



1 **Fine particle pH and sensitivity to NH₃ and HNO₃ over summertime South Korea during**
2 **KORUS-AQ**

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27



28 **Abstract**

29 Using a new approach that constrains thermodynamic modeling of aerosol composition with
30 measured gas-to-particle partitioning of inorganic nitrate, we estimate the acidity levels for aerosol
31 sampled in the South Korean planetary boundary layer during the NASA/NIER KORUS-AQ field
32 campaign. The pH (mean $\pm 1\sigma = 2.43 \pm 0.68$) and aerosol liquid water content determined were then
33 used to determine the “chemical regime” of the inorganic fraction of particulate matter (PM)
34 sensitivity to ammonia and nitrate availability. We found that the aerosol formation is always
35 sensitive to HNO_3 levels, especially in highly polluted regions, while it is only exclusively
36 sensitive to NH_3 in some rural/remote regions. Nitrate levels are further promoted because dry
37 deposition velocity is low and allows its accumulation in the boundary layer. Because of this,
38 HNO_3 reductions achieved by NO_x controls prove to be the most effective approach for all
39 conditions examined, and that NH_3 emissions can only partially affect PM reduction for the
40 specific season and region. Despite the benefits of controlling PM formation to reduce ammonium-
41 nitrate aerosol and PM mass, changes in the acidity domain can significantly affect other processes
42 and sources of aerosol toxicity (such as e.g., solubilization of Fe, Cu and other metals) as well as
43 the deposition patterns of these trace species and reactive nitrate.

44

45 **1. Introduction**

46 Poor air quality from high concentrations of fine particulate matter over South Korea has its origin
47 in domestic emissions from vehicles, industry, and biomass burning, combined with long-range
48 transport of pollutants from mainland China (Kim et al., 2018; Nault et al., 2018). Because air
49 quality in South Korea is a mixture of factors such as regional and local emissions from both
50 anthropogenic and natural (e.g. dust) sources, as well as meteorological (e.g. wind, relative



51 humidity) and chemical interactions (e.g. photochemistry), assessing possible air pollution control
52 strategies in this region is challenging (RSSR, 2016).

53 To improve our understanding of poor air quality in South Korea, the Korean National Institute of
54 Environmental Research (NIER) and the United States National Aeronautics and Space
55 Administration (NASA), conducted a field study in South Korea from 26 April to 18 June 2016
56 (RSSR, 2016). The aim of the Korea-United States Air Quality Study (KORUS-AQ), was to
57 determine the factors that contribute to the poor regional air quality to aid the development of
58 effective air quality mitigation strategies.

59 During the KORUS-AQ campaign, secondary production of particulate matter (PM) – organic and
60 inorganic – constitute a significant fraction of PM pollution, with significant contribution from
61 local sources. Since sulfate and nitrate comprised nearly half of the mass of PM at sizes smaller
62 than 1 μm (PM_{1}) (Jordan et al., in review; Nault et al., 2018), aerosol acidity, liquid water content
63 and temperature will significantly affect aerosol properties including its mass through the gas-
64 particle partitioning of semi-volatile species (Nenes et al., 2020a). More broadly, fine aerosol
65 particle acidity in South Korea affects air quality and human health and therefore requires
66 knowledge of aerosol pH levels (e.g., Nenes et al., 2020a).

67 Given that direct measurement of atmospheric aerosol pH remains a challenge, there is large
68 uncertainty in this important parameter (Pye et al., 2020). Of all approaches used to date for
69 constraining aerosol pH, use of thermodynamic analysis of gas-aerosol observations together with
70 a model provides the most robust estimates of aerosol acidity (Hennigan et al., 2015; Song et al.,
71 2018; Pye et al., 2020). To achieve the most robust pH predictions, thermodynamic analysis
72 requires observations of all species, both gas and particle phase, that affect pH. Knowledge of gas-
73 phase ammonia and particulate ammonium is especially important, owing to its role as the



74 dominant cation in fine-mode aerosol and pH-sensitive partitioning. Accurate measurement of
75 ammonia, especially at low concentrations, is far from trivial (Zhu et al., 2015; Wang et al., 2015)
76 and subject to biases from adsorption of NH_3 in instrument inlets and NH_4^+ volatilization (Dawson
77 et al., 2014; Osada 2011; Yokelson et al., 2003; Guo et al., 2018a). Although measurements of
78 aerosol NH_4^+ are common, gas-phase measurements of NH_3 are often missing in field studies –
79 including for KORUS-AQ. There is a need to accurately infer aerosol pH in the absence of gas-
80 phase NH_3 data.

81 This study aims to accurately determine fine particle pH for aerosol in South Korea, and use this
82 inferred pH to understand the sensitivity of PM mass to the availability of ammonia and nitric acid
83 (which are two major aerosol precursors). Focusing only on the inorganic fraction of the aerosol,
84 we develop a computationally rigorous approach to estimate aerosol pH and particle liquid water
85 despite missing gas-phase ammonia data and apply it to the KORUS-AQ dataset. We then utilize
86 these data to assess optimal strategies to control inorganic aerosol mass through reduction of either
87 available nitric acid or ammonia, based on the new approach of Nenes et al. (2020a).

88

89 **2. Methods**

90 **2.1 Instrumentation**

91 *Non-refractory PM_1 composition*

92 The CU-Boulder highly-customized Aerodyne high-resolution time-of-flight aerosol mass
93 spectrometer (HR-ToF-AMS) measured non-refractory (NR) PM_1 composition, including
94 ammonium, nitrate, sulfate, chloride, and organic aerosol. The basic concept, operation, and
95 aircraft deployment of the AMS has been described elsewhere (DeCarlo et al., 2006, 2008; Dunlea
96 et al., 2009; Kimmel et al., 2011) and the deployment for KORUS-AQ is discussed in detail by



97 Nault et al. (2018). Observations are reported in units of $\mu\text{g sm}^{-3}$ (under standard conditions of
98 $T=273.15\text{ K}$, $p=1013\text{ hPa}$). For the following analysis, the data were converted to concentrations
99 at ambient conditions for the thermodynamic calculations.

100 The AMS also separately measures the contribution of amines, organonitrates and organosulfates
101 (Fry et al., 2019; Farmer et al., 2010; Chen et al., 2019). During the campaign, organic nitrates
102 comprised roughly 8% of the AMS particulate NO_3 signal and were only an important contribution
103 to the signal when NO_3 was below $0.50\ \mu\text{g sm}^{-3}$ ($0.45\ \mu\text{g m}^{-3}$) (Nault et al., 2018). Average
104 concentration of nitrate during the campaign was $8.09 \pm 6.16\ \mu\text{g m}^{-3}$, so we consider AMS nitrate
105 to be approximately equal to inorganic aerosol NO_3^- . Accuracy (2σ) for AMS detection of
106 inorganic species is estimated to be 35% (Bahreini et al, 2009).

107 *Gas-phase HCl and HNO₃*

108 HCl and HNO₃ measurements were made using the Georgia Institute of Technology Chemical
109 Ionization Mass Spectrometer (GT CIMS) and the California Institute of Technology CIMS (CIT
110 CIMS), respectively. The GT CIMS technique utilizes a low-pressure ion source/reactor and SF_6^-
111 reagent ion chemistry to detect HCl (Huey et al., 2004; Slusher et al., 2004, Kim et al., 2007). The
112 CIT CIMS utilizes CF_3O^- chemistry to detect HNO₃ by way of fluoride transfer (Huey et al., 1996;
113 Amelynck et al., 2000; Crouse et al., 2006). Both methods are configured to allow for ion
114 chemistry in a heated Teflon-coated flow tube at low pressure to selectively cluster the reagent
115 ions with HNO₃ (HCl) in ambient air (Huey et al., 2004). The ions from the flow tube enter a mass
116 spectrometer, where they are mass-analyzed and quantified. Gas-phase concentrations were
117 reported in ppbv (parts-per-billion, by volume), and were subsequently converted to $\mu\text{g m}^{-3}$ at
118 ambient conditions prior to thermodynamic analysis. The estimated CIMS measurement
119 uncertainty is 40%.



