

Response to review comments on "Global carbon budgets estimated from atmospheric O₂/N₂ and CO₂ observations in the western Pacific region over a 15-year period"

Anonymous Referee #1:

We would like to thank the anonymous referee #1 for his/her helpful comments and suggestions on our paper. We have revised the manuscript as is described in the following. The referee's comments are in *blue italics*, and the modifications are shown in **red**.

First, we would like to mention about the GCP-reported data used in this paper. In the original manuscript, we used fossil fuel emissions, atmospheric accumulation, and global sink estimates taken from Global Carbon Budget 2017 reported by GCP (Le Quéré, et al., 2018). However, the updated data of Global Carbon Budget 2018 (Le Quéré, et al., 2018) is now available. So, we have used the updated GCP data for recalculating the global carbon budgets in the revised manuscript. Since the fossil fuel-derived CO₂ emission rates have been slightly downwardly revised, the ocean and land sinks based on the APO data have been slightly decreased. But the changes are at most 0.1 PgC yr⁻¹. Consequently, this change has affected very little the conclusion of the original manuscript.

Reply to minor comments:

1) The uncertainties of F and Z_{eff} should be presented explicitly somewhere in the main text and/or Tables. Did the authors assume the uncertainty of 100% for Z_{eff} following Keeling and Manning (2014)? Also, for those readers not familiar with O₂/N₂ studies, it would be better to present the representative values of α_F for the period 1998-2016 or the respective values for the periods in Table 2.

In response to the suggestions about the uncertainties from both Referee #1 and #2, to clarify how we calculated the uncertainties associated with the sink estimations, we have added the following paragraph after the first paragraph in section 3.2.:” **The uncertainties in the parameters used for the carbon budget calculation (Eqs. (6) and (7)), which are also listed in Table 2, are briefly discussed here. Note that in this study the estimated uncertainties are $\pm 1\sigma$. Since the ocean outgassing effect is rather speculative, we assumed that the values of Z_{eff} for the individual periods had $\pm 100\%$ uncertainties in accordance with previous studies (e.g. Manning and Keeling, 2006; Tohjima et al., 2008). We adopted uncertainties of $\pm 5\%$ for the fossil fuel-derived CO₂ emission rate and ± 0.2 PgC yr⁻¹ for the atmospheric CO₂ increasing rate from Le Quéré, et al. (2018). As for the uncertainties of the observed APO changing rates, we adopted the standard deviations among the sites shown in Fig. 7 (± 0.37 per meg yr⁻¹ for longer than 10 years and ± 0.54 per meg yr⁻¹ for 5 years). The estimated uncertainty of the O₂/N₂ scale stability (± 0.45 per meg yr⁻¹) discussed in Section 2. 3, the uncertainty of the O₂/N₂ span sensitivity ($\pm 3\%$), and the uncertainty in the global averaged APO associated with the limited atmospheric sampling (± 0.2 PgC yr⁻¹) discussed in Nevison et al. (2008) were also included in the calculation of the uncertainties in ΔAPO . The uncertainties of α_B and α_F were ± 0.10 (Keeling and Manning, 2014) and ± 0.04 (Tohjima et al., 2008), respectively. Finally, these uncertainties were propagated to the ocean and land sink uncertainties in accordance with Eqs. (6) and (7).”** Additionally, we have added a column for the values of α_F in Table 2 and a description of the uncertainties in the footnote. Following these changes, Table 2 has been modified as follows:

Table 2. Comparison of global carbon budgets based on APO with those from GCP^{a,b}

Atm.	Fossil	Sink of this study	Sink of GCP ^d
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Period	ΔAPO^c	CO_2^d	fuel ^d	α_F^e	$Z_{\text{eff}}/1.1^f$	Ocean	Land	Ocean	Land	Imb.
1998-2016	-10.3(0.91)	4.45	8.28	1.38	0.52	2.57(0.71)	1.26(0.89)	2.24	1.46	0.13
2000-2016	-10.3(0.91)	4.45	8.48	1.38	0.54	2.55(0.73)	1.48(0.91)	2.27	1.48	0.29
2003-2016	-9.9(0.91)	4.58	8.83	1.38	0.52	2.35(0.73)	1.90(0.93)	2.34	1.55	0.36
2000-2004	-8.8(0.94)	3.93	7.11	1.40	0.59	2.23(0.76)	0.94(0.90)	2.01	1.30	-0.14
2004-2008	-9.2(0.96)	4.08	8.21	1.38	0.33	1.97(0.62)	2.17(0.82)	2.18	1.74	0.22
2008-2012	-10.4(0.98)	4.19	9.05	1.37	0.54	2.54(0.77)	2.31(0.97)	2.32	1.85	0.68
2012-2016	-11.6(1.06)	5.36	9.65	1.37	0.71	3.05(0.90)	1.25(1.09)	2.55	1.26	0.49

^aFigures are given in units of per meg yr⁻¹ for ΔAPO , mol mol⁻¹ for α_F , and PgC yr⁻¹ for the others.

^bFigures in parentheses represent the uncertainties.

^c ΔAPO is based on the data from HAT, COI, and cargo ships (40°S-30°N). The uncertainty in parentheses includes the uncertainty associated with the observations, stability in the O₂/N₂ scale, uncertainty derived from limited sampling, and uncertainty in the O₂/N₂ span sensitivity (see text).

^dThese figures are computed from the dataset summarized by the Global Carbon Project (GCP). The uncertainties are ± 0.2 PgC yr⁻¹ for the atmospheric CO₂ and $\pm 5\%$ for the fossil fuel emissions, ± 0.5 PgC yr⁻¹ for the ocean sinks, and ± 0.9 PgC yr⁻¹ for the land sinks (Le Quéré, et al., 2018).

^eThe uncertainties for α_F are ± 0.04 mol mol⁻¹ (Tohjima et al., 2008).

^fThe values of Z_{eff} include both global ocean warming and anthropogenic nitrogen deposition effects, and the uncertainties are assumed to be $\pm 100\%$ (see text).

2) *The discussion on the evaluation of the needed interval to suppress the temporal variability in Z_{eff} is useful in deriving reasonable interannual variations in CO₂ sinks from O₂/N₂ observations. The needed interval was estimated to be 5 years in the paper, and the authors used the ocean heat storage of 0-2000 m layer to estimate Z_{eff} based on the gas flux / heat flux ratio reported by Keeling and Garcia (2002). However, I think the circulation time of ocean deep layer water is much longer than 5 years. Please explain why the authors have considered the use of heat storage of 0-2000 m to be more reasonable than that of 0-700 m. I suppose there is an implicit assumption in the analysis that the ocean circulation and oxygen concentration are in steady-state from the surface to the deep layer. However, temporal variations found in the 5-years average Z_{eff} in Table2 suggest that the ocean is not in steady-state.*

There are several evidences of an increase in the ocean heat content, suggesting that the present ocean is not in steady-state. The changes in both the heat and O₂ contents of the ocean depend on net air-sea exchanges of the heat and O₂. Considering the similarity between heat and gas regarding the air-sea exchange, we can assume that the air-sea O₂ flux is proportional to the air-sea heat flux (O₂ flux/heat flux = constant). Under this assumption, we can evaluate the O₂ outgassing flux if we know the total increase in the ocean heat content. In previous studies, ocean heat content data for the 0-700 m layer were used to evaluate the O₂ outgassing fluxes. However, Levitus et al. (2012) revealed that a much deeper layer significantly contributed to the ocean heat storage: about one third of the heat storage of the 0-2000 m layer is stored in the 700-2000 m layer. Therefore, we used the data for the 0-2000 m layer in this study following the study of Keeling and Manning (2014). These explanations are given in the fourth paragraph of Section 2.5.

In addition to the gradual ocean warming, there are large inter-annual variations in the air-sea heat exchange, which are attributed to an imbalance of the large seasonality in the air-sea heat fluxes. These inter-annual variations in the air-sea heat fluxes are considered to cause rather large inter-annual variations in the air-sea gas exchanges. In this study, we conclude that the 5-year average would to some extent, but not completely, suppress such interannual variability. These explanations are also given in the fourth paragraph of Introduction.

3) *I think it may be helpful for the reader to note the differences in land and ocean CO₂ uptakes expected from the 3% difference in the span sensitivity between NIES and SIO. Has the conclusion about the comparison of the CO₂ uptake reported by the present study with those by GCP changed significantly due to the difference in the span sensitivity?*

A 3% higher span sensitivity, which corresponds to the SIO oxygen span sensitivity, results in an approx. 0.3 PgC yr^{-1} increase and decrease in the ocean and land sinks, respectively. These changes, although not negligible, have not changed significantly the conclusion about the comparison of the CO_2 sinks between this study and those of GCP. To clearly state the influence of the difference in the span sensitivity, we have added the following sentences after the last sentence of the second paragraph (the third paragraph in the revised manuscript) in section 3.2.: “A 3% higher span sensitivity of the O_2/N_2 measurements, which corresponds to the difference in the span sensitivity between SIO and NIES (section 3.1), would result in an increase and decrease of 0.27 PgC yr^{-1} in the ocean and land sinks, respectively, for the three long periods. Although these changes would enlarge the differences in sinks between GCP and this study, the differences are still not significant given the uncertainties of both this study and GCP.”

4) The authors have compared land and ocean CO_2 sinks estimated in the present study with those obtained by GCP, with and without the imbalance sinks added. It seems to me that the authors conclude that the differences between the present study and GCP are reduced, both for land and ocean CO_2 sinks, by adding the “total” imbalance to the respective sinks. However, I actually think we can only add the imbalance to the land and ocean CO_2 sinks based on an appropriate differential distribution. I understand it would be difficult to suggest the best distribution due to uncertainties of the estimated CO_2 sinks, but I would like to hear the authors’ thoughts on this.

We cannot draw any certain conclusion about how to partition the budget imbalance between ocean and land sinks because of the large uncertainties associated with the sink estimation of this study, as Referee #1 also pointed out. Nevertheless, the differences in the sinks between the estimates of this study and those of GCP correspond to the best estimation of the partitioning of the budget imbalance. Since the budget imbalances seem to increase after 2007 in the pentad averages, we have added a discussion about the partitioning of the imbalances in the revised manuscript. Consequently, we have added one paragraph after the fourth paragraph in section 3.3. to read as: “Shown as discrepancies between the pentad sinks of GCP with and without budget imbalances (Fig. 8), the magnitude of the budget imbalances seems to increase after 2007. For the pentad sinks centered on 2007, 2008 and 2009, the ocean and land sinks of this study agree with those of GCP without and with the budget imbalances, respectively. For the pentad sinks between 2010 and 2014, both the ocean and land sinks of this study are plotted between those of GCP with and without the budget imbalances. Although we cannot show any definitive partitioning of the budget imbalance between ocean and land sinks because of a rather large uncertainty associated with the sink estimations, the above results seem to suggest that the budget imbalances are allocated to land sinks for the former period and to both sinks for the latter period.”

5) P3, line 32: A literal error “: : heat content,.” should be corrected.

The literal errors have been corrected.

6) P7, line 19: I think the unit of Z_{eff} is not TgC yr^{-1} but PgCyr^{-1} in this context.

The unit of Z_{eff} has been corrected to “ PgC yr^{-1} ”.

7) References: Please consolidate the format of references. For example, some journal titles are written in Italic and the others are not.

We have consolidated the format of references.

8) *Caption in Fig.7: The phrase “changing ratio” should be changed to “changing rate”.*

“changing ratio” has been changed to “changing rate”.

Response to review comments on "Global carbon budgets estimated from atmospheric O₂/N₂ and CO₂ observations in the western Pacific region over a 15-year period"

Anonymous Referee #2:

We would like to thank the anonymous referee for his/her helpful comments and suggestions on our paper. We have revised the manuscript as is described in the following. The referee's comments are in *blue italics*, and the modifications are shown in **red**.

Reply to main comments:

1) It's unclear how uncertainty in estimating Z_{eff} comes into play in the carbon budget estimates, especially given its relevance for the shorter timescales considered. The authors do show trends without Z_{eff} (figure 8) but it's not evident if this is incorporated in the carbon sinks and trends calculation (i.e. unclear if incorporated in grey shading in figure 8 or error estimates in carbon sinks column in Table 2). Perhaps, the authors could evaluate uncertainty in Z_{eff} from using the upper/lower bounds with and without Z_{eff}? Further, it's unclear how correcting for Z_{eff} in the carbon budgets plays out in the 5 year timescales described (as also discussed in Nevison et al. 2008 and elsewhere). It seems, as referred to by the authors, that the ventilation events of 1999-2001 could impact the pentad trends, and thus similar variability during other years probably could have similar effects on other pentad periods (e.g. 2004-2005 dip in pentad ocean sink seems to co-occur with inter-annual variation that don't seem to be fully suppressed?). Finally, an additional and not insignificant component of Z_{eff} not included in this study is the atmospheric deposition effect, as detailed in Keeling and Manning 2014, which adds about 0.1 (+/-0.1) Pg C/yr, and which should raise Z_{eff} from 0.1-0.9 Pg C/yr, to 0.2-1.0 Pg C/yr. Overall, I feel the treatment of Z_{eff} uncertainty within the shorter timescales considered here merits further clarification.

