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# Regularisation model study for the least squares retrieval of aerosol extinction time series from UV/VIS MAX-DOAS observations for a ground layer profile parametrisation and westward viewing direction

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## Abstract

The retrieval of tropospheric aerosol extinctions from MAX-DOAS observations of O<sub>4</sub> using a small number of three or four extinction profile parameters suitable for boundary layer reconstruction is investigated with respect to the following questions. First, to what extent does this nominally over-constrained pure least squares problem for the inversion of the radiative transfer equation require regularisation and how should parameters of the regularisation be chosen. Second, how can a lack of information in the under-constrained case be best compensated by using the information contained in a sequence of observations and by explicitly including intensities into the fit.

The forward model parametrises the optical properties of the boundary layer aerosol by its extinction profile, single scattering albedo and a Henyey–Greenstein phase function. Forward calculations are carried out online, i.e. without look-up tables. The retrieval uses a Tikhonov regularisation combined with an approximate L-curve criterion and empirical a priori information from the retrieval sequence based on previous valid solutions. The consistency of the approach is demonstrated in selected model case studies assuming a polluted urban scenario and westward viewing direction of the instrument. It is shown that a dynamic choice of the regularisation parameter is crucial for high aerosol load and large diurnal variations. The quality of the retrieval can be improved significantly, if the retrieval sequence and thus the a priori is chosen according to the information content of the measurement series. Additional intensities improve the solution for all solar angles if suitably weighted. This flexible retrieval algorithm allows reconstruction of aerosol profiles in the boundary layer for a wide range of viewing directions and extinctions. It can thus be applied to observational geometries scanning the sky in two angular dimensions and to retrieve further aerosol optical parameters in the boundary layer.

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## 1 Introduction

The experimental method to obtain tropospheric aerosol extinction and trace gas concentration profiles by applying the differential optical absorption spectroscopy (DOAS, Platt and Stutz, 2008) to ground based observation of sun light intensities under different viewing directions has been an area of considerable activity in the past years – in field measurements in urban and remote areas (e.g. Wittrock et al., 2004; Irie et al., 2009; Lee et al., 2009; Li et al., 2010; Clémer et al., 2010; Frieß et al., 2011; Halla et al., 2011; Shaiganfar et al., 2011), within intensive measurement campaigns (e.g. Heckel et al., 2005; Sinreich et al., 2007; Brinksma et al., 2008; Roscoe et al., 2010; Irie et al., 2011; Vlemmix et al., 2011; Wagner et al., 2011; Zieger et al., 2011), as well as in comparative radiative model studies (Hendrick et al., 2006; Wagner et al., 2007). While the basic idea of this multi-axis DOAS, or MAX-DOAS, technique, namely to infer altitude information on an atmospheric absorber from its absorption signal along several light paths, is fairly straightforward and the experimental setup relatively inexpensive, the actual conversion requires inversion of the underlying radiative transfer equation. It strongly depends on tropospheric aerosol and, in general, the equation cannot be linearised. Evolving in a series of studies (Wagner et al., 2002; Hönninger et al., 2004; Sinreich et al., 2005; Wagner et al., 2004; Frieß et al., 2006), it has therefore been suggested to fit forward modelled absorption signals of an absorber with known concentration profile, the Oxygen dimer complex  $O_4$ , to the measured data in order to retrieve aerosol extinction profile parameters.

This method has been used in several of the field measurements referred to above with different experimental setups and different retrieval procedures. Instruments, amongst other factors, may differ in the number of  $O_4$ -wavelength absorptions bands their spectrometers cover and in the number and orientation of viewing angles under which their telescopes scan the sky. Apart from details of the DOAS-fit itself – here the  $O_4$  absorption cross section is a source of uncertainty (Wagner et al., 2009; Clémer et al., 2010), retrieval methods mainly differ in the way they parametrise the aerosol

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extinction and other aerosol parameters, in which radiative transfer model they use and in their fit algorithms. Currently there are mainly two kinds of approaches. One combines a linear parametrisation of the extinction profile by discrete layers and a statistical parameter estimation which in the area of satellite profile retrieval is commonly referred to as optimal estimation method (Rodgers, 2000) to result in a nonlinear least squares problem. This approach bears the advantage of providing an estimate for the uncertainty of the retrieval, as well as diagnostic tools for the information content of the measurement. Frieß et al. (2006) employed this approach in their comprehensive model studies to show, among other things, that the retrieval improves with increasing number of  $O_4$ -wavelength bands and if (relative) intensities are explicitly added to the fit quantities. Clémer et al. (2010) retrieved monthly aerosol extinction profiles in Beijing at four  $O_4$ -wavelengths separately using this method. Both studies come to the conclusion that the number of extinction profile parameters constrained by a MAX-DOAS measurement is rather limited and although the nonlinearity of the problem makes it hard to give a generally valid number, it is even for ideal conditions somewhere below four. The second kind of approach may be characterised by making use of empirically motivated profile shapes like linear or exponential functions and combinations of them, thus leading to a smaller set of parameters like ground extinction, layer height, aerosol optical depth etc. or, alternatively, scaling factors. Li et al. (2010) used this kind of parametrisation assuming a well-mixed ground layer for their measurements in Guangzhou, South of China. Wagner et al. (2011) proceeded similarly, but also accommodated a linear decrease from ground for their retrieval of aerosol and trace gas profiles in Milano. In both studies a pure (i.e. unregularised) nonlinear least squares fit is used to determine the four to six (for the former) or three (for the latter) profile parameters. Irie et al. (2008), on the other hand, introduced a set of three scaling factors for the total aerosol optical depth to parametrise partial optical depths and combined this with the optimal estimation method to retrieve four profile parameters in total. The forward radiative transfer model in the least squares fit uses pre-calculated “offline” look-up tables. All studies using profile shape parametrisation with a small set of parameters

employ look-up tables, whereas those mentioned using the discrete layer parametrisation carry out their calculation “online”.

The inversion of the atmospheric radiative transfer equation as an integral equation potentially constitutes an ill-posed inverse problem (e.g. Groetsch, 1993; Twomey, 1997; Rodgers, 2000), meaning, for example, that even if all fit parameters are well constrained by the observation, they might be extremely sensitive to measurement errors. This study first addresses the question to what extent a nominally “over-constrained” least squares problem for the inversion of ground based MAX-DOAS measurements of  $O_4$  to aerosol extinction coefficients requires modification to stabilise its solutions (regularisation). We use a simple parametrisation that represents the lower part of the profile linearly and the upper part exponentially to arrive at this formulation. The regularisation is studied using a basic Tikhonov scheme by comparing ad hoc choices of the regularisation parameter to an approximated L-curve criterion. We secondly examine the “under-constrained” problem and how to provide further (“a priori”) information from the context of a measurement series and by adding intensities to the fit (both similar to Frieß et al. (2006), but with slightly different conclusions for our case). By way of addressing these questions in selected, detailed case studies building on each other in successive sections of this paper, we finally arrive at a robust retrieval algorithm that can be used for any observation geometry and level of aerosol optical depth.

The measurement scenario is the same for all case studies. We assume a situation in urban air pollution monitoring in the UV/VIS spectral range and choose the day and location of the one-day time series arbitrarily as the 5 June 2010 in Shanghai. Aerosol optical properties other than the extinction coefficient are fix and chosen to be representative for air pollution in a Chinese megacity (see Sect. 3). All radiative transfer calculations are carried out online using the model SCIATRAN (Rozanov et al., 2005, <http://www.iup.uni-bremen.de/sciattran>). The retrieval combines data of the two  $O_4$  absorption bands around 360 and 477 nm.

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## 2 Method

### 2.1 Principle

The DOAS-technique makes use of the fact that atmospheric scattering and absorption processes have different and distinct wavelength dependencies which can be analysed in a spectrum by separating broad and narrow wavelength bands to retrieve trace gases through their unique narrow-band absorption structures. Radiative calibration, in other words absolute intensities, are dispensed with by dividing intensities  $I$  of a spectrum by those of a reference. Defining the optical depth of a certain absorber at wavelength  $\lambda$  through

$$\tau(\lambda) = -\ln\left(\frac{I(\lambda)}{I'(\lambda)}\right), \quad (1)$$

where  $I'$  are the intensities without this absorber, the DOAS-analysis ultimately yields differential optical depths of an absorber as

$$\Delta\tau(\lambda) = \tau(\lambda) - \tau_{\text{ref}}(\lambda). \quad (2)$$

We refer the reader to Platt and Stutz (2008) for details of this procedure. Slant column densities (SCDs)  $S$  of an absorber are defined by dividing optical densities by the absorption cross section  $\sigma_{\text{abs}}$ , i.e.  $S = \tau/\sigma_{\text{abs}}$ . Differential slant column densities (DSCDs)  $\Delta S$  are understood here in terms of the above differential optical densities as  $\Delta S = \Delta\tau/\sigma_{\text{abs}}$ .

We consider an experimental setup where the instrument records intensities  $I(\alpha)$  for a certain wavelength range in a cycle of elevation angles  $\alpha$  between  $0^\circ$  (horizon) and  $90^\circ$  (zenith) in a fix azimuthal plane  $\phi = \phi_{\text{obs}}$  and take the aforementioned reference spectrum to be recorded for each cycle in zenith direction  $\alpha_{\text{ref}} = 90^\circ$ . Through this choice, the stratospheric contributions in Eq. (2) cancel and one is left with the desired tropospheric part to the differential absorption optical depth. DOAS-evaluation

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of the measured intensities in this way with respect to the absorber  $O_4$  at one or several wavelength bands makes it possible to infer tropospheric aerosol extinctions by inverse modelling because the  $O_4$  profile is known. Details of this reasoning can be found in Wagner et al. (2002); Hönninger et al. (2004); Wagner et al. (2004); Sinreich et al. (2005); Frieß et al. (2006).