120

121 *Non-volatile aerosol cations and meteorological variables*

122 Non-volatile cations, such as Na^+ , K^+ , Ca^{2+} , and Mg^{2+} , were measured using the soluble acidic
123 gases and aerosols (SAGA) instrumentation. Bulk aerosol (nominally $< 4.1 \mu\text{m}$ aerodynamic
124 diameter) are collected onto filters to quantify soluble ions (Dibb et al., 2003). Detection limits are
125 5 pptv for Mg^{2+} and better than 20 pptv for Na^+ , K^+ , and Ca^{2+}
126 (https://cloud1.arc.nasa.gov/docs/intex-b/SAGA_Dibb.pdf). Filter sampling times were
127 approximately 5 min or less. Units of NVCs were reported in $\mu\text{g m}^{-3}$ under ambient conditions
128 when used for thermodynamic calculations. Water vapor content was measured using a diode laser
129 hygrometer (Diskin et al., 2002) and then converted to relative humidity based on the measurement
130 of temperature measured onboard.

131

132 **2.2 Thermodynamic analysis of observations and PM sensitivity to aerosol precursors**

133 ISORROPIA II (Fountoukis and Nenes, 2007; <http://isorro피아.epfl.ch>), was used to calculate the
134 equilibrium phase state and composition of inorganic aerosol systems containing NH_4^+ , SO_4^{2-} ,
135 NO_3^- , Cl^- , Na^+ , Ca^{2+} , K^+ , Mg^{2+} , aerosol liquid water content, and partitioning of semi-volatile
136 species (e.g. particle nitrate, NO_3^- , and gas phase nitric acid, HNO_3). ISORROPIA was run in
137 forward mode, which results in more robust and accurate predictions of pH owing to the sensitivity
138 of reverse mode calculations to measurement uncertainty (Hennigan et al., 2015; Song et al., 2018).
139 The model assumes that fine particles and their corresponding volatile counterparts are in
140 thermodynamic equilibrium, which is a good assumption for submicron particles that are not
141 kinetically limited by size (i.e., in the fine mode) (Dassios and Pandis 1999; Cruz and Pandis 2000;
142 Fountoukis et al., 2009) and slow diffusivity in the particle phase. For the selected KORUS-AQ



143 dataset (<1 km ASL, 45% < RH < 95%, average RH: 62±12%; average temperature: 20.2±2.5 C)
144 the aerosol tends to be in a liquid metastable state (i.e. no precipitation of salts under supersaturated
145 conditions; Fountoukis and Nenes 2007; Day and Malm 2001; Seinfeld and Pandis, 2016).
146 Efflorescence typically occurs for RH between 10-30%, however, such humidity points were not
147 considered in our analysis.

148 While ISORROPIA II can handle systems containing non-volatile cations (NVCs), such as Na⁺,
149 K⁺, Mg²⁺, and Ca²⁺, they were not considered in the analysis owing to their very low concentration
150 (Figure 1) and minor impact on aerosol pH (Guo et al., 2018a); as well as, the fact that much of
151 the NVCs were in the supermicron mode (Saide et al., 2019; Heim et al., 2020). A sensitivity
152 calculation, using SAGA measurements as an upper limit of submicron NVCs pH confirms this
153 (Table 1).

154 ISORROPIA II is used to calculate the pH of aerosols. pH is defined as:

$$155 \quad pH = -\log_{10}(\gamma_{H^+} H_{aq}^+) = -\log_{10}\left(\frac{1000 \gamma_{H^+} H_{air}^+}{W_i + W_o}\right) \approx -\log_{10}\left(\frac{1000 \gamma_{H^+} H_{air}^+}{W_i}\right) \quad (1)$$

156 where γ_{H^+} is the hydronium ion activity coefficient – here assumed to be unity, H_{aq}^+ (mol L⁻¹)
157 hydronium ion concentration in aerosol liquid water, H_{air}^+ (μg m⁻³) hydronium ion concentration
158 per volume of air, and W_i (μg m⁻³) and W_o (μg m⁻³) are the particle liquid water concentrations
159 associated with inorganic and organic species, respectively. Although W_o can be estimated, e.g.,
160 through the hygroscopicity parameter (e.g., Guo et al., 2015), its effect on pH, together with
161 organic effects on the activity coefficient, are secondary - introducing somewhere between a 0.15
162 and 0.30 pH units change (Guo et al., 2015; Song et al., 2018; Vasilakos et al., 2018; Battaglia Jr.
163 et al., 2019). Note Equation 1 is consistent with the “pH_F” definition of Pye et al. (2020).



164 As applied here, the thermodynamic analysis provides pH consistent with “bulk” pH, which
165 assumes that particles are internally mixed. This assumption tends to provide good estimates of
166 pH for submicron particles, given that the equilibrium assumption is largely satisfied and semi-
167 volatile partitioning of gases such as NH_3 and HNO_3 is well captured (Guo et al., 2018b, Pye et
168 al., 2020). Internal mixing is achieved in a few hours in polluted areas due to rapid secondary
169 aerosol production (Wang et al., 2010; Zhu et al., 2016; Riemer et al., 2019). Both the ground
170 AMS (Kim et al., 2018) and the aircraft AMS (Nault et al., 2018) saw no evidence of external
171 mixing based on the mass size distributions of the individual components, so the internal mixing
172 assumption is appropriate.

173 The results of the thermodynamic analysis are then combined with the conceptual framework of
174 Nenes et al. (2020a) to identify the chemical domains of PM mass sensitivity to HNO_3 and NH_3
175 availability for this observational dataset. These sensitivity domains are characterized as *i*)
176 primarily NH_3 -sensitive, *ii*) primarily HNO_3 -sensitive, *iii*) combined NH_3 and HNO_3 sensitive,
177 and, *iv*) HNO_3 / NH_3 insensitive. This thermodynamically consistent approach enables us to
178 directly determine how change in gas phase species will elicit a change in PM mass. As it pertains
179 to the KORUS-AQ campaign, we adopt the Nenes et al. (2020a) threshold value for nitrate and
180 ammonia partitioning that separates each regime at the 10% level. For example, for nitrate this
181 means that when more than 10% of the total nitrate is in the particle phase, we would expect PM
182 responses to NO_3^- precursors to be significant.

183

184 **2.4 Nitrate partitioning constrained pH (NPC-pH)**

185 When the thermodynamic model is run in forward mode, input of semi-volatile species (i.e., those
186 that can exist in particle- or gas-phase) are assumed to be total gas and particle concentrations. As



187 in this study, when NH_3 data is unavailable, the total NH_3 will be lower than the true value, as
188 some fraction of the total NH_3 is actually in the gas phase. In these cases, particle acidity is likely
189 to be overestimated. Past studies have proposed an iterative approach to calculate pH and gas phase
190 ammonia data (e.g., Guo et al., 2016) that involves running ISORROPIA (without gas-phase
191 ammonia) to retrieve the predicted equilibrium gas phase ammonia concentration from the model
192 output. This equilibrium NH_3 along with the measured NH_4^+ is then used as total ammonia input
193 for the next ISORROPIA iteration (which eventually adds some mass to the system), holding all
194 other input values constant. After each such iteration, convergence is checked by examining
195 whether the values of gas-phase NH_3 agree to within a predefined criterion. This method, however,
196 is unconditionally unstable, i.e., the method does not converge when an increasingly strict criterion
197 is used (see Figure 2 for a schematic). To demonstrate this, Figure 3 presents how pH and gas-
198 phase NH_3 change with iteration; a stable algorithm would eventually converge to values that do
199 not change with iteration; in reality NH_3 and pH increase monotonically without a bound. This
200 instability is not unique to ISORROPIA, but inherent to the algorithm and should apply to any
201 thermodynamic model, as the amount of total ammonia with every iteration increases – without
202 any bound. Applying a less strict upper bound would result in an arbitrary estimate of total
203 ammonia, which requires some prior knowledge on NH_3 levels to provide realistic values of pH
204 (which is what Guo et al., 2016 adopted).

205 As an alternative to the Guo et al. (2016) algorithm, we propose a method to infer pH calculations
206 from thermodynamic analysis of observations when NH_3 data are lacking. This approach involves
207 using an algorithm that is tied to the observed nitrate/nitric acid partitioning (hereafter referred to
208 as “nitrate partitioning constrained pH”, NPC-pH) to infer gas-phase NH_3 concentrations required
209 for plausible pH predictions. Figure 4 summarizes the methodology behind this approach. The



210 observed NH_4^+ , total nitrate, and nitrate partitioning fraction, $\varepsilon(\text{NO}_3)$ —defined as the fraction of
211 total nitrate (gas + aerosol) present in the aerosol phase—are used as input to an iterative algorithm
212 that determines the value of gas-phase NH_3 that, together with the observed value of NH_4^+ ,
213 temperature, and relative humidity reproduces the observed $\varepsilon(\text{NO}_3)$ to some predetermined level
214 of accuracy. The thermodynamic calculations required for each step in the iterative procedure are
215 done with ISORROPIA-II. Upon convergence (here, to within 10^{-6} between iterations), particle
216 liquid water, H^+ _{air} concentration, and particle pH are obtained and used for further analysis. In the
217 following, we assess the predicted pH with NPC-pH, which together with aerosol liquid water
218 content is needed to determine an effective PM control strategy.

