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Molecular and seasonal characteristics of organic vapors in urban Beijing: insights from Vocus-PTR measurements

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Abstract. Understanding the composition and evolution of atmospheric organic vapors is crucial for exploring their impact on air quality. However, the molecular and seasonal characteristics of organic vapors in urban areas, with complex anthropogenic emissions and high variability, remain inadequately understood. In this study, we conducted measurements in urban Beijing during 2021–2022, covering four seasons, using improved proton-transfer-reaction mass spectrometry (Vocus-PTR-MS). During the measurement period, a total of 895 peaks were observed, and 512 of them can be assigned to formulas. The contribution of $C_xH_yO_z$ species is most significant, composing up to 54 % of the number and 74 % of the mixing ratio of total organics. With enhanced sensitivity and mass resolution, various species with multiple oxygens (≥ 3) or at a level of sub-parts per trillion (ppt) were observed, with 44 % of the number measured at sub-ppt level and 38 % of the number containing three–eight oxygen atoms. Organic vapors with multiple oxygens mainly consist of intermediate/semi-volatile compounds, and many formulae detected were reported to be the oxidation products of various volatile organic precursors. In summer, the fast photooxidation process generated organic vapors with multiple oxygens and led to an increase in both their mixing ratio and proportion, while, in other seasons, the variations in organic vapors with multiple oxygens were closely correlated with those of organic vapors with 1–2 oxygens, which could be substantially influenced by primary emissions. Organic vapors with low oxygen content (≤ 2 oxygens) are comparable to the results obtained by traditional PTR-MS measurements in both urban Beijing and neighboring regions.

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1 Introduction

Volatile organic compounds (VOCs) play a crucial role in the formation of ozone and fine particulate matter ($PM_{2.5}$) in the atmosphere, subsequently affecting air quality, climate, and human health (Carter, 1994; Williams and Koppmann, 2007; Jimenez et al., 2009; Hallquist et al., 2009). The sources and atmospheric evolution of VOCs in the atmosphere are complex due to the coexistence of compounds from both primary emissions and secondary formation (Gentner et al., 2013; Gilman et al., 2015; Millet et al., 2015). Understanding their molecular characteristics is essential for studying their hydroxyl radical (OH) reactivities, ozone, and secondary organic aerosol (SOA) formation potentials. However, the diverse range of species and wide distribution of oxidation products of atmospheric VOCs make it challenging to unravel their molecular properties (Goldstein and Galbally, 2007).

Instrumental advances have allowed us to improve the understanding of the compositions and variations in VOCs at the molecular level, especially for oxygenated VOCs (OVOCs). Gas chromatography, or multidimensional gas chromatography coupled with mass spectrometry, is the most commonly used technology for VOC measurement, capable of detecting major non-methane hydrocarbons and select OVOCs (Lewis et al., 2000; Xu et al., 2003; Noziere et al., 2015). Proton-transfer-reaction mass spectrometry (PTR-MS) enables real-time detection of VOCs without preconcentration and separation, greatly enriching the molecular understanding of OVOCs due to its high sensitivity to oxygen-containing compounds (Hansel et al., 1995; de Gouw and Warneke, 2007; Yuan et al., 2017). Hundreds of OVOCs have been detected and characterized in different areas using PTR-MS, e.g., urban (Wu et al., 2020), suburban (He et al., 2022), and forest areas (Pugliese et al., 2023). Recent developments in the ion-molecule reactor (IMR) configuration have greatly increased sensitivities and concurrently lowered the limits of detection of PTR-MS by several orders of magnitude by incorporating radio frequency electric fields to focus ions (Breitenlechner et al., 2017; Krechmer et al., 2018; Reinecke et al., 2023). A consequential issue is that these advanced PTR-MS instruments typically need to eliminate lighter ions to protect the detector from overload, and, similarly to traditional PTR-MS instruments, they are incapable of obtaining molecular structure information.

These improvements have expanded the detection capabilities of PTR-MS, particularly for organic vapors with lower volatility and multiple oxygens (≥ 3) (Riva et al., 2019), which enables the simultaneous measurement of VOC precursors and their primary, secondary, and higher-level oxidation products using a single instrument (Li et al., 2020). Despite their low concentrations, these vapors may condense on pre-existing aerosols and make a significant contribution to secondary aerosol growth and cloud condensation nuclei (Bianchi et al., 2019; Pospisilova et al., 2020; Nie et al., 2022). Organic vapors with multiple oxygens are likely to be simultaneously detected by other chemical ionization mass spectrometry (CIMS), e.g., nitrate (NO_3^-), iodide (I^-), bromide (Br^-), and ammonium (NH_4^+) (Riva et al., 2019; Huang et al., 2021), which are widely used for measuring oxygenated organic compounds in the atmosphere (Bianchi et al., 2019; Ye et al., 2021; Huang et al., 2021). Therefore, using these improved PTR-MS instruments can supplement our understanding of oxygenated organic vapors and facilitate the study of the atmospheric chemical evolution of organics (M. Wang et al., 2020).

The improved PTR-MS systems have gradually gained traction in research applications over the past few years, including measuring organics in controlled lab studies (Zaytsev et al., 2019a, b; Riva et al., 2019; H. Li et al., 2022; K. Li et al., 2024), emission sources (Sreeram et al., 2022; Yu et al., 2022; Yacovitch et al., 2023; Wohl et al., 2023; Jahn et al., 2023), and ambient air. For ambient measurements, observations in forested regions have been extensively conducted to study the compositions, variations, fluxes, and emissions of organics from different plants (Li et al., 2020, 2021; Huang et al., 2021; Fischer et al., 2021; Thomas et al., 2022; Vettikkat et al., 2023; Vermeuel et al., 2023). Terpenes and their oxidation products with oxygen numbers up to 6 have been detected (Li et al., 2020). Diterpenes have been directly observed in the ambient air for the first time owing to the substantial improvement in sensitivity of Vocus-PTR (Li et al., 2020). Ambient measurements have also been conducted on a mountain in China and found that terpenes and their oxidation products dominate the detected organic compounds, while the influence of industrial emissions can also be observed (Zhang et al., 2024).

In urban atmospheres, the sources and evolution of VOCs are considerably complex, potentially exhibiting distinct characteristics compared to forested areas. Several studies have carried out measurements in urban air using these improved PTR-MS instruments. Jensen et al. (2023) conducted a 1-month observation to address the production of reliable measurements. Coggon et al. (2024) evaluated the fragmentation and interferences of a series of urban VOCs. Pfannerstill et al. (2023a, 2024) measured hundreds of VOCs to calculate their emission fluxes in Los Angeles. A few low-signal species, including dimethylamine, icosanal, dimethyl disulfide, and siloxanes, emitted from diverse emission sources have been detected as a result of the enhanced sensitivity (Y. Wang et al., 2020a; Chang et al., 2022; Jensen et al., 2023). However, the understanding of organic vapors with multiple oxygens in urban air, including their species, mixing ratios, diurnal profiles, and seasonal variations, remains inadequate.

In this study, we conduct measurements of organic vapors using a Vocus-PTR in urban Beijing during 2021–2022, covering four seasons. We present general characteristics of measured organic vapors and compare them with traditional PTR-MS and previous Vocus-PTR measurements. We focus

on organic vapors with multiple oxygens (3 or more), which have rarely been individually analyzed in previous studies due to their low mixing ratios. Their chemical compositions, atmospheric mixing ratios, and diurnal and seasonal variations are reported. Cluster analysis is further conducted to resolve the main driving factors of their variations.

2 Methods

2.1 Measurements

The observation site is located in the central area of Tsinghua University, Beijing (40°0′ N, 116°20′ E). It is an urban site with no significant direct influence from industrial activities or heavy-traffic arteries (Fig. S1 in the Supplement). Details of this site can be found in a previous study (Cai and Jiang, 2017). Organic vapors were measured by a Vocus-2R PTR-TOF-MS (referred to as Vocus-PTR hereinafter; Tofwerk AG and Aerodyne Research Inc.), which is situated on top of a fourth-floor tower building, with its sampling inlet positioned approximately 20 m above the ground. The observation period is from 1 May 2021 to 10 March 2022, covering four seasons. Detailed information about observation periods and their corresponding seasons is shown in Table S1 in the Supplement.