The observational data  $\mathbf{d} = \mathbf{d}(\theta, \phi)$  for a measurement at solar zenith and azimuth angles  $\theta$  and  $\phi$  is thus given by an  $m$ -dimensional vector

$$\mathbf{d} = (\Delta S_1, \dots, \Delta S_m)^T, \quad m = m_\alpha m_{O_4}, \quad (3)$$

where  $m_\alpha$  is the number of elevation angles (exclusive the reference direction) and  $m_{O_4}$  is the number of  $O_4$ -wavelengths bands used simultaneously in the profile retrieval. Frieß et al. (2006) suggested to use both  $O_4$ -optical depths (or SCDs) and relative intensities (RIs)  $\tilde{I} = I/I_{\text{ref}}$  as fit quantities. In this case  $\mathbf{d}$  becomes

$$\mathbf{d} = (\Delta S_1, \dots, \Delta S_{m/2}, \tilde{I}_1, \dots, \tilde{I}_{m/2})^T, \quad m = 2 m_\alpha m_{O_4}. \quad (4)$$

We assume that the aerosol extinction profile  $k(\lambda, z)$  at a certain wavelength  $\lambda$  ( $\lambda_{\text{fm}}$ , specified in Sect. 2.3) is given by  $n$  parameters  $x_j$

$$\mathbf{x} = (x_1, \dots, x_n)^T \mapsto k(\lambda, z, \mathbf{x}) \quad (5)$$

and that for this wavelength the forward model  $F(\mathbf{x})$  produces the quantities corresponding to the measurement data  $\mathbf{d}$ , that is

$$\mathbf{d} + \boldsymbol{\epsilon} = F(\mathbf{x}, \mathbf{b}) + \boldsymbol{\delta}, \quad (6)$$

with  $\boldsymbol{\epsilon}$ ,  $\boldsymbol{\delta}$  being the data and model errors, respectively, and  $\mathbf{b}$  including all forward model parameters but  $\mathbf{x}$ . Then  $\mathbf{x}$  can be retrieved by fitting the forward modelled quantities  $F(\mathbf{x})$  to the observational data using some cost function of  $\mathbf{d}$  and  $F$ .

## 2.2 Least squares fit – regularisation, a priori and information content

For parametrisations with (a small number of) parameters such that  $n \ll m$  the most straightforward approach to the inverse problem of retrieving  $\mathbf{x}$  in Eq. (6) is a nonlinear least squares principle, i.e. a quadratic cost function  $\chi^2(\mathbf{x})$  of the form

$$\min_{\mathbf{x} \in \mathbb{R}^n} (\mathbf{d} - F(\mathbf{x}))^T \mathbf{S}_e^{-1} (\mathbf{d} - F(\mathbf{x})), \quad (7)$$

where  $\mathbf{S}_e$  is the covariance matrix of the measurement data and  $\mathbf{b}$  is dropped for the moment. This approach requires that the underlying inverse problem is well-posed in the sense that the desired parameters are actually constrained by the measurement and that they are reasonably stable for a given level of noise. To examine whether the problem is ill-posed in this sense, we first linearise it, assuming that  $\mathbf{y}$  is a point in parameter space in the vicinity of the true optimum  $\hat{\mathbf{x}}$  so that

$$\mathbf{d} + \mathbf{e} = F(\mathbf{y}) + F_{\mathbf{y}}(\mathbf{y})(\hat{\mathbf{x}} - \mathbf{y}) + \dots \quad (8)$$

(derivatives are written throughout as  $F_x = \partial F / \partial \mathbf{x}$  etc.). We now apply the singular value decomposition (SVD) to the  $m \times n$  Jacobian matrix  $F_{\mathbf{y}}$ . For an arbitrary  $m \times n$  matrix  $\mathbf{M}$  of rank  $r$  it can be written as

$$\mathbf{M} = \mathbf{U} \mathbf{\Sigma} \mathbf{V}^T = \sum_{i=1}^r \sigma_i \mathbf{u}_i \mathbf{v}_i^T, \quad (9)$$

where the  $m \times n$  matrix  $\mathbf{\Sigma}$  is defined as

$$\mathbf{\Sigma} = \begin{bmatrix} \mathbf{\Sigma}_r & \mathbf{0} \\ \mathbf{0} & \mathbf{0} \end{bmatrix} \quad (10)$$

and  $\mathbf{\Sigma}_r = \text{diag}(\sigma_1, \sigma_2, \dots, \sigma_r)$  contains the  $r$  non-negative singular values by convention ordered as  $\sigma_1 \geq \sigma_2 \geq \dots \geq \sigma_r$ .  $\mathbf{U}$  and  $\mathbf{V}$  are orthonormal matrices of dimension  $m \times m$

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and  $n \times n$ , respectively, and  $\mathbf{u}_i, \mathbf{v}_j$  their column vectors. Inserting the singular vector expansion for  $F_y$ , Eq. (9), into Eq. (8) allows to draw conclusions similar to the linear case (e.g. Twomey, 1997; Hansen, 1998). Only if the contribution  $\sigma_i \mathbf{v}_i^T \hat{\mathbf{x}}$  of a singular value is significantly larger than the error components  $\epsilon_i$  (and  $\delta_i$ ) it adds information to the retrieval process, otherwise the inversion of the normal equation of the least squares problem may magnify errors up to the order of the condition number  $\sigma_r/\sigma_1$ . This is irrespective of the rank of the Jacobian matrix.

We will show later that for our aerosol parametrisation the pure least squares solution can indeed suffer from instability in some of its parameters and consider here the following modified least squares principle to regularise the solution of Eq. (7)

$$\min_{\mathbf{x} \in \mathbb{R}^n} (\mathbf{d} - F(\mathbf{x}))^T \mathbf{S}_e^{-1} (\mathbf{d} - F(\mathbf{x})) + \gamma^2 (\mathbf{x} - \mathbf{x}_a)^T (\mathbf{x} - \mathbf{x}_a). \quad (11)$$

This formulation of the so-called Tikhonov regularisation implies that the regularisation parameter  $\gamma$  is kept constant when solving the minimisation problem. The vector  $\mathbf{x}_a$  is usually referred to as a priori of  $\mathbf{x}$ . In the context of statistical parameter estimation for Gaussian probability distributions (e.g. Tarantola, 1987, 2005; Rodgers, 2000) Eq. (11) corresponds to the assumption of uncorrelated a priori distributions with mean  $\mathbf{x}_a$  and the same variance  $\sigma_a = \gamma^{-1}$ . Without going into any details, and assuming uncorrelated errors with variances of the same size  $\sigma_e$  for the weighted least squares problem, we note for the discussion of the local linearisation following Eq. (7) that, as a consequence of this regularisation, a singular value of the inverse for the weighted linear least squares problem is changed from  $(\sigma_i/\sigma_e)^{-1}$  to  $f_i (\sigma_i/\sigma_e)^{-1}$ , where the filter factor  $f_i = (\sigma_i/\sigma_e)^2 / ((\sigma_i/\sigma_e)^2 + \gamma^2)$  “filters out” contributions with  $\sigma_i/\sigma_e < \gamma$ . Contributions with  $\sigma_i/\sigma_e \gg \gamma$  are more or less unaffected by the regularisation. We note that substitution of  $\sigma_a^{-1}$  for  $\gamma$  yields the definition of independent measurements given in (Rodgers, 2000, chap. 2.4).

The right choice of the regularisation parameter  $\gamma$  (or  $\sigma_a$ ) is evidently important, but even in the linear case an intricate matter. Simple regularisation schemes can be

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based on an estimate  $\epsilon$  of the data error, e.g. a choice of  $\gamma$  such that  $\|d - F(x)\| > \|\epsilon\|$  to avoid “fitting to noise” (discrepancy principle). Other methods do not use an estimate of the noise. The empirically developed L-curve criterion, for example, relies on the observation that the optimal regularisation parameter appears close to a  $\gamma$  that balances the data residual and the regularisation norm  $\|x - x_a\|$  in Eq. (11). For details on these and more sophisticated regularisation schemes in the linear case we refer to (Hansen, 1998) and the discussion in Sect. 4. The statistical inversion of Eq. (11) using fix a priori uncertainties  $\sigma_a$  of the fit parameters as regularisation parameters implicitly assumes that the resulting problem is well-posed. Iterative solution of the nonlinear (or linear) least squares problem can, of course, also involve regularisation parameters (or even schemes) which depend on the state of the iteration (see, e.g. Doicu et al. (2010) for a comprehensive overview of atmospheric applications), but we would like to postpone the related discussion to Sect. 4 as part of our conclusions. In the present study we firstly want to investigate the relevance of regularisation for an “over-determined” formulation of the inverse problem at hand and discuss the efficiency of a fix, ad hoc choice of the regularisation parameter by comparing it to a very simple scheme related to the L-curve criterion for the linear case. It uses the fact, that if  $e$  and  $x_{\text{est}}$  are good estimates for  $\|\epsilon\|$  and  $x$  at the true values, then the choice  $\gamma = e/\|x_{\text{est}}\|$  is similar to that of the L-curve method (for the unweighted least squares problem) (Hansen, 1998, Sect. 7.5.3). Assuming that the expected value of the data residual is a reasonable estimate, this choice here becomes

$$\gamma = \frac{\sqrt{m}}{\|x_{\text{est}}\|}, \quad (12)$$

where the estimate  $x_{\text{est}}$  of  $x$  will be specified in Sect. 2.3. In the statistical framework this may be expressed as  $\sigma_a = \sqrt{n}/\sqrt{m} \cdot x_{\text{RMS}} \sim x_{\text{RMS}}$  with  $x_{\text{RMS}}$  being the RMS of  $x_{\text{RMS}}$ .