219 Nitric acid can condense onto fine mode and coarse particles containing non-volatile cations (from
220 seasalt or dust). Because of this, nitrate may over time volatilize and recondense onto coarse mode
221 cations forming nonvolatile species (in the form of $\text{Ca}(\text{NO}_3)_2$, NaNO_3 and other salts). This
222 disequilibrium will lead to a gas phase HNO_3 concentration which is lower than the equilibrium
223 value expected from the PM_{10} composition and higher than the corresponding value based on the
224 coarse mode composition. Given that the equilibration timescale of submicron aerosol is much
225 smaller than for coarse mode particles, the degree of disequilibrium between the PM_{10} nitrate and
226 gas-phase HNO_3 is much smaller than that for coarse mode nitrate. We therefore assume that the
227 PM_{10} semi-volatile inorganic species (NO_3^- , NH_4^+) are in equilibrium with their gas phase
228 components (HNO_3 , NH_3). Analysis carried out by Guo et al. (2017) supports this approach.

229

230 **2.5 Data selection and analysis**

231 Data points were filtered for conditions where ambient relative humidity (RH) fell within the range
232 of 45-95%, and flight altitude below 1 km which is often within the boundary layer and most



233 relevant from an air quality perspective (the sampling RH was often significantly lower, but
234 exposure to these low levels of humidity is too short to have a significant impact on semi-volatile
235 nitrate and ammonium; Shingler et al, 2016; Guo et al., 2017). The RH range was chosen to ensure
236 robust estimations from the thermodynamic model. Data for $RH > 95\%$ were excluded owing to
237 the exponential growth in particle liquid water with RH, which leads to high W_i and subsequently
238 large pH uncertainty owing to propagation of RH uncertainties (Guo et al., 2015). Guo et al. (2016)
239 suggests that below RH of 40%, pH estimations are subject to considerable uncertainty owing to
240 the low aerosol liquid water and other uncertainties, which limits the ability to capture the observed
241 partitioning of nitrate and other species. From this filtering process, a total of 11 of the total flights
242 were analyzed and are summarized in Table S1.

243

244 **3. Results**

245 **3.1 Sensitivity studies to evaluate the new NPC-pH algorithm**

246 The NPC-pH algorithm for predicting pH without NH_3 data was assessed based on a synthetic
247 dataset for which aerosol is in perfect thermodynamic equilibrium. To this dataset, random
248 variability is added to the concentration of semi-volatile species (within a predefined but relatively
249 wide range). The NPC-pH algorithm was applied to the original and noisy synthetic data sets, to
250 predict pH. The aerosol acidity obtained from NPC-pH was then compared to the pH of the initial
251 data where the aerosol was in equilibrium with all species. The robustness of the acidity to noise
252 level is important, especially given that aerosol semivolatile species could be in disequilibrium.

253 The synthetic data was constructed of aerosol precursor values relevant for KORUS-AQ
254 conditions: total NO_3 concentration ranged from 0.2 to $110 \mu g m^{-3}$ allowing for $\epsilon(NO_3)$ to range



255 from 0-0.95. Total SO_4 and total NH_4^+ concentrations ranged from $0.1\text{-}10\ \mu\text{g m}^{-3}$ and $0.2\text{-}110\ \mu\text{g}$
256 m^{-3} , respectively. $\epsilon(\text{NH}_4)$ ranged from 0-1. The total Cl concentration was kept constant at 0.50
257 $\mu\text{g m}^{-3}$ and NVC concentrations were set to zero. Temperature was kept constant at 298K, and the
258 RH ranged from 45-95%. Comparing the equilibrium partitioning retrieved from the synthetic data
259 (without added noise) to that generated from the algorithm resulted in nearly perfect agreement
260 between the two quantities, when $\epsilon(\text{NO}_3)$ was greater than about 40% (Figure 5). Higher sensitivity
261 of low $\epsilon(\text{NO}_3)$ values to pH and gas-phase NH_3 results in more scatter in $\epsilon(\text{NO}_3)$ generated from
262 the algorithm for $\epsilon(\text{NO}_3) < 0.4$. Average pH for the whole synthetic dataset for equilibrium and
263 NPC-pH method are 2.14 ± 1.33 and 2.26 ± 1.25 , respectively while the average LWC for
264 equilibrium and NPC-pH method are $42.50 \pm 137.85\ \mu\text{g m}^{-3}$ and $42.25 \pm 137.93\ \mu\text{g m}^{-3}$,
265 respectively. The large standard deviation is a result of a wide range of conditions tested during
266 the analysis.

267 To ascertain how uncertainties in nitrate partitioning (i.e., deviations from thermodynamic
268 equilibrium) would impact pH inferences from NPC-pH, random noise at the 1-50% level is added
269 to the original $\epsilon(\text{NO}_3)$ values from the synthetic dataset. Reapplication of NPC-pH to the noisy
270 datasets then quantify the effect of this noise to the inferred pH. The method used to add noise is
271 shown in equations 2 and 3

$$272 \quad X = 2\Phi\left(-\frac{1}{2} + Rnd\right)\epsilon(\text{NO}_3) \quad (2)$$

$$273 \quad \epsilon(\text{NO}_3)_{Rnd} = \epsilon(\text{NO}_3) + X \quad (3)$$

274 where $\epsilon(\text{NO}_3)_{Rnd}$ is $\epsilon(\text{NO}_3)$ with added random noise X that is symmetrical about zero and scaled
275 to the magnitude of $\epsilon(\text{NO}_3)$; Φ is the maximum fractional noise level (ranging from 0.01-0.5 to
276 express a noise level of 1-50%) and Rnd is a random number between 0 and 1, generated by the
277 "rand()" pseudorandom number generator available in the Matlab® environment.



278 Results from this sensitivity analysis reveal that a 50% relative error in $\varepsilon(\text{NO}_3)$ resulted in an
279 average absolute error in pH of 0.28 ± 0.45 units, and explained as follows. Figure 6 presents
280 $\varepsilon(\text{NO}_3)$, $\varepsilon(\text{NH}_4)$ and pH for $T = 288$ K, and with average ISORROPIA-predicted liquid water
281 content ($13.78 \pm 10.52 \mu\text{g m}^{-3}$) and activity coefficients (0.125 and 1.794 for γ_{NO_3} – γ_{H^+} and
282 $\gamma_{\text{H}^+}/\gamma_{\text{NH}_4^+}$, respectively) derived from the KORUS-AQ flight data analysis (Table 1). In the blue
283 region, where $\varepsilon(\text{NO}_3)$ approaches 0 or 1, we observe that a small uncertainty in $\varepsilon(\text{NO}_3)$ can result
284 in significant changes to pH. In the pink region, however, we would expect that even large changes
285 in $\varepsilon(\text{NO}_3)$ would only result in minor changes to pH. Evaluating $0.2 < \varepsilon(\text{NO}_3) < 0.8$ from the
286 synthetic data resulted in an average absolute error in pH of 0.21 ± 0.15 units. Therefore, pH
287 predictions using this method are reasonably accurate, especially when considering the inherent
288 uncertainty of pH inferences using thermodynamic models (Pye et al. 2020) and the low sensitivity
289 of pH to NH_3 inferences error (about a factor of 10 error in NH_3 provides a pH error of 1 unit,
290 regardless of acidity regime; Guo et al., 2017; Song et al., 2019).

291

292 3.2 Aerosol acidity during KORUS-AQ

293 KORUS-AQ integrated aircraft and ground-based measurements, and satellite observations. The
294 campaign was conducted over South Korea (33 and 39 °N, 124 and 130E °E) and the Yellow Sea
295 during the months of May and June in 2016. Flight tracks for the DC-8 during KORUS-AQ are
296 shown in Figure 7, colored by measured concentrations of NO_3^- . Most flights focused on the
297 western region (35 and 38 °N, 126 and 127°E) of South Korea extending into the Yellow Sea.
298 Major emitters of NO_x are the Taean coal power plant (36.90 °N, 126.23 °E), Dangjin power plant
299 (37.06°N, 126.51°E), and Yeongheung power station (37.24°N, 126.44°E) – which are all within
300 the region of observed high nitrate concentrations (Hong et al. 2019; Kafle et al. 2017; Kim et al.