As both Referee #1 and #2 indicated, the description of the uncertainty in the ocean outgassing effect was unclear in the original manuscript. We assumed a $\pm 100\%$ uncertainty for Z_{eff} for the corresponding budget calculation period. To make it clear, we have added a sentence in the second paragraph in the revised manuscript (see reply (2) for more details): "**Since the ocean outgassing effect is rather speculative, we assumed that the values of Z_{eff} for the individual periods had $\pm 100\%$ uncertainties in accordance with previous studies (e.g. Manning and Keeling, 2006; Tohjima et al., 2008).**"

As Referee#2 suspected, it seems that the five-year average cannot fully suppress the APO variations associated with the anomalous air-sea gas exchanges. As we discussed in Section 3.1 and as is shown in Fig. 7, the pentad APO trends have a temporal variability of about +1.2 per meg yr⁻¹ or +0.5 PgC yr⁻¹, which is comparable to the 2004-2005 dip in the pentad ocean sink shown in Fig. 8. To emphasize the limitations of the pentad averaging, we added the following sentences after the first sentence of the second paragraph in Section 3.3: "**Nevison et al. (2008) suggested that a decadal or longer period is needed to suppress the influence of the interannual variation in the ocean O₂ flux on the carbon sink estimation within ± 0.1 PgC yr⁻¹ based on an ocean ecosystem model and an atmospheric transport model. In addition, the pentad APO changing rate still contains an uncertainty corresponding to ± 0.5 PgC yr⁻¹ as is discussed in Section 3.1. Therefore, the anomalous dip in the ocean sink for 2004-2005 might be an error caused by the anomalous ocean O₂ flux variations.**"

In response to the suggestion of Referee #2, we have included the anthropogenic N deposition effect in the Z_{eff} estimation. To explain the anthropogenic nitrogen deposition effect (Z_{anthN}), we have separated Z_{eff} into Z_{anthN} and the global ocean warming component (Z_{gow}), and added the following paragraph after the fourth paragraph in Section 2.5:

“In addition to the ocean warming effect, Keeling and Manning (2014) introduced recently another ocean outgassing effect caused by atmospheric deposition of excess anthropogenic nitrogen to the open ocean. The excess nitrogen is considered to enhance the ocean biotic production of organic matter, which is associated with the O₂ production. Keeling and Manning (2014) evaluated the anthropogenic nitrogen-induced outgassing as being about 0.1×10^{14} mol O₂ yr⁻¹. Since the outgassing effect caused by the anthropogenic nitrogen deposition is small but rather significant, we adopted the effect as Z_{anthN} , with a magnitude of 0.12 PgC yr^{-1} ($=0.1 \times 10^{14}$ mol O₂ yr⁻¹ $\times 12.01 \text{ gC mol}^{-1}$). Eventually, the total outgassing effect, Z_{eff} , is expressed as the summation of Z_{gow} and Z_{anthN} :

$$Z_{eff} = Z_{gow} + Z_{anthN}. \quad (9)”$$

2) I see the need perhaps for a section dedicated to clarifying and detailing the sources, contribution, and methods for calculating carbon budget uncertainty, as it can help clarify the confidence in the pentad trends and conclusions presented here. Table 2, for instance, could incorporate uncertainty due to Z_{eff} in the carbon sinks column, and in Figure 8 (gray shading). Uncertainty due to undersampling the global signal has also been shown by Nevison et al. (2008) to contribute to uncertainty in estimating budgets on the shorter timescales evaluated here. What is impact of sampling over the western Pacific (40S-40N) vs. full global sampling on the carbon sink trends? How does uncertainty in α_B in using values of 1.1 vs 1.05 affect the uncertainty in carbon sink budgets? Finally, it’s unclear how the contribution of measurement uncertainty, due to span calibration of the gas chromatographic technique and potential longterm drift shared across all cylinders, is incorporated in the uncertainty analysis.

In the original manuscript, we didn’t incorporate the anthropogenic N deposition effect on Z_{eff} , the uncertainty associated with the global average of APO from limited samples, uncertainty due to span sensitivity of the gas chromatographic technique, and the potential long-term drift among the O₂/N₂ reference cylinders. In addition, the uncertainty of ± 0.05 for α_B in the original manuscript should be increased to ± 0.10 . Considering these uncertainties, we have carefully reevaluated the uncertainties. To make it clear, we have modified the original manuscript as follows:

To clarify the long-term stability of the O₂/N₂ scale, we have added the following sentences at the end of Section 2.3: “However, this stability test cannot exclude the possibility that the O₂/N₂ ratios of the reference gases drift across all the cylinders rather uniformly. There are several mechanisms that affect the O₂/N₂ ratios of the gases within the high-pressure cylinders, including corrosion of the inner surface, leakage, thermal diffusion and gravitational fractionation. Keeling et al. (2007) assessed carefully and comprehensively the influences of those potential mechanisms on the long-term stability of the O₂/N₂ ratio of the reference gases and obtained an estimated uncertainty of ± 0.4 per meg yr⁻¹. We also treated the reference cylinders which were kept horizontally in a thermally insulated box, with the greatest care (Tohjima et al., 2008). Therefore, we adopted the value of ± 0.4 per meg yr⁻¹ as the long-term drift of the reference gases caused by the above degradation effects. Consequently, we assumed that the total uncertainty of the long-term stability of the O₂/N₂ reference scale was ± 0.45 per meg yr⁻¹ ($= (0.2^2 + 0.4^2)^{1/2}$) in this study.”

As for the uncertainty of α_B , we have added the following sentences at the end of Section 2.4: “Considering the recent reports about the global net -O₂/CO₂ exchange ratio, Keeling and Manning (2014) revised the uncertainty of α_B upward from ± 0.05 (Severinghaus, 1995) to ± 0.10 . Thus, we also adopted ± 0.10 for the uncertainty of α_B in this study.”

Finally, to clarify how we computed the total uncertainties associated with the global sink estimations, we have added the following paragraph after the first paragraph in Section 3.2:

“The uncertainties in the parameters used for the carbon budget calculation (Eqs. (6) and (7)), which are also listed in Table 2, are briefly discussed here. Note that in this study the estimated uncertainties are $\pm 1\sigma$. Since the ocean outgassing effect is rather speculative, we assumed that the values of Z_{eff} for the individual periods had $\pm 100\%$ uncertainties in accordance with previous studies (e.g. Manning and Keeling, 2006; Tohjima et al., 2008). We adopted uncertainties of $\pm 5\%$ for the fossil fuel-derived CO_2 emission rate and $\pm 0.2 \text{ PgC yr}^{-1}$ for the atmospheric CO_2 increasing rate from Le Quéré, et al. (2018). As for the uncertainties of the observed APO changing rates, we adopted the standard deviations among the sites shown in Fig. 7 ($\pm 0.37 \text{ per meg yr}^{-1}$ for longer than 10 years and $\pm 0.54 \text{ per meg yr}^{-1}$ for 5 years). The estimated uncertainty of the O_2/N_2 scale stability ($\pm 0.45 \text{ per meg yr}^{-1}$) discussed in Section 2. 3, the uncertainty of the O_2/N_2 span sensitivity ($\pm 3\%$), and the uncertainty in the global averaged APO associated with the limited atmospheric sampling ($\pm 0.2 \text{ PgC yr}^{-1}$) discussed in Nevison et al. (2008) were also included in the calculation of the uncertainties in ΔAPO . The uncertainties of α_B and α_F were ± 0.10 (Keeling and Manning, 2014) and ± 0.04 (Tohjima et al., 2008), respectively. Finally, these uncertainties were propagated to the ocean and land sink uncertainties in accordance with Eqs. (6) and (7).”

In accordance with the uncertainty revision, we have also revised the uncertainties of the O_2/N_2 and APO changing rates listed in Table 1. The uncertainties of ± 0.2 or $\pm 0.3 \text{ per meg yr}^{-1}$ have increased to $\pm 0.8 \text{ per meg yr}^{-1}$. In addition, we have also redrawn Fig. 8 as follows:

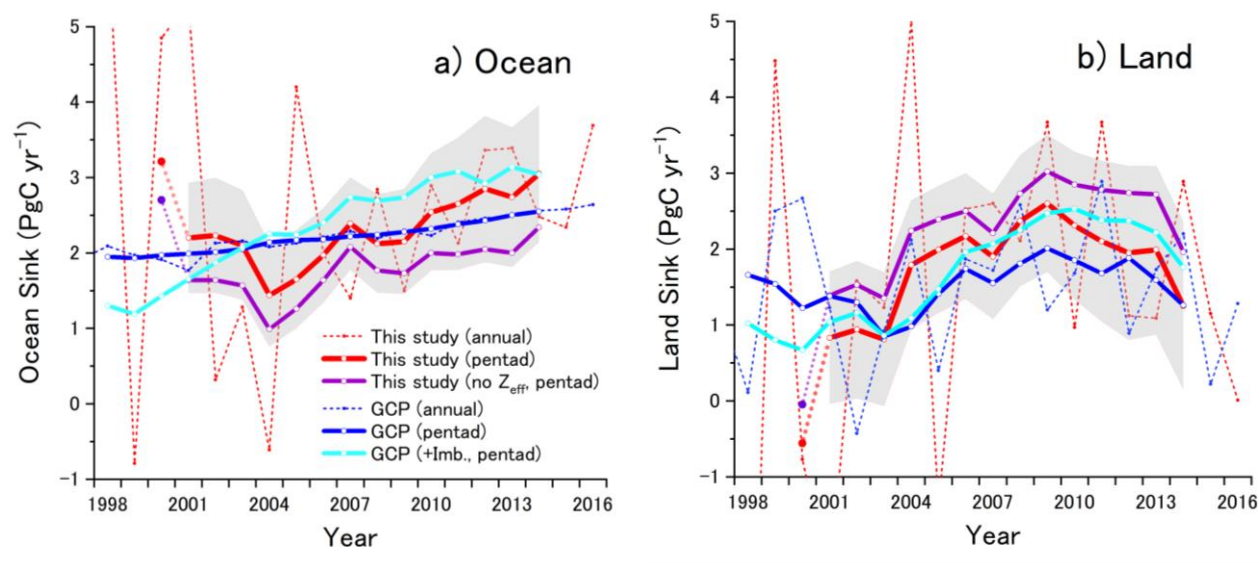


Fig. 8. Temporal variations in (a) ocean and (b) land biospheric sinks estimated from APO variations of this study (red) and process-based models of GCP (blue). The thin broken lines represent the annual sinks and the thick lines represent the pentad sinks. The purple lines represent the pentad sinks based on APO without ocean outgassing correction (Z_{eff}) and the light blue lines represent the sinks of GCP with the imbalance sinks added. The uncertainty associated with the pentad sinks with Z_{eff} corrections are shown as shaded area.

3) The comparison against the GCP could be elaborated on a bit more, as it raises important issues in the field. The authors could elaborate further (through existing or new figure/table) how different estimates reported by GCP compare to the APO method, including hindcast ocean models and ocean observation based

products, all of which are readily available in the GCP product as globally integrated fluxes: <https://www.icoscp.eu/GCP/2018>. It is interesting that the comparison to the GCP mean showcases similarities in magnitude and in temporal evolution of pentads. The point that the uptake of carbon by the ocean is larger than expected from atmospheric increase alone is very interesting. How do the decadal trends (2000-2016) in this study compare to the pCO₂ based air-sea flux timeseries by Landschutzer et al (2016) and Rodenbeck et al (2013), as both of these estimates seem to show larger decadal variability than the ocean models? These items may be beyond the current scope of this study, but could substantially improve the impact of this paper with (hopefully?) relatively minor figure/text additions.

We have carefully compared our pentad ocean sinks based on APO with those of GCP and the pCO₂-based estimations and found that the increasing trend of the ocean sinks based on APO was close to those based on the pCO₂ observations. To explain this clearly, we have added the following sentences after the second to the last sentence of the second paragraph in Section 3.3: “For a detailed comparison, the global ocean sinks based on pCO₂ observations and interpolation techniques (Landschützer et al., 2016; Rödenbeck et al., 2014) for the period of 1990-2017 are plotted in Fig. 9 together with the ocean sinks of this study and GCP. Note that the extended pCO₂-derived ocean sinks were given as supplementary data of Le Quéré, et al. (2018) and those sinks were uniformly inflated by 0.78 PgC yr⁻¹ to compensate for the pre-industrial steady state source of CO₂ derived from riverine input of carbon to the ocean (Resplandy et al., 2018). As you can see, both the GCP and pCO₂-derived ocean sinks show changes in the trends between before and after 2001 while the magnitude of the changes in the pCO₂-derived sinks are larger. The increasing rates determined by a linear regression during 2001-2014 are 0.08 ± 0.01 PgC yr⁻² in Landschützer et al. (2016) and 0.07 ± 0.02 PgC yr⁻² in Rödenbeck et al. (2014), which are more consistent with the rate found in this study. Therefore, our result seems to support a previous conclusion that the recent increase in the ocean sinks exceeds the increasing trend of ocean sink expected only from the atmospheric CO₂ increase (Landschützer et al., 2015; DeVries et al., 2017).” In accordance of this change, we have modified the third sentence of Conclusion 3) as “The pentad ocean sinks showed an overall increasing trend for the entire period (2001-2014) with a linear increasing rate of 0.08 ± 0.02 PgC yr⁻². This increasing rate was about two times larger than that for the GCP ocean sinks (0.04 ± 0.01 PgC yr⁻²) but was consistent with those for the global ocean sinks based on pCO₂ observations and interpolation techniques (Landschützer et al., 2016; Rödenbeck et al., 2014).” We have also added Resplandy et al. (2018) to Reference and Fig. 9 showing ocean sinks based on the APO data, process-based models (GCP), and pCO₂ observations as follows:

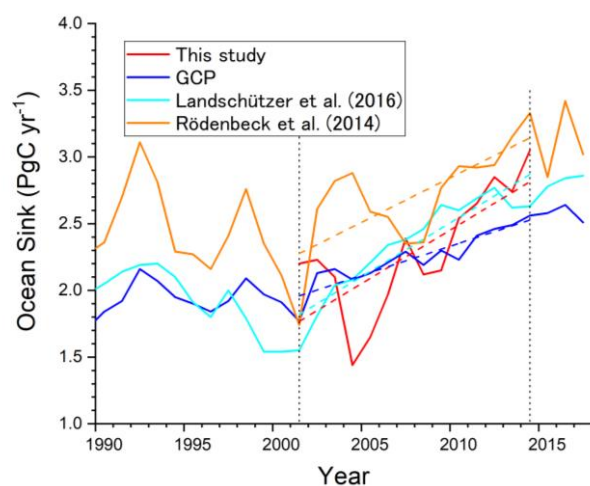


Fig. 9. Comparison of the temporal variations of the ocean sinks based on the APO data of this study (red), global ocean biogeochemistry models (GOBMs) of GCP (blue), and pCO₂ data of Landschützer et al. (2016) (light blue) and Rödenbeck et al. (2014) (orange). The broken lines represent the regression lines for the corresponding data during 2001-2014. Note that the pCO₂-based ocean sinks are adjusted for the pre-industrial ocean CO₂ emissions ($\pm 0.78 \text{PgC yr}^{-1}$) caused by riverine CO₂ input to the ocean (Resplandy et al., 2018).

In addition to the above modifications, we have changed the GCP-reported data (fossil fuel emissions, atmospheric accumulation, and global sinks) from Global Carbon Budget 2017 to the data from Global Carbon Budget 2018 (Le Quéré, et al., 2018). We have used the updated GCP data for recalculating the global carbon budgets in the revised manuscript. Because the fossil fuel-derived CO₂ emission rates have been slightly downwardly revised, the ocean and land sinks based on the APO data have been slightly decreased. But the changes are at most 0.1 PgC yr⁻¹. Consequently, this change has affected very little the conclusion of the original manuscript.

Reply to minor issues:

Pg2 L27: “The estimated value for α_F is about 1.10 ± 0.05 (Severinghaus, 1995) and that for α_B is about 1.4 (Keeling, 1988).” Should be the other way around: α_B is 1.10 and α_F is 1.4.