The operating parameters of the Vocus-PTR used in this study are briefly described here. In PTR-MS, VOCs are ionized via proton transfer by hydronium ions (H_3O^+) in the IMR (Hansel et al., 1995; Yuan et al., 2016). The sensitivity can be quantified based on the proton transfer reaction rate while simultaneously considering ion transmission, detector efficiency, etc. (Cappellin et al., 2012; Jensen et al., 2023). The ion source was supplied with a water vapor flow of 20 sccm. The IMR was operated at 100 °C and 2 mbar with an axial voltage of 600 V and a quadrupole amplitude voltage of 450 V. The IMR operating parameters were optimized to minimize the formation of water clusters. Mass spectra were collected from m/z 11 to m/z 398 with a time resolution of 5 s, achieving a mass resolution of ∼ 10 000 for $C_7H_9^+$ throughout the measurement period. Ambient air was sampled via a tetrafluoroethylene (PTFE) tube (1.35 m long, 1/4 in. OD) at a flow rate of 3 L min^{-1} to reduce wall losses, with only 150 sccm flow entering the Vocus-PTR. The sampling tube was heated to 50 ± 5 °C during the measurement. A regularly replaced Teflon filter (every 7 d) was used in front of the sampling line to prevent the orifice from clogging. The data within 30 min after membrane replacement were excluded. Measurements were made on a 2 h cycle with 110 min for ambient air, 5 min for zero gas, and 5 min for fast calibration. The fast calibrations involved the use of mixed-calibration gases, with detailed information available in Table S2.

The ambient $PM_{2.5}$, NO_2, and O_3 data are from a state-operated air quality station (Wanliu station), located approximately 3.6 km away from our observation site. The meteorological parameters, including temperature (T), relative humidity (RH), wind speed, and wind direction, are also from Wanliu station. The diurnal variations in $PM_{2.5}$, O_3, NO_x, RH, and T in four seasons are shown in Fig. S2.

2.2 Data processing

Data analysis of Vocus-PTR mass spectra, including mass calibration, baseline subtraction, and high-resolution peak fitting, was conducted using Tofware (v3.2.3; Tofwerk AG and Aerodyne Research Inc.) within the Igor Pro 8 platform (WaveMetrics, OR, USA). The ambient mass spectra were averaged over 1 min for subsequent processing in Tofware. The peaklist used for high-resolution peak fitting was manually made based on mass spectra of both clean days ($PM_{2.5} < 75\,\mu g\,m^3$) and polluted days ($PM_{2.5} \geq 75\,\mu g\,m^3$). The maximum mass error allowed for identifying peaks is 5–10 ppm, which is consistent with the error of mass calibration. When there were multiple options of formulas meeting the error limit under a peak with oxygen numbers ≤ 8 and carbon numbers ≤ 20, especially at high molecular weights, a lower degree of unsaturation was selected; otherwise, the peak would be classified as unknown peak. The maximum peak area residual for each unit mass resolution was 5 %. Subsequent analysis was performed in MATLAB R2022a (The MathWorks Inc., USA).

In PTR-MS, the sensitivities of organic vapors are typically determined through their direct linear correlation with their PTR rate constant (k_{PTR}). Vocus-PTR utilizes a big segmented quadrupole with a high-pass band filter, which detects ions < 35 m/z with reduced transmission efficiency (Krechmer et al., 2018). Consequently, determining sensitivities in Vocus-PTR involves consideration of both reaction efficiency and transmission efficiency. Figure S3a shows the measured sensitivities of mixed-calibration gases and their corresponding k_{PTR} values. The linear regression between k_{PTR} and sensitivities was obtained based on sensitivities of $C_7H_9^+$, $C_8H_{11}^+$, $C_9H_{13}^+$, $C_{10}H_9^+$, and $C_5H_9O_2^+$ with an R^2 of 0.87. Sensitivities of other ions in mixed-calibration gases may be influenced by transmission (ions labeled as gray) and fragmentation ($C_5H_9^+$, $C_{10}H_{17}^+$, and $C_{11}H_{11}^+$). The transmission efficiency of mixed-calibration gases was calculated using sensitivities of mixed-calibration gases, as shown in Fig. S3b. The transmission efficiency of mixed-calibration gases aligns well with the fitted transmission efficiency curve, except for $C_5H_9^+$, $C_{10}H_{17}^+$, and $C_{11}H_{11}^+$, which potentially experience fragmentation (fragmentation of measured ions is discussed below). For organic vapors without standards, their theoretical k_{PTR} were used to constrain sensitivities, while, for organic vapors with no theoretical k_{PTR}, an average k_{PTR} of known species, 2.5×10^{-9} cm^3 molec.$^{-1}$ s^{-1}, was used to constrain their sensitivities. The theoretical k_{PTR} of organic vapors is from previous studies (Zhao and Zhang, 2004; Cappellin et al., 2012; Sekimoto et al., 2017). Average limits of detection (LODs; 1 min) of the measured com-

pounds were determined using zero-gas background measurements taken every 2 h during the observation periods, as shown in Fig. S4. The LODs were calculated as 3 times the standard deviation of the zero-gas background divided by the obtained sensitivity. The LODs show a correlation with masses; as masses increase, instrument backgrounds decrease, leading to lower LODs. This trend was observed for species with different oxygen content, with LODs around 0.03 ± 0.03 pptv at m/z 200. Note that LODs in this study are 1 min averages, with raw 1 s data averaged to 1 min before Tofware analysis, as mentioned before, which may account for the lower LODs compared to those in Jensen et al. (2023).

The fragmentation, water cluster, and interferences for calibrated and uncalibrated species were corrected. The ratio of the electric field strength (E) to the buffer gas number density (N) used in our study was 146.9 Td, and the gradient between BSQ skimmer 1 and skimmer 2 was 9.8 V, which in this case limited the formation of water clusters, promoted the simple reaction kinetics, and improved the sensitivity but may have led to stronger fragmentation. For α-pinene, we identified its fragments based on GC chromatograms. The Vocus-PTR was calibrated in GC mode before atmospheric measurement. A total of four species were tested in GC mode, including severely fragmented α-pinene. The spectrum of α-pinene showed that the main fragment was $C_6H_9^+$. Several long-chain aldehydes and cycloalkanes may fragment on $C_5H_8H^+$, the ion typically attributed to isoprene in PTR-MS (Gueneron et al., 2015; Pfannerstill et al., 2023b; Coggon et al., 2024). We corrected isoprene signals following an approach by Coggon et al. (2024). The correction was calculated as follows:

$$m/z\ 69.07_{\text{Corrected}} = S_{69.07} - S_{111.12+125.13} \cdot f_{69.07/(111.12+125.13)}. \quad (1)$$

$S_{69.07}$ is the signal measured at $C_5H_9^+$. $S_{111.12+125.13}$ is the signal of the isoprene interferences, referring to $C_8H_{15}^+$ (m/z 111.12) and $C_9H_{17}^+$ (m/z 125.13), which are dehydrated products from octanal and nonanal, respectively. $f_{69.07/(111.12+125.13)}$ was determined from nighttime data (00:00–04:00 local time) of each period. Similarly, acetaldehyde was corrected for ethanol fragments. We also checked the fragments and water cluster list in Pfannerstill et al. (2023b) and Jensen et al. (2023). When the Pearson correlation coefficient r is greater than 0.95, the ions were regarded as fragments or water clusters of the parent ion. We also tried to exclude the effects of unknown fragments and water clusters based on correlations of times series. Similar to Pfannerstill et al. (2023b), any ion showing a correlation with another ion with $r^2 > 0.97$ (if chemically reasonable) was analyzed for possible water clustering or fragmentation effects and added up with its parent ion. The ions corrected are specified in Table S3.

Here, we discuss the uncertainties of quantification for calibrated and uncalibrated compounds. The uncertainty of calibrated ions ranges from 2 % to 16 % determined from the standard deviations of the fast calibrations during the measurement periods. The quantification was conducted for uncalibrated compounds with their sensitivities constrained by the k_{PTR} linear relationship and transmission efficiency. The uncertainty of these uncalibrated compounds arising from linear fitting and transmission efficiency fitting is 20 % using Monte Carlo simulation. Additionally, undetermined fragmentations and water clusters also contribute to the uncertainty, though we identified some potential fragments and water clusters through the strength of correlations as previously indicated. We acknowledge that this method cannot identify all fragments and clusters and that fragments and clusters may still be present in the measured VOCs and OVOCs. Further research is needed to explore the impact of fragments and clusters on the measurements, particularly concerning OVOCs with multiple oxygens.