The role of the a priori  $x_a$  in the regularised framework is commonly discussed for a linearised approximation of  $F$  and the retrieval function  $R$ , so that the retrieved  $\hat{x} =$

$R(\mathbf{d})$  and the true  $\mathbf{x}$  are related through

$$\hat{\mathbf{x}} = \mathbf{A}\mathbf{x} + (\mathbf{I}_n - \mathbf{A})\mathbf{x}_a + \text{errors.} \quad (13)$$

The explicit form of the averaging kernel matrix (or resolution matrix)  $\mathbf{A} = R_d F_x$  can be found, e.g. in Rodgers (2000, chap. 3) (following from setting the a priori covariance matrix  $\mathbf{S}_a$  to  $\gamma^{-2}\mathbf{I}_n$  or  $\sigma_a^2\mathbf{I}_n$ , respectively). The “errors” include all contributions from measurement and forward model errors, see Sect. 2.4 for details. The first expression on the right hand side represents the contribution from the true state to the retrieval, the complementary second one describes how the “missing” components are provided by the a priori. For least squares solutions  $\gamma \rightarrow 0$  ( $\sigma_a \rightarrow \infty$ ) of full rank  $r = n < m$  the a priori has no influence. In the rank-deficient case the solution becomes ambiguous, algorithms producing the generalised inverse yield the minimum-norm solution. In the statistical context the trace of  $\mathbf{A}$  is often used to quantify how many independent degrees of freedom  $d_s$  out of the maximally  $m$  the measurement has (see, e.g. Rodgers (2000, chap. 3) or the discussion after Eq. 11). For uncorrelated errors  $\mathbf{e}$  with same variance  $\sigma_e$  it here takes the form

$$d_s = \sum_{i=1}^m \frac{(\sigma_i/\sigma_e)^2}{(\sigma_i/\sigma_e)^2 + \sigma_a^{-2}}, \quad (14)$$

where  $\sigma_i$  are the singular values of  $F_x(\hat{\mathbf{x}})$  (assumed to have full rank), so that  $m$  in the sum can here be replaced by  $n$ .

### 2.3 Aerosol parametrisation, forward model and retrieval of a time series

The aerosol extinction profile  $k(z)$  is divided with respect to altitude into a lower part  $k(z, \mathbf{x}, \mathbf{b})$  ( $z \leq z_{\text{TOR}}$ ) given mainly by the parameters  $\mathbf{x}$  to be retrieved and an upper part  $k(z, \mathbf{b})$  ( $z > z_{\text{TOR}}$ ) given by some of the forward model parameters  $\mathbf{b}$  not subject to the retrieval (with TOR standing for “top of retrieval”). The profile below  $z_{\text{TOR}}$  consists of a linear part starting at ground and an exponential part above with continuous

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transition at height  $H_L$ , see Fig. 1. The linear part is parametrised by extinction coefficients  $k_{L_1}$  at ground and  $k_{L_2}$  at height  $H_L$  and the partial aerosol optical depth (AOD)  $\tau_L$  between. The exponential part is given by its optical depth  $\tau_E$ . For this four-parameter representation we set

$$5 \quad \mathbf{x} = (x_1, x_2, x_3, x_4)^T = (k_{L_1}, k_{L_2}, \tau_L, \tau_E)^T. \quad (15)$$

For a well-mixed layer ( $k_{L_1} \equiv k_{L_2}$ ) the three parameters are

$$\mathbf{x} = (x_1, x_2, x_3)^T = (k_L, \tau_L, \tau_E)^T. \quad (16)$$

These aerosol profiles for  $\mathbf{x}$  are mapped to the forward quantities  $F(\mathbf{x})$  using the numerical radiative transport model SCIATRAN (version 2.2.2) (Rozanov et al., 2000, 2001, 2005, <http://www.iup.uni-bremen.de/sciattran>). In this study we are not particularly concerned with the details of the numerical model other than the fact that it has to be suitable for this application and the way it parametrises aerosol optical properties. SCIATRAN describes the atmosphere in terms of altitude layers for profiles and their derivatives (weighting functions) accounting for multiple scattering processes in a fully spherical geometry including refraction. It offers two methods to solve the radiative transfer equation, a combined differential-integral (CDI) method and a discrete ordinate method (DOM), where the choice of solver also depends on the output quantities desired. The former allows iterative improvement of the multiple scattering contribution with respect to sphericity, but not computation of aerosol weighting functions, whereas the latter produces these quantities, but does not allow refinement of the multiple scattering calculation. The model has been used in other MAX-DOAS studies (e.g. Wittrock et al., 2004; Frieß et al., 2006, 2011) and has been validated for this application in comparison with other radiative transfer models in Hendrick et al. (2006); Wagner et al. (2007).

25 Aerosol optical properties can be described in several ways in SCIATRAN, we use the mode that parametrises the aerosol by a Henyey–Greenstein phase function  $P_{HG}$

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with asymmetry parameter  $g(z)$  and its absorption and scattering coefficient profiles  $k_{a,s}(z)$  in discrete layers (the dependency on  $\lambda$  for all quantities is suppressed for the moment). Within this study, we assume that  $g$  and the aerosol single scattering albedo  $\omega_0$  are constant below  $z_{\text{TOR}}$ , so that the set of optical aerosol parameters  $k_a(z)$ ,  $k_s(z)$  and  $P_{\text{HG}}(z)$  becomes  $k(z, \mathbf{x})$ ,  $\omega_0$ ,  $g$  for  $z \leq z_{\text{TOR}}$  and  $k_a(z)$ ,  $k_s(z)$  and  $g(z)$  for  $z > z_{\text{TOR}}$ , respectively. The profile parametrisations involve a variable height  $H_L$  and have to be accommodated to the fix model grid given by altitude levels  $z_j$ . We do this by linear interpolation between the forward quantities  $F(z_j)$  and  $F(z_{j+1})$ , where  $z_j \leq H_L < z_{j+1}$ .

The retrieval algorithm based on the regularised least square fit in Sect. 2.2 and the forward model described above can be carried out with or without explicit use of intensities, Eq. (4), and for wavelengths  $\lambda_i$ ,  $i = 1, \dots, m_{O_4}$  representing the  $O_4$ -bands either individually or simultaneously. In the latter case profile parameters are calculated at one reference wavelength  $\lambda_{\text{fm}}$  which, together with Ångstrom exponents  $a_i$  for the conversion to extinction coefficients at  $\lambda_i$ , have to be specified as part of the forward model parameters  $\mathbf{b}$ .

The optimisation problem Eq. (11) is solved using an implementation of the well established Levenberg–Marquardt (LM) algorithm for this kind of nonlinear least squares problems (see, for example Nocedal and Wright (2006) and references therein) provided by the MINPACK library (Moré et al., 1980; Cowell, 1984, <http://www.netlib.org/minpack>). This library contains a thrust-region implementation of the Levenberg–Marquardt algorithm and was chosen here because it includes routines both for the case where derivatives are available and for the case where these are not directly available. In the latter case the Jacobian matrix  $\mathbf{J}$  used to calculate the next step correcting  $\mathbf{x}$  in each iteration is estimated using forward-differences. In the statistical framework the output at the optimum can be used to calculate the covariance matrix  $\hat{\mathbf{S}}$  of the result. All radiative transfer calculations during the iterative fit are carried out “online”, i.e. without using look-up tables.

The retrieval described so far refers to individual data  $\mathbf{d}$  and – unless other sources are available – both the a priori  $\mathbf{x}_a$  and the regularisation parameter in Eq. (12) have to

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be chosen more or less ad hoc. But since a MAX-DOAS measurement usually consists of a series of diurnal data, such an ad hoc choice is in fact only necessary for the very first retrieval within a day. The retrieval algorithm for a time series of data  $\mathbf{d}(T_j)$  for one day implemented here proceeds as follows: it starts with some non-committal choice of the initial parameters  $\mathbf{x}_{a,k=0}$  and  $\gamma_{k=0}$  (or  $\sigma_{a,k=0}$ ), preferably at a point in the series considered to represent the best measurement conditions/highest information content. The result  $\hat{\mathbf{x}}_k$  of the LM-iteration (starting throughout with  $\mathbf{x}_0 = \mathbf{x}_{a,k}$ ) is regarded here as valid, if the weighted data residual  $\|\mathbf{r}_d\|$  lies within a certain range of its expected value

$$10 \quad \mathbf{r}_{d_k}^T \mathbf{r}_{d_k} = (\mathbf{d}_k - \mathbf{F}(\hat{\mathbf{x}}_k))^T \mathbf{S}_{e,k}^{-1} (\mathbf{d}_k - \mathbf{F}(\hat{\mathbf{x}}_k)) \leq \Delta^2 \cdot m, \quad (17)$$

where  $\Delta = 3$  in this model study. This choice represents a 3- $\sigma$  threshold, but the exact value is not very important in this model study, where the only errors are random measurement errors. If the retrieval fails, a second fit uses the initial a priori parameters. A valid retrieval  $\hat{\mathbf{x}}_k$  is used as a priori for the next data point  $\mathbf{d}_{k+1}$  such that

$$15 \quad \mathbf{x}_{a,k+1} = \hat{\mathbf{x}}_k \quad (18)$$

$$\gamma_{k+1} = \frac{\sqrt{m}}{\|\hat{\mathbf{x}}_k\|}, \quad (19)$$

where the latter equation again defines  $\sigma_{a,k+1}$ . This procedure is somewhat similar to the retrieval scheme using a Kalman filter in Frieß et al. (2006), but without the explicit requirements of the Bayesian interpretation and a model for the evolution of  $\mathbf{x}$  with  $k$  or time  $T_j$ .