301 2013). Increased nitrate levels observed in Table 1 for Asia are most likely the result of emissions
302 from vehicular traffic and power plants, and active photochemistry (Nault et al., 2018) and
303 nighttime nitrate formation from N_2O_5 hydrolysis (Jordan et al., 2019). More details on KORUS-
304 AQ flights can be found in Nault et al. (2018).

305 Measurements of HNO_3 and NO_3^- indicate that $\epsilon(\text{NO}_3)$ most of the time (approximately 92%;
306 Figure S3) falls between 0.1 and 0.9, indicating that aerosol nitrate levels will be sensitive to pH
307 (i.e., reside in the “sensitivity window” of Vasilakos et al., 2018; Figure 6). Of the remaining 8%,
308 approximately 6% of these partitioning values fall below 0.1, while 2% of all nitrate partitioning
309 measurements were greater than 0.9, where nitrate partitioning is expected to be less sensitive to
310 pH changes. Given this and the sensitivity analysis of Section 3.1, the NPC-pH algorithm can
311 provide robust estimates of aerosol pH for the majority of the KORUS-AQ data. We find that the
312 average pH for all flights was 2.43 ± 0.68 (Table 1), while the average acidity between flights varied
313 between 1.74 and 2.90 (Table S1). These pH levels are similar to what has been reported by others
314 in the region for summertime conditions; Spring in Beijing, China (pH range= 1.8-4.3) (Ding et
315 al., 2019; Tan et al., 2018; Wang et al., 2019; Pye et al. 2020).

316 The inorganic fraction of the KORUS-AQ aerosol is dominated by sulfate, ammonium, and nitrate
317 (Figure 1). Since as already mentioned in the methods the relative mass concentrations of NVCs
318 to other major inorganic ion components are low, we assume the formation of nonvolatile salts
319 (e.g. $\text{Ca}(\text{NO}_3)_2$ and $\text{Na}(\text{NO}_3)$) negligibly impact the pH derived from nitrate partitioning. This
320 assumption is valid as the pH calculated using SAGA measurements of NVCs is within 1% of that
321 predicted, consistent with the impact suggested by Guo et al. (2018a). We also do not consider the
322 contribution of organic nitrates to the total amount of nitrates (which constitute less than 10% of



323 the total nitrate, hence have a minimal impact on liquid water content and pH, based on the analysis
324 in Section 3.1).

325 Particle pH is affected by several coupled variables such as particle nitrate and nitrate partitioning
326 levels. Higher ambient particle pH is often associated with higher concentrations of particle nitrate
327 (Guo et al., 2018b). This occurs when nitrate aerosol (usually in the form of ammonium nitrate)
328 dominates the aerosol liquid water content; in such situations, the aerosol can contain considerable
329 amounts of aerosol water but maintain small amounts of H^+ in solution – as it tends to combine
330 with NO_3^- to form volatile HNO_3 . In contrast, when aerosol liquid water is controlled by
331 hygroscopic sulfates, significant amounts of H^+ can exist in solution (e.g., from the semi-volatile
332 partitioning of NH_4^+ to form NH_3) as its tendency to associate with HSO_4^- , SO_4^{2-} is relatively weak
333 and sulfates are involatile. This general effect of higher (lower) pH associated with higher (lower)
334 NO_3 has been seen in a number of other field campaigns (see comparisons in Table 1). In Cabauw,
335 Netherlands during the summer, Guo et al. (2018b) reported an $\epsilon(NO_3)$ of 88%, with a
336 corresponding pH of 3.3 ± 0.5 . Both quantities are higher than the values found for the KORUS-
337 AQ campaign. The pH of aerosol in Beijing, China in the summer was found to be 3.9 ± 1.3 for
338 nitrate levels that are higher than those measured for the South Korean data (Table 1).

339

340 **3.3 Acidity and PM sensitivity regimes to NH_3 and HNO_3 during KORUS-AQ**

341 Nenes et al. (2020a) developed a framework that allows PM sensitivity to NH_3 and HNO_3
342 availability to be determined from aerosol acidity and liquid water content. This framework
343 directly determines effective PM reduction policies – which is important given recent work
344 identifying NH_3 reductions over other policies (e.g., NO_x and SO_x reductions) (Pozzer et al., 2017;



345 Xu et al., 2019) as the most effective for PM reductions, and the dominance of PM in the region
346 by ammonium and nitrate (22.2 and 36.6%, respectively; Figure 1)

347 In its simplest form, the Nenes et al. (2020a) framework is expressed in terms of a “policy map”
348 (Figure 8) characterized by four distinct regimes: one, shaded pink, where $\epsilon(\text{NO}_3)$ is small and
349 $\epsilon(\text{NH}_4)$ is large (i.e., the majority of nitrate resides in the gaseous phase and ammonia in the particle
350 phase, defined by a relevant threshold); here PM mass responds proportionally to changes in the
351 total ammonia but tends to be insensitive to total nitrate changes. For this reason, Nenes et al.
352 (2020a) characterize PM in this regime as being “NH₃ sensitive”. The opposite is seen in the blue-
353 shaded regime, as the majority of nitrate resides in the aerosol phase and ammonia in the gas phase.
354 For this reason, Nenes et al. (2020a) characterize PM in this region as being “HNO₃ sensitive”. In
355 both acidity regimes, partitioning may not be strongly affected by pH changes, therefore
356 uncertainties in its exact value carry minor implications for PM sensitivity to available ammonia
357 and nitrate. In the purple acidity domain, however, which Vasilakos et al. (2018) terms “sensitivity
358 window”, PM tends to respond to both HNO₃ and NH₃ emissions, as an important fraction of both
359 species is in the aerosol phase (Nenes et al., 2020a). Here, rather precise knowledge of aerosol pH
360 is important – as variations to within one unit usually imply a very large change in the partitioning
361 fraction for each semi-volatile species, hence PM sensitivity. The fourth domain, colored white, is
362 characterized by low nitrate and ammonium partitioning fraction, and PM is then relatively
363 insensitive to changes in NH₃ and HNO₃ availability.

364 From the histogram of observed nitrate partitioning (Figure S3), we expect a large fraction of the
365 data to lie in the HNO₃/NH₃ sensitive domain. As distribution of nitrate partitioning is toward the
366 high end of $\epsilon(\text{NO}_3)$ for the study (average nitrate partitioning value of $58 \pm 24\%$), PM₁ reductions
367 would be sensitive to reductions in HNO₃ – hence NO_x reductions. Indeed, if the pH and aerosol



368 liquid water content for the KORUS-AQ data are plotted on the policy map (Figure 8), the data
369 reside in the domain where HNO_3 or a mix of HNO_3/NH_3 controls are generally the most effective
370 routes for reduction of inorganic PM_1 mass for the time period of the study.

371 Focusing on flights that span the nitrate $\varepsilon(\text{NO}_3)$ range with high HNO_3 availability provides
372 additional insights on effective PM controls: during flight 19 the highest levels of nitrate and
373 ammonium of the mission were measured; and in contrast during flight 15 approximately 30% of
374 its ambient nitrate partitioning values fell below 0.1 (Figure S4; other per-flight characteristics can
375 be found in table S1). High levels of nitrate availability are expected for flight 19 because of its
376 proximity to a number of power plants (and associated NO_x emissions) along the northwest coast.
377 The DC-8 aircraft also flew very close to Seoul, South Korea for this flight (see Figure S1) – an
378 area where nitrate availability from vehicular NO_x emissions are expected to be large. Plotting the
379 data from both flights on the Nenes et al. (2020a) policy map indicates where the high HNO_3
380 availability can lead to high NO_3^- levels. The results (Figure 9) suggest that data from each flight
381 reside mostly in the $\text{NH}_3\text{-HNO}_3$ sensitivity domain, which implies similar emissions control
382 strategies are effective, despite the very different aerosol characteristics in both flights.