The double bond equivalent (DBE), carbon oxidation state ($\overline{OS}_C$), and volatility of organic vapors were calculated to address the chemical and physical properties of detected organic vapors (see Sect. S1 in the Supplement). The condensational growth rates contributed by detected organic vapors were simulated using a kinetic partitioning method, as detailed in Y. Li et al. (2024). For comparison, the condensational growth rates of low volatile and extremely low volatile organic compounds measured by nitrate-CIMS were also simulated (Y. Li et al., 2024). The OH reactivities of detected organic vapors were calculated, and the rate constants are from Data S1 in Pfannerstill et al. (2024) and Table S4 in Wu et al. (2020). For species with unreported rate constants, we calculated the OH reactivities for hydrocarbons and OVOCs using the reported median rate constants of hydrocarbons and OVOCs, respectively.

Quantified mixing ratios were further processed by cluster analysis to investigate their characteristics. The intraclass correlation coefficient (ICC) is a suitable method for assessing the consistency of trends in unbalanced data. It quantifies the stability of differences between two sets of measurement results, enabling the evaluation of their consistency. The ICC combined with k-means cluster analysis was used. ICC(C, 1) was selected among several typical consistency evaluation parameters for its evaluation results, which exhibit the highest level of differentiation based on factual evidence (Qiao et al., 2021). ICC(C, 1) was calculated as follows:

$$\text{ICC(C, 1)} = (D(X+Y) - D(X-Y))/(D(X+Y) + D(X-Y)), \quad (2)$$

where $D(\cdot)$ is the arithmetic operators of variance. X and Y are two sets of measurement data, in this case referring to the mixing ratios of any organic vapors we are concerned about. The ICC matrices of various organic vapors were subsequently utilized as input for k-means analysis. The square Euclidean distance was selected to calculate the distances between different organic vapors.

3 Results and discussion

3.1 General characteristics of organic vapors

During the measurement period, a total of 895 peaks were observed, and 512 of them can be assigned to formulae, divided into C_xH_y, $C_xH_yO_z$, $C_xH_yN_i$, and $C_xH_yO_zN_i$ categories based on their elemental compositions (Fig. 1a). $C_xH_yO_z$ comprises up to 54 % of the total number of formulae followed by $C_xH_yO_zN_i$, C_xH_y, and $C_xH_yN_i$, with proportions of 26 %, 14 %, and 6 %, respectively (Fig. 1b). $C_xH_yO_z$ dominates, contributing 74 % of the annual median mixing ratios of total organics, followed by C_xH_y, $C_xH_yO_zN_i$, and $C_xH_yN_i$, with proportions of 22 %, 2 %, and 2 %, respectively (Fig. 1c). In addition to these resolved formulae, we also detect 18 peaks containing other elements, such as S, Cl, and Si, and 79 CH(O)(N) peaks that do not comply with nitrogen rules, which we regard as fragments or free radicals. Others are unknown peaks for which formulae cannot be assigned or water clusters/fragments excluded from analysis. The mixing ratios of organic vapors vary substantially in urban Beijing, ranging from 0.01 parts per trillion (ppt) to 10 parts per billion (ppb) in volume at a time resolution of 1 min, with many species detected at sub-ppt levels in particular (Fig. 1d). The units of the mixing ratio in the following text are all volume fractions. As the molecular masses of organics increase, their annual median mixing ratios decrease. The mixing ratios of $C_xH_yO_z$ and $C_xH_yO_zN_i$ categories start to decrease below the ppt level above molecular weights of 160 and 125, respectively.

With enhanced sensitivity and mass resolution, an increased number of formulae have been identified compared to traditional PTR-MS measurements in urban Beijing, especially formulae with lower mixing ratios and higher oxygen contents. Note that most organics with low mixing ratios have high oxygen content. Of the formulae measured in this study, 44 % are at sub-ppt level, while 31 % of the formulae are between 1 and 10 ppt (Fig. 1e). Only compounds detected above ppt levels were previously reported in urban sites within Beijing (Sheng et al., 2018; Li et al., 2019) and at a suburban site located 100 km southwest of Beijing (He et al., 2022). Simultaneously, organic vapors with multiple oxygens ($C_xH_yO_{\geq 3}$ and $C_xH_yO_{\geq 3}N_i$ species) have successfully been detected in this study in the urban atmosphere. Traditionally, they have often been recognized as total $C_xH_yO_{\geq 3}$ species, with no individual analysis in traditional PTR-MS (Yuan et al., 2023; X.-B. Li et al., 2022; He et al., 2022). Many other studies only focus on reporting OVOCs containing up to 2–3 oxygens or omit to address the presence of nitrogen-containing OVOCs (L. Wang et al., 2021; Liu et al., 2022). The low mixing ratios and high wall losses of organic vapors with multiple oxygens impact the detection in traditional PTR-MS (Breitenlechner et al., 2017). Figure 2a reinterprets the mass defect plot of measured organics with a focus on oxygen numbers, ranging from 0 to 8. The analysis of mixing ratio levels and variations in organic vapors with multiple oxygens (≥ 3) are shown in Sect. 3.2. Organic vapors with low oxygen content (≤ 2) are reported in Sect. 3.3. Subsequent comparisons of Vocus-PTR and traditional PTR in urban Beijing and Vocus-PTR measurements in both urban Beijing and European forests are also shown in Sect. 3.3.

3.2 Organic vapors with high oxygen content

A total of 195 observed organics with multiple oxygen atoms account for 38 % of the total organics, including 136 species of $C_xH_yO_{\geq 3}$ and 59 species of $C_xH_yO_{\geq 3}N_i$. Organics with oxygen numbers 3 and 4 dominate within the $C_xH_yO_{\geq 3}$ and $C_xH_yO_{\geq 3}N_i$ species (Fig. 2b and c). Organics with oxygen numbers of 3, 4, 5, and ≥ 6 comprise 15 %, 11 %, 7 %, and 6 % of the total species number of $C_xH_yO_z$ compounds, respectively, while compounds with oxygen numbers of 3, 4, 5, and ≥ 6 comprise 15 %, 12 %, 7 %, and 2 % of the total species number of $C_xH_yO_zN_i$ compounds, respectively.

The measured organic vapors with multiple oxygens are mainly intermediate volatile organic compounds (IVOCs) and semi-volatile organic compounds (SVOCs). The dominant carbon numbers range from 5 to 9, and the DBE ranges between 1 and 5, accounting for over three-quarters of the total species number of organic vapors with multiple oxygens (Fig. 3a and b). The maximum occurrence of organic vapors with 3 or 4 oxygen atoms is observed within the carbon range of 7–8 and a DBE value of 2. For organic species with 5 or more oxygens, they reach their peak at a smaller carbon number of 4–5 and a higher DBE value of 3. Aromatic VOCs have DBE values no smaller than 4, while aliphatic VOCs usually have DBE values smaller than 2. For organic vapors with a DBE between 2 and 3, they are likely oxidation products of aliphatic and aromatic VOCs (Z. Wang et al., 2021; Nie et al., 2022). For the same number of carbon atoms, organic vapors with a higher number of oxygen atoms exhibit a higher carbon oxidation state (as shown in Fig. S5). Compared to organic vapors with 3 or 4 oxygen atoms, organic vapors with 5 or more oxygens have undergone more extensive atmospheric oxidation and functionalization processes (Kroll et al., 2011; Isaacman-VanWertz et al., 2018). Based on calculated volatility, 81 % of the species are IVOCs, and the remaining 19 % are SVOCs (Fig. 3c). With the increase in oxygen number, the volatility of the compounds gradually decreases, while the potential partitioning to aerosols increases, manifested by a gradual reduction in the peak values of the $\log_{10}C_0$. Compounds containing nitrogen, referred to as shaded bars with white stripes in Fig. 3c, have a lower volatility compared to non-nitrogen species.

