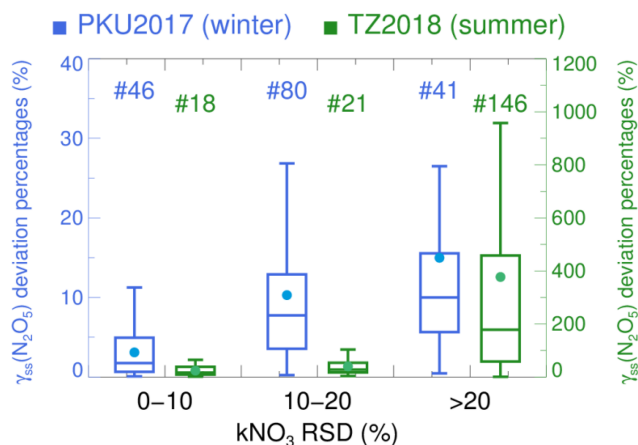


Figure 3. Relationship between γ (N_2O_5) derivation through steady-state approximation and $k\text{NO}_3$ relative standard deviation (RSD) in box-whisker plot. The blue and green colors represent datasets from PKU2017 and TZ2018 respectively, binned according to $k\text{NO}_3$ RSD. The dots are the mean deviation of γ_{ss} (N_2O_5). The number above the box-whisker plot represents the valid data points in each bin.

boosts the approach to steady state (Sect. S5 and Fig. S8), higher levels of $k\text{NO}_3$ amplify the bias of γ_{ss} (N_2O_5), contrary to $k\text{N}_2\text{O}_5$, with the same relative variation in $k\text{NO}_3$ (Sect. S6 and Fig. S10). This indicates that the region with plural emissions (e.g., strong biogenic or vehicular emission) might not be suited for steady-state fit due to the high $k\text{NO}_3$. Therefore, a trade-off between the variation in $k\text{NO}_3$ and the high level of $k\text{NO}_3$ (fast approach to steady state) should be made when deriving γ_{ss} (N_2O_5).

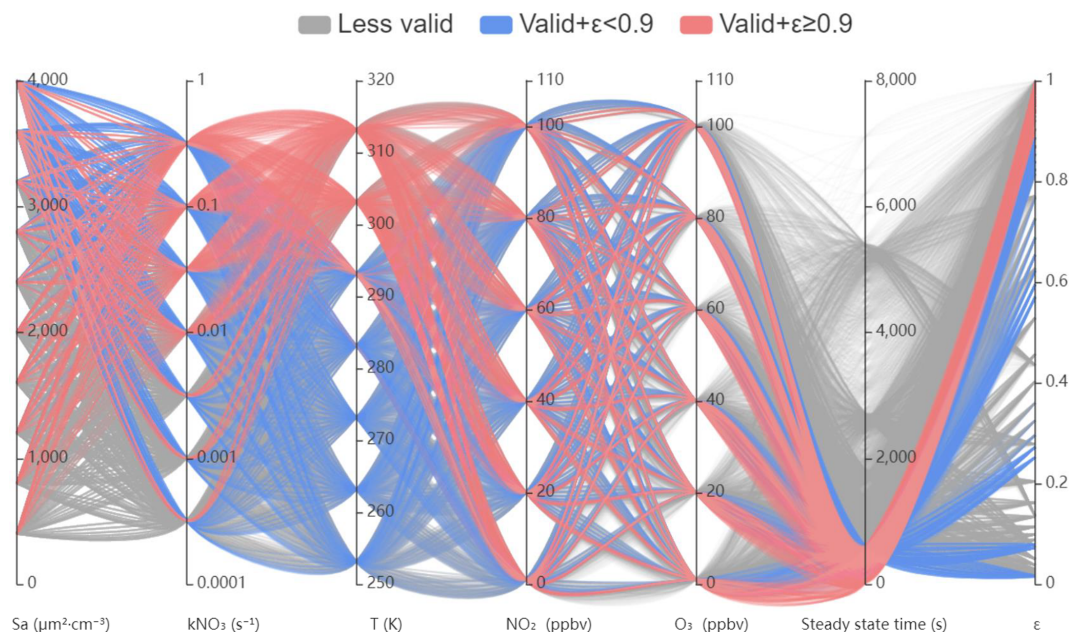


Figure 4. Numerical simulations for determining conditions available for the steady-state approximation method in a parallel-axis plot. Each line simply represents a simulation associated with different parameters in different vertical axes. The first five axes from the left represent initial variables used for constraining the simulations. The last two axes represent the time required for achieving steady state and the ϵ value calculated from the simulated results. The gray lines show cases approaching steady state for longer than 600 s (less valid). The blue lines show cases approaching steady-state cases within 600 s with ϵ less than 0.9, which is also inappropriate for steady-state analysis. The pink lines show cases approaching steady-state cases within 600 s with ϵ higher than 0.9, which is suited for steady-state analysis.