383 The above analysis can be further expanded to consider only PM levels above a regulatory mass
384 threshold, as PM control policies are most effective when relevant for the high PM levels seen in
385 pollution events. To assess this, we plot all the available data on the policy map, but with points
386 color-coded with PM level (Figure 10). For PM levels exceeding e.g., $30 \mu\text{g m}^{-3}$, HNO_3 is always
387 an effective control strategy, while NH_3 is effective for about half the points. The relatively fewer
388 cases that fall into the NH_3 -sensitive domain are characterized by low PM levels, hence less
389 important to control. Given that dry deposition dominates the loss of boundary layer PM during
390 haze episodes, and the expected low deposition rate of nitrate when $\varepsilon(\text{NO}_3)$ is relatively large



391 (Nenes et al., 2020b) further emphasizes the need to control NO_x , as such conditions favor the
392 rapid accumulation of available HNO_3 – and buildup of the high levels of NO_3^- seen in the
393 observations (Figure 1).

394 **4. Summary and broader implications**

395 Accurate estimates of atmospheric aerosol acidity are important for understanding a number of
396 atmospheric processes sensitive to pH. Here, we present a method – called NPC-pH - for
397 estimating aerosol pH, through thermodynamic analysis of observations that lack gas-phase
398 ammonia measurements. NPC-pH is based on the observed gas-to-particle partitioning of nitrate
399 in the absence of ammonia measurements, and is shown to perform much better than a previously
400 proposed algorithm that iterated for total ammonia (using aerosol ammonium as an initial guess),
401 as the latter is shown to be unconditionally unstable. NPC-pH is also shown to provide robust pH
402 levels that are relatively insensitive to nitrate partitioning errors.

403 Applying NPC-pH to airborne observations collected from the NASA/NIER KORUS-AQ field
404 campaign in South Korea resulted in pH predictions ($\text{pH}=2.43 \pm 0.68$) that are consistent with
405 published estimates in this region and season. The pH and LWC calculated from our
406 thermodynamic analysis and the approaches of Nenes et al. (2020ab) determine the “chemical
407 regime” of PM sensitivity to ammonia and nitrate availability, and, “dry deposition velocity
408 regime” of inorganic nitrogen (which controls the lifetime, hence accumulation, of nitrate in the
409 boundary layer during haze episodes). For KORUS-AQ, we found that the aerosol formation is
410 often in the NH_3 and HNO_3 -sensitive or HNO_3 -sensitive zone, while a small fraction (4%) of the
411 points fall in NH_3 -limited region near the Yellow Sea, Jeju Island, Busan and Eastern Sea.
412 Nevertheless, when PM levels are high, the data always lies in the HNO -sensitive or HNO_3/NH_3
413 – sensitive region. Under these conditions, we conclude that HNO_3 reductions prove to be the most



414 effective for all conditions examined, and that NH_3 emissions would only partially be effective in
415 reducing PM levels – especially given that during pollution episodes, the pH and LWC levels
416 promote rapid accumulation of nitrate aerosol in the boundary layer owing to its slow dry
417 deposition (Nenes et al., 2020b). A complete in-depth analysis of the complex chemistry and
418 contributions of different sources however is required to fully assess the most effective NO_x
419 emission controls to reduce HNO_3 production. Source attribution information can also be
420 represented on the policy maps shown here to understand their role in shaping the acidity, liquid
421 water, and PM sensitivity/deposition regimes (e.g., Zang et al., in review) to further refine types
422 of policies that could be effective during pollution episodes.

423 Despite the benefits of controlling PM formation to reduce ammonium-nitrate aerosol and PM
424 mass, we must consider that the acidity domain can significantly affect other processes and sources
425 of aerosol toxicity. Fang et al. (2017) and Wong et al. (2020) showed that acid-driven dissolution
426 of transition metals (e.g., Fe, Cu) can potentiate health effects such as cardiovascular morbidity
427 and mortality through oxidative stress (Bates et al., 2015; Ghio et al., 2012). If emissions controls,
428 in an attempt to reduce PM levels also lead to reduction in pH, may unintentionally increase aerosol
429 toxicity with adverse health effects in humans. Lastly, increased aerosol acidity can impact the
430 deposition pattern of reactive nitrogen (Nenes et al., 2020b) and bioavailability of micronutrients
431 (e.g. Fe, P) with both synergistic and/or antagonistic effects on remote ecosystems (e.g.,
432 Meskhidze et al., 2003; Nenes et al., 2011; Ito et al., 2016).

433

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440

441 **Author contributions**

442 II was responsible for the thermodynamic analysis of the ambient data, contributed to the ammonia
443 estimation algorithm and wrote the initial draft of the manuscript with significant contributions by
444 AN and RW. AN is a core developer of ISORROPIA-II, developed the conceptual framework used
445 here for understanding the sensitivity of PM_{2.5} to total nitrate and ammonia, the tools to include
446 the ambient data on the acidity maps and also demonstrated the unconditional instability of the
447 original NH₃ iteration algorithm of Guo et al. (2016). AN and RW were involved in planning and
448 supervision of the work. All authors provided feedback on the analysis approach and extensively
449 contributed to the manuscript text.

450

451 **Code and Data availability**

452 The ISORROPIA-II thermodynamic equilibrium code is available at <http://isorropia.epfl.ch>.
453 KORUS-AQ data is available at <https://www-air.larc.nasa.gov/cgi-bin/ArcView/korusaq>.

454

455 **Competing interests**

456 The authors declare that they have no conflicts of interest.



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- 723



724 **Table 1:** Summary of selected KORUS and other campaign data.

Campaign	KORUS-AQ	Institute of Urban Meteorology	SEARCH	CalNex	SOAS	SENEX
Type of Measurement	Air	Ground	Ground	Ground	Air	Air
Year	2016	2017	2010	2013	2013	2013
Season	Summer	Spring, Summer, Fall, Winter	Late Summer Early Fall	Summer	Summer	Summer
Location	South Korea	Beijing, China	SE U.S.	SW U.S.	SE U.S.	SE U.S.
Average RH (%)	62±12	-	69±18	79±17	74±16	72±9
Avg. HNO₃ (µg m⁻³)	4.61±2.72	-	0.50±0.26	6.65 ± 7.03	0.36 ± 0.14	1.35 ± 0.66
Avg. NO₃ concentration (µg m⁻³)	8.09± 6.16	12.6 ± 14.2 (spring), 13.7 ± 21.0 (winter), 9.5 ± 9.5 (summer), 18.5 ± 19.5 (fall)	0.2±0.1	PM1: 3.58 ± 3.65	0.08 ± 0.08	0.28 ± 0.09
Average ε(NO₃)	58±24%	-	26±15%	39±16%	22±16%	18 ± 6 %
pH	2.43 ±0.68 (no NVCs) 2.45±0.96 (with NVCs)	4.3±1.6 (spring), 4.5±1.1 (winter), 3.9±1.3 (summer), 4.1±1.0 (fall)	2.2 ±0.6	1.9 ±0.5	0.9 ± 0.6	1.1 ± 0.4
Reference	This study	Ding et al. (2019)	Nah et al. (2018)	Guo et al. (2017)	Guo et al. (2015)	Xu et al. (2016)

725



726 **Figure Captions**

727 **Figure 1:** Average inorganic PM₁ mass composition throughout the entire study and for altitudes
728 below 1km. Average total mass is 22 $\mu\text{g m}^{-3}$.

729 **Figure 2:** Guo et al. (2016) method for constraining aerosol pH in the absence of NH₃ data.

730 **Figure 3:** Total ammonia (blue) and pH (red) as a function of iteration number using the Guo et
731 al. (2016) algorithm.

732 **Figure 4:** The nitrate partitioning constrained pH (NPC-pH) method to obtain aerosol pH from
733 nitrate partitioning observations.

734 **Figure 5:** Comparison of equilibrium nitrate partitioning retrieved from synthetic data vs. the
735 value from the NPC-pH method. Meteorological conditions in the synthetic dataset cover a RH
736 of 45-95%, Temperature 298K and SO₄, NO₃^T, NH₄^T of 0.1-10 $\mu\text{g m}^{-3}$, 0.2-110 $\mu\text{g m}^{-3}$, and 0.2-
737 110 $\mu\text{g m}^{-3}$, respectively. Non-volatile cation concentrations were set to zero, and Cl-
738 concentration was kept constant at 0.5 $\mu\text{g m}^{-3}$.