The annual median mixing ratio of measured organic vapors with multiple oxygens in median $\pm$ standard deviation is 2.0 ± 0.9 ppb, accounting for 4 % of the total $C_xH_yO_z$ and $C_xH_yO_zN_i$ mixing ratios. For the $C_xH_yO_z$ category, the annual median mixing ratios of species with 3, 4, 5, and ≥ 6 oxygens are 1.4 ppb, 186.0 ppt, 17.8 ppt, and 5.9 ppt, respec-

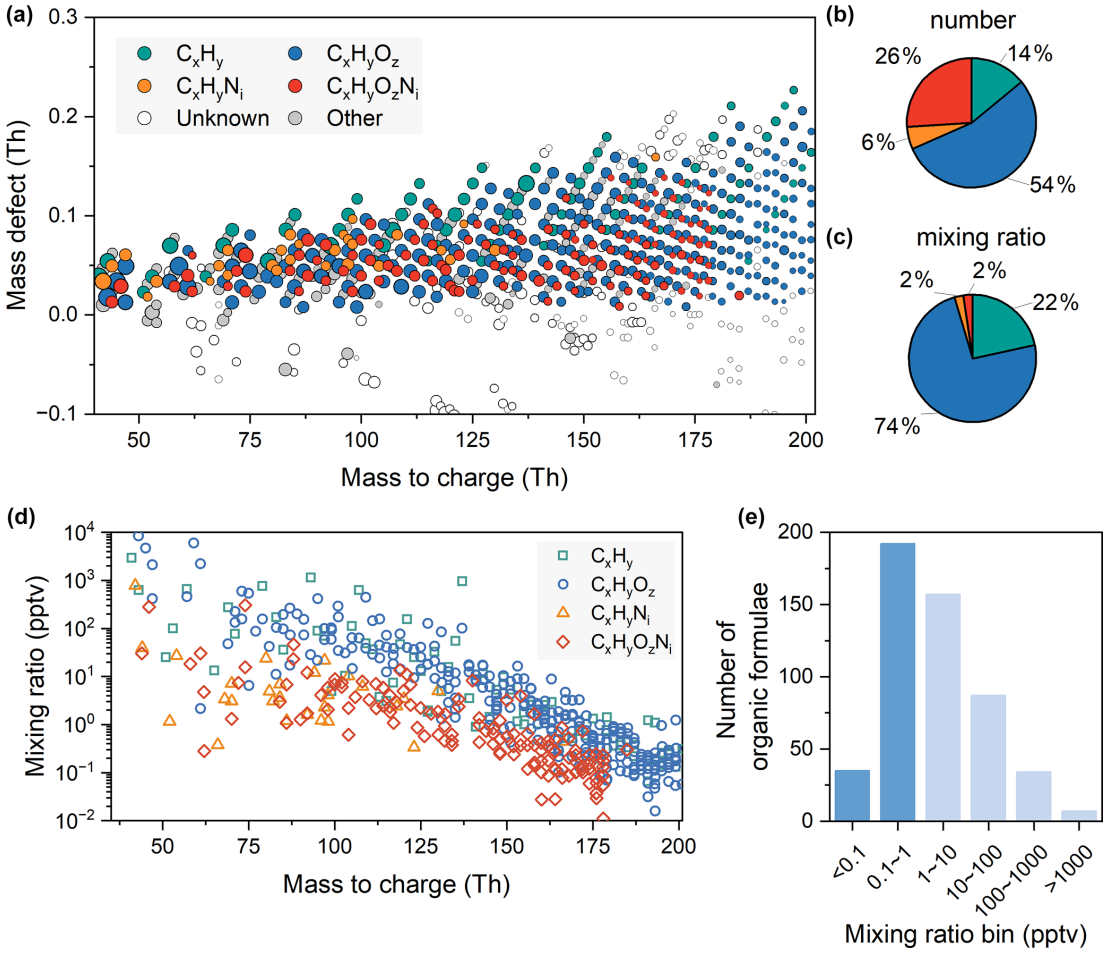

Figure 1. Identified formulae in urban Beijing using Vocus-PTR. **(a)** Mass defect plot. The sizes of the bubbles represent the annual median mixing ratios. The bubbles are colored by different elemental compositions as labeled in the legend. The "unknown" refers to fitted peaks without matched formulae. The "other" refers to peaks containing elements other than C, H, O, and N or fragment peaks (or radicals). **(b)** Pie chart of the number of identified formulae. **(c)** Pie chart of the annual median mixing ratios of identified formulae. The color scheme of the pie charts is the same as that of the mass defect plot. **(d)** The annual median mixing ratios of identified formulae versus their masses. **(e)** Histogram of annual mixing ratios of identified formulae. Bins with values less than 1 ppt are emphasized in dark blue.

tively. For the $C_xH_yO_zN_i$ category, the annual median mixing ratios of species with 3, 4, 5, and ≥ 6 oxygens are 49.6, 24.5, 2.4, and 0.5 ppt, respectively (Fig. 2d and e). Organic vapors with 3 oxygens constitute the overwhelming majority of the mixing ratio of measured organic vapors with more than 3 oxygens. As a result, the mixing-ratio-weighted carbon number and DBE distributions (Fig. 3d and e) are significantly different from those of species number distributions for organic vapors with multiple oxygens. The mixing ratios of species with carbon numbers ranging from 2 to 6 are significantly higher, with those containing 4 carbons exhibiting the highest mixing ratios. Similarly, the mixing ratios of species with a DBE ranging from 0–4 are notably higher than those of other DBE values. As compounds containing 3 oxygens dominate the mixing ratio, IVOCs nearly entirely contribute to the mixing-ratio-weighted volatility of organic vapors with multiple oxygens (Fig. 3f). The mixing ratios of organic vapors with multiple oxygens measured in this study are higher than in other studies and will be detailed in Sect. 3.3.

Though the contribution of the measured IVOCs and SVOCs to the overall VOC mixing ratio is low, their contribution to the condensational growth rates is non-negligible and may influence the growth of new particles (Ehn et al., 2014), SOA formation (Jimenez et al., 2009), and haze (Nie et al., 2022). The condensational growth rates of total organic vapors are calculated, including extremely low, low, and semi-volatile organic compounds detected by nitrate-CIMS and I/SVOCs detected by Vocus-PTR. The contribution to the condensational growth rate from I/SVOCs detected by Vocus-PTR increases with particle size and decreases with temperature. For 8 nm particles, the contribution of SVOCs detected by Vocus-PTR is 9 %, while IVOCs contribute 1 %. For 40 nm particles, the contribution of SVOCs

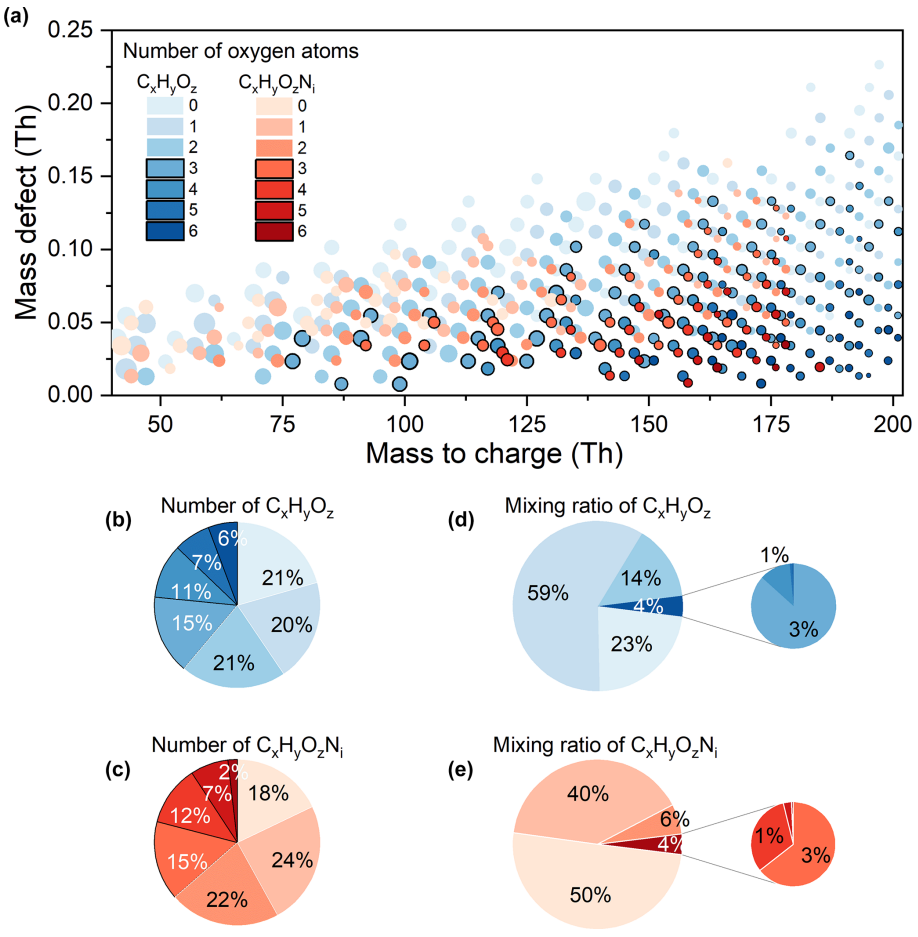

Figure 2. Organic vapors of different oxygen content. **(a)** Mass defect plot. The sizes of the bubbles represent the annual median mixing ratios. The bubbles are colored according to different oxygen numbers as labeled in the legend. Bubbles representing organic vapors with 3 or more oxygens are highlighted with black borders. Bars labeled as 6 refer to organic vapors with oxygen numbers equal to or larger than 6. **(b)** Pie chart of the number of $C_xH_yO_z$ species. **(c)** Pie chart of the number of $C_xH_yO_zN_i$ species. **(d)** Pie chart of the mixing ratio of $C_xH_yO_z$ species. **(e)** Pie chart of the mixing ratio of $C_xH_yO_zN_i$ species. The color scheme of the pie charts is the same as that of the mass defect plot.

increases to 13 % and IVOCs rise to 4 %. At sub-zero temperatures for 8 nm particles, the SVOC contribution detected by Vocus-PTR can reach up to 21 %, with IVOCs contributing 10 %.