Finally, we introduce a modification of the data weights  $\sigma_e$  for the case that relative intensities  $\tilde{I}$  are used as additional fit quantities. Choosing fix errors for DSCDs and RIs here similar to Frieß et al. (2006) (the latter model study is for a northward viewing geometry), we observed that, in the presence of noise, for small relative azimuth angles the object function is dominated by contributions from  $\tilde{I}$  in such a way that the profile

retrieval suffers in terms of height information. We found that including intensities can still improve the retrieval, if their contribution is weighted (penalised) relative to the one from the  $O_4$ -slant columns. In the following, they are scaled with weights  $w_i$  such that the relative error for the  $\tilde{I}_i$  is not smaller than the one for the corresponding  $\Delta S_i$ . I.e., in the notation of Eq. (4)

$$\sigma_{\tilde{I}_i}^w = w_i^{-1} \sigma_{\tilde{I}_i} \quad \text{with} \quad w_i = \left( \frac{\sigma_{\Delta S_i}}{\Delta S_i} \right)^{-1} \frac{\sigma_{\tilde{I}_i}}{\tilde{I}_i}. \quad (20)$$

## 2.4 Error of the retrieval

We first address the error of the retrieval in parameter space, where we have to take into account that the simple profile parametrisation in the form of  $k(z, \mathbf{x})$ , Eqs. (5), (15) and (16), might not provide an adequate representation of the true extinction profile  $k(z)$  (e.g. for an elevated layer or two distinct layers etc.). We here define the parameters  $\mathbf{x}$  of the best possible representation as the ones that minimise  $k(\lambda_{\text{fm}}, z) - k(\lambda_{\text{fm}}, z, \mathbf{x})$  in a least square sense, which thus defines  $h(\lambda_{\text{fm}}, z)$  in

$$k(\lambda_{\text{fm}}, z) = k(\lambda_{\text{fm}}, z, \mathbf{x}) + h(\lambda_{\text{fm}}, z). \quad (21)$$

Assuming that the altitude grid of the forward model given by  $z_j$  has a sufficient spatial resolution,  $h(\lambda_{\text{fm}}, z)$  may be expressed by discrete parameters  $h_j$ . We now follow the error analysis in Rodgers (2000, chap. 3), treating  $\mathbf{h}$  in the same way as the other forward model parameters  $\mathbf{b}$ , so that the error in the retrieval  $\hat{\mathbf{x}}$  can be written as

$$\hat{\mathbf{x}} - \mathbf{x} = (\mathbf{I}_n - \mathbf{A})(\mathbf{x}_a - \mathbf{x}) + \mathbf{R}_d \mathbf{e} + \mathbf{R}_d \mathbf{F}_h \mathbf{h} + \mathbf{R}_d \Delta \mathbf{F}(\mathbf{x}, \mathbf{b}) + \mathbf{R}_d \mathbf{F}_b (\mathbf{b} - \hat{\mathbf{b}}) \quad (22)$$

with derivatives of  $\mathbf{F}$  evaluated at  $\mathbf{x} = \mathbf{x}_a$ ,  $\mathbf{b} = \hat{\mathbf{b}}$ ,  $\mathbf{h} = \mathbf{0}$  and  $\mathbf{R}_d = \mathbf{R}_d(\mathbf{d})$ . The first expression on the right hand side, usually referred to as smoothing error (also regularisation error), was discussed in Sect. 2.2. The second one (perturbation error) describes the propagation of the data error and the last two expressions are errors due to shortcomings in the formulation of the forward model itself and the choice of its parameters

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deviating from the optimal ones  $\hat{\mathbf{b}}$ . The third contribution represents the parametrisation error of  $k$ . Equation (22) results from various linearisation of both the forward and inverse model  $\hat{\mathbf{x}} = \mathbf{R}(\mathbf{d})$ , respectively, and therefore might not be accurate, for example, if  $\mathbf{x}_a$  is very different from  $\hat{\mathbf{x}}$  or if the representation error  $\mathbf{h}$  is large. Expectation values of individual contributions to the covariance-variance matrix for Eq. (22) in the statistical framework are discussed in Rodgers (2000, chap. 3). If  $\mathbf{x}_a$  and  $\mathbf{S}_a$  are good estimates for the true mean and its variances and covariances, then the covariance matrix of  $\hat{\mathbf{x}}$  is

$$\hat{\mathbf{S}} = \left( \mathbf{R}_d^T \mathbf{S}_e^{-1} \mathbf{R}_d + \mathbf{S}_a^{-1} \right)^{-1}, \quad (23)$$

if all forward model errors are negligible.

The error for a function  $f(\mathbf{x})$  at  $\hat{\mathbf{x}}$  is in first order of  $\hat{\mathbf{x}} - \mathbf{x}$  given by  $f_x(\hat{\mathbf{x}})(\hat{\mathbf{x}} - \mathbf{x})$ . In particular, the variance of the profile  $k(z, \hat{\mathbf{x}})$  becomes

$$\sigma_k^2(z, \hat{\mathbf{x}}) = [k_x(z, \hat{\mathbf{x}})]^T \hat{\mathbf{S}} k_x(z, \hat{\mathbf{x}}) \quad (24)$$

under the same circumstances as for Eq. (23). The variance of  $H_L(\hat{\mathbf{x}})$  follows correspondingly. The tropospheric (meaning here below  $z_{\text{TOR}}$ ) AOD  $\tau = \tau_L + \tau_E$  can be expressed as the special case of a linear function  $\mathbf{t}^T \mathbf{x}$ , so that

$$\sigma_\tau^2(\hat{\mathbf{x}}) = \mathbf{t}^T \hat{\mathbf{S}} \mathbf{t}, \quad (25)$$

with  $\mathbf{t} = (0, 0, 1, 1)^T$  for the four-parameter representation, Eq. (15), and  $\mathbf{t} = (0, 0, 1)^T$  for the three parameters in Eq. (16).

We would like to conclude the error discussion with a general remark on the error in data space. Using a linearisation around the optimum with assumptions similar to those leading to Eq. (22), one can derive the following expression for the error of the forward modelled quantities  $\hat{\mathbf{d}} = \mathbf{F}(\hat{\mathbf{x}})$

$$\hat{\mathbf{d}} - \mathbf{d} = (\mathbf{I}_m - \mathbf{A}_d)(\mathbf{F}(\mathbf{x}_a) - \mathbf{d}) + \mathbf{A}_d \mathbf{e}'. \quad (26)$$



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Here  $\mathbf{d}$  is the ideal, error-free data,  $\mathbf{A}_d = \mathbf{F}_x \mathbf{R}_d$  the data averaging kernel matrix or influence matrix and  $\mathbf{e}'$  represents all measurement and forward model errors. Using the picture of the linear case for the sake of clarity, one essential difference between the errors in Eqs. (22, 26) lies in the null-spaces of the maps  $\mathbf{F}_d$  and  $\mathbf{R}_d$ . For the least squares retrieval,  $\mathbf{I}_n - \mathbf{A}$  and  $\mathbf{I}_m - \mathbf{A}_d$  project onto the data and parameter null-spaces, respectively. In the over-determined, full rank case the former is empty, whereas the latter is not, so that the potential of any discrepancy between the model and data leading to large data residuals is higher than in the under-determined case. In this sense the data residual, Eq. (26), can be used as a criterion for the consistency of model and error assumptions of the retrieval.

### 3 Results

#### 3.1 Model and retrieval settings

In this model case study we consider mainly the retrieval of aerosol extinctions for the situation of urban air pollution monitoring, i.e. medium to high tropospheric aerosol optical depths. Unless stated otherwise in the following, the most important settings for the model experiment, the forward model and the retrieval are contained in Table 1. Representing restrictions of the view for a location in an area of high buildings, we choose the lowest elevation angle as  $2^\circ$  and the orientation of the telescope as west (north). Solar zenith and relative solar azimuth angles  $\theta$ ,  $\phi_{\text{rel}}$  shown in Fig. 2 are arbitrarily chosen corresponding to a midsummer day (5 June 2010) in Shanghai ( $31^\circ 12' \text{ N}$ ,  $121^\circ 30' \text{ E}$ ). Apart from the extinction profile below the retrieval height  $Z_{\text{TOR}}$ , all other aerosol optical properties are fix throughout as specified in Table 1. The upper tropospheric/stratospheric aerosol follows a standard background aerosol taken from the LOWTRAN database (Kneizys et al., 1988) with extinction coefficient  $\sim 0.0115 \text{ km}^{-1}$  at 5 km altitude and a single scattering albedo  $\omega_0$  of 0.99 (at 477 nm). The single scattering albedo  $\omega_0$  and the Henyey–Greenstein parameter  $g_{\text{HG}}$  below

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$z_{\text{TOR}}$  are set to values that are typical for air pollution containing absorbing aerosol (e.g. Lee et al., 2007; Garland et al., 2008, 2009). The Ångström parameter used here in the simultaneous fit of the first two  $O_4$ -bands to convert extinctions from the reference wavelength  $\lambda_{\text{fm}} = 477 \text{ nm}$  to  $360 \text{ nm}$  is set to 1.25. The number  $n$  of fit parameters is three or four. The initial a priori  $x_{a,0}$  is an exponential with ground extinction coefficient  $k(0) = 0.1 \text{ km}^{-1}$  and scale height  $\zeta = 1 \text{ km}$ .

We use the  $O_4$ -absorption cross section according to Greenblatt et al. (1990), but do not take into account the discrepancy between measured and modelled  $O_4$ -slant column densities reported in Clémer et al. (2010) (the investigation of this issue using results of this study will be part of our future work). Measurement errors are assumed to take the form of uncorrelated, Gaussian distributed noise with standard deviation  $\sigma_{e_0}$  as in Frieß et al. (2006).