“The estimated value for α_F is about 1.10 ± 0.05 (Severinghaus, 1995) and that for α_B is about 1.4 (Keeling, 1988)” has been changed to “The estimated value for α_B is about 1.1 (Severinghaus, 1995) and that for α_F is about 1.4 (Keeling, 1988).”

Perhaps add citations for Equations (1), (2), and (3)?

Citation for Eq. (1), (2), and (3): We have added the citation, Manning and Keeling (2006), for these equations.

P3 L1, this paragraph could use a brief explanation of APO concept as a tracer for those not familiar with APO, i.e. cancellation of terrestrial influence, etc

In accordance with the suggestion, we have added the following sentence after the first sentence of the paragraph: “Since the APO is defined to be invariant with respect to the land biotic exchange, the secular trend in the APO is determined by fossil fuel combustions which cause a gradually decreasing trend in APO, and the air-sea gas exchange.”

Pg 7 L29, shouldn't Z_{eff} be in PgC/yr?

The unit “TgC yr⁻¹” has been altered to “PgC yr⁻¹”.

Pg 9 Line 20, the ENSO topic deserves a bit more clarification here. It would be good to preface the ENSO sentence with the findings of Rodenbeck et al 2008, who suggest anomalous outgassing of APO during El Niño, while Tohjima et al 2015 show a suppressed peak instead, and clarify that Eddebbbar et al (2017) reconcile this apparent discrepancy through a model-simulated zonal dipole-like ENSO response in the equatorial Pacific, and that enhanced observational zonal coverage in this region is needed to constrain the full basin ENSO response.

In response to the Referee’s suggestion, we have modified the relevant part as “Conducting atmospheric inversion analyses based on the APO data from the Scripps observation network,

Rödenbeck et al. (2008) suggested anomalous outgassing of APO from the equatorial region during El Niño periods, while Tohjima et al. (2015) found a suppressed equatorial peak during El Niño periods based on the western Pacific observations. Eddebbar et al. (2017) reconciled these conflicting results by predicting the existence of a zonal dipole-like ENSO response in the equatorial Pacific based on several ocean process-based models and an atmospheric transport model. These results suggest that an enhanced zonal coverage of the atmospheric observations in the equatorial Pacific is needed to constrain the full basin-scale ENSO response. We can see a considerable suppression of the equatorial peak during the strong 2015/2016 El Niño event in Fig. 6c, which was not reported in Tohjima et al. (2015). Any detailed discussion about the temporal variation of the equatorial peak during the 2015/2016 El Niño event is, however, beyond the scope of this study and will be given elsewhere.”

Suggested editing notes:

Pg 2 Line 5: remove “still”, and add year by which emissions rose to 10 Pg C/yr?

“...the global fossil fuel-derived CO₂ emissions in recent years still increased gradually and rose toward 10 PgC yr⁻¹ (Boden et al. 2017)” has been changed to “...the global fossil fuel-derived CO₂ emissions in recent years still increased gradually and rose to 9.9 PgC yr⁻¹ by 2014 (Boden et al. 2017)”.

Pg 2 Line 6: “Paris Agreement . . . aimed to balance the anthropogenic greenhouse gas emissions and natural removals in the second half of this century. . .”, I suggest editing to: “. . . aimed to reduce anthropogenic greenhouse gas emissions to maintain the increase in global mean surface temperature well below 2_C by 2100, . . .”?

The relevant part has been modified to “..., the Paris Agreement adopted at COP21 in 2015 aimed to **reduce** the anthropogenic greenhouse gas emissions to maintain the increase in global **mean surface** temperatures well below 2°C **by 2100**, ...”.

Pg3 L8, suggest deleting “In these days”.

“In these days” has been deleted.

Pg3 L21, “which reduces the ventilation of the seawater.”, suggest instead: “which reduces the ventilation of interior water masses.”

We have changed the relevant sentence from “the ventilation of the seawater” to “the ventilation of interior water masses”.

Pg3 L26: replace “huge” with “large”

“huge” has been replaced with “large”.

Pg 14 L 20: Not sure I understand this sentence: “This means that the changing trends of carbon budgets may be evaluated by the at least decadal APO data.” Suggest rephrasing and/or elaborating further?

The ambiguous fourth conclusion has been deleted. We have also modified the first two sentences of the last paragraph in Section 3.3, “From the above discussions, we feel ...in the temporal resolution.”, to “From the above discussions, we feel that a five-year duration effectively suppresses **to some extent** the anomalous variations in the carbon budget estimations

based on APO, which are considered to be caused by the imbalance of the seasonal air-sea O₂ exchange. **Probably, the five-year average suppresses the variability of Z_{eff} to a level of ±0.5 PgC yr⁻¹ as is discussed in Section 3.1.**”

Pg 12 L 32. Replace “stagnant” with “stagnancy”

“stagnant” has been replaced with “stagnancy”.

Pg13 L 1: replace “in spite of” with “despite”

“in spite of” has been replaced with “despite”.

Global carbon budgets estimated from atmospheric O₂/N₂ and CO₂ observations in the western Pacific region over a 15-year period

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Abstract. Time series of atmospheric O₂/N₂ ratio and CO₂ mole fraction of flask samples obtained from NIES's flask sampling network are presented. The network includes two ground sites, Hateruma Island (HAT, 24.05N, 123.81°E) and Cape Ochiishi (COI, 43.17°N, 145.50°E), and cargo ships regularly sailing in the western Pacific. Based on temporal changes in fossil fuel-derived CO₂ emissions, global atmospheric CO₂ burden, and atmospheric potential oxygen (APO), which were calculated from the observed O₂/N₂ ratio and CO₂ mole fraction according to $APO = O_2 + 1.1 \times CO_2$, we estimated the global carbon sinks of the ocean and land biosphere for a period of more than 15 years. In this carbon budget calculation, we adopted a correction for the time-varying ocean O₂ outgassing effect with an average of 0.54 PgC yr⁻¹ for 2000-2016. The outgassing effect, attributed mainly to global ocean warming, was evaluated under the assumption that the net ocean gas flux is proportional to the change in the ocean heat content for the 0-2000 m layer. The resulting oceanic and land biotic carbon sinks were 2.6 ± 0.7 PgC yr⁻¹ and 1.5 ± 0.9 PgC yr⁻¹, respectively, for a 17-year period (2000-2016) and 2.4 ± 0.7 PgC yr⁻¹ and 1.9 ± 0.9 PgC yr⁻¹, respectively, for a 14-year period (2003-2016). Despite the independent approaches, these sink values of this study agreed with those estimated by the Global Carbon Project (GCP) within a difference of about ± 0.4 PgC yr⁻¹. We examined the carbon sinks for an interval of five years to assess the temporal trends. The pentad (5-year) ocean sinks showed an increasing trend at a rate of 0.08 ± 0.02 PgC yr⁻² during 2001-2014, while the pentad land sinks showed an increasing trend at a rate of 0.23 ± 0.04 PgC yr⁻² for 2001-2009 and a decreasing trend at a rate of -0.22 ± 0.04 PgC yr⁻² during 2009-2014. Although there is good agreement in the trends of the pentad sinks between this study and that of GCP, the increasing rate of the pentad ocean sinks of this study was about two times larger than that of GCP.

25

1. Introduction

In spite of various international efforts to reduce anthropogenic carbon dioxide (CO₂) emissions, the atmospheric CO₂ levels observed around the world have shown a steady increase and exceeded the benchmark of 400 parts per million mole fraction (ppm) in past years (Betts et al., 2016). The Carbon Dioxide Information Analysis Center (CDIAC) reported that global fossil fuel-derived CO₂ emissions in recent years still increased gradually and rose to 9.9 PgC yr⁻¹ by 2014 (Boden et al. 2017). Under these circumstances, the Paris Agreement adopted at COP21 in 2015 aimed to reduce the anthropogenic greenhouse gas emissions to maintain the increase in global mean surface temperatures well below 2°C by 2100, and if possible, to limit the increase to 1.5°C. To achieve this goal, it is crucially important to quantitatively understand the natural sink strengths or land biosphere and ocean sinks. A variety of approaches have so far been applied to the quantification of ocean or land sinks or both, including process-based land and ocean models, bottom-up emission estimates based on flux measurements, and top-down estimates based on atmospheric measurements. Developing process-based models to enhance the accuracy of the global carbon budget is crucially important because they are expected to predict the future global carbon cycle in a warmer world. However, carbon budget estimates based on observations are still important to validate and improve the process-based models.

The budget estimation based on atmospheric CO₂ and O₂ observations is a simple and straightforward approach, and it has historically settled the controversy whether the land biosphere is a net carbon sink or source (Keeling and Shertz, 1992). Although several techniques based on interferometer (Keeling, 1988), mass spectrometry (Bender et al., 1994), paramagnetic analyzer (Manning et al., 1999), fuel cell analyzer (Stephens et al., 2007), vacuum ultraviolet absorption photometer (Stephens et al, 2003) and so on, have been developed to detect the ppm level changes in the atmospheric O₂ concentration, the accurate quantification of the O₂ change is still challenging. The carbon budget is evaluated by simultaneously solving the mass balance equations of the atmospheric CO₂ and O₂ as follows (Manning and Keeling, 2006):

$$\Delta CO_2 = F - B - O, \quad (1)$$

$$\Delta O_2 = -\alpha_f F + \alpha_B B + Z, \quad (2)$$

where ΔCO_2 and ΔO_2 represent the changes in the atmospheric CO₂ and O₂ burdens based on atmospheric observations, respectively, F represents the fossil fuel-derived emissions, and B and O represent the uptake by the land biosphere and the ocean, respectively. α_f and α_B are the -O₂/C exchange ratio for the globally averaged fossil fuel combustions and land biotic processes, respectively. The estimated value for α_B is about 1.1 (Severinghaus, 1995) and that for α_f is about 1.4 (Keeling, 1988). These equations mean that the CO₂ and O₂ fluxes associated with fossil fuel combustion and land biotic processes are tightly coupled. In contrast, the ocean CO₂ uptake, O , and ocean O₂ emissions, denoted as Z , are decoupled because the ocean acts as a carbon sink by physicochemically dissolving the CO₂. Since the values of F and α_f can be evaluated from energy statistics (Keeling, 1988), we can evaluate the ocean and land uptake by solving the above equations if we could evaluate the value of Z .

The global carbon budget can also be related to tracer atmospheric potential oxygen (APO), which is defined by the equation of $APO = O_2 + \alpha_B \times CO_2$ (Stephens et al., 1998). Since the APO is defined to be invariant with respect to the land biotic exchange, the secular trend in the APO is determined by fossil fuel combustions which cause a gradually decreasing trend in APO, and the air-sea gas exchange. Combining Eq. (1) multiplied by α_B and Eq. (2) in accordance with the APO definition, results in the following equation for APO budget (Manning and Keeling, 2006):

$$\Delta APO = -(\alpha_f - \alpha_B)F - \alpha_B O + Z. \quad (3)$$

Since observation sites for atmospheric O₂ are still limited compared with those for atmospheric CO₂, Manning and Keeling (2006) proposed an alternative approach that the global carbon budgets could be obtained by simultaneously solving Eqs. (1) and (3) and using globally averaged CO₂ data based on NOAA/ESRL/GMD's measurements. This approach, making

maximum use of the available data, is expected to give the most reliable estimation. This APO approach has been adopted for the estimation of global carbon budget based on atmospheric O₂ and CO₂ measurements (e.g. Manning and Keeling, 2006; Tohjima et al., 2008; Ishidoya et al., 2012; Goto et al., 2017).

To evaluate the carbon budgets based on O₂ and CO₂ measurements, we need to quantify the magnitude of Z and its temporal variation, if possible. It is considered that Z has a large interannual variability because observed trends of APO generally show large interannual variations which would result in unrealistic variations in the ocean uptake if the variability in Z were rather small (e.g. Bender et al., 2005). Probably, an imbalance of the air-sea seasonal O₂ exchanges, outgassing flux associated with primary production in spring and summer and ingassing flux associated with ocean ventilation in autumn and winter, cause the interannual variations in Z. The results of ocean model simulations also support this mechanism (e.g. McKinley et al., 2003; Nevison et al. 2008). Therefore, it is difficult to estimate the short-term carbon budgets unless the temporal variations in Z are accurately evaluated. Additionally, as for long timescales, it is considered that the present ocean acts as an O₂ source because of the global ocean warming (Keeling and Garcia, 2002). The increase in surface ocean temperature not only reduces the solubility of gases in seawater but also strengthens the ocean stratification, which reduces the ventilation of interior water masses. The reduction of ventilation reduces the ingassing flux of O₂. In contrast, the reduction of ventilation also causes a reduction of the nutrient supply from deep water, which might decrease the primary production and O₂ outgassing in summer. Therefore, the influence of the ocean warming on the net air-sea gas exchange is rather complicated.

Unfortunately, there is little observational evidence to quantify the magnitude of Z and its temporal variations. The long-term average values of Z are inferred under the assumption that Z is proportional to the air-to-sea heat flux (Keeling and Garcia, 2002). The change in the global ocean heat content has been evaluated based on the large data set of ocean observations (Levitus et al. 2012). Keeling and Garcia (2002) estimated the O₂ flux / heat flux ratio from the relationship between the dissolved O₂ corrected for the mineralization effect and the potential temperature. This approach was basically adopted by most of the studies to evaluate the long-term global carbon budgets (e.g. Bender et al., 2005; Manning and Keeling, 2006; Tohjima et al., 2008). On the other hand, Ishidoya et al. (2012) evaluated the instantaneous variations in the land and ocean sinks based on the APO data at Ny-Ålesund, Svalbard, and Syowa, Antarctica for the period 2001-2009 by calculating the interannual variation in Z from the temporal variation in the ocean heat content. They concluded that the above-mentioned Z values adequately suppressed artifacts caused by the imbalance of the seasonal air-sea O₂ exchange.

We have been conducting air sampling into glass flasks for the measurement of the atmospheric O₂/N₂ ratio and CO₂ mole fraction at two ground sites in Japan since the late 1990s (Tohjima et al., 2003), and have been evaluating the global carbon budgets for up to 7 years (1999-2005) based on the APO data from the flask observations (Tohjima et al., 2008). To extend the observation area, we started additional flask sampling aboard commercial cargo ships regularly sailing in the Pacific region in 2002 (Tohjima et al., 2005b, 2012). About a decade has passed since we previously reported the global carbon budgets, and now we have more than a 15-year long record of atmospheric O₂/N₂ and CO₂ of the flask samples. In this study, we estimated the ocean and land biotic carbon sinks for over a decade by using the temporal changes in the APO based on these flask data. In addition, we sequentially computed the ocean and land sinks for an interval of five years and examined the changing trends of both sinks. In these budget calculations, we estimated the values of Z for the corresponding period by using the temporal changes in the global ocean heat content. Finally, the estimated ocean and land carbon sinks of this study were compared with those of the Global Carbon Project (GCP).