The molecular formulae of the measured organic vapors with multiple oxygens are displayed in the mass spectra, categorized by carbon numbers ranging from 2–11 (Fig. 4 and Table S4). Many of the formulae are reported as oxidation products of various VOC precursors in previous studies. Take isoprene as an example: detected formulae are reported as various oxidation products of isoprene, including $C_5H_{10}O_3$ and subsequent oxidation products in C5 species, e.g., $C_5H_8O_6$ and $C_5H_9NO_4$. (Wennberg et al., 2018). For several C4 species, such as $C_4H_7NO_4$ and $C_4H_4O_3$, they are reported as oxidation products of two additional important oxidation products of isoprene, methacrolein (MACR) and methyl vinyl ketone (MVK). We also see formulae reported as oxidation products of precursors such as benzene (C6) (Priestley et al., 2021), alkyl-substituted benzenes (C7–C9) (Pan and Wang, 2014; Y. Wang et al., 2020b; Cheng et al., 2021), and monoterpenes (C10) (Rolletter et al., 2019). Besides, we can also detect some organic vapors with a relatively low DBE (≤ 3), which may originate from the oxidation of aliphatic precursors. For example, $C_5H_8O_4$ observed is reported as one of the oxidation products of C5 aldehyde, the photolysis of which releases OH radicals. This mechanism may explain the source gap of OH radicals between simulations and observations in low-nitrogen-oxide and high-VOC regimes (Yang et al., 2024). Note that these species may be oxidation products as reported by previous studies; however, confirming this would require additional techniques such as GC.

Measured molecular formulae may react with OH radicals, contributing to OH reactivity. The calculated OH reac-

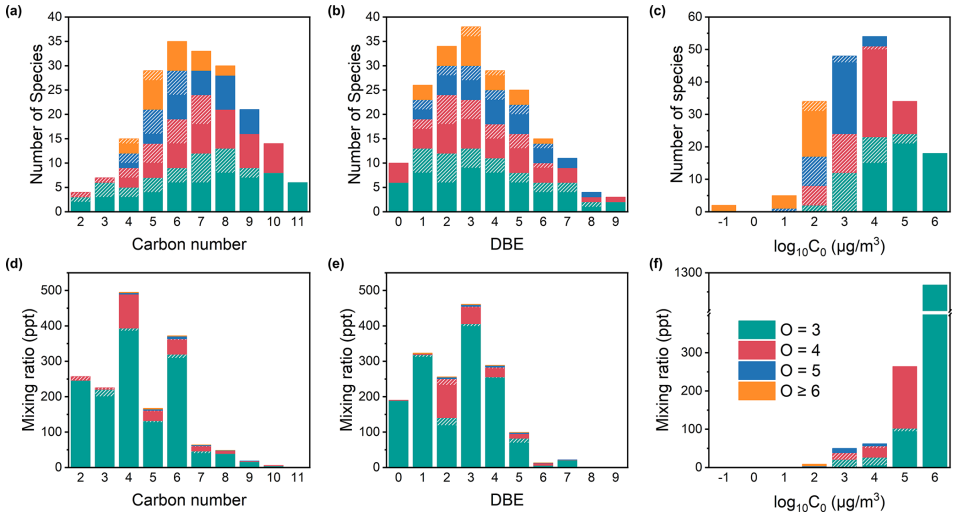

Figure 3. Distribution of carbon number, double bond equivalent (DBE), and volatility of organic vapors with multiple oxygens. Panels **(a)**–**(c)** represent species number distributions of carbon number, DBE, and volatility, respectively. Panels **(d)**–**(f)** represent mixing ratio distributions of carbon number, DBE, and volatility, respectively. Different bar colors refer to compounds with different oxygen content. Bars without white stripes represent $C_xH_yO_{\geq 3}$, while shaded bars with white stripes represent $C_xH_yO_{\geq 3}N$. The y axes refer to annual median mixing ratios.

tivity of organic vapors with multiple oxygens accounts for 6 % of the total detected VOCs, with an average annual value of $1.2\,s^{-1}$. Previous studies show differences between measured and calculated or modeled OH reactivity (Hansen et al., 2014), and unmeasured species from photochemical oxidation likely explain this gap (Ferracci et al., 2018). Therefore, the OH reactivity contributed by detected organic vapors with multiple oxygens in this study may potentially reduce this gap, thereby improving the accuracy of diagnosis of sensitivity regimes for ozone formation (Wang et al., 2024). Using Vocus-PTR has the potential to simultaneously measure both precursors and multi-generational oxygenated products, which is beneficial for studying the evolution process of organic compounds in the atmosphere.

As for the seasonal variations, the overall mixing ratio of organic vapors with multiple oxygens is the highest in winter, followed by summer and spring, and the lowest in autumn (Fig. 5a). The mixing ratios expressed in median ± standard deviation (ppb ± ppb) are 1.9 ± 0.5, 1.9 ± 0.9, 1.4 ± 1.2, and 2.2 ± 0.8 for spring, summer, autumn, and winter, respectively. Compounds with different oxygens exhibit different seasonal variations, as shown in Fig. 5b and c and Table S5. For $C_xH_yO_z$ with 3 or 4 oxygens, the mixing ratios are higher in winter than in other seasons, while, for compounds containing 5 or more oxygens, the mixing ratios are highest in summer. For $C_xH_yO_zN_i$ with 3 or 4 oxygens, the mixing ratios are high in both summer and winter, while, for compounds containing 5 or more oxygens, the mixing ratios are high in summer and spring. As the oxygen number increases, the contribution from secondary sources becomes greater and the high mixing ratio of oxidants in summer intensifies this process. Thus, the fraction of the mixing ratio of compounds with multiple oxygens increases with the oxygen number in summer (Fig. 5d). In winter, the mixing ratios of compounds containing five or more oxygens are substantially suppressed, which may be due to reduced generation. Alternatively, it could be that these compounds belong to SVOCs, with a majority being partitioned onto particulate matter at low temperatures.

The seasonal variations in organic vapors with multiple oxygens differ from those of total OVOCs (Fig. S6), with the latter's mixing ratio being primarily influenced by organic vapors containing 1–2 oxygen atoms. The mixing ratio of total OVOCs in winter is substantially higher than in the other three seasons, followed by autumn and summer, with the lowest mixing ratio observed in spring. The seasonal variations in OVOCs are partly caused by the variation in mixing layer height (Li et al., 2023), which is lowest in winter. Cluster analysis is performed to further explore the dominated driving factors of the seasonal variations in organic vapors with multiple oxygens. Three clusters are identified in each season based on the diurnal profiles of each compound. To increase the interpretability of the clusters, two of them are merged. Figures 6 and S7 show the cluster results for organic vapors with multiple oxygens. For comparison, cluster analysis is performed on organic vapors with 1–2 oxygens as well (Figs. S8 and S9).