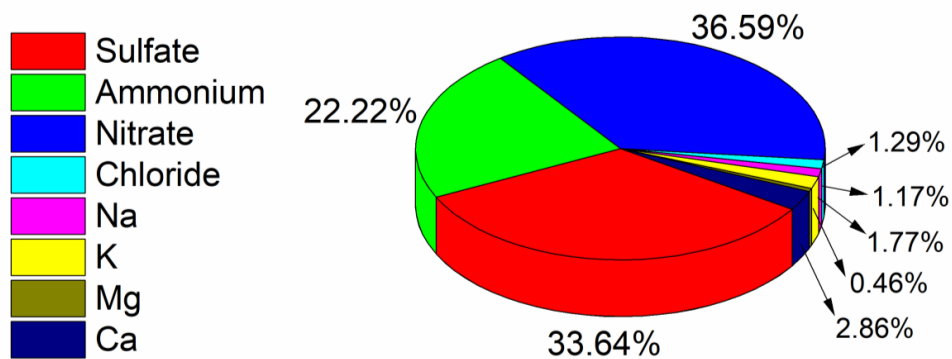
739 **Figure 6:** Particle phase fraction of total nitrate, $\epsilon(\text{NO}_3^-)$ (blue curve) and total ammonium,
740 $\epsilon(\text{NH}_4^+)$ (red curve) versus pH for a temperature of 288 K and an aerosol liquid water content of
741 10 $\mu\text{g m}^{-3}$. The pink zone is a region where PM is sensitive to both HNO₃ and NH₃. Following the
742 approach of Nenes et al. (2020a), the dotted black line represents a pre-defined threshold, below
743 which the aerosol is deemed insensitive to changes in NH₃ and/or HNO₃.

744 **Figure 7:** (a) Flight trajectories for altitudes below 1km during KORUS-AQ, color mapped by
745 NO₃⁻ levels. Highest concentrations of NO₃⁻ observed near Seoul, for which NO_x emissions are
746 dominated by the transportation and energy industries. (b) NO_x emission sources in Korea labelled
747 by industry.

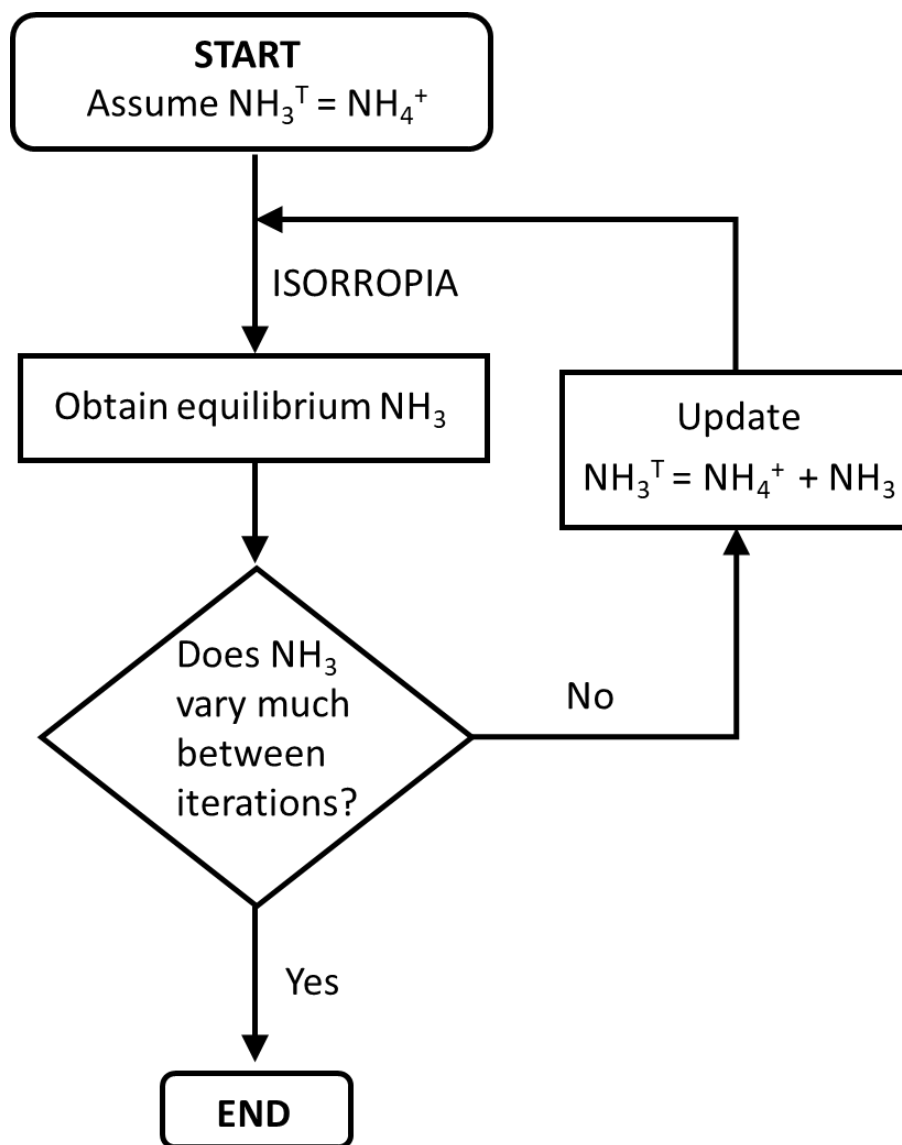
748 **Figure 8:** Chemical domains for entire KORUS-AQ study data. Larger fraction of data falls in
749 HNO₃ sensitive region as a result of moderate-high values of nitrate partitioning.

750 **Figure 9:** Chemical domains for (a) flight 15 and (b) flight 19. For Flight 15, a significant
751 number of its data points are characterized by low nitrate partitioning values (approx. 30%)^{**}.
752 Flight 19 is characterized by high levels of particle nitrate, and moderate to high levels of nitrate
753 partitioning. For these reasons, PM would be for response to NH₃ and HNO₃ for flights 15 and
754 19 respectively.

755 **Figure 10:** Chemical domains for entire KORUS-AQ study data, with symbols colored by the
756 value of PM.