### 3.2 Example for SVD-analysis of the forward model

To examine the conditions under which a measurement of the type just described allows a least squares fit of the three (four) parameters without regularisation, we first consider the singular value decomposition of the linearised model, Eq. (8), for two profile shapes and levels for the optical depths of about 1.2 and 0.12, respectively. The first profile type is a well-mixed ground layer of height  $H_L = 1.5 \text{ km}$  ( $O_4$ -DSCDs for these profiles are shown in Fig. 2), the second one is an exponential.

#### 3.2.1 Singular values

Figure 3 shows the three nonzero singular values for the three-parameter representation, Eq. (16), for viewing direction west (Fig. 3a) and north (Fig. 3b). For both levels of aerosol optical depths the range of singular values comprises several orders, as is characteristic of ill-posed problems and especially for the higher aerosol extinction, the smaller modes show a distinct variation with solar angles, indicating changes in the information content related to different profile parameters. Only singular values whose

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contribution according to Eq. (8) is above the measurement error represent a degree of freedom of  $x$  that can be distinguished from noise. For the profiles with higher aerosol extinction (top panel in Fig. 3), where Fourier-coefficients  $v_i^T x$  are of the order 0.1 to 1, a measurement error of about  $0.1 \times 10^{43} \text{ molecules}^2 \text{ cm}^{-5}$  would mean that the contribution of the smallest singular value is comparable to the size of the noise. Singular values for the lower optical depths (bottom panel in Fig. 3) are about ten times larger, but the Fourier-coefficients are reduced by a similar factor. To avoid that the smallest mode affects the retrieval through error magnification, the least squares formulation thus needs regularisation – at least for the errors and profiles assumed here. This will be further substantiated by model retrievals below, where it will be furthermore shown that the seemingly small variations of the second singular value for the ground layer profile have in fact significant influence on the retrieval.

### 3.2.2 Singular vectors in state space

To interpret the three modes distinct in all singular value decompositions of Fig. 3, the corresponding singular vectors for the layer profile in the top panel of this figure are shown in Fig. 4. The vector for the dominant first mode is almost completely in direction of the layer extinction coefficient  $k_L$ , whereas it does not contribute to the lower modes. This means that this parameter is both robust with respect to noise and the choice of the regularisation parameters. The second mode is modulated similarly to the differential slant columns. For westward viewing direction, it represents more or less the sum of  $\tau_L$  and  $\tau_E$ , i.e. the optical depth, except for times when the solar zenith (and the relative azimuth) angle is small. This means that the total optical depth  $\tau_L + \tau_E$  is a relatively well determined parameter, unless the noise is so large or the signal so low that the second mode becomes affected (as shown for this example in Sect. 3.4). In the same way it becomes clear that for the northward viewing direction the optical depth for this layer profile is less well determined for low solar elevation, illustrating the advantage of observation in different azimuthal directions.

### 3.3 Pure least squares retrieval of a polluted ground layer

We now consider the unregularised least squares problem Eq. (7) for the polluted ground layer with optical density  $\tau \sim 1.2$ .

#### 3.3.1 $\chi^2$ -landscape

5 In the linear case the object function  $\chi^2(\mathbf{x})$  in Eq. (7) describes ellipsoids in parameter space for a given value  $c^2 = \chi^2$ . Figure 5 illustrates how the nonlinear case behaves, again for the ground layer profile of the previous section with  $\tau \sim 1.2$ ,  $k_L = 0.5 \text{ km}^{-1}$ ,  $H_L = 1.5 \text{ km}$  (Fig. 2b), this time represented by the four parameters  $k_{L_1} = k_{L_2} = 0.5 \text{ km}^{-1}$ ,  $\tau_L = 0.75$  and  $\tau_E = 0.48$ . Figure 5a shows the dependency of  $\chi^2$  on the layer extinctions at the minimum  $\hat{\mathbf{x}}$ , Fig. 5b its variation with optical depths. Both graphs illustrate that only in a relatively close neighbourhood of  $\hat{\mathbf{x}}$  the linearisation is a sufficient approximation. In agreement with the discussion at the end of Sect. 3.2, the minimum is quite sharp in the sense that the residual function  $\chi^2$  grows rapidly with distance from  $\hat{\mathbf{x}}$  and thus a sufficiently fine resolution in  $x_j$  is needed in the forward model calculation in order to achieve an accurate value of the minimum. The function  $\chi^2$  shows the strongest variation in the ground extinction  $k_{L_1}$ , the penalising effect of the ground layer parameters is considerably larger than the exponential one.

#### 3.3.2 Monte Carlo simulations

20 The previous discussion did not involve any errors. To show how even moderate measurement errors affect the solution of the unregularised least squares problem, Eq. (7), we consider the same profile – i.e. a relatively high aerosol optical depth – for westward viewing direction (the corresponding example for low optical depth will appear in the next Sect. 3.4). Without measurement errors the inverse problem for both parametrizations Eqs. (16) and (15) can be solved almost exactly for all solar positions, as shown for the three-parameter representation with its parameters  $k_L$ ,  $\tau_L$  and  $\tau_E$  in Fig. 6. A Monte

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Carlo simulation of the retrieval with uncorrelated Gaussian distributed errors of variance  $\frac{1}{2}\sigma_{e_0}$  (about 5%) shows, as anticipated, that the layer extinction coefficient  $k_L$  corresponding to the first mode of the singular value decomposition is relatively stable against noise:  $\sim 5\%$  error on the data have an effect of less than 10% on the retrieval, except for solar positions with small  $\theta$  and  $\phi_{\text{rel}}$  (between about 13:00 and 15:00 LT) pointed out in the previous discussion of the singular vectors. The largest effect on any of the retrieval parameters occurs for the optical depth of the exponential  $\tau_E$ . Here the magnification in the unregularised inversion of an error of only 5% can lead to errors in the retrieval of more than 100%, which merely reflects the fact that it is the least well determined parameter for the MAX-DOAS retrieval. The error of the totally retrieved AOD  $\tau = \tau_L + \tau_E$  is considerably smaller than the individual errors, reflecting a strong correlation of  $\tau_E$  and  $\tau_L$ , see Eq. (25). As anticipated in the discussion of the singular value decomposition, the error of  $\tau$  is largest (up to 50%) when the second singular value is smallest (Fig. 3a bottom). The  $1\text{-}\sigma$  variation obtained from the Monte Carlo calculation generally agrees well with the variance of  $\hat{\mathbf{S}}$  calculated for error-free data at the true minimum. The reason for their discrepancy around 14:00 LT is that the retrieval can produce exponential profiles as acceptable solutions for which the estimate  $\hat{S}$  obtained from linearisation does not hold. Acceptable here refers to the data residuum in Eq. (17), shown in Fig. 7 (red). Taking into account the limited number of random samples ( $N = 100$ ), its mean value agrees well with the expected value of  $\sqrt{m}$  for all solar positions. To illustrate that this is not necessarily the case if explicit or implicit assumptions on the data or model error are incorrect, we also show  $\chi$  for the error-free retrieval under the assumption that  $\sigma_e = 10^{-3}\sigma_{e_0}$ . The residual could in principle be zero. But for small  $\phi_{\text{rel}}$  it lies outside the  $2\text{-}\sigma$  range, indicating that numerical errors in the forward model are in fact larger than the assumed data error.

### 3.4 Regularised least squares retrieval of ground layer profiles

In this section we demonstrate that the regularisation criterion according to Eqs. (12) and (19) is indeed a useful choice of the regularisation parameter and that the update

of the a priori according to Eq. (18) works consistently in the presence of noise. We consider again the retrieval of  $\mathbf{x} = (k_L, \tau_L, \tau_E)^T$  for the two ground layer profiles with AOD  $\tau \sim 0.12$  and  $\tau \sim 1.2$  and westward viewing direction. The retrieval procedure is as described at the end of Sect. 2.3 except that here we consider different ad hoc choices of a constant regularisation parameter.

### 3.4.1 Information content and averaging kernels

To get a rough idea of the potential influence of the regularisation parameter on the retrieval for arbitrary (but appropriate) a priori  $\mathbf{x}_a$  in the present scenario, we first look at the information content in terms of the degrees of freedom  $d_s$  and at the averaging kernels  $\mathbf{A}$ . Figure 8 shows  $d_s$  for both profiles and two levels of noise calculated for the error-free retrieval at  $\hat{\mathbf{x}}$ . The range of the regularisation parameters chosen corresponds to the size of the true parameters  $x_j$  ( $\sigma_a = 0.05$  for  $\tau \sim 0.1$  and  $\sigma_a = 0.5$  for  $\tau \sim 1$ ). The general tendencies of  $d_s$  in  $\sigma_a$  and  $\sigma_e$  can be understood from the special case in Eq. (14). The values of  $d_s$  become larger if either the signal-to-noise ratio increases or the regularisation becomes weaker ( $\sigma_a$  larger). Consequently  $d_s$  is considerably smaller for the profile with larger AOD  $\tau \sim 1$ . Depending on the value of  $\sigma_a$ , a measurement at AOD around  $\tau \sim 0.1$  can be expected to hold between 2 to almost 3 DOFs for a noise level given by  $\sigma_e \sim 10\%$ , whereas for high AODs around  $\tau \sim 1$  this number is 1 to 2.5. The temporal variation of  $d_s$  with solar angles is similar to that of the condition number except when the second singular value varies significantly. Thus the minimum of this singular value for  $\tau \sim 1$  at small  $\theta$  and  $\phi_{\text{rel}}$  (around 14:00 LT in Fig. 3a bottom) occurs again as a distinct reduction in information content. As argued before, the second singular value is related to the AOD  $\tau = \tau_L + \tau_E$  and the corresponding lack of information on  $\tau$  clearly appears in the averaging kernels in Fig. 9.