Daytime clusters, where the peak occurs during the daytime, were identified across the four seasons for organic vapors with multiple oxygens (shown as cluster 1 in Fig. 6). Daytime clusters start to rise at 06:00–07:00 (06:00 for summer and 07:00 for other seasons), peak at 11:00–14:00, and

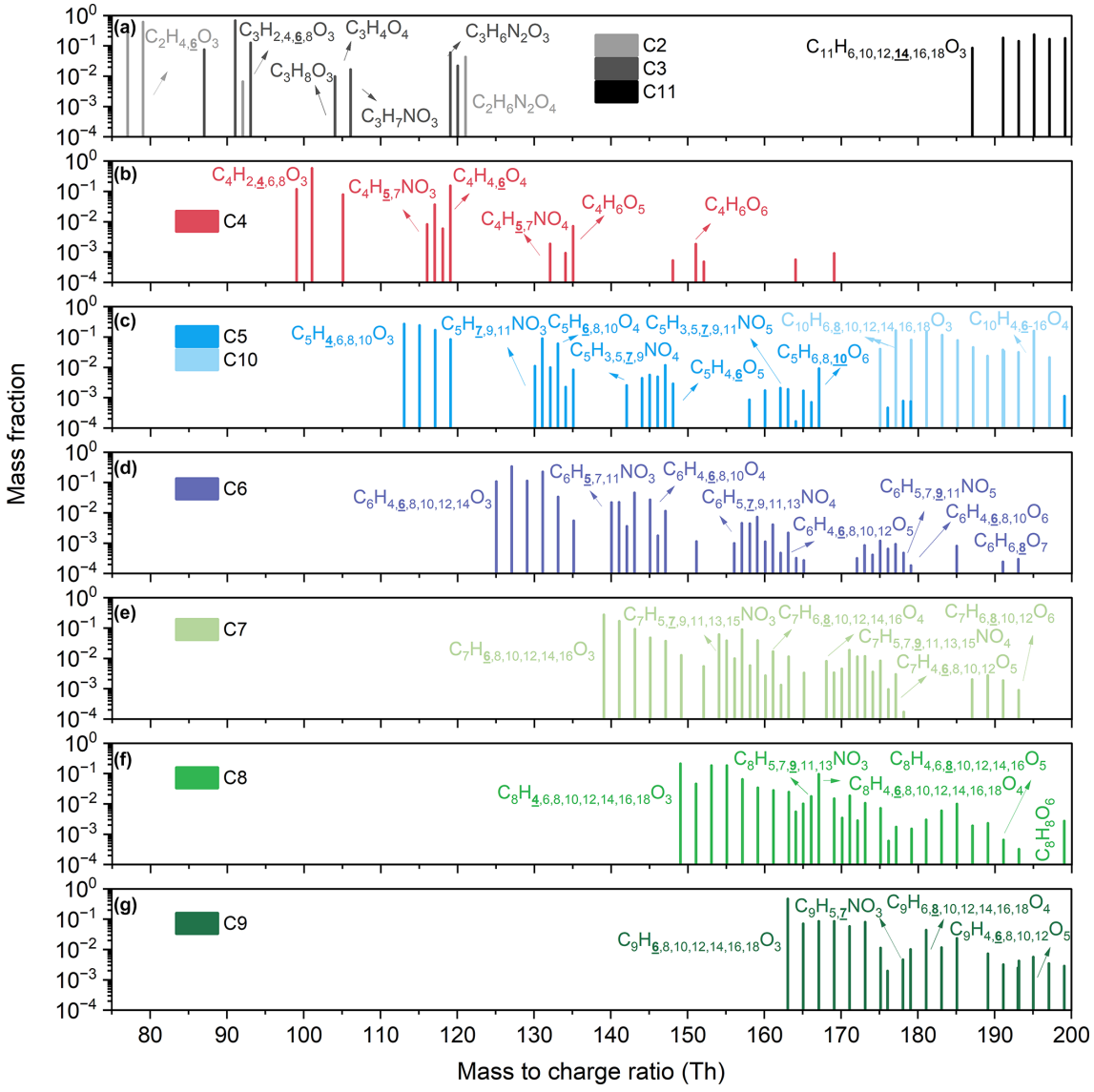

Figure 4. Mass spectra of organic vapors with multiple oxygens with different carbon numbers: **(a)** C2, C3, and C11; **(b)** C4; **(c)** C5 and C10; **(d)** C6; **(e)** C7; **(f)** C8; and **(g)** C9. The *y* axis shows the annual median mixing ratio fraction of organic vapors for each carbon number, which means that, for different organic vapors with the same carbon number, the sum of the mixing ratio fractions is 1. The unprotonated formulae of organic vapors with multiple oxygens are labeled. In molecular formulae with the same number of carbons and oxygens, the hydrogen content in the organic vapors with the highest intensity is emphasized by bold and underlined formatting.

then slowly decrease, following the diurnal variation in solar radiation (Li et al., 2023), ozone, and temperature (Fig. S2). Figure S10 further demonstrates the dependence of daytime clusters on temperature. The mixing ratio of daytime clusters shows an apparent increase in summer (when temperature is higher than 15 °C), which indicates that higher temperatures accompanied by an increase in solar radiation and ozone favor the formation of daytime clusters. The number and corresponding mixing ratios of species allocated to the daytime clusters vary in four seasons. In summer, the vast majority of species (76 %) exhibit daytime characteristics, with a mixing ratio percentage as high as 82 %, which may be related to the strongest solar radiation (Li et al., 2023) and lowest NO_x concentrations (Fig. S2). The contribution of daytime clusters in autumn is also significant, with 68 % and 61 % of the species and mixing ratios being accounted for. The noon peaks of daytime clusters in winter and spring are relatively less pronounced, with the species and day / night mixing ratios also being comparatively lower. The afternoon peak of daytime clusters in autumn and winter is accompanied by a decrease in mixing layer height (Li et al., 2023). For organic vapors with 1 or 2 oxygens, a significant daytime cluster was observed only in summer (Fig. S8d–f).