2. Data and analysis

2. 1. Flask sampling locations

We started air samplings for the measurement of the atmospheric O₂/N₂ ratio and CO₂ mole fraction at two monitoring stations

located on Hateruma Island (HAT, 24.05N, 123.81°E) in July 1997 and at Cape Ochiishi (COI, 43.17°N, 145.50°E) in December 1998 (Tohjima et al., 2003). In addition, we have been collecting air samples from the Pacific regions by using commercial cargo vessels equipped with automated flask sampling systems (Tohjima et al., 2005b; 2012). The shipboard flask samplings were started between Japan and North America in December 2001, between Japan and Australia/New Zealand in December 2001, and between Japan and Southeast Asia in September 2007. The flask sampling sites are depicted in Figure 1. Unfortunately, the shipboard data in Southeast Asia, the northern North Pacific (north of 30°N), and the eastern North Pacific are spatiotemporally rather sporadic. (See the inserted figure in Fig. 1 showing time-latitude plots of the shipboard flask samples.) Thus, in the following analysis, we only used the data set obtained at HAT, COI and the western Pacific region between 40°S and 30°N and between 130°E and 180°E.

Air samples were collected in glass flasks hermetically sealed by two glass valves with Viton O-rings. The volumes of the flasks were 2 liters for the samplings at HAT and COI and 2.5 liters for the shipboard samplings. It should be noted that glass flasks with a volume of 1 liter were also used in the early period from the start to March 2006 and only 1-liter flasks were used from the start to January 1999 at HAT.

2. 2. O₂/N₂ and CO₂ analytical methods

In this study, we used a gas chromatograph (GC) equipped with a thermal conductivity detector (TCD) for the measurements of atmospheric O₂ (Tohjima 2000). In this GC/TCD method, O₂/N₂ ratios of sample air and working reference air were alternately measured and the atmospheric O₂ change was determined as the relative difference in the O₂/N₂ ratio from an arbitrary reference. We used the delta notation according to Keeling and Shertz (1992) to express the relatively small difference in the O₂/N₂ ratio as follows:

$$\delta(O_2/N_2) = \frac{(O_2/N_2)_{sam}}{(O_2/N_2)_{ref}} - 1, \quad (4)$$

where subscripts “*sam*” and “*ref*” refer to sample and reference, respectively, and the $\delta(O_2/N_2)$ value multiplied by 10⁶ is expressed in “per meg” units. The change in 1 μ mol of O₂ per mole of dry air changed the O₂/N₂ ratio by 4.77 per meg, which corresponds to 1 ppm change in the atmospheric trace gas abundance. APO was calculated from the CO₂ mole fraction (X_{CO_2}) in ppm and $\delta(O_2/N_2)$ in per meg according to

$$\delta APO = \delta(O_2/N_2) + \alpha_B X_{CO_2} / S_{O_2} - 1850, \quad (5)$$

where S_{O_2} is the mole fraction of O₂ in the air ($S_{O_2}=0.2094$, Tohjima et al., 2005a) and the value of 1850 is an arbitrary reference point of δAPO in per meg. The values of $\delta(O_2/N_2)$ were determined against the NIES O₂/N₂ scale (Tohjima et al. 2008). Its temporal stability is examined in the following section.

A nondispersive infrared (NDIR) analyzer (LI-COR, Lincoln, Ne., model LI-6252) was used for the CO₂ measurement of the flask samples. The CO₂ mole fractions were determined against the NIES 09 scale, which is based on a set of gravitationally prepared CO₂-in-air standard gases (Machida et al., 2011). The relationship between the NIES 09 scale and the NOAA scale were repeatedly compared through the WMO Round-robin inter-comparison program. The results showed that the differences of the NIES 09 scale from the NOAA scale were kept within ± 0.15 ppm during the period from 1996 to 2014 (https://www.esrl.noaa.gov/gmd/ccgg/wmorr/wmorr_results.php?).

2. 3. O₂/N₂ scale stability

As details of the NIES O₂/N₂ scale are given elsewhere (Tohjima et al., 2008), here we describe only briefly the outline of the scale and add some new information below. The zero point of NIES O₂/N₂ scale had been related to an ambient dry air stored in a high-pressure cylinder (HDA-1). The O₂/N₂ scale was maintained by three cylinders during the first four years (1997-

2001) of our O₂/N₂ measurement program. In 2001, the NIES O₂/N₂ scale was transferred to 11 other high-pressure cylinders (five 10-liter cylinders and six 48-liter cylinders), of which the δ(O₂/N₂) values were carefully determined against the original O₂/N₂ scale. Another primary reference gas (48-liter cylinder, CQB-07080) was added in 2002, and now 12 primary reference gases keep the NIES O₂/N₂ scale. The air samples delivered from glass flasks or high-pressure cylinders were measured against the working reference airs stored in 48-liter aluminum cylinders, of which the δ(O₂/N₂) values were repeatedly determined against the individual primary gas cylinders at intervals of a few months. The working reference gas cylinders were replaced by new ones every one to two years.

Fig. 2 is the extended version of the previously reported figure (Fig. 1 in Tohjima et al., 2008), showing the temporal changes in the O₂/N₂ ratio of primary reference gases relative to the NIES O₂/N₂ scale. In the figure, the deviations of the O₂/N₂ ratio from the average value for HDA-2 and from the initially determined values for the second set of 12 cylinders are plotted. The averages and the standard deviations (1σ) of the differences for the individual 12 cylinders range from -4.2 per meg to 3.3 per meg and from 3.1 per meg to 5.2 per meg, respectively. The changing rates of the deviations for the 12 reference gases during 2001-2017, determined by least square linear regression, range from -0.34 per meg yr⁻¹ to 0.2 per meg yr⁻¹. Solid and broken horizontal bars in the bottom of the figure indicate the durations of use of the individual working reference gases.

To assess the stability of the NIES O₂/N₂ scales, we have continued to measure the reference gases in two 48-L aluminum cylinders (CQB-15645 and CQB-15649) since 2003, which are independent from the reference gases for the NIES scale. The results are shown in Figure 3. The average changing rates for the whole period, evaluated by linear regression analysis, are -0.14 ± 0.06 per meg yr⁻¹ for CQB-15645 and -0.05 ± 0.06 per meg yr⁻¹ for CQB-15649. Therefore, we conclude that the stability of the NIES O₂/N₂ scale has been maintained within ±0.2 per meg yr⁻¹ at least during the period of 2003-2016. However, this stability test cannot exclude the possibility that the O₂/N₂ ratios of the reference gases drift across all the cylinders rather uniformly. There are several mechanisms that affect the O₂/N₂ ratios of the gases within the high-pressure cylinders, including corrosion of the inner surface, leakage, thermal diffusion and gravitational fractionation. Keeling et al. (2007) assessed carefully and comprehensively the influences of those potential mechanisms on the long-term stability of the O₂/N₂ ratio of the reference gases and obtained an estimated uncertainty of ±0.4 per meg yr⁻¹. We also treated the reference cylinders, which were kept horizontally in a thermally insulated box, with the greatest care (Tohjima et al., 2008). Therefore, we adopted the value of ±0.4 per meg yr⁻¹ as the long-term drift of the reference gases caused by the above degradation effects. Consequently, we assumed that the total uncertainty of the long-term stability of the O₂/N₂ reference scale was ±0.45 per meg yr⁻¹ (= (0.2²+0.4²)^{1/2}) in this study.

2. 4. Carbon budget calculation

The ocean and land uptake, *O* and *B*, are given by the following equations (Manning and Keeling, 2006; Tohjima et al., 2008):

$$O = \left[-(\alpha_F - \alpha_B)F - \left(\frac{SO_2}{\beta} \right) \times \Delta APO + Z_{eff} \right] \times \frac{1}{\alpha_B}, \quad (6)$$

$$B = \left[\alpha_F F + \left(\frac{SO_2}{\beta} \right) \times \Delta APO - \left(\frac{\alpha_B}{\beta} \right) \times \Delta X_{CO_2} - Z_{eff} \right] \times \frac{1}{\alpha_B}, \quad (7)$$

where β is the coefficient converting PgC to ppm CO₂ in the atmosphere (β=0.470 ppm/PgC, Tohjima et al., 2008), and Z_{eff} represents the net effect of the air-sea O₂ and N₂ exchange on the atmospheric O₂/N₂ ratio. In these equations, *O*, *B*, *F* and Z_{eff} are given in units of PgC yr⁻¹, Δ*AP*O in units of per meg yr⁻¹, and Δ*X*_{CO₂} in units of ppm yr⁻¹. Note that *F* and α_{*F*} include the CO₂ emissions associated with cement manufacturing. The values of α_{*F*} were calculated from the CO₂ emission amounts and the -O₂/CO₂ molar exchange ratios of the individual fuel types (Keeling 1988). Since α_{*F*} slightly varies year by year, the value of α_{*F*} for the relevant period was used for the carbon budget calculations. The values of Z_{eff} were calculated based on the effects of global ocean warming and anthropogenic nitrogen deposition in accordance with the approach of Keeling and Manning

(2014). The details of the Z_{eff} calculation is discussed in the following section.

We used the same data set of the fossil fuel-derived CO₂ emissions and the global average of the atmospheric CO₂ mole fractions as the Global Carbon Project (GCP) used for the **estimation of global carbon budget 2018** (Le Quéré et al. 2018). The fossil CO₂ emissions from fossil fuel combustion and cement production were basically based on the dataset from CDIAC and
5 other energy statistics (Boden et al, 2017). The change in the atmospheric CO₂ burden was calculated based on the global observation by the US National Oceanic and Atmospheric Administration Earth System Research Laboratory (NOAA/ESRL; Dlugokencky and Tans, 2018). The temporal variations in the fossil CO₂ emissions and the atmospheric CO₂ accumulation rate are depicted in Fig. 4.

Annual means of APO centered on January 1 were computed by using the same procedure as Tohjima et al. (2008). First,
10 smooth curve fits to the data were computed in accordance with the methods of Thoning et al. (1998) with a cut-off frequency of 4.6 cycles yr⁻¹. Then the flask APO data were modified to represent the values at the center of the individual months by shifting them in parallel with the smooth curve fits. This procedure aimed to reduce the influence from biases of the sampling timings within the individual months. The monthly averages were calculated from the modified APO data. When there were
15 no flask data in the monthly time frame, the monthly average of the smooth curve was used. The annual means were calculated from the consecutive 12 monthly averages from July to June of the following year. The standard errors of the differences between the flask data and the smooth curve fits for the corresponding annual periods were adopted as the uncertainties for the annual averages. The averages and ranges (minimum ~ maximum) of the errors for the annual means of APO were 0.8 per meg (0.6~1.3 per meg) for HAT and 1.1 per meg (0.9 ~ 1.4 per meg) for COI.

In this study, we adopted the value of 1.10 of Severinghaus (1995) for α_B in accordance with a series of previous studies (e.g.
20 Bender et al., 2005; Manning and Keeling, 2006; Ishidoaya et al., 2012; Keeling and Manning, 2014; Goto et al., 2017). However, several studies (e.g. Randerson et al., 2006; Worrall et al., 2013), investigating the elemental compositions of organic matters in soil and plants, indicated that the value of 1.1 is rather large for the globally averaged net -O₂/CO₂ exchange ratio for the terrestrial biosphere. These studies suggest that the value of 1.05 is much more appropriate for α_B . Although it's beyond
25 the scope of this study to discuss which value is better for α_B , it is useful to mention that the use of 1.05 for α_B results in larger decreasing rates of APO by about 5% and an increase in land sinks and a decrease in ocean sinks by about 0.06 PgC yr⁻¹ on average in the following results. **Considering recent reports about the global net -O₂/CO₂ exchange ratio, Keeling and Manning (2014) revised the uncertainty of α_B upward from ± 0.05 (Severinghaus, 1995) to ± 0.10 . Thus, we also adopted ± 0.10 for the uncertainty of α_B in this study.**

2. 5. Evaluation of outgassing effect (Z_{eff})

30 As is discussed in the Introduction, today's ocean is considered to act as a net source of atmospheric O₂ because of global ocean warming, which also affects the air-sea N₂ exchange. Since the atmospheric O₂ change is measured as the change in the atmospheric O₂/N₂ ratio, the outgassing effect **caused by the global ocean warming, which is denoted as Z_{gow}** , should include the influences from not only ocean O₂ outgassing but also ocean N₂ outgassing. Assuming the relationship is proportional between the gas fluxes and heat fluxes across the air-sea interface, **Manning and Keeling (2006)** gave the equation for Z_{gow} as:

$$35 \quad Z_{gow} = \left(\gamma_{O_2} - \frac{S_{O_2}}{S_{N_2}} \gamma_{N_2} \right) \times Q \times m_c \times 10^{-15}, \quad (8)$$

where Q represents the changing rate of the global ocean heat storage in units of J yr⁻¹, γ_{O_2} and γ_{N_2} are gas flux / heat flux ratios between the air and sea in units of mol J⁻¹, S_{N_2} is the mole fraction of atmospheric nitrogen ($S_{N_2}=0.7809$, Tohjima et al., 2005b) and m_c is the atomic mass of carbon ($m_c=12.01$). Z_{gow} is given in units of **PgC yr⁻¹**.

The primary mechanism that affects the air-sea gas exchange is a reduction of gas solubility caused by the increase in the ocean

temperature. Therefore, the gas flux / heat flux ratio derived from the above thermal effect can be evaluated from the temperature dependence of gas solubility in the seawater and the specific heat of the seawater. Since the air-sea N₂ exchange is predominantly driven by the thermal effect, we adopted the estimated γ_{N_2} of 2.2 nmol J⁻¹ in this study in accordance with previous studies (Keeling and Garcia, 2002; Manning and Keeling, 2006).