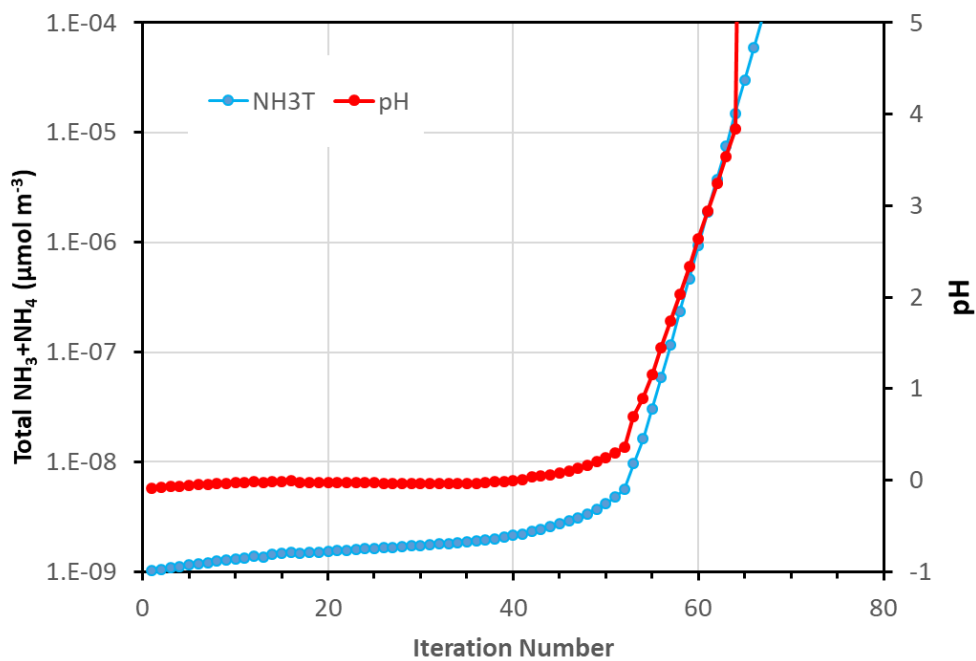


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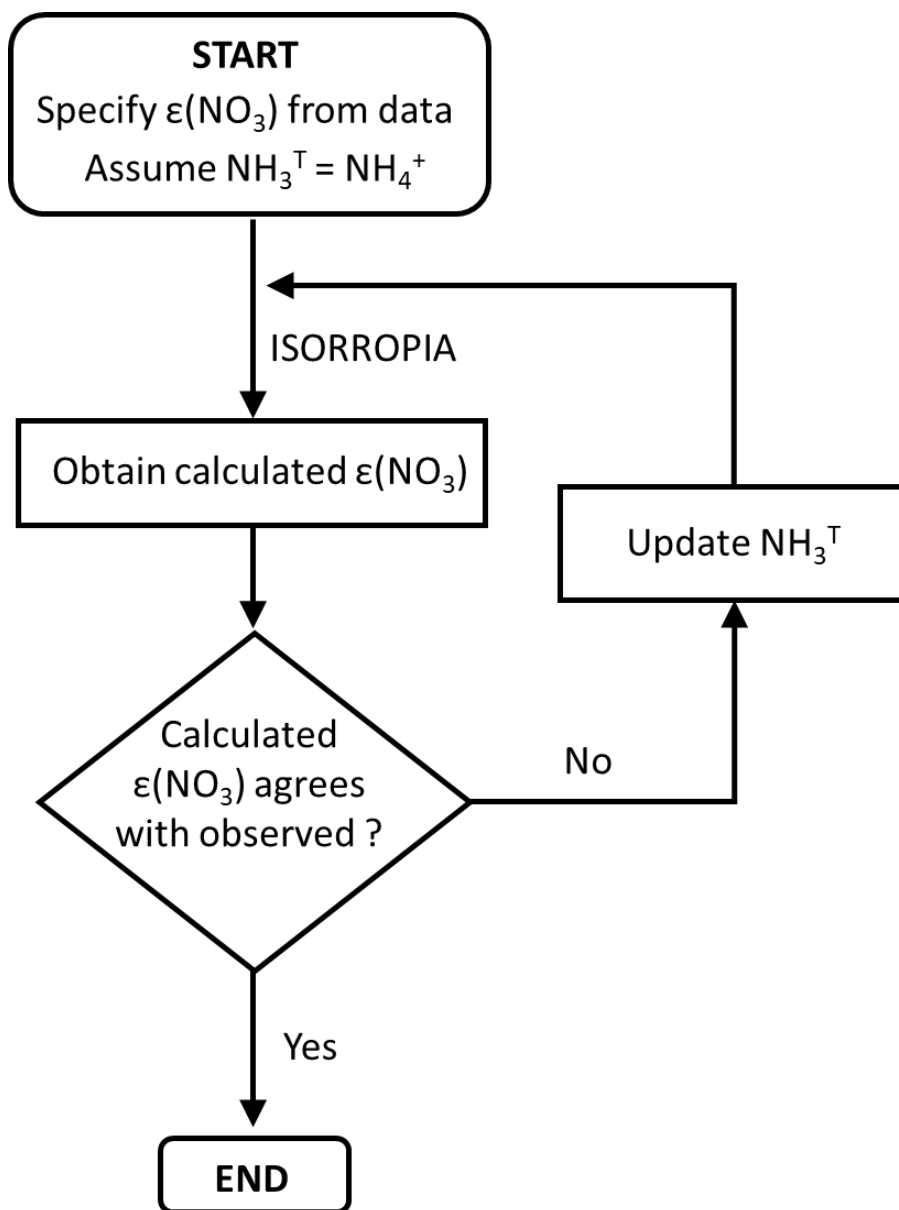


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Figure 2

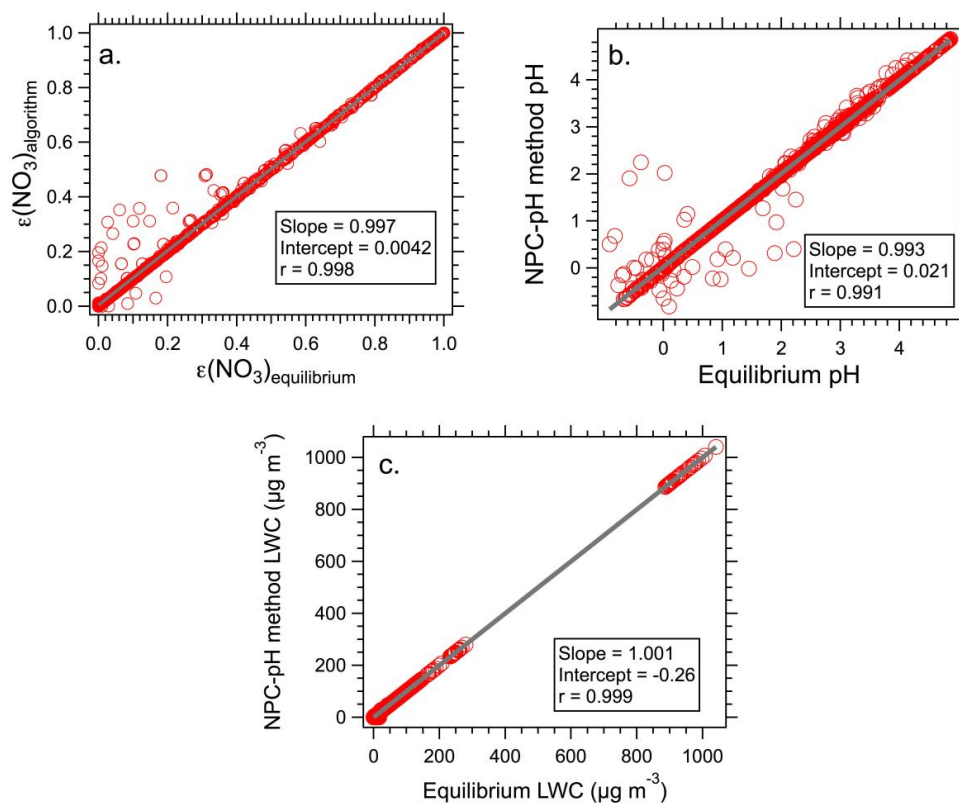


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765 **Figure 3**
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Figure 4