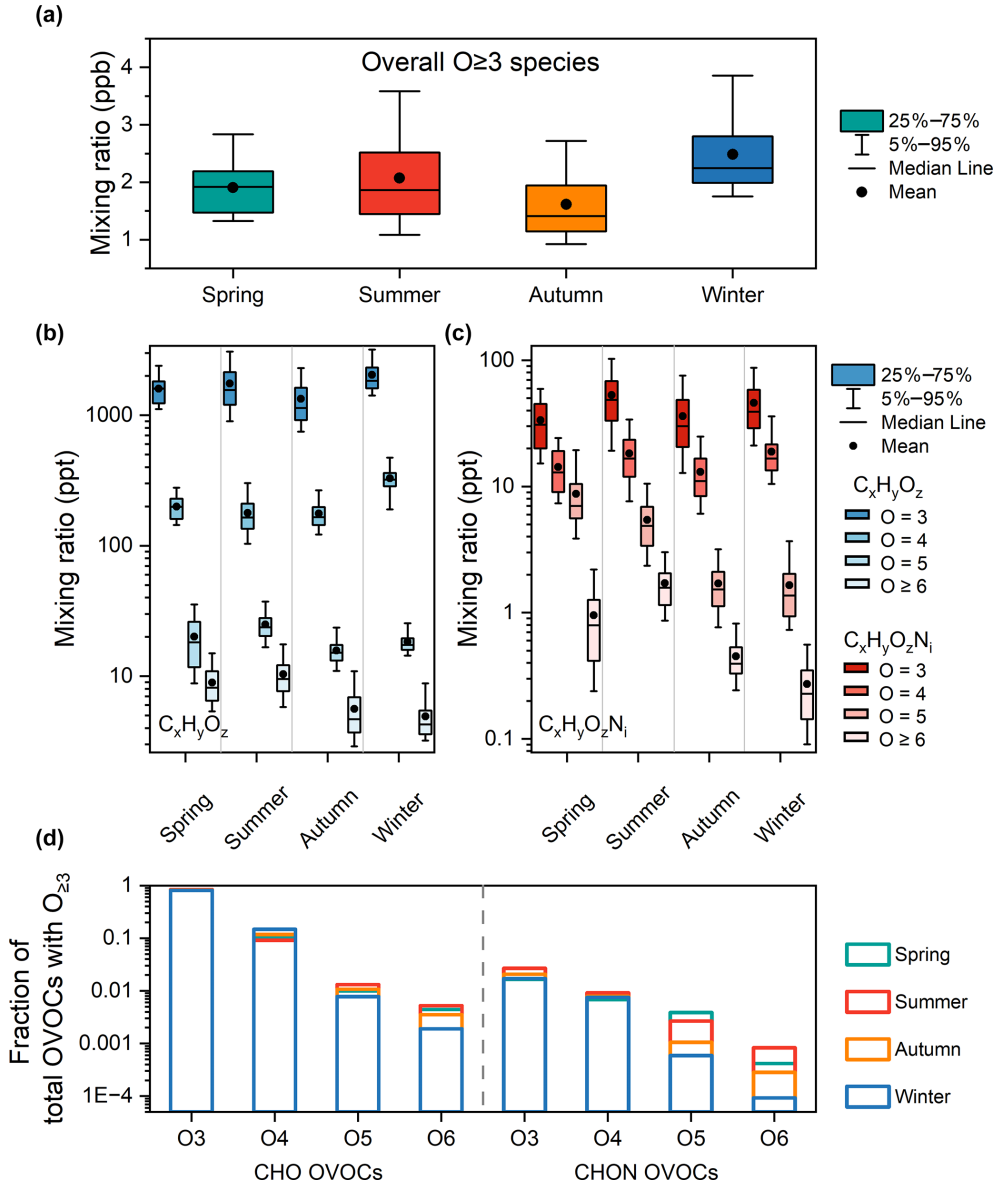

Figure 5. Seasonal variations in organic vapors with multiple oxygens in urban Beijing. **(a)** Total organic vapors with multiple oxygens. **(b)** $C_xH_yO_z$ with different oxygens. **(c)** $C_xH_yO_zN_i$ with different oxygens. **(d)** Fractions of organic vapors with different oxygens of total organic vapors with multiple oxygens.

Another cluster type considered is nighttime clusters, as the corresponding species have their highest mixing ratios at night. Unlike the daytime cluster, the diurnal variations in nighttime clusters are different in four seasons (Fig. 6). In spring, the nighttime cluster comprises over 87 % of species and 75 % of mixing ratios, and it peaks at 04:00 with low daytime values. The nighttime clusters in winter and autumn show bimodal diurnal variations, with the highest peak occurring during the night from 19:00 to 23:00 and the second peak appearing during the day from 08:00 to 12:00. Altogether, 45 % and 32 % of species exhibit the characteristics of the nighttime cluster in winter and autumn, constituting 59 % and 39 % of the mixing ratio, respectively. The contribution of the nighttime cluster is minimal in summer, reaching its peak at midnight. We found that each nighttime cluster of organic vapors with multiple oxygens shows good consistency with the corresponding major clusters of organic vapors containing 1–2 oxygens (Figs. S8 and S11), while the mixing ratios during midday differ. Nighttime clusters also show better consistency with $PM_{2.5}$ compared to daytime clusters (Fig. S2), which may be related to mixed sources.

Most organic vapors with multiple oxygens could be assigned to different clusters in different seasons (Fig. S12). Only a small number of species can be categorized into the

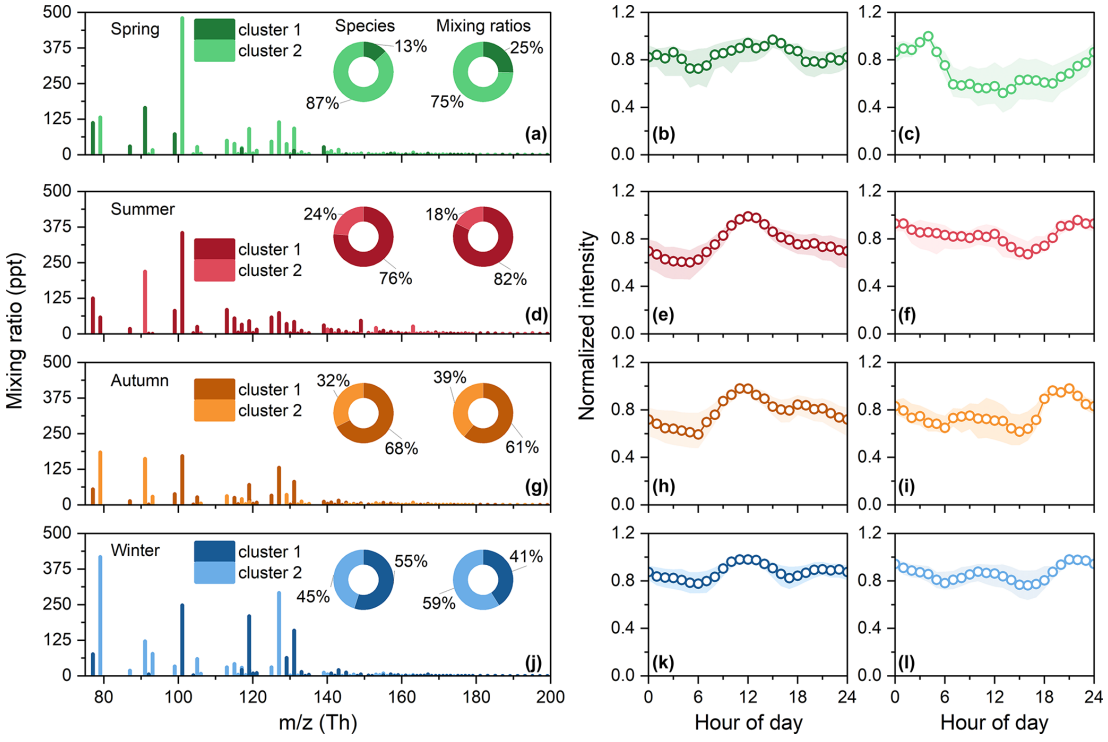

Figure 6. Cluster results of organic vapors with multiple oxygens in four seasons. **(a–c)** Cluster results for spring. **(a)** Mass spectra of organic vapors with multiple oxygens in spring. The *y* axis is the median mixing ratio of each compound. Two different shades are used to distinguish between two clusters. Two pie charts represent the distribution of species numbers and mixing ratios of organic vapors for two clusters. **(b)** Normalized median diurnal variation in cluster 1, daytime cluster. **(c)** Normalized median diurnal variation in cluster 2, nighttime cluster. The shaded areas in panels **(b)** and **(c)** represent the 25th and 75th percentiles. **(d–f)** Cluster results for summer. **(g–i)** Cluster results for autumn. **(j–l)** Cluster results for winter.

same cluster in four seasons. Figure S13 shows the average C, H, O, and N number of species assigned to daytime cluster zero to four times during the four seasons. As compounds exhibit more characteristics associated with the daytime cluster, there is no significant change in the carbon number, but there is an increase in hydrogen and oxygen number and a decrease in nitrogen number. This may be due to multi-step oxidation reactions in the atmosphere, causing an increase in the oxygen number and DBE of species (Kroll et al., 2011; Isaacman-VanWertz et al., 2018), with diurnal variations peaking at noon as a result of the strongest photochemistry. The decreasing trend in the number of nitrogen atoms in Fig. S13 indicates that nitrogen-containing compounds measured in this study are more likely to come from nocturnal production or emissions. Regarding the average elemental composition (C, H, O, and N) of species assigned to two clusters (see Fig. S14), daytime clusters typically exhibit higher oxygen content and lower H/C compared to nighttime clusters, providing further evidence supporting the atmospheric photochemical origin of daytime clusters. The nighttime clusters have higher nitrogen contents than daytime clusters, indicating more of the impacts of nocturnal sources.

3.3 Organic vapors with low oxygen content

In addition to multiple oxygens, organic vapors with low oxygen content were also measured in urban Beijing in this study. Here we primarily discuss comparisons between the results of this study and those of previous studies. The mixing ratios and variations in typical VOCs measured in this study are comparable to the results obtained by traditional PTR-MS measurements in both urban Beijing and neighboring regions. Figure S15 shows the diurnal profiles of 12 representative VOCs in four seasons. OVOCs of C_2H_4O, C_3H_6O, and C_4H_4O, usually identified as acetaldehyde, acetone, and furan, are mainly from anthropogenic sources, as reported by previous studies (Qian et al., 2019). Their diurnal variations exhibit a characteristic of being higher at night and lower during the day, similar to other studies reported in Beijing during the winter (Sheng et al., 2018; He et al., 2022). The mixing ratios of acetaldehyde, methyl ethyl ketone (MEK), and furan in winter are slightly lower than those observed in Beijing in the winters of 2016 and 2018 (Sheng et al., 2018; He et al., 2022). The winter mixing ratios of acetone are higher than in other seasons, and this is observed in other studies, indicating an unknown emission source during winter. The mixing ratios of benzene (C_6H_6), toluene (C_7H_8),

and naphthalene ($C_{10}H_8$) in winter are slightly lower than reported in winter in Beijing during the past few years (Sheng et al., 2018; Li et al., 2019; He et al., 2022), possibly due to improvements in air pollution policies, especially those targeting emissions from residential combustion and motor vehicles (Liu et al., 2023). As for phenols, the mixing ratios of C_6H_6O are similar to measurement at a background site in the North China Plain in winter, while the mixing ratios of C_7H_8O are much lower than that (He et al., 2022). High mixing ratios of biogenic emissions in summer are observed; for example, isoprene (C_5H_8) and the sum of its oxidation products MACR and MVK (Apel et al., 2002) have peak mixing ratios of 2.6 and 0.6 ppb, respectively. Their mixing ratios in winter are lower and consistent with other studies (Sheng et al., 2018; He et al., 2022).