5 In contrast to the air-sea N₂ exchange, the changes in the ocean circulation and ocean primary production also affect the air-sea O₂ exchange as is mentioned in the Introduction. Examining the ratio of the seasonal ocean outgassing of O₂ to the seasonal ocean heating and the negative linear relationship between the dissolved O₂ concentrations corrected for ocean biological processes and the potential temperature in the main thermocline based on archived global observation data, Keeling and Garcia (2002) obtained the estimate of 4.9 nmol J⁻¹ for γ_{O_2} . The value of γ_{O_2} was also investigated by using ocean biogeochemical models to revise the global carbon budgets based on O₂ observations (e.g. Plattner et al., 2002; Bopp et al., 2002). Keeling et al. (2010) summarized the model-based values of γ_{O_2} ranging from 5.9 to 6.7 nmol J⁻¹. On the other hand, Stendardo and Gruber (2012) examined a huge archived dataset of observations in the North Antarctic Ocean during the past five decades and obtained changing ratios of O₂ inventory to heat content of -4.3 ± 2.4 nmol J⁻¹ in the upper 700m and -1.6 ± 1.9 nmol J⁻¹ between 700 and 2750 m. These basin-scale ocean O₂/heat changing ratios seem to suggest that the global ocean acts as a net
15 O₂ source due to global ocean warming. Therefore, in this study, we used the value of 4.9 nmol J⁻¹ for γ_{O_2} according to previous studies.

To compute Q, we used the estimates of the world ocean heat content (OHC) based on a variety of oceanographic data (Levitus et al., 2000; Levitus et al., 2012). Time series of OHC for the 0-700 m and 0-2000 m layers are available from the NOAA's web site (https://www.nodc.noaa.gov/OC5/3M_HEAT_CONTENT/). In previous carbon budget estimations based on
20 atmospheric O₂/N₂ measurements, the values of Q were estimated from the OHC for the 0-700 m layer (e.g. Manning and Keeling, 2006; Tohjima et al., 2008; Ishidoya et al. 2012). Levitus et al. (2012) showed, however, that the ocean heat storage of the 700-2000 m layer contributes to about one third of the total heat storage of the 0-2000 m layer. Keeling and Manning (2014) estimated the value of Q by considering not only Q for the depths above 700 m but also Q for the depths below 700, which contributed to 30% of the total value of Q. Therefore, the time series of OHC for the 0-2000 m layer was used in this
25 study. Note that since the annual average of OHC for the 0-2000 m layer is available only after 2005, we used the pentad (5-year) averages before 2005.

In addition to the ocean warming effect, Keeling and Manning (2014) introduced recently another ocean outgassing effect caused by atmospheric deposition of excess anthropogenic nitrogen to the open ocean. The excess nitrogen is considered to enhance the ocean biotic production of organic matter, which is associated with the O₂ production. Keeling and Manning
30 (2014) evaluated the anthropogenic nitrogen-induced outgassing as being about 0.1×10^{14} mol O₂ yr⁻¹. Since the outgassing effect caused by the anthropogenic nitrogen deposition is small but rather significant, we adopted the effect as Z_{anthN} , with a magnitude of 0.12 PgC yr⁻¹ ($=0.1 \times 10^{14}$ mol O₂ yr⁻¹ $\times 12.01$ gC mol⁻¹). Eventually, the total outgassing effect, Z_{eff} , is expressed as the summation of Z_{gow} and Z_{anthN} :

$$Z_{eff} = Z_{gow} + Z_{anthN}. \quad (9)$$

35 The time series of the annual Z_{eff} divided by α_B are depicted in Fig. 4 as purple lines. The value of Z_{eff}/α_B ranges from 0.2 Pg yr⁻¹ to 1.0 Pg yr⁻¹, and the 19-year average for 1998-2016 is 0.52 Pg yr⁻¹. There are not much differences in Z_{eff} between this study and previous studies (e.g. Manning and Keeling, 2006; Keeling and Manning, 2014; Tohjima et al., 2008; Ishidoya et al., 2012; Goto et al., 2017).

3. Results and discussion

40 3.1. Observed O₂/N₂, CO₂, and APO

The time series of the atmospheric CO₂ mole fraction, O₂/N₂ ratio, and APO of the air samples collected at HAT and COI and onboard cargo ships sailing between 40°S and 30°N in the western Pacific are depicted in Fig. 5 together with the smooth-curve fits. The ship data were binned into 10-degree latitudinal bands (40-30°S, 30-20°S, ..., 20-30°N). Note that there are no data gaps with more than 50 days in the time series at HAT and COI while the time series of TF5 have data gaps during the 7-month period from October 2006 to April 2007. The ship data during the 7-month period were significantly contaminated by the inboard air due to the failure of diaphragm of the sampling pump.

The temporal variations in the annual means of the atmospheric CO₂, O₂/N₂ and APO, centered on January 1st, are shown in Figs. 6a, 6b, and 6c, respectively, where linear trends obtained from least square fitting to the data of HAT were subtracted from the individual time series to emphasize the interannual variations. The standard errors of the annual means for HAT and 10-0°S bin are depicted as vertical bars for typical examples. Note that the annual means of the atmospheric CO₂, O₂/N₂ and APO and the corresponding standard errors for HAT, COI and the 10-degree bins are summarized in Table S1-S9 in the Supplement. The annual means of CO₂ and O₂/N₂ show latitudinal gradients of northward increase and southward increase, respectively, because the fossil fuel-derived CO₂ emissions and O₂ consumptions occur predominantly in the northern mid-latitudes. In contrast, the highest values of the annual mean APO were generally observed around the equator as previously reported (Battle et al., 2006; Tohjima et al., 2005b, 2012). This equatorial peak is mainly attributed to large-scale air-sea gas exchanges: ingassing in the mid- and high-latitudes and outgassing in the equatorial region. Conducting atmospheric inversion analyses based on the APO data from the Scripps observation network, Rödenbeck et al. (2008) suggested anomalous outgassing of APO from the equatorial region during El Niño periods, while Tohjima et al. (2015) found a suppressed equatorial peak during El Niño periods based on the western Pacific observations. Eddebbar et al. (2017) reconciled these conflicting results by predicting the existence of a zonal dipole-like ENSO response in the equatorial Pacific based on several ocean process-based models and an atmospheric transport model. These results suggest that an enhanced zonal coverage of the atmospheric observations in the equatorial Pacific is needed to constrain the full basin-scale ENSO response. We can see a considerable suppression of the equatorial peak during the strong 2015/2016 El Niño event in Fig. 6c, which was not reported in Tohjima et al. (2015). Any detailed discussion about the temporal variation of the equatorial peak during the 2015/2016 El Niño event is, however, beyond the scope of this study and will be given elsewhere.

Fig. 6d shows the time series of the annual changes in the annual mean APO, which are the annual changing rates of APO for a one-year interval ($\Delta t = 1$ year). As you can see, there are considerable differences in the annual changing rates among the observation sites in the same years; the standard deviations range from 1.6 to 4.4 per meg and the average is 2.8 per meg. We also depict the averages of the annual changing rates of APO of HAT and COI and of all the shipboard data as thick grey line in Fig. 6d. Note that these average annual changing rates of APO were used for calculation of the global carbon budget in the following sections. The average annual changing rates show also a large interannual variability with a standard deviation of 4.7 per meg yr⁻¹ for the entire observation period.

The differences in the changing rate of APO among the sites in the same years decrease with increase in the interval for the calculation (Δt) as shown in Fig. 7, where the average (red circles) and the minimum and maximum (red broken lines) standard deviations of the changing rates are plotted against Δt . The differences among the sites decrease almost inversely with Δt ; the average standard deviation for $\Delta t = 5$ year is 0.54 per meg yr⁻¹. The temporal variability in the changing rate also decreases inversely with Δt as depicted in Fig. 7 (blue circles); the standard deviation is reduced to 1.2 per meg yr⁻¹ for $\Delta t = 5$ years. The above results seem to suggest that the temporal variability in the APO fluxes exceeds the spatial variability. As is indicated by Eq. (3), the temporal variability in the APO changing rate should be attributed mostly to those in O and Z_{eff} . Therefore, the above results also indicate that an interval of 5 years could suppress the temporal variability in Z_{eff} to the level of ± 1.2 per meg yr⁻¹, which corresponds to a carbon budget of about ± 0.5 PgC yr⁻¹.

The changing rates of the atmospheric CO₂, O₂/N₂, and APO for several combinations of time periods and the observed data

(HAT, COI and shipboard) are summarized in Table 1. Here, the uncertainties of the changing rates were computed from the uncertainties of the corresponding annual means at both ends of the periods, the estimated uncertainty of the O_2/N_2 scale stability (± 0.45 per meg yr^{-1} , Section 2. 3), and the uncertainty in the O_2/N_2 span sensitivity ($\pm 3\%$, see below). The time periods of 2000-2010, 2001-2010, and 2001-2014 were selected to compare the observational results of this study with those of Scripps
5 Institution of Oceanography (SIO) and Tohoku University (TU) (Keeling and Manning, 2014; Ishidoya et al., 2012; Goto et al., 2017). As is discussed in the above section, the differences in the long-term changing rates of APO between HAT, COI and shipboard data are less than 0.3 per meg yr^{-1} , while the increasing rates of CO_2 and the decreasing rates of O_2/N_2 for HAT are slightly larger than those for COI and other sites. The monitoring station of HAT is located at the marginal region of continental East Asia, and the anthropogenic CO_2 emissions from China often influence the observations at HAT during winter due to the
10 East Asian monsoon (c.f. Minejima et al., 2012; Tohjima et al., 2010, 2014). Additionally, for the period of 2000-2014, the fossil fuel-derived CO_2 emissions from China show a rapid increase in association with the unprecedented economic growth. These situations may explain the rather large increase in CO_2 and decrease in O_2/N_2 at HAT. In contrast to CO_2 and O_2/N_2 , the emissions from fossil fuel combustion and land biotic processes contribute less to the APO variations, resulting in relatively small differences in the long-term APO changing rates among the sites.

15 It should be noted that the decreasing rates of APO of our study are $0.5 \sim 1.1$ per meg yr^{-1} smaller than those of SIO and TU. Except for the differences of the observation sites, we can offer two explanations for the discrepancy. First, the calculation methods of the changing rate adopted by Goto et al. (2017) are different from those adopted in this study, which might partially explain the discrepancy. This is understandable when comparing the changing rates of CO_2 , O_2/N_2 , and APO for the individual studies. In this study, the APO changing rates are almost consistent with those calculated from the CO_2 and O_2/N_2 changing
20 rates according to the APO definition. However, in the study of Goto et al. (2017), the CO_2 and O_2/N_2 changing rates of Ny-Ålesund give APO decreasing rates of -9.4 per meg yr^{-1} , which is 0.7 per meg yr^{-1} smaller than the originally reported values. Second, inter-laboratory comparison of flask samples and high-pressure cylinders suggests a possibility that the span sensitivity of the O_2/N_2 measurements of NIES is about 3% lower than that of SIO, which can almost explain the differences in the APO decreasing rates. However, to obtain an accurate conclusion, we need much more studies.

25 3. 2. Calculation of global carbon budgets

The rates of the global carbon uptake by the ocean and land biosphere were calculated from the average changing rate of APO based on observations at COI, HAT and on cargo ships ($40^{\circ}S-30^{\circ}N$). The results for several time periods are summarized in Table 2 together with the average changing rates of APO, globally averaged atmospheric CO_2 accumulation rates, fossil fuel-derived CO_2 emission rates, and the ocean outgassing effect divided by α_B . The 19-year (1998-2016), 17-year (2000-2016),
30 and 14-year (2003-2016) periods correspond to the individual maximum observation periods for HAT, COI, and the western Pacific, respectively. For example, the estimated ocean and land sinks for 2000-2016 were found to be 2.6 ± 0.7 PgC yr^{-1} and 1.5 ± 0.9 PgC yr^{-1} , respectively.

The uncertainties in the parameters used for the carbon budget calculation (Eqs. (6) and (7)), which are also listed in Table 2, are briefly discussed here. Note that in this study the estimated uncertainties are $\pm 1\sigma$. Since the ocean outgassing effect is rather
35 speculative, we assumed that the values of Z_{eff} for the individual periods had $\pm 100\%$ uncertainties in accordance with previous studies (e.g. Manning and Keeling, 2006; Tohjima et al., 2008). We adopted uncertainties of $\pm 5\%$ for the fossil fuel-derived CO_2 emission rate and ± 0.2 PgC yr^{-1} for the atmospheric CO_2 increasing rate from Le Quéré, et al. (2018). As for the uncertainties of the observed APO changing rates, we adopted the standard deviations among the sites shown in Fig. 7 (± 0.37 per meg yr^{-1} for longer than 10 years and ± 0.54 per meg yr^{-1} for 5 years). The estimated uncertainty of the O_2/N_2 scale stability
40 (± 0.45 per meg yr^{-1}) discussed in Section 2. 3, the uncertainty of the O_2/N_2 span sensitivity ($\pm 3\%$), and the uncertainty in the global averaged APO associated with the limited atmospheric sampling (± 0.2 PgC yr^{-1}) discussed in Nevison et al. (2008)

were also included in the calculation of the uncertainties in ΔAPO . The uncertainties of α_B and α_F were ± 0.10 (Keeling and Manning, 2014) and ± 0.04 (Tohjma et al., 2008), respectively. Finally, these uncertainties were propagated to the ocean and land sink uncertainties in accordance with Eqs. (6) and (7).

We compared our global carbon budget estimations with those of GCP (Global Carbon Budget 2018) updated by Le Quéré, et al. (2018). In the GCP carbon budget assessment, the ocean and land sinks were estimated by combining multiple results from a variety of models including global ocean biogeochemistry models (GOBMs) and dynamic global vegetation models (DGVMs). Since the sum of the model-based ocean and land sinks was not necessarily balanced with the difference between fossil fuel emissions and atmospheric accumulation, Le Quéré, et al. (2018) listed the discrepancies as budget imbalances. The ocean sinks, the land sinks which are ‘net’ land sinks computed as the differences between land uptake and emissions associated with land-use change, and the budget imbalances for the corresponding periods are listed in Table 2. Note that the uncertainties of the sinks of GCP are $\pm 0.5 \text{ PgC yr}^{-1}$ for ocean and $\pm 0.9 \text{ PgC yr}^{-1}$ for land. The carbon sinks of this study and GCP for the three long periods are consistent with each other; the largest difference in sink strength is 0.35 PgC yr^{-1} , which is smaller than the uncertainties associated with the individual estimations. A 3% higher span sensitivity of the O_2/N_2 measurements, which corresponds to the difference in the span sensitivity between SIO and NIES (section 3.1), would result in an increase and decrease of 0.27 PgC yr^{-1} in the ocean and land sinks, respectively, for the three long periods. Although these changes would enlarge the differences in sinks between GCP and this study, the differences are still not significant given the uncertainties of both this study and GCP.

The carbon budgets for four pentad periods (2000–2004, 2004–2008, 2008–2012, 2012–2016) are also listed in Table 2. Here, we consider that the 5-year interval effectively reduced the apparent errors caused by the imbalance of the seasonal ocean O_2 fluxes, as is discussed in Section 3.1. Again, the discrepancies of the pentad ocean and land sinks between this study and GCP are within $\pm 0.5 \text{ PgC yr}^{-1}$, which is also less than the estimated uncertainties. The land sink during 2008–2012 and the ocean sink during 2012–2016 of this study are about 0.5 PgC yr^{-1} larger than those of GCP. These discrepancies in the carbon sinks, partly explained by the rather large values of the carbon budget imbalances of the GCP estimation, might give a clue about how to partition the imbalance values between the land and ocean sinks.