This figure shows the rows of  $\mathbf{A}$  for  $\sigma_a = 0.1$  in the case of the ground layer profile with  $\tau \sim 0.12$  and the range  $\sigma_a = 0.1 \dots 0.5$  for  $\tau \sim 1.2$  (this choice of  $\sigma_a$  will become clear shortly. For  $\tau \sim 0.1$  only one value of the regularisation parameter  $\sigma_a$  is shown because  $\mathbf{A}$  is less sensitive to  $\sigma_a$  than in the case of higher AOD  $\tau \sim 1$ ). As anticipated in the

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discussion of the singular values in Sect. 3.2 for the profile with  $\tau \sim 1$ , the layer extinction coefficient  $k_L$  is a well-resolved parameter almost unaffected by regularisation. For low AOD  $\tau \sim 0.1$  the averaging kernels for the optical depths  $\tau_L$  and  $\tau_E$  are “most diagonal” for large  $\theta$ . Vice versa the profile information is lowest for small  $\theta$ . Apart from the layer extinction  $k_L$ , the averaging kernels for higher AOD strongly depend on the regularisation level. The same regularisation parameter  $\sigma_a$  as before leaves almost only one DOF, namely  $k_L$ .

### 3.4.2 Monte Carlo simulations

We now turn to the retrieval results of simulated measurements with errors for the above fix values of the regularisation parameter. Averages and standard deviations are again obtained from Monte Carlo simulations for the measurement errors given in Table 1. Starting with the ground layer profile with low AOD  $\tau \sim 0.1$ ,  $(k_L, \tau_L, \tau_E) = (0.05 \text{ km}^{-1}, 0.075, 0.048)$  and singular values as in Fig. 3 (a top), it can be expected from the discussion following Eq. (11) that for errors of the DSCD around  $0.1 \times 10^{43} \text{ molecules}^2 \text{ cm}^{-5}$  the regularisation parameter  $\gamma$  has to be larger than around 1 to be effective for the smallest singular value. Figure 10 contains the result of modelled retrievals for different levels of regularisation. The pure least squares fit suffers from large errors in the retrieved optical depths. Regularisation parameters  $\sigma_a$  larger than about 0.1 ( $\gamma$  less than 10) have small impact on the retrieval. Values smaller than around 0.01 affect the lowest two singular values, forcing corresponding parameters  $x_j$  strongly towards the a priori (“over-regularisation”). A choice of  $\gamma$  around 20 appears to be a good compromise between minimising the perturbation error and the bias introduced by regularisation (In this case these are the only contributions to the retrieval error in Eq. 22). In the statistical framework this choice of  $\sigma_a = \gamma^{-1} = 0.05$  would correspond to an a priori variance being of the order of the true, unknown parameters.

Turning to the profile with high AOD  $\tau \sim 1$  and parameters  $(k_L, \tau_L, \tau_E) = (0.5 \text{ km}^{-1}, 0.75, 0.48)$  and assuming the same (or similar) measurement errors as before, the previous argument on the basis of the smallest singular value would again lead

to choose  $\sigma_a$  smaller than about 1. Results for the retrieval of Monte Carlo simulations in Fig. 11 suggest an optimal value of  $\sigma_a \sim 0.2$ , larger values increase the bias without real benefit for the variance of the perturbation error. In terms of a priori variances, this choice would again be similar to the size of the  $x_j$ .

5 The “optimal” regularisation parameters  $\sigma_a \sim 0.05$  ( $\gamma \sim 20$ ) for  $\tau \sim 0.1$  and  $\sigma_a \sim 0.2$  ( $\gamma \sim 5$ ) for  $\tau \sim 1$  found in the modelled retrievals agree reasonably with the values of  $\gamma \sim 33$  (for  $\tau \sim 0.1$ ) and  $\gamma \sim 3.3$  (for  $\tau \sim 1$ ) according to Eqs. (12, 19) if one inserts the true parameters  $x$ . Since this criterion for the choice of the regularisation parameter is not very sophisticated, we are not too concerned about precise values of the  
10 “best” parameter. However, we want to point out that for the range of aerosol extinctions considered here the choice of a constant regularisation parameter is not a good compromise (e.g. setting  $\gamma = 10$  in both examples would hardly have any regularising effect for the profile with  $\tau \sim 0.1$  and over-regularise the solution for  $\tau \sim 1$ ). The same conclusion holds, of course, for the a priori uncertainties  $\sigma_a$ , so that the approach to  
15 choose a non-committal  $\sigma_a$  reflecting the natural uncertainty of the aerosol extinctions is similarly inadequate.

Comparing the “optimal” regularisation parameters with the corresponding  $d_s$  in Fig. 8 shows that, depending on the solar angles, for low  $\tau \sim 0.1$  between 20–30 % are not fixed by the measurement, for high  $\tau \sim 1$  between 30–60 %. These numbers  
20 are mostly larger than the relative errors and here do not seem to allow any conclusion back on the choice of the regularisation parameter.

### 3.5 Use of relative intensities as fit quantities

The advantage of including (relative) intensities into the optimal estimate of aerosol extinction profiles has been demonstrated for the parametrisation by discrete layers  
25 and northward orientation of the telescope by Frieß et al. (2006). It can be expected that the retrieval of three (four) profile shape parameters benefits in a similar way.

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### 3.5.1 Information content and averaging kernels

For ideal, error-free measurement data for the polluted layer with  $\tau \sim 1.2$  this is indeed confirmed by the degrees of freedom for the measurement  $d_s$  shown in Fig. 12. Compared to the case without intensities (open diamonds in this figure)  $d_s$  enhances by 0.5 to 1.5 degrees of freedom when (unscaled) intensities are added to the fit (filled diamonds and squares in this figure). The highest increase of information occurs for large relative solar azimuth angles  $\phi_{\text{rel}}$  in the morning, the lowest for small solar zenith angles  $\theta$  around noon. The corresponding averaging kernels in Fig. 13 illustrate which profile parameters contribute to this increase. Comparing again the retrieval without and with (unscaled)  $\tilde{I}$ , Fig. 13b top indicates that for large and medium  $\phi_{\text{rel}}$  especially the upper, exponential part of the profile given by  $\tau_E$  can be retrieved more accurately, whereas for small  $\phi_{\text{rel}}$  the lower ground layer part,  $\tau_L$ , profits most from the additional information provided by  $\tilde{I}$  (Fig. 13a bottom). The change of the averaging kernels is less pronounced for small  $\theta$ , here mainly  $\tau_E$  is better resolved. The total tropospheric AOD  $\tau$  can now be retrieved reliably for all solar positions.

### 3.5.2 Monte Carlo simulations with relative intensities

Taking into account measurement errors for the simulated data of slant column densities and relative intensities according to Table 1 changes the situation as summarised in Fig. 14. The graphs are result of Monte Carlo simulations for the ground layer profile with  $\tau \sim 1.2$  similar to the previous ones, but now including relative intensities in the retrieval. The retrieval of the aerosol optical depth  $\tau = \tau_L + \tau_E$  is indeed greatly improved both with respect to its mean value and its variance for all solar positions when compared to the case without relative intensities in Fig. 11b bottom. However, the retrieved profile parameters  $k_L$ ,  $\tau_L$  and  $\tau_E$  do not show the behaviour expected from the above analysis of the information content for the error-free data. For very large relative azimuth angles ( $\phi_{\text{rel}} > \sim 150^\circ$  in this example) all parameters are more or less insensitive to the choice of the regularisation parameter (within the range of  $\sigma_a = 0.1 \dots 0.5$ ),

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showing similar variance and bias, which for  $\tau_E$  is actually larger than without intensities. With  $\phi_{\text{rel}}$  getting larger the parameters become increasingly sensitive to regularisation if the weights of  $\tilde{I}$  are not modified, that is  $\sigma_{\tilde{I}}$  not scaled. Irrespective of the regularisation parameter or weighting scheme for  $\tilde{I}$ , the AOD  $\tau_L$  and the height  $H_L$  of the layer are underestimated for  $\phi_{\text{rel}} < \sim 90^\circ$ , whereas they become overestimated for  $\phi_{\text{rel}} > \sim 90^\circ$  if  $\tilde{I}$  is not specially weighted. Scaling  $\sigma_{\tilde{I}}$  as proposed in Eq. (20) has only slight effect on the profile parameters for large  $\phi_{\text{rel}}$ , but improves their retrieval significantly for small  $\phi_{\text{rel}}$  and in this particular example of high AOD is even more accurate than for large relative azimuth angles.

The dependency of the error afflicted retrieval on the solar azimuth angle  $\phi_{\text{rel}}$  in this example is, of course, related to a strong increase in the signal-to-noise ratio of the relative intensities for smaller  $\phi_{\text{rel}}$  when looking into the direction of the sun. The fact that the contribution from the relative intensities to the cost function in Eq. (11) may outweigh the one from the differential slant columns, which carry most of the profile information, does not pose a problem in itself, if measurement errors can be neglected. If measurement errors cannot be neglected, one first has to take into account that the two data sets of differential column densities and relative intensities are assumed to be completely uncorrelated within themselves and with respect to each other. In order to see to what extent the neglect of correlations between DSCDs and RIs in their relative weighting in the object function becomes relevant, we look at the statistics of the data residual in Fig. 15. If all model assumptions were correct, the average data residual for the Monte Carlo runs would be around its expected value  $\sqrt{m} \sim 5.3$ , similarly to Fig. 7. With the exception of small solar zenith angles  $\theta$ , the actual mean residuals are significantly larger and for small relative solar azimuth angles  $\phi_{\text{rel}}$  frequently exceed the threshold for valid retrievals in Eq. (17). The sometimes high number of invalid results distorts the statistics and causes the jumps in some of the curves in Fig. 15.