The mixing ratio fractions of organic categories in urban Beijing using Vocus-PTR differ from the results obtained using traditional PTR-MS. Previous studies in Beijing have only reported a few selected VOCs up to around 100 species, resulting in limited results on systematic characterizations of VOCs using PTR-MS in Beijing (Sheng et al., 2018; Li et al., 2019; L. Wang et al., 2021; Liu et al., 2022). Therefore, we compare with a suburban site, Gucheng, which is located 100 km southwest of our site. The two sites (urban Beijing and Gucheng) are both located in the North China Plain and are subject to regional air pollution simultaneously. Figure S16 shows the comparison results of five categories, including C_xH_y, C_xH_yO, $C_xH_yO_2$, $C_xH_yO_{\geq 3}$, and N/S-containing compounds. The first difference is that the mixing ratio fraction of species containing 2 or more oxygens measured by Vocus-PTR is higher than that measured by traditional PTR-MS. The mixing ratio fractions of $C_xH_yO_2$ and $C_xH_yO_{\geq 3}$ in Vocus-PTR are 12 % and 4 %, respectively, whereas they are 6 % and 1 % for traditional PTR-MS. In terms of mixing ratios, the mixing ratio of $C_xH_yO_{\geq 3}$ is approximately double in Vocus-PTR compared to traditional PTR-MS, while the mixing ratio of C_xH_yO is half compared to traditional PTR-MS measurements. The mixing ratio of $C_xH_yO_2$ remains similar. This is because Vocus-PTR can detect more OVOCs with multiple oxygens due to its high sensitivity and mass resolution, whereas, due to its low transmission efficiency for low masses, it is difficult to detect high mixing ratios of OVOCs such as methanol and formaldehyde. The other difference is that the mixing ratio and the corresponding fraction of C_xH_y species measured by Vocus-PTR are much lower than those measured by traditional PTR. For several major C_xH_y compounds, such as benzene, C7, C8, and C9 aromatics, their mixing ratios are comparable between the two methods. The main difference between the two methods lies in the mixing ratio of low-mass hydrocarbons. Overall, when applied to the urban atmosphere, Vocus-PTR has advantages in measuring oxygenated VOCs, especially with multiple oxygens. However, it has limitations in measuring low-molecular-weight VOCs due to the low-mass cutoff in the transmission efficiency.

The molecular characteristics of organic vapors measured by Vocus-PTR in urban Beijing show several differences from those in forested areas (Li et al., 2020; Huang et al., 2021; Li et al., 2021). Firstly, organics up to 300 m/z can be observed in forested areas, while organics up to 230 m/z are observed (Fig. 1a). Two main reasons are responsible for this. The complexity of the species introduces challenges in interpreting mass spectra, which is evidenced by the total number of species being similar to existing atmospheric measurements using Vocus-PTR, despite a narrower mass range in this study. The higher particulate matter concentrations in urban areas provide a larger sink for organic vapors (Deng et al., 2020), and this loss effect is especially pronounced for compounds with high molecular weights due to their lower volatility. The second difference is that $C_xH_yO_z$ and $C_xH_yO_zN_i$ species are the dominant organics in both urban and forested areas, whilst $C_xH_yN_i$ species are more common and abundant in urban areas, which may come from biomass burning emissions (Laskin et al., 2009). Thirdly, VOCs with low carbon and oxygen numbers play a more significant role in the total organic mixing ratio compared to results from forested regions. As shown in Fig. S17a, C_2 and C_3 organics contribute 79 % of the total organic mixing ratio in this study, while C_4–C_6 organics contribute approximately 75 % in forested regions. In contrast to forested areas, where VOC and IVOC mixing ratios are comparable, the majority of the total organic mixing ratio is attributed to VOCs in this study (Fig. S17b). Typical C_2 and C_3 organics, such as C_3H_6O, C_2H_4O, and $C_2H_4O_2$, contribute 14 %, 11 %, and 5 %, respectively, to the total organic mixing ratio and mainly originate from anthropogenic emissions, including industrial and vehicular activities, solvent utilization, and other sources (Qian et al., 2019).

4 Conclusions

In this study, we explore the molecular and seasonal characteristics of organic vapors in urban Beijing using a Vocus-PTR over four seasons. A total of 895 peaks are observed, and 512 of them can be assigned to formulae. The contribution of $C_xH_yO_z$ species is most significant, as they compose up to 54 % of the number and 74 % of the mixing ratios of total organics. With enhanced sensitivity and mass resolution, an increased number of species was observed compared to traditional PTR-MS measurements in urban Beijing, especially compounds with lower mixing ratios and higher oxygen content. A total of 44 % of the species measured in this study are at sub-ppt level, and 38 % of the species contain 3–8 oxygens, resulting in a higher fraction of species containing 3 or more oxygens compared to traditional PTR-MS measurements. Organic vapors with low oxygen content are comparable to those obtained in both urban Beijing and neighboring regions, and they exert a more substantial influence on the overall organic mixing in forested areas.

The mixing ratio of organic vapors with multiple oxygens accounts for 4 % of the total VOC mixing ratio, with the highest levels observed in winter, followed by summer and spring, and the lowest in autumn. These vapors also make a non-negligible contribution to condensational growth and OH reactivity. In summer, the majority of species are aligned to daytime cluster (peaking at noon), primarily originating from the photooxidation process. As the oxygen number increases, the impact of the photooxidation process becomes more pronounced, leading to an increase in both the mixing ratio and the proportion of organic vapors with multiple oxygens during summer. In spring and winter, when the nighttime cluster (peaking at night) dominated, the variations in organic vapors with multiple oxygens are strongly correlated with organic vapors with one or two oxygens. The measured compositions and seasonal variabilities in organic vapors with multiple oxygens emphasize the importance of high-sensitivity and high-mass-resolution measurements in the urban atmosphere, suggesting the potential for future research.

Data availability. Data are available upon request from the corresponding author.

Supplement. The content of the Supplement includes the map of the observation site (Fig. S1); the diurnal variations in $PM_{2.5}$, O_3, NO_x, RH, and T in four seasons (Fig. S2); the calibration results of mixed-calibration gases (Fig. S3); the average limits of detection (1 min) for detected compounds (Fig. S4); the carbon oxidation state of organic vapors with different oxygens (Fig. S5); a boxplot of total OVOC mixing ratios in four seasons (Fig. S6); the diurnal variation cluster results of organic vapors with multiple oxygens (Fig. S7); the cluster results of organic vapors with 1–2 oxygens (Figs. S8–S9); the dependence of daytime clusters on temperature (Fig. S10); the dependence of nighttime clusters on major clusters of organic vapors with 1–2 oxygens (Fig. S11); the distribution of organic vapors with multiple oxygens across different clusters (Fig. S12); the average C, H, O, and N numbers of organic vapors containing multiple oxygens with different diurnal patterns (Fig. S13); the average C, H, O, and N numbers of organic vapors containing multiple oxygens in two clusters (Fig. S14); the diurnal profiles of representative VOCs in four seasons (Fig. S15); the comparison results with the Gucheng site (Fig. S16); the molecular characteristics of the total measured organic vapors by Vocus-PTR (Fig. S17); the observation periods of Vocus-PTR (Table S1); information about calibration gases (Table S2); corrected fragments and water clusters (Table S3); the main $C_xH_yO_{\geq 3}$ and $C_xH_yO_{\geq 3}N$ species measured in this study (Table S4); and the seasonal mixing ratios of OVOCs with multiple oxygens (Table S5). The supplement related to this article is available online at: https://doi.org/10.5194/acp-24-13793-2024-supplement.

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Competing interests. At least one of the (co-)authors is a member of the editorial board of *Atmospheric Chemistry and Physics*. The peer-review process was guided by an independent editor, and the authors also have no other competing interests to declare.

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