Examining the temporal variations in the pentad sink strengths of this study, we found a gradual increase in the ocean sinks for the latter three pentad periods and a rapid increase and decrease in the land sinks for the former and latter two pentad periods, respectively. The pentad averages of the GCP sinks seem to show similar temporal variations: a steady increase in the ocean sinks for the whole period and a rapid increase and decrease in the land sinks for the former and latter two pentad periods, respectively. These results suggest that the carbon sinks for the pentad periods can be used to evaluate the temporal changes. In the following section, we will examine the temporal change in the carbon sinks in more detail.

3. 3. Temporal change in the carbon sinks

Fig. 8 shows the temporal variations in the ocean and land sinks for the annual (broken red lines) and pentad (red lines) intervals calculated from the average of the APO changing rates based on the observations from HAT, COI, and cargo ships in the western Pacific. The uncertainties for the pentad sinks ($\pm 1\sigma$), which were calculated as described in the previous section, are shown as gray shading. To clearly understand the effect of Z_{eff} correction, the pentad sinks without corrections are also depicted as purple lines in the figure. Additionally, the annual and pentad sinks of GCP are also depicted in the figures for comparison. Although the annual sinks show considerable variability especially for the first several years with peak-to-peak differences of more than 10 PgC yr^{-1} , the variability of the pentad sinks is effectively suppressed. Only the pentad budgets for 2000 show a rather large ocean uptake (3.21 PgC yr^{-1}) and a rather weak land emission (0.56 PgC yr^{-1}), which are depicted as dotted lines. These anomalous values may be explained by the fact that the influences from the considerable drawdown of APO in 2000–2001 cannot be compensated for in the pentad APO changing rate for 2000. Hamme and Keeling (2008) reported that the APO

drawdown in 2000-2001, which was also observed in the SIO observations, may be attributed to deep ventilation associated with the unprecedented cooling of the western Pacific, and the variations in the ocean heat content exerted only secondary influence. Therefore, we don't use the anomalous pentad ocean and land sinks for 2000 in the following discussions.

The pentad ocean sinks show an overall increasing trend although there is a dip in the ocean sink centered on 2004-2005 by about 0.6 PgC yr⁻¹. Nevison et al. (2008) suggested that a decadal or longer period is needed to suppress the influence of the interannual variation in the ocean O₂ flux on the carbon sink estimation within ±0.1 PgC yr⁻¹ based on an ocean ecosystem model and an atmospheric transport model. In addition, the pentad APO changing rate still contains an uncertainty corresponding to ±0.5 PgC yr⁻¹ as is discussed in Section 3.1. Therefore, the anomalous dip in the ocean sink for 2004-2005 might be an error caused by the anomalous ocean O₂ flux variations. The increasing rate of the ocean sink during 2001-2014, determined by linear regression, is 0.08 ± 0.02 PgC yr⁻², which is larger than that of GCP which was 0.04 ± 0.01 PgC yr⁻². Although the temporal variability in the ocean sink in the GCP study is rather suppressed, which is attributed to the rather coarse resolution of the GOBMs (Le Quéré, et al., 2018), a much larger decadal and sub-decadal variability has been reported in the ocean sink estimations based on archived data of the observed surface partial pressure of CO₂ (pCO₂) (Landschützer et al., 2015; DeVries et al., 2017). Results from the Surface Ocean pCO₂ Mapping (SOCOM) initiative show that the decadal linear trend of the global ocean sink enhancement over 2001-2011 based on pCO₂ data and selected mapping methods is about 0.8 PgC yr⁻¹ per decade (Rödenbeck et al., 2015). For a detailed comparison, the global ocean sinks based on pCO₂ observations and interpolation techniques (Landschützer et al., 2016; Rödenbeck et al., 2014) for the period of 1990-2017 are plotted in Fig. 9 together with the ocean sinks of both this study and GCP. Note that the extended pCO₂-derived ocean sinks were given as supplementary data of Le Quéré, et al. (2018) and those sinks were uniformly inflated by 0.78 PgC yr⁻¹ to compensate for the pre-industrial steady state source of CO₂ derived from riverine input of carbon to the ocean (Resplandy et al., 2018). As you can see, both the GCP and pCO₂-derived ocean sinks show changes in the trends before and after 2001, while the magnitude of the changes in the pCO₂-derived sinks are larger. The increasing rates determined by linear regression during 2001-2014 are 0.08 ± 0.01 PgC yr⁻² for Landschützer et al. (2016) and 0.07 ± 0.02 PgC yr⁻² for Rödenbeck et al. (2014), which are more consistent with the rate found in this study. Therefore, our result seems to support a previous conclusion that the recent increase in the ocean sinks exceeds the increasing trend of ocean sink expected only from the atmospheric CO₂ increase (Landschützer et al., 2015; DeVries et al., 2017).

In contrast, the pentad land sinks of both this study and the GCP study show an increasing trend during 2001-2009 followed by a decreasing trend during 2009-2014, although the range of variations of this study is about two times larger than that of GCP. The linear trends for the former period are 0.23 ± 0.04 PgC yr⁻² for this study and 0.10 ± 0.03 PgC yr⁻² for GCP, and those for the latter period are -0.22 ± 0.04 PgC yr⁻² for this study and -0.12 ± 0.04 PgC yr⁻² for GCP. An enhancement of the land uptake during the 2000s has been reported recently by several studies based on atmospheric inversions and biosphere models (Keenan, et al., 2016; Kondo et al., 2018; Piao et al., 2018). Although there is an ongoing discussion about the detailed mechanisms of the enhanced net land uptake, the accelerated land uptake may partially explain the stagnancy of the growth rate of the atmospheric CO₂ in the 2000s despite the increasing anthropogenic CO₂ emissions. Examining the atmospheric inversion studies and the previous version of GCP (Le Quéré, et al., 2015), in which the net land uptake were computed as residuals among the other carbon budget components, Piao et al. (2018) found that the linear increasing trend of the net land carbon sink during 1998-2012 was 0.17 ± 0.05 PgC yr⁻². The linear trend of this study during 2001-2009 is close to the above value within the uncertainty. Although the corresponding linear trend of the latest GCP estimation is about half of that of the present study, the sum of the net land sink and the budget imbalances of GCP, plotted as light blue lines in Fig. 8, shows a much larger increasing trend, 0.20 ± 0.03 PgC yr⁻², which is almost identical to our trend.

The land sinks of both this study and the GCP study exhibit decreasing trends for the period 2009-2014, which are partially compensated for by the steady increase in the ocean uptake. The atmospheric accumulation rate of CO₂ significantly increased

in 2015 and 2016 (see Figure 5), when one of the strongest El Niño events occurred. Studies based on atmospheric CO₂ observations from stations and by satellite indicated that a reduction in biospheric uptake and an increase in biomass burning contributed to the CO₂ increase during the El Niño event (Chatterjee, et al., 2017; Patra et al., 2017). The decreasing trend of the pentad land uptake also reflects the change in the global carbon cycle associated with the El Niño event.

- 5 Shown as discrepancies between the pentad sinks of GCP with and without budget imbalances (Fig. 8), the magnitude of the budget imbalances seems to increase after 2007. For the pentad sinks centered on 2007, 2008 and 2009, the ocean and land sinks of this study agree with those of GCP without and with the budget imbalances, respectively. For the pentad sinks between 2010 and 2014, both the ocean and land sinks of this study are plotted between those of GCP with and without the budget imbalances. Although we cannot show any definitive partitioning of the budget imbalance between ocean and land sinks
- 10 because of a rather large uncertainty associated with the sink estimations, the above results seem to suggest that the budget imbalances are allocated to land sinks for the former period and to both sinks for the latter period.

From the above discussions, we feel that a five-year duration effectively suppresses to some extent the anomalous variations in the carbon budget estimations based on APO, which are considered to be caused by the imbalance of the seasonal air-sea O₂ exchange. Probably, the five-year average suppresses the variability of Z_{eff} to a level of ± 0.5 PgC yr⁻¹, as is discussed in

15 Section 3.1. To reduce uncertainty in the carbon budget estimation, we need more effort to improve the quantification of the net O₂ outgassing associated with global ocean warming because the quantification of the Z_{eff} at this state is still very speculative. Applying the approach of Stendardo and Gruber (2012), who examined the long-term changes in dissolved O₂ and heat content by using archived oceanographic data of the Atlantic Ocean, to other ocean basins would improve our understanding of the long-term net ocean O₂ flux/heat flux ratio.

20 4. Conclusion

We have evaluated the global carbon budgets based on the APO data computed from the O₂/N₂ and CO₂ of the flask samples collected in the Pacific region since 1997. In the carbon budget calculation, we corrected the ocean and land sinks with the ocean O₂ outgassing effect, Z_{eff} , based on the ocean heat increment for the 0-2000 m layer. Eventually, we have obtained the following conclusions:

- 25 1) The long-term oceanic and land biotic carbon sinks were 2.6 ± 0.7 PgC yr⁻¹ and 1.5 ± 0.9 PgC yr⁻¹, respectively, for a 17-year period (2000-2016), and 2.4 ± 0.7 PgC yr⁻¹ and 1.9 ± 0.9 PgC yr⁻¹, respectively, for a 14-year period (2003-2016). These long-term carbon sinks agreed well with those of the latest GCP estimation (Le Quéré, et al., 2018); the differences of the individual estimations are less than ± 0.35 PgC yr⁻¹.
- 30 2) The ocean and land sinks for the four pentad (five-year) periods (2000-2004, 2004-2008, 2008-2012, 2012-2016) of this study also showed good agreement with those of GCP within a difference of ± 0.5 PgC yr⁻¹. The land and ocean sinks of this study showed larger values by about 0.5 PgC yr⁻¹ than those of GCP for 2008-2012 and 2012-2016, respectively, when rather large carbon budget imbalances (>0.5 PgC yr⁻¹) were found. Therefore, the discrepancies in the sinks between this study and GCP might give a clue about how to partition the imbalance values between the land and ocean sinks.
- 35 3) Calculating the carbon budgets for the pentad periods consecutively, we examined the changing trend of the ocean and land sinks during a 14-year period (2001-2014). In general, the changing trends of both land and ocean sinks of this study agreed well with those of GCP, although the range of variations of this study was about two times larger than that of the GCP study. The pentad ocean sinks showed an overall increasing trend for the entire period (2001-2014) with a linear increasing rate of 0.08 ± 0.02 PgC yr⁻². This increasing rate was about two times larger than that
- 40 for the GCP ocean sinks (0.04 ± 0.01 PgC yr⁻²) but was consistent with those for the global ocean sinks based on

pCO₂ observations and interpolation techniques (Landschützer et al., 2016; Rödenbeck et al., 2014). In contrast, the pentad land sinks showed an increasing trend for 2001-2009 and a decreasing trend for 2009-2014. The linear trends of the land sinks for this study and the GCP (in parentheses) were 0.23 ± 0.04 PgC yr⁻² (0.10 ± 0.03 PgC yr⁻²) for the former period and -0.22 ± 0.04 PgC yr⁻² (-0.12 ± 0.04 PgC yr⁻²) for the latter period. Enhancement of the land carbon uptake was reported also by previous studies (Keenan, et al., 2016; Kondo et al., 2018; Piao et al., 2018). In addition, the recent decreasing trend of the land uptake was found to be partially related to the global carbon cycle variation associated with the strong El Nino event in 2015 and 2016.

5. Author contributions

YT designed the study and drafted the manuscript. YT and YH conducted the O₂/N₂ measurements and analyzed the observed data. HM and TM organized the flask sampling at the monitoring stations and conducted the CO₂ measurements. SN managed the shipboard flask sampling. HM, TM, SN, and YH examined the results and provided feedback on the manuscript. All the authors approved the final manuscript.

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Tables

Table 1. Comparison of changing rate of the atmospheric CO₂, O₂, and APO

Period	Site	Average changing rate			Ref.
		CO ₂ (ppm yr ⁻¹)	O ₂ (per meg yr ⁻¹)	APO (per meg yr ⁻¹)	
1998-2016	HAT	2.19 ± 0.01	-21.8 ± 0.8	-10.4 ± 0.8	This study
2000-2016	HAT	2.21 ± 0.02	-22.0 ± 0.8	-10.4 ± 0.8	This study
2000-2016	COI	2.22 ± 0.02	-21.9 ± 0.8	-10.2 ± 0.8	This study
2003-2016	HAT	2.25 ± 0.02	-21.8 ± 0.8	-10.1 ± 0.8	This study
2003-2016	COI	2.26 ± 0.02	-21.6 ± 0.8	-9.8 ± 0.8	This study
2003-2016	W-Pacific	2.15 ± 0.06	-21.2 ± 0.8	-10.0 ± 0.8	This study
2000-2009	HAT	2.04 ± 0.02	-20.4 ± 0.8	-9.7 ± 0.8	This study
2000-2009	COI	1.91 ± 0.03	-19.9 ± 0.8	-9.7 ± 0.8	This study
2000-2009	Global	1.90 ± 0.02	-	-10.4 ± 0.5	Keeling & Manning (2014)
2001-2009	HAT	2.08 ± 0.03	-20.0 ± 0.8	-8.9 ± 0.8	This study
2001-2009	COI	1.87 ± 0.03	-19.0 ± 0.8	-9.2 ± 0.7	This study
2001-2009	Ny-Ålesund	2.00 ± 0.08	-21.2 ± 0.8	-	Ishidoya et al. (2012)
2001-2009	Showa	1.99 ± 0.06	-22.0 ± 0.8	-	Ishidoya et al. (2012)
2001-2013	HAT	2.19 ± 0.02	-21.3 ± 0.8	-9.7 ± 0.8	This study
2001-2013	COI	2.07 ± 0.03	-20.4 ± 0.3	-9.6 ± 0.8	This study
2001-2013	Ny-Ålesund	1.99 ± 0.02	-19.9 ± 0.3	-10.1 ± 0.3	Goto et al. (2017)
2001-2013	ALT, MLO, SPO	1.98 ± 0.03 ^a	-20.5 ± 0.3 ^a	-10.8 ± 0.1 ^a	Goto et al. (2017)

^aAverage and standard deviation of the changing rates for the three sites (ALT, MLO and SPO) listed in Table 1 of Goto et al. (2017) are given in this table.