The second, less obvious implication of adding relative intensities to the fit vector is that the latter are in some form functions of  $\tau$ , and therefore the fit parameters  $\tau_L$ ,  $\tau_E$  themselves (in particular if  $\tau$  is chosen as fit parameter). Instead of regarding the

addition of intensities as an augmentation of the data vector, one might look at it as an additional contribution to the a priori or regularisation term with covariance matrix depending on the solar position. In this sense we empirically weight this contribution according to Eq. (20).

### 5 3.6 Retrieval of profile times series

The retrieval algorithm of the previous section for static, well-mixed ground layer profiles is now applied in two examples of profile time series with Gaussian functions serving as model profiles. In the first example the series of profiles shown in Fig. 16a top start with low AOD  $\tau \sim 0.063$  at large relative solar azimuth  $\phi_{\text{rel}}$ . The maximum extinction  
10  $k_0 = 0.05 \text{ km}^{-1}$ , its height  $z_0 = 1.5 \text{ km}$  and the half width  $\sigma_0 = 0.5 \text{ km}$  of the Gaussian profile evolve linearly in time such that at the other end of the time series for small  $\phi_{\text{rel}}$  these parameters are  $k_0 = 1.5 \text{ km}^{-1}$ ,  $z_0 = 0 \text{ km}$  and  $\sigma_0 = 1 \text{ km}$ . The AOD increases by a factor of about 30 to  $\tau \sim 1.86$ . The second example of profile time series consists of the same profiles in reverse order in time, see Fig. 18a top, i.e. from high AOD for  
15 long light paths to low AOD for short light paths. The profiles in the forward model are represented now by the four parameters in  $\mathbf{x} = (k_{L_1}, k_{L_2}, \tau_L, \tau_E)^T$  (see Fig. 1) to be retrieved and Monte Carlo simulations of measurements are carried out in the same manner as before.

#### 3.6.1 Example for $d_s$ increasing with time

20 Statistics for the retrieval of the first profile series are illustrated in Fig. 16 for four different settings of the retrieval process. The first setting uses a fix regularisation parameter  $\sigma_a = 0.1$  – this value was found to be a reasonable choice for the range of profile parameters in Sect. 3.4 – and no relative intensities in the data vector. The second retrieval uses a variable regularisation parameter according to Eqs. (12, 19) and the third and  
25 forth additionally include relative intensities with unscaled and scaled weights. Values for  $\sigma_{a_k}$  range from less than about 0.02 for the lowest AOD to about 0.4–0.6 for highest

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AODs. Depending on the choice of the regularisation scheme one obtains different numbers for the degrees of freedom of the signal  $d_s$ , but in either case the fact that extinctions increase and the lengths of the light paths generally decrease within the time series results in decreasing signal-to-noise ratios (assuming constant measurement errors). For the first retrieval scenario with constant  $\sigma_a = 0.1$  (green colour in this figure)  $d_s$  goes down continuously from around 3.8 for large  $\phi_{\text{rel}}$  in the morning to only 0.8 for small  $\phi_{\text{rel}}$  in the evening. For the second retrieval with variable  $\sigma_{a_k}$  (orange) the corresponding numbers are about 3.8 and 2.3.

In this case, it therefore makes sense to start the retrieval algorithm at the first data point with largest  $\theta$  and  $\phi_{\text{rel}}$  and since the different choices of regularisation and weighting perform in general similarly to the previous constant profile examples, we here concentrate on some aspects we consider most noteworthy. The first thing to notice is that the reconstruction of the first two or three profiles consisting of elevated layers with low extinctions does not work well for the regularisation with  $\sigma_{a_k}$ , especially when including  $\tilde{l}$ . The small values of  $\sigma_{a_k}$  at this point possibly over-regularise the problem and result in exponential (for scaled  $\sigma_{\tilde{l}}$ , red) or no (for unscaled  $\sigma_{\tilde{l}}$ , blue) solutions, which might suggest that for very low AODs the regularisation criterion in Eqs. (12, 19) is not appropriate. Another explanation is that the true profile  $k(z)$  cannot be well represented by  $\mathbf{x}$ . The subsequent retrievals, however, confirm the advantage of the variable regularisation over the fix  $\sigma_a = 0.1$  as illustrated by the examples in Fig. 17. The latter now over-regularises and produces profiles that, in terms of standard deviations, are partly inconsistent with the true profile parameters. Without intensities the AOD is under-estimated for small  $\phi_{\text{rel}}$ , as before, due to the lack of information, i.e. low signal-to-noise ratio. A tendency to overestimate the extinction coefficient  $k_{L_2}$  evident in Figs. 18a top and 17 occurred frequently, also for retrievals of well-mixed ground layers parametrised by  $\mathbf{x} = (k_{L_1}, k_{L_2}, \tau_L, \tau_E)^T$  – usually accompanied by an under-estimation of the ground value. The data residuum  $\|\mathbf{r}_{\text{d}}\|$  behaves similarly to the retrievals shown before and, except for some violations of the validity criterion Eq. (17) for the setting

with  $\sigma_{a_k}$  and unscaled  $\sigma_{\tilde{I}}$ , in practice does not allow to discriminate between good or bad retrievals.

### 3.6.2 Example for $d_s$ decreasing with time

While for all profile series so far starting point and sequence of the retrieval process coincided with the chronological order of the time series, the second example in Fig. 18a top will now underline how important the choice of the sequence of retrievals is in order to make optimal use of varying information content of the measurement and thus quality of the empirical a priori. Analysing the information content expressed in  $d_s$  for the variable regularisation with  $\sigma_{a_k}$  (without  $\tilde{I}$ ) both in forward and backward direction, it turns out that  $d_s$  has a maximum value of almost 4 when starting with the last data point in time. It decreases to a value of about 2 at the first data when going through the sequence in reverse chronological order. The decrease is not strictly monotonic, but the effect on the retrieval quality is clearly obvious from Fig. 18. Monte Carlo results for simulated retrievals were carried out for the regularisation parameter  $\sigma_{a_k}$  with and without relative intensities in forward direction starting with the first data point ( $\rightarrow$ ) and backward direction starting with the last data point ( $\leftarrow$ ).

In the discussion of the results in this figure we restrict ourselves again to the most important observations. Starting the retrieval in forward direction, the lower information content at the beginning of the profile time series shows in the “warm-up” character of the first retrieval needed to adjust both  $x_a$  and  $\sigma_a$ . The ground extinction coefficient  $k_{L_1}$  (corresponding to the largest singular value for the polluted ground layer profile) can be estimated relatively well. This is less the case in this direction of the retrieval sequence for the “height”  $H_L$  of the ground layer and the extinction coefficient  $k_{L_2}$  at its top. The height  $H_L$  is overestimated throughout and for  $\phi_{\text{rel}} < \sim 90^\circ$  the profile information with respect to  $H_L$  and  $k_{L_2}$  is rather limited for the high AODs in this case. Including relative intensities makes little difference here. For  $\phi_{\text{rel}} < \sim 90^\circ$  and small  $\theta$ , however, they improve notably the AOD  $\tau$ . But for small  $\phi_{\text{rel}}$  at the end of the series the last two retrievals fail regardless. This time not due to over-regularisation (the other direction works), but

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either again due to the inadequacy of the profile representation or due to an interplay of regularisation and a priori. The effect of the latter, carried through from each retrieval to the next, is most obvious when comparing the retrievals in forward and backward direction, especially the profile parameters  $H_L$  and  $k_{L_2}$ . The retrievals (without  $\tilde{I}$ ) for small relative solar azimuth illustrate that depending on the a priori the AOD can be both persistently over-estimated ( $\rightarrow$ , green in Fig. 18b bottom) and under-estimated ( $\leftarrow$ , yellow in Fig. 18b bottom). Similar to the first example – but not as evident here – it would not improve the retrieval of the profile time series, if forward and backward direction were to be combined.

## 4 Conclusions and outlook

In this model study we address the questions how the information content in the least-squares retrieval of profile shape parameters for aerosol optical extinction profiles  $k(z)$  from MAX-DOAS observations can be used more efficiently by choosing appropriate parameters for noise filtering (regularisation) and by exploiting the context of a time series of measurements (empirical a priori). The simulated measurements for two  $O_4$ -wavelengths in the UV/VIS (360 and 477 nm) are assumed to take place in an urban, polluted environment with the instrument pointing to the west. Forward calculations are carried out at 470 nm and the retrieval combines both wavelengths using a fix Ångström exponent  $a$ . Other (equally fix) aerosol optical parameters are expressed in terms of the single scattering albedo  $\omega_0$ , the asymmetry parameter  $g$  and a Henyey–Greenstein phase function. The aerosol profile  $k(z)$  is parametrised linearly up to a variable height  $H_L$  and exponentially above  $H_L$ . The either three or four profile parameters are given by extinction coefficients and optical depth for the linear part ( $k_L$  and  $\tau_L$  or  $k_{L_1}$ ,  $k_{L_2}$  and  $\tau_L$ ) and the optical depth of the exponential ( $\tau_E$ ). The retrieval uses either differential slant column densities  $\Delta S$  or  $\tilde{\Delta S}$  plus relative intensities  $\tilde{I}$  for all elevations as fit quantities. Errors for these data are assumed to be completely uncorrelated. The forward model for the fit is involved “online”, i.e. without look up tables. We use a Tikhonov regularisation

scheme with one regularisation parameter chosen to be fix or to approximate the L-curve criterion.