3.3 Implication for accurate steady-state analysis of $\text{NO}_3\text{--N}_2\text{O}_5$

While a few studies have examined the validity of steady state under certain conditions via numerical modeling when interpreting the ambient data (Brown et al., 2009, 2003), a clear range well suited to steady-state analysis of $\text{NO}_3\text{--N}_2\text{O}_5$, taking both K_{eq} and validity of steady state into consideration, has not been determined to date.

Here almost 20 000 simulations are displayed in the parallel plot of Fig. 4, where each line connects five constraint parameters to the calculated steady-state time and ϵ (the correction factor for K_{eq} parameterization to match the exact ratio of $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$, detailed in Eq. 6). The gray traces represent the simulations that could not match steady state within 600 s and were defined as less valid cases here. By this definition, we intend to indicate that it is also viable to apply steady-state approximation on air mass, which requires more than 600 s to match steady state, whereas the uncertainty caused therefrom could increase to some extent. The pink and blue traces together represent the simulations that could match valid steady state within 600 s without consideration of K_{eq} deviation (in other words the value of ϵ). Furthermore, the criteria to apply steady-state approximation appropriately is approach to steady state within 600 s and ϵ larger than 0.9, which are indicated as pink traces. While the levels of T , NO_2 , and O_3 have a minor effect on the approach

to steady state, simultaneous low kN_2O_5 (indicated as low S_a in the plot) and kNO_3 prevent the $\text{NO}_3\text{--N}_2\text{O}_5$ system from developing steady state. For example, when kNO_3 is lower than 0.01 s^{-1} , the air mass will be valid only if S_a increases to at least $3000 \mu\text{m}^2 \text{ cm}^{-3}$ with $\gamma(\text{N}_2\text{O}_5)$ of 0.02. This implies that clean air mass is not suited for steady state in any cases, whereas the high-aerosol condition provides more possibilities to approach steady state even with low kNO_3 . However, in order to interpret $\text{NO}_3\text{--N}_2\text{O}_5$ chemistry with an accurate K_{eq} coefficient, the ϵ larger than 0.9 is additionally taken into consideration, which excludes 50 % of valid steady-state cases mainly with high aerosol and lower than 10°C . These cases could bias $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ from original K_{eq} (also indicated in Fig. 2), leading to inaccurate results of calculation based on K_{eq} .

4 Conclusions

In this study, we found that the parameterized K_{eq} coefficient deviates much from the ratio of $[\text{N}_2\text{O}_5]/([\text{NO}_2] \times [\text{NO}_3])$ in some cases where steady state is valid. The indicator of the deviation, ϵ , is relatively sensitive to N_2O_5 reactivity and ambient temperature. It implies that conditions suited for steady-state analysis should be determined according to not only the validity of steady state but also K_{eq} , especially under high-aerosol conditions, like some regions in India, China, Europe, and the US (Baasandorj et al., 2017; Cesari

et al., 2018; Huang et al., 2014; Mogno et al., 2021; Petit et al., 2017; Wang et al., 2017b). Considering that a high level of kNO_3 might amplify the bias of γ_{ss} (N_2O_5) yield from steady-state fit and appears to be accompanied by fast variations, air mass of kNO_3 less than 0.01 s^{-1} with high aerosol and T higher than 10°C are therefore the best suited for steady-state analysis of $\text{NO}_3\text{--N}_2\text{O}_5$ chemistry, which indicates that this method would be more applicable in polluted regions with high aerosol loading during summertime. If the restriction of ε is relaxed to 30 %, some winter conditions will also be applicable. Our results provide an insight into improving the accuracy of the steady-state approximation method and finding suited areas to interpret nighttime chemistry. Further improvement of in situ $\text{NO}_3\text{--N}_2\text{O}_5$ budget quantification might rely on the direct measurements via flow tube system or machine learning prediction based on ancillary parameters.

Code and data availability. The datasets used in this study are available from the corresponding author upon request (wanghch27@mail.sysu.edu.cn; k.lu@pku.edu.cn).

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