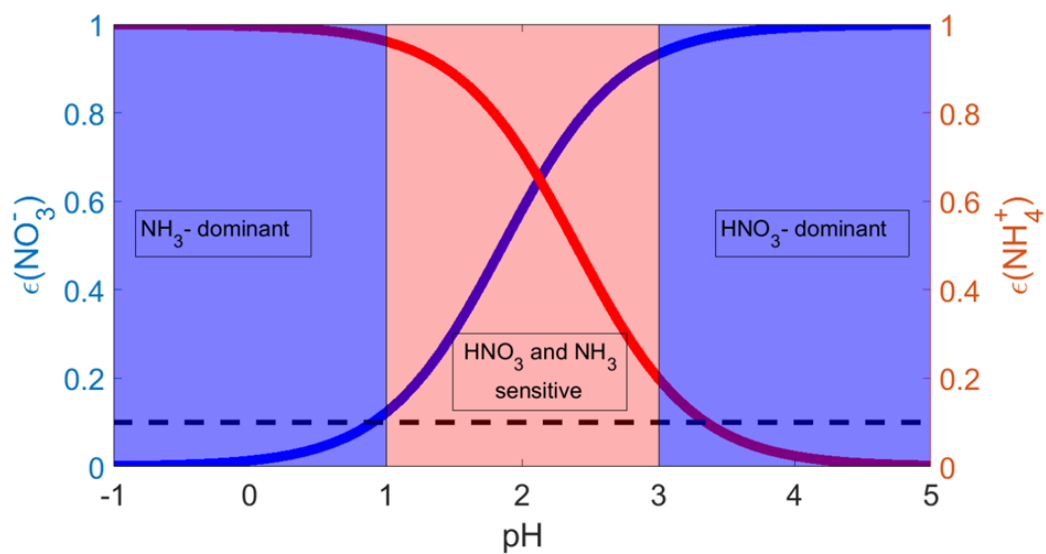


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Figure 5

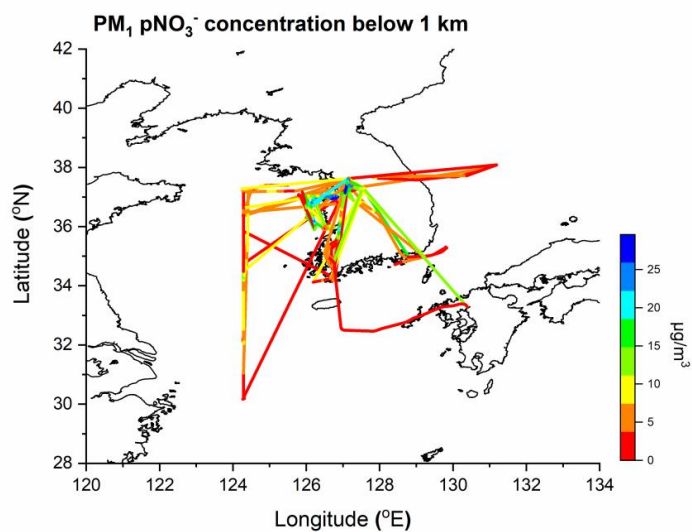


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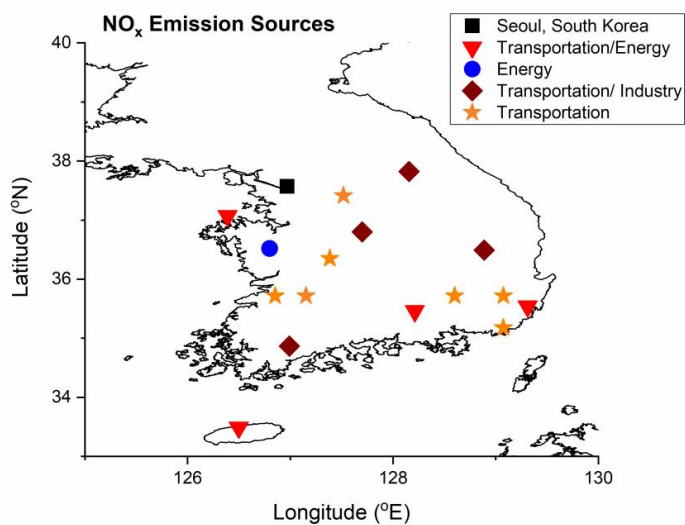
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Figure 6



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(a)



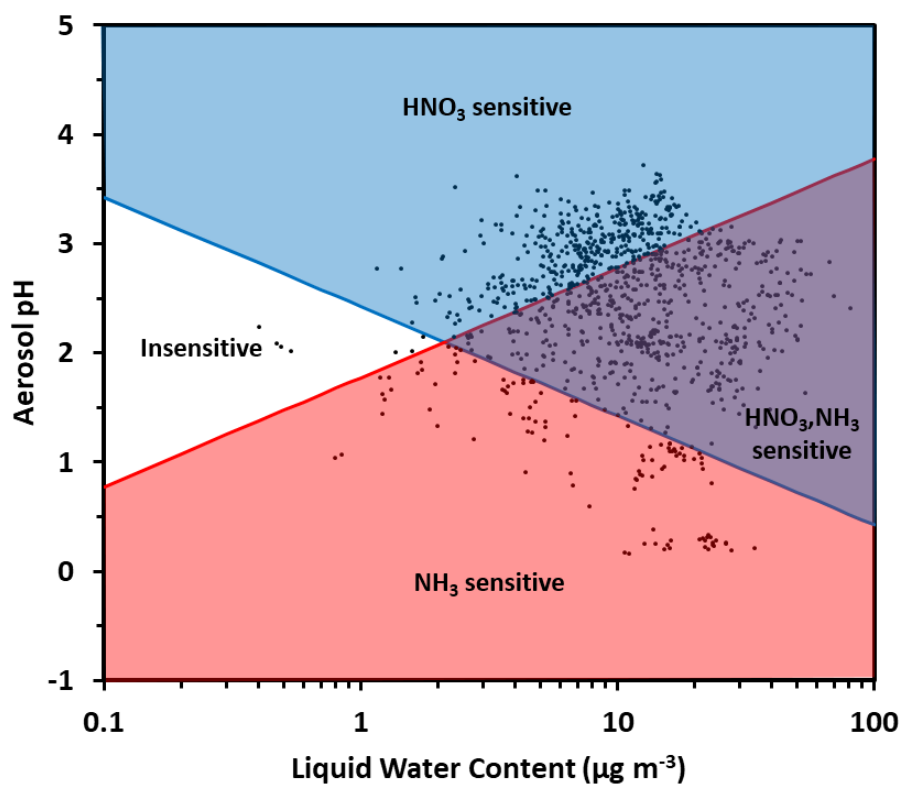
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(b)

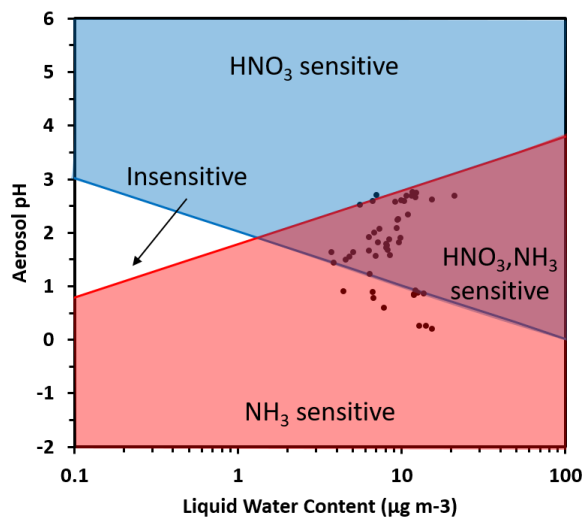
Figure 7



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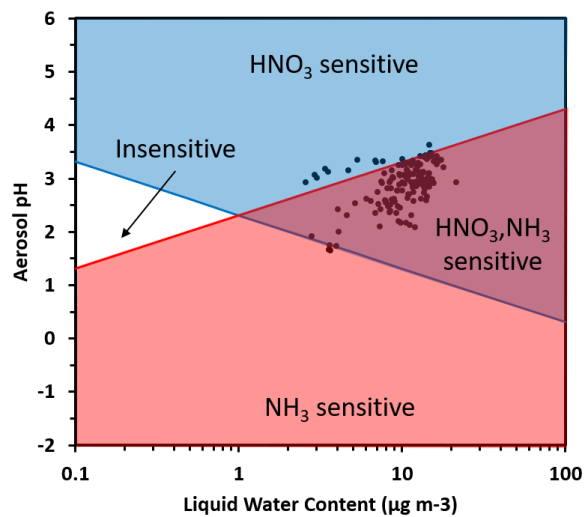


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789 **Figure 8**
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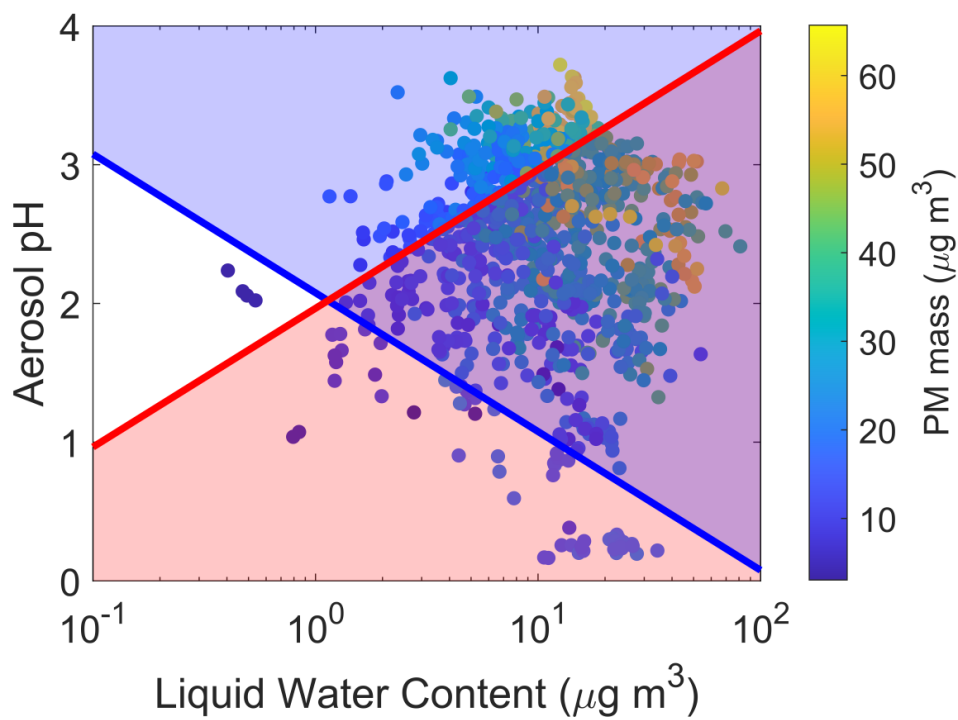
(a)



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Figure 9

(b)



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Figure 10