Table 2. Comparison of global carbon budgets based on APO with those from GCP^{a,b}

Period	ΔAPO^c	Atm.	Fossil	α_F^e	$Z_{\text{eff}}/1.1^f$	Sink of this study		Sink of GCP ^d		
		CO_2^d	fuel ^d			Ocean	Land	Ocean	Land	Imb.
1998-2016	-10.3(0.91)	4.45	8.28	1.38	0.52	2.57(0.71)	1.26(0.89)	2.24	1.46	0.13
2000-2016	-10.3(0.91)	4.45	8.48	1.38	0.54	2.55(0.73)	1.48(0.91)	2.27	1.48	0.29
2003-2016	-9.9(0.91)	4.58	8.83	1.38	0.52	2.35(0.73)	1.90(0.93)	2.34	1.55	0.36
2000-2004	-8.8(0.94)	3.93	7.11	1.40	0.59	2.23(0.76)	0.94(0.90)	2.01	1.30	-0.14
2004-2008	-9.2(0.96)	4.08	8.21	1.38	0.33	1.97(0.62)	2.17(0.82)	2.18	1.74	0.22
2008-2012	-10.4(0.98)	4.19	9.05	1.37	0.54	2.54(0.77)	2.31(0.97)	2.32	1.85	0.68
2012-2016	-11.6(1.06)	5.36	9.65	1.37	0.71	3.05(0.90)	1.25(1.09)	2.55	1.26	0.49

^aFigures are given in units of per meg yr⁻¹ for ΔAPO , mol mol⁻¹ for α_F , and PgC yr⁻¹ for the others.

^bFigures in parentheses represent the uncertainties.

5 ^c ΔAPO is based on the data from HAT, COI, and cargo ships (40°S-30°N). The uncertainty in parentheses includes the uncertainty associated with the observations, stability in the O₂/N₂ scale, uncertainty derived from limited sampling, and uncertainty in the O₂/N₂ span sensitivity (see text).

^dThese figures are computed from the dataset summarized by the Global Carbon Project (GCP). The uncertainties are ± 0.2 PgC yr⁻¹ for the atmospheric CO₂ and $\pm 5\%$ for the fossil fuel emissions, ± 0.5 PgC yr⁻¹ for the ocean sinks, and ± 0.9 PgC yr⁻¹ for the land sinks (Le Quéré, et al., 2018).

10 ^eThe uncertainties for α_F are ± 0.04 mol mol⁻¹ (Tohjima et al., 2008).

^fThe values of Z_{eff} include both global ocean warming and anthropogenic nitrogen deposition effects, and the uncertainties are assumed to be $\pm 100\%$ (see text).

Figure captions

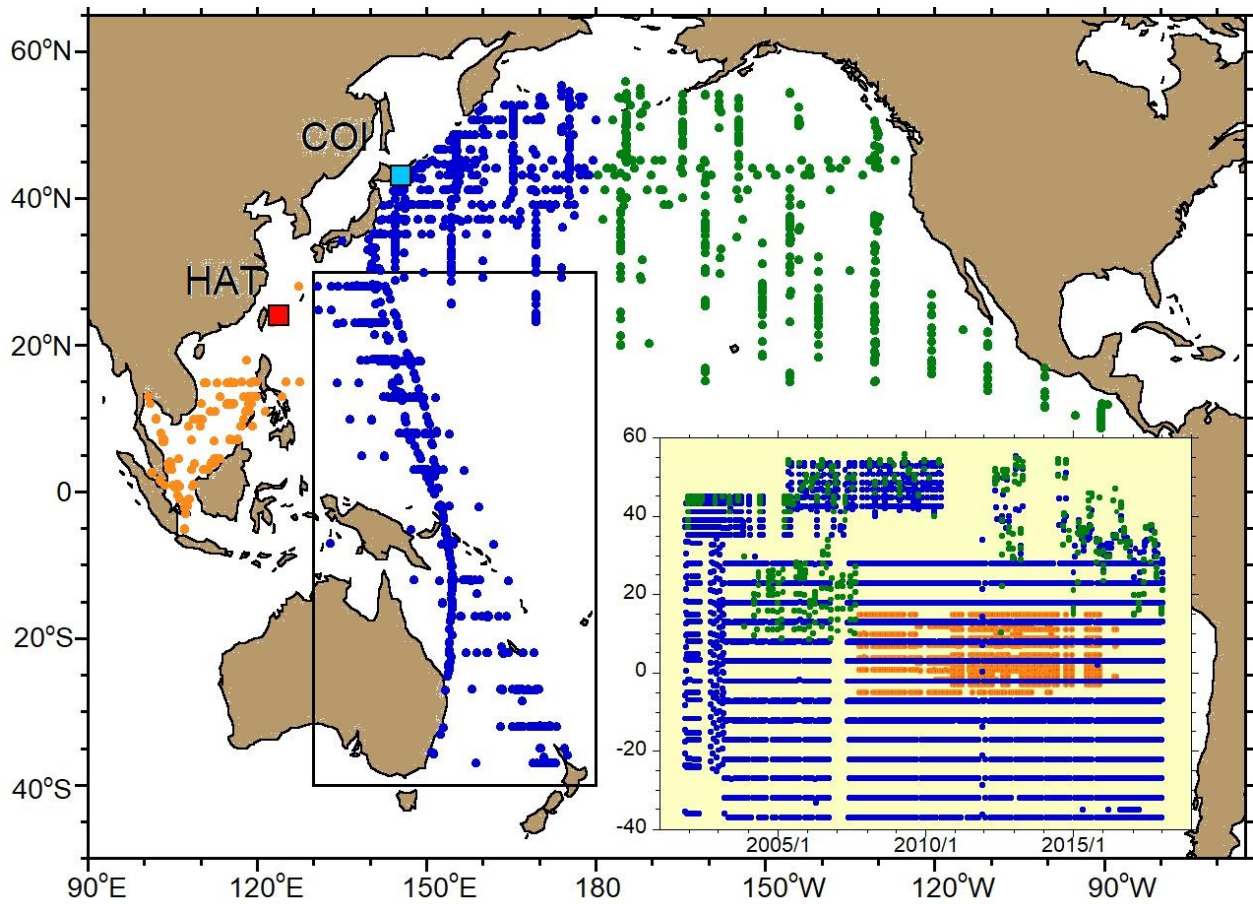


Fig. 1. Map showing the air sampling locations in the Pacific region. The light blue and red squares represent the monitoring stations of COI and HAT, respectively. The orange, blue, and green circles correspond to the positions where flask samplings were taken onboard cargo ships in South East Asia, western Pacific, and eastern Pacific, respectively. The inserted figure shows the time latitude distribution of onboard flask samples. The flask data from COI, HAT and the regions in the black rectangle (130-180°E, 40°S-30°N) were used in the budget calculations.

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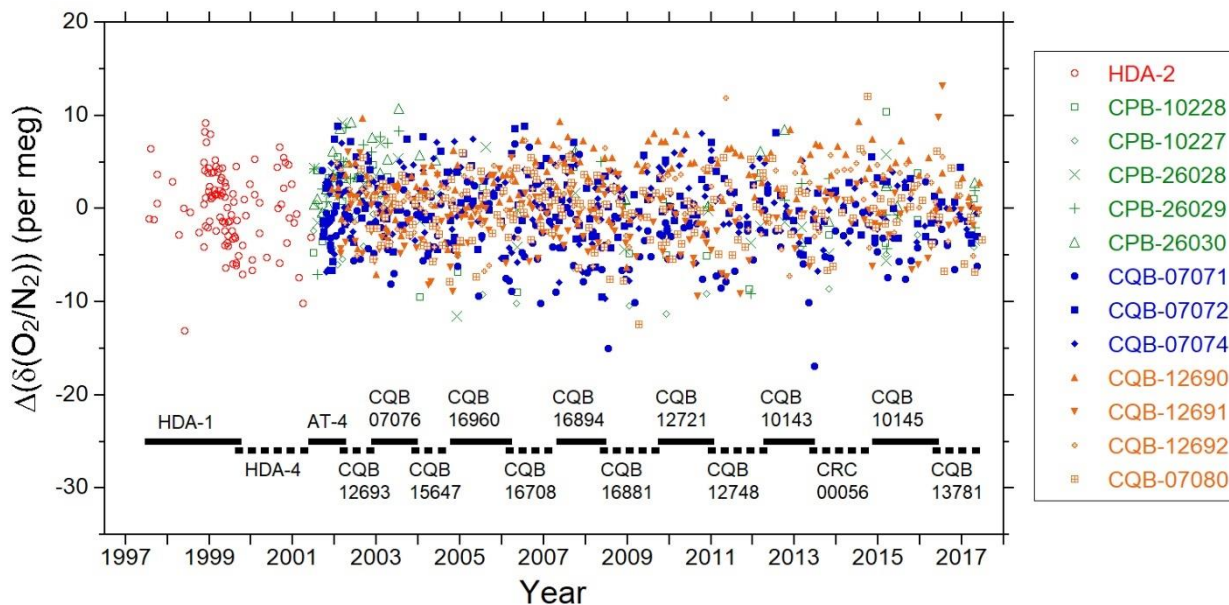


Fig. 2. Temporal changes in the O_2/N_2 ratio of primary reference gases relative to the NIES O_2/N_2 scale. The differences of the O_2/N_2 ratio from the average are plotted for HDA-2 along with the differences of the O_2/N_2 ratio from the initial values for the individual cylinders except HDA-2. Solid and broken horizontal bars in the lower part of the figure indicate the periods when the working reference gases were used.

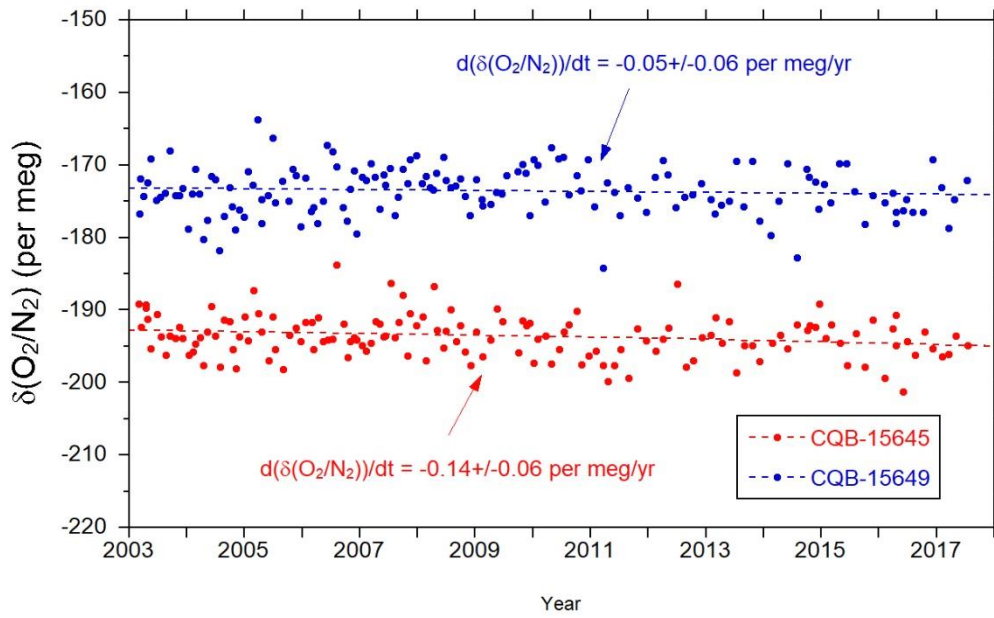


Fig. 3. Temporal changes in the O_2/N_2 ratio of reference airs in two aluminum cylinders which are independent from the NIES primary reference gases, relative to the NIES O_2/N_2 scale. The broken lines represent the linear regression lines.

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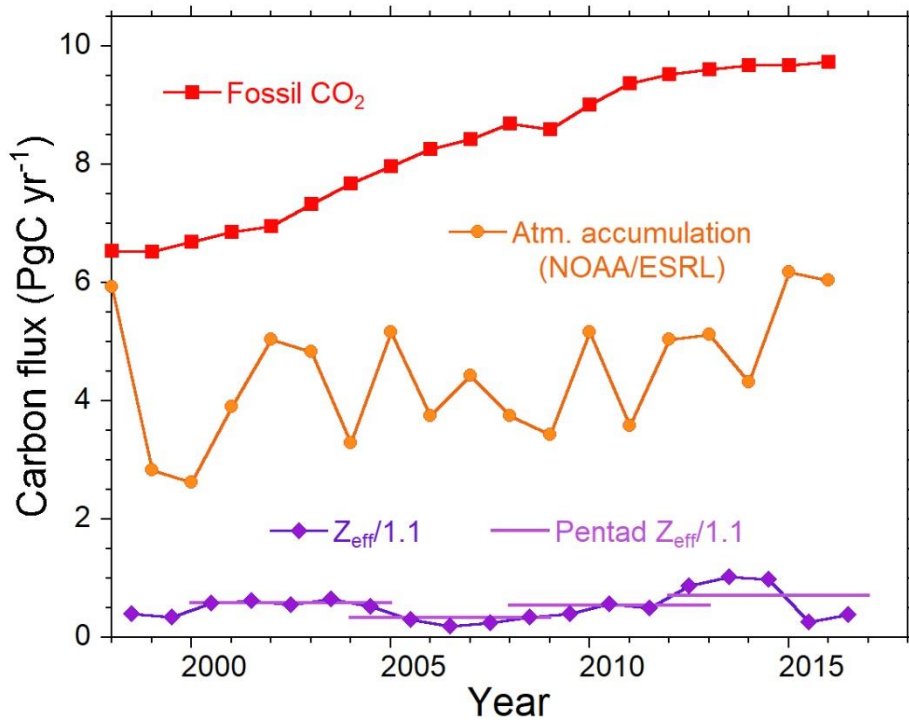


Fig. 4. Temporal changes in fossil fuel CO₂ emissions (red squares), atmospheric CO₂ accumulation rate (orange circles), and ocean outgassing effect Z_{eff} divided by land biotic $-O_2/CO_2$ exchange rate (1.1) (purple diamonds). The 5-year averages of $Z_{\text{eff}}/1.1$ used for the pentad (five-year) carbon sink calculations are also depicted as purple lines.

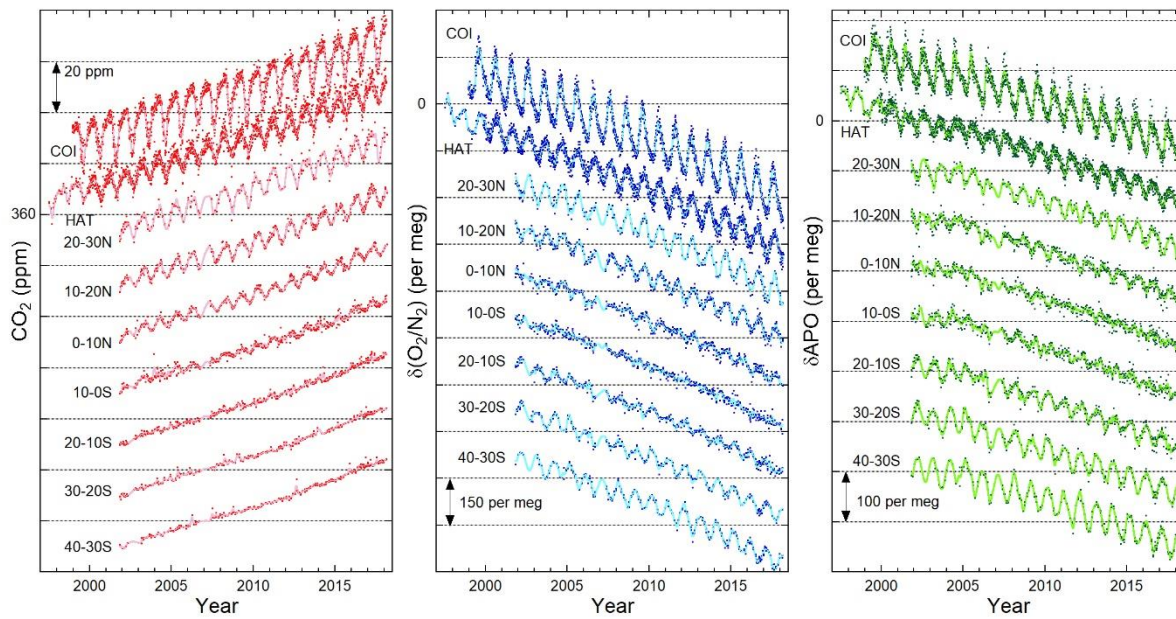


Fig. 5. Time series of the atmospheric CO₂ mole fraction (left), O₂/N₂ ratio (middle), and APO (right) of the flask samples obtained from the NIES flask sampling network shown in **Fig. 1**. Observed data from COI, HAT, and cargo ships operating between 40°S and 30°N were used for the global carbon budget calculation.

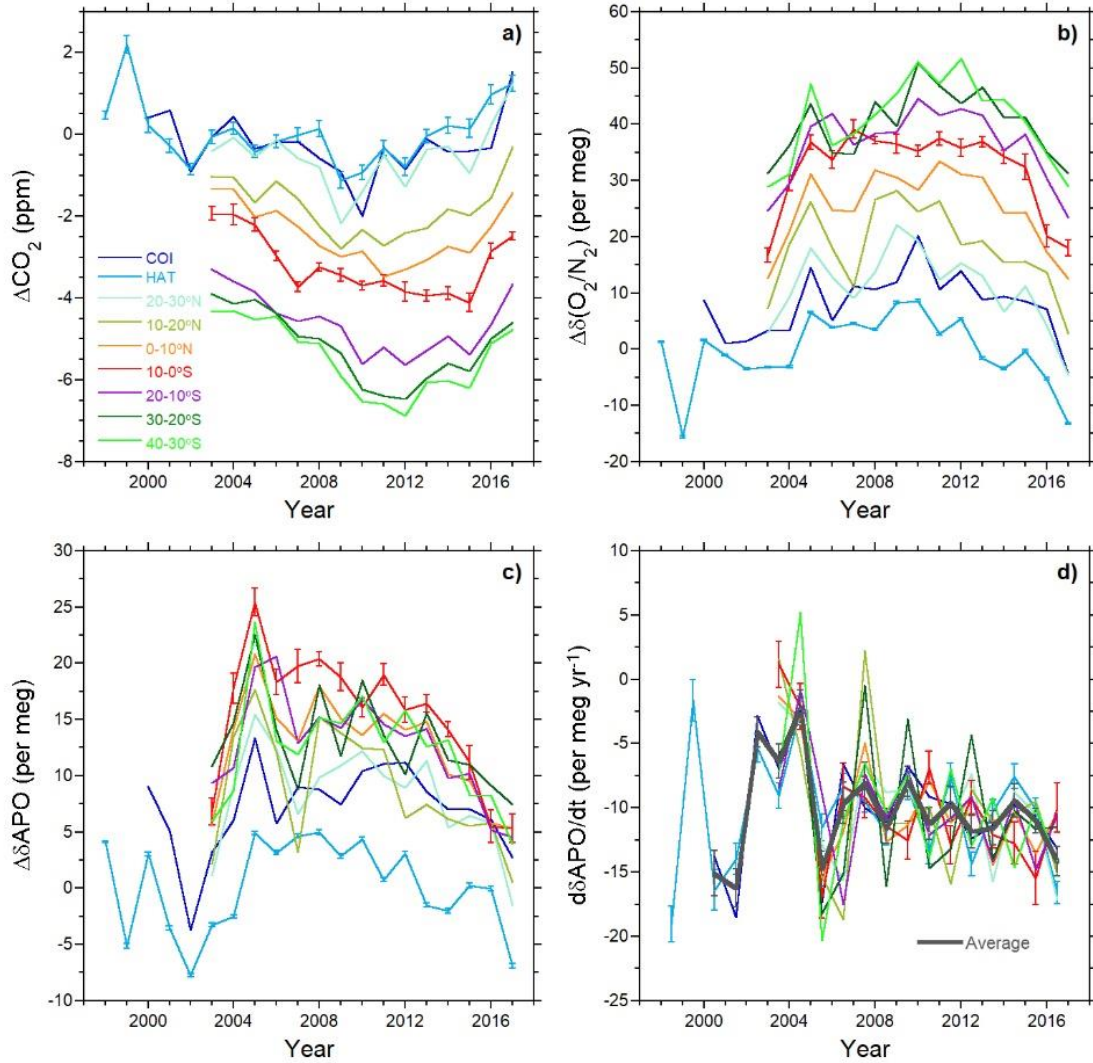


Fig. 6. Temporal variations of the (a) annual mean CO_2 , (b) annual mean O_2/N_2 , (c) annual mean APO, and (d) annual changing rate of APO based on the flask samples collected from HAT, COI, and cargo ships in the western Pacific (40°S - 30°N). The differences in the annual means from the linear trends fitted to the data at HAT are depicted in the figures to emphasize the interannual variations. Vertical bars for the plots of HAT and $10\text{-}0^\circ\text{S}$ bin correspond to the standard errors of the annual means.

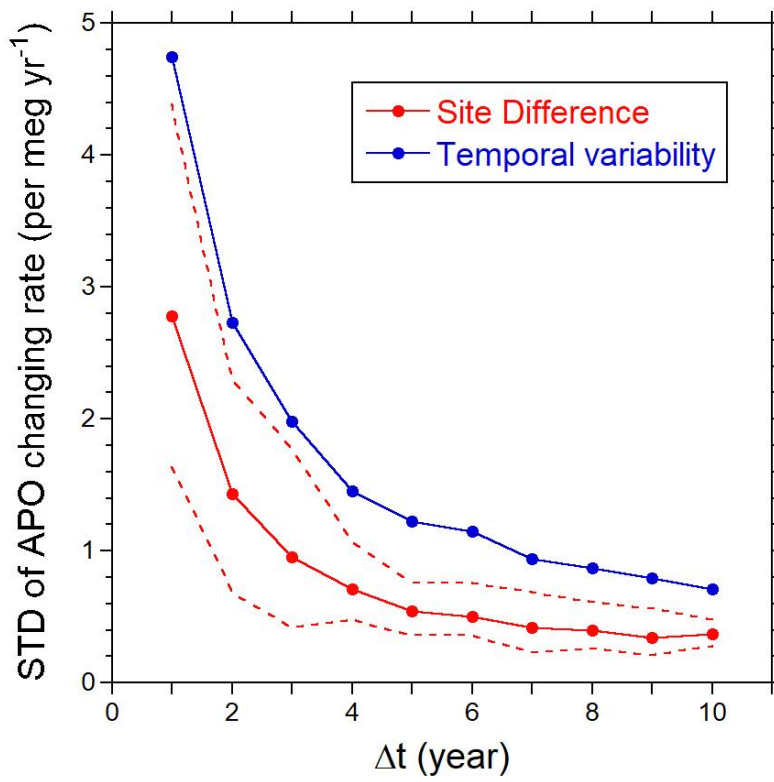


Fig. 7. Relationship between the standard deviation of the APO changing rate and the time interval to calculate the changing ratio. The red circles represent the averages of the standard deviations of the APO changing ratios from the different sites for the same year. The broken lines represent the minimum and maximum of the standard deviations. The blue circles represent the temporal variability of the average APO changing ratio of the different sites.

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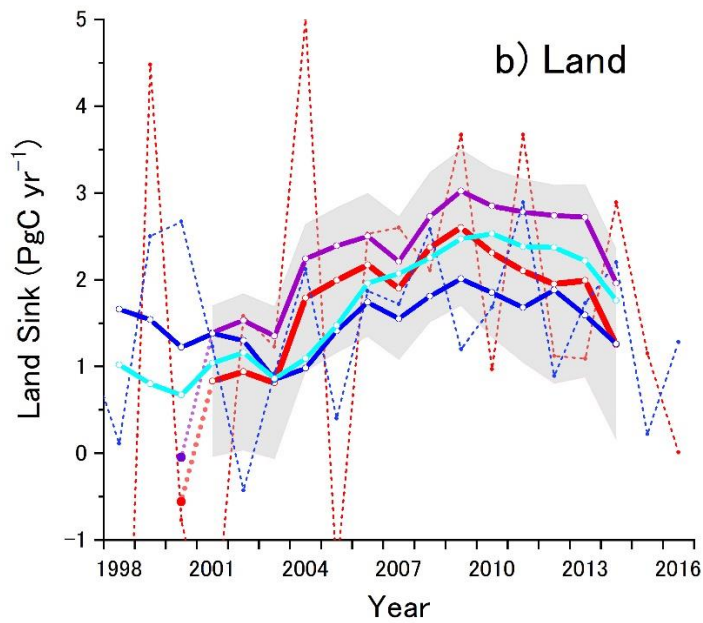
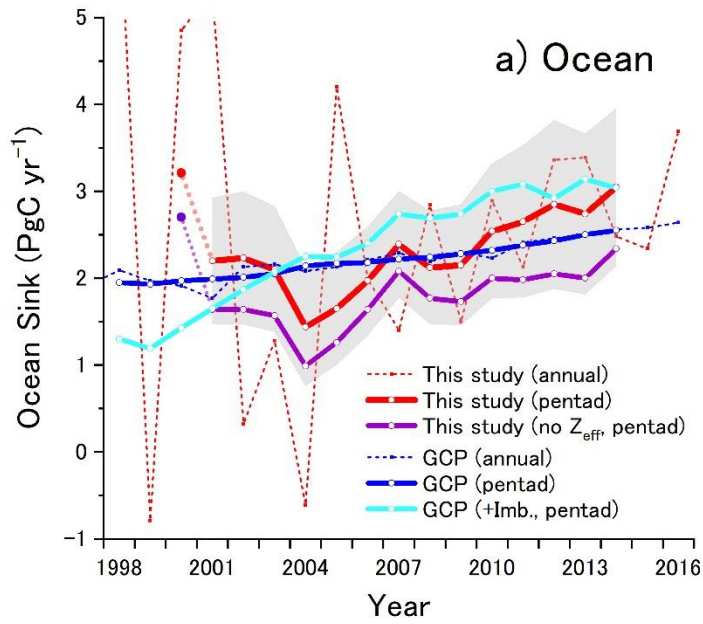


Fig. 8. Temporal variations in (a) ocean and (b) land biospheric sinks estimated from APO variations of this study (red) and process-based models of GCP (blue). The thin broken lines represent the annual sinks and the thick lines represent the pentad sinks. The purple lines represent the pentad sinks based on APO without ocean outgassing correction (Z_{eff}) and the light blue lines represent the sinks of GCP with the imbalance sinks added. The uncertainty associated with the pentad sinks with Z_{eff} corrections are shown as shaded area.

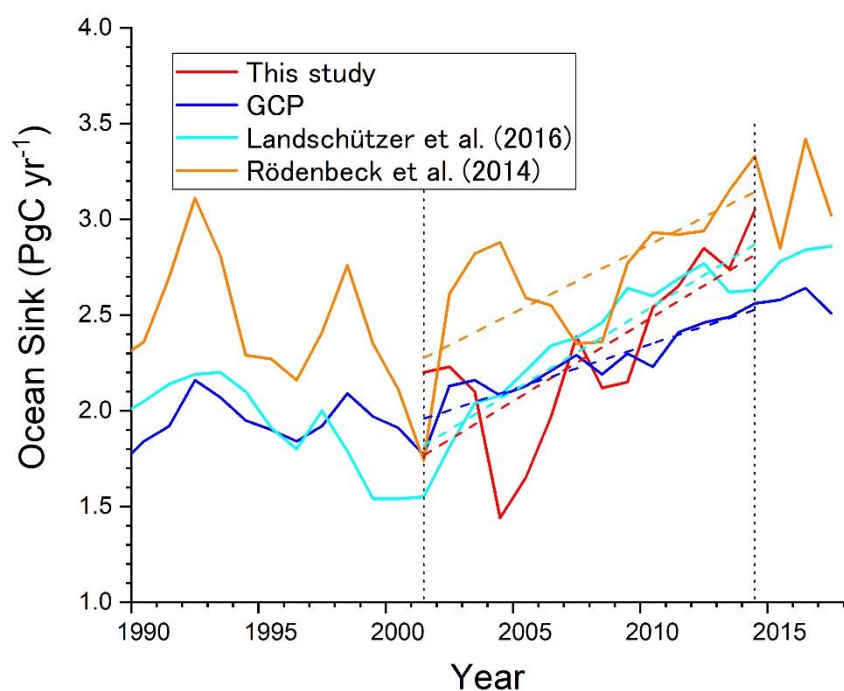


Fig. 9. Comparison of the temporal variations of the ocean sinks based on the APO data of this study (red), global ocean biogeochemistry models (GOBMs) of GCP (blue), and pCO₂ data of Landschützer et al. (2016) (light blue) and Rödenbeck et al. (2014) (orange). The broken lines represent the regression lines for the corresponding data during 2001-2014. Note that the pCO₂-based ocean sinks are adjusted for the pre-industrial ocean CO₂ emissions ($\pm 0.78 \text{ PgC yr}^{-1}$) caused by riverine CO₂ input to the ocean (Resplandy et al., 2018).