#### 4.1 Conclusions and further discussion

Summarising our results from both the singular value decomposition (SVD) and Monte Carlo simulations, we first conclude that the inverse MAX-DOAS problem is ill-posed in the sense that some of its fit parameters cannot be determined if measurement errors are present. By this we mean that even under ideal conditions (no model errors, low noise, long light paths etc.) the relative error in the retrieved parameter may exceed relative errors of the measurement by one or two orders of magnitude. For the examples of ground layer profiles containing absorbing aerosols shown here, these parameters correspond to the profile information incorporated in  $\tau_E$  (and  $\tau_L$ ), while the layer extinction coefficient  $k_L$  and to lesser extent the totally retrieved AOD  $\tau$  respond moderately. This holds for low and high AODs, but depends, of course, on the signal-to-noise ratio.

Regularisation in form of the Tikhonov scheme with parameter  $\gamma$  (or  $\sigma_a^{-1}$ ) proves effective, but the right choice of the regularisation parameter is crucial, as it is shown to depend on the aerosol extinctions themselves. A fix value of  $\gamma$  does not accommodate a realistic range of AODs. The approximate L-curve criterion employed here works well for low to high AODs, but possibly over-regularises for very small AODs and might just be replaced by an upper bound (for  $\gamma$ ) in these cases. Such an upper bound can, in principle, be given by a smallest “possible” measurement error (see the discussion after Eq. 9), but we rather suggest a more refined parameter choice as proposed in the outlook below.

The dynamic update of both the regularisation parameter  $\sigma_{a_k}$  and the a priori  $\mathbf{x}_{a_k}$  based on the last valid retrieval within a daytime series of data was demonstrated to work consistently for random measurement noise and strong temporal variation of the aerosol profile. While for the present “over-determined” formulation of the inverse problem we did not find the information content embodied in the signal degrees of freedom  $d_s$  helpful for the choice of the regularisation parameter, variations of  $d_s$  within the

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time series (of a day) can be used to establish the starting point  $x_{a_0}$  and order of the retrieval sequence. Choosing  $x_{a_0}$  according to maximal information content allows to reduce the smoothing error for subsequent retrievals suffering from a lack of information, e.g. for short light paths and high AODs. Results from different retrieval sequences (for example in forward/backward direction) can, in principle, be combined and used to estimate the smoothing error, but this should only be considered if these sequences are regarded as equally valid or good.

Explicit addition of relative intensities to the measurement data improves the retrieval of  $\tau$  in all cases for all viewing geometries. For the reconstruction of profile features, however, we find a certain discrepancy between the error-free analysis using averaging kernels and results from MC simulations. While the former suggest better retrieval of profile shape parameters like  $\tau_L$ ,  $\tau_E$  as well, the latter indicate that these retrieval qualities depend to a certain extent on the weighting of  $\tilde{I}$ , especially for small relative solar azimuth when the signal-to-noise ratios of  $\tilde{I}$  are large. This disagreement is most likely due to neglected correlations between the  $\Delta S$  and  $\tilde{I}$  and is here accounted for by giving relative errors of  $\tilde{I}$  weights not larger than the ones of  $\Delta S$ . Alternative approaches will be outlined below.

An immediate consequence of the inverse MAX-DOAS problem being ill-posed is the fact that all data errors should be kept as small as possible. On the one hand, large errors do not necessarily mean that all retrieved parameters are equally affected. In our examples of ground layer profiles the layer extinction is a largely decoupled and robust parameter. This implies that the experimental validation of such “safe” parameters related to the largest singular mode in the SVD of the forward model’s Jacobian does not necessarily mean a validation of the complete profile. On the other hand, the benefit of accurate data can partly be undone by an unsuitable choice of the retrieval parameters, for example by giving the a priori too much weight. Not only the minimisation of measurement errors but also the correct estimate of their (and all the errors’) actual size and statistics is of great importance to be able to use absolute values of the data residuals  $\|r_d\|$  as indicators for a valid retrieval – obviously vital for the way the



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empirical a priori is chosen here. Large values of the residual suggest a problem with the underlying model assumptions. One example discussed was the expected value of  $\|r_d\|$  being too high when  $\tilde{T}$  is included into the fit. Another simple example would be a true extinction profile that can only be represented with large errors ( $h$  in Eq. 21) by the profile parametrisation used in this study, resulting in high residuals in Eq. (26). Low or consistent residuals, however, do by no means automatically guarantee a good retrieval, see the example of the pure least squares fit. This might be the point to raise the question whether coarse look-up tables are able to achieve consistent residuals in a similar manner (compare Fig. 5). As a side remark, we note that the (linear) interpolation involved by look-up tables is one way of implicit regularisation (Groetsch, 1993).

## 4.2 Outlook

This first attempt to investigate the regularisation behaviour of the inverse MAX-DOAS problem was naturally simplified. An obvious generalisation of the regularisation scheme would be to adjust the parameter  $\gamma$  (or  $\sigma_a$ ) during the Levenberg–Marquardt (LM) (or Gauss–Newton (GN)) iteration such that at each step it is chosen to minimise the object function. However, this is computationally expensive and, for example in Doicu et al. (2004), replaced by an empirical relaxation scheme of the regularisation parameter. It might even be possible to carry out the decisive choice of the regularisation parameter at the last step of the LM/GN iteration, that is, for a linear inverse problem. Replacing the regularising contribution in the object function  $\gamma\|x - x_a\|$  by  $\gamma\|\mathbf{L}(x - x_a)\|$  with  $\gamma^2\mathbf{L}^T\mathbf{L}$  corresponding to  $\mathbf{S}_a^{-1}$  in the optimal estimate (see Sect. 1) allows to give parameters different relative weights (e.g. regularise  $\tau_E$  stronger than  $\tau_L$ ) and to introduce correlations (e.g. for  $\tau_E$  and  $\tau_L$ ). An alternative viewpoint is that this more general regularisation enables to choose the basis vectors  $\mathbf{v}_j$  of the now generalised SVD (GSVD), i.e. to control how individual modes are regularised (Hansen, 1998). If the regularisation can be carried out for the linearised problem, a promising method for the choice of the overall regularisation parameter is the Generalised Cross-Validation (GCV, e.g.

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Hansen, 1998 and references therein) developed for statistical parameter estimation. It is designed to regularise the problem according to the different information content of individual data and would automatically account for correlation between them. While the Tikhonov regularisation and consequently its filtering behaviour is not intended for any specific class of ill-posed problems, certain algorithms originally designed to solve large tomographic problems have been proven quite successful for regularising the inverse problem arising from this kind of remote-sensing measurement. These so-called row acting methods or algebraic reconstruction techniques regularise the least squares problem iteratively and have been applied to active DOAS measurements in Laepple et al. (2004); Hartl et al. (2006) and to stratospheric ozone profile retrieval from limb scatter satellite observations in Degenstein et al. (2009). Since MAX-DOAS observations are essentially tomographic measurements, we believe that their analysis might benefit from this more specialised approach.

The time series retrieval presented in this study requires strictly speaking an a posteriori analysis of the information content ( $d_s$ ), i.e. two retrieval runs. While, in principle, this is not a problem, it would be desirable to perform the time series analysis in a more rigorous probabilistic framework and, for example, to estimate relaxation parameters introduced ad hoc in Frieß et al. (2006) to mimic a model for the temporal evolution used in the application of a Kalman filter.

Relative intensities are used here and in the original work Frieß et al. (2006) to constrain certain parameters or combinations thereof, namely the integrated extinction coefficients, rather than to add independent profile information. Taking  $\tilde{I}$  into account in form of inequalities like  $\tilde{I}_{i_{\min}} \leq f_i(\mathbf{x}) \leq \tilde{I}_{i_{\max}}$  may turn the resulting equality bound least-squares problem less sensitive to uncertainties in the true  $\sigma_j$  and avoid the problem arising from very different weights of slant column densities and relative intensities in the object function altogether. Alternatively, the forward and inverse problem may be expressed in terms of (relative) intensities alone, which might be a more appropriate formulation to exploit the full wavelength-intensity spectrum measured, for example in a simultaneous retrieval of aerosol and trace gas profiles, or to retrieve other optical

parameters like the single scattering albedo  $\omega_0$ , the asymmetry parameter  $g$  etc. Regardless of the formulation, regularisation of the underlying inverse problem remains important in order to avoid losing valuable information.

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**Table 1.** Settings for model experiments and retrieval unless stated otherwise in the text. The model errors  $\sigma_{e_0}$  for the differential optical depths and the relative intensities are adopted from Frieß et al. (2006) (for the wavelength 550 nm).

Model experiment	
$\alpha(m_\alpha)$	2, 4, 6, 8, 10, 12, 15, 90° (7)
$\phi_{\text{obs}}$	West (North)
$\theta, \phi$	Shanghai, 5 June 2010 (see Fig. 2)
$\lambda_{\text{O}_4}, (m_{\text{O}_4})$	360, 477 nm (2)
$\sigma_{e_0}(\Delta\tau)$	$5 \times 10^{-4}$
$\sigma_{e_0}(\tilde{I})$	$5 \times 10^{-4}$
Forward model	
$\lambda_{\text{fm}}$	477 nm
$z_{\text{TOR}}$	5 km
$z_i (< z_{\text{TOR}})$	$\Delta z = 0.1$ km
$\text{O}_4$ -cross section	Greenblatt et al. (1990)
Model profiles	
aerosol $z < z_{\text{TOR}}$	$\omega_0 = 0.95, g_{\text{HG}} = 0.68, \text{\AA} = 1.25$
aerosol $z > z_{\text{TOR}}$	LOWTRAN background
$\rho, T$ , trace gases	standard atmosphere
Retrieval	
$m = m_\alpha m_{\text{O}_4} (\times 2)$	14 (28)
$n$	3 or 4
$x_{a,0}$ (initial)	$0.1 \text{ km}^{-1} \cdot \exp(-z/1 \text{ km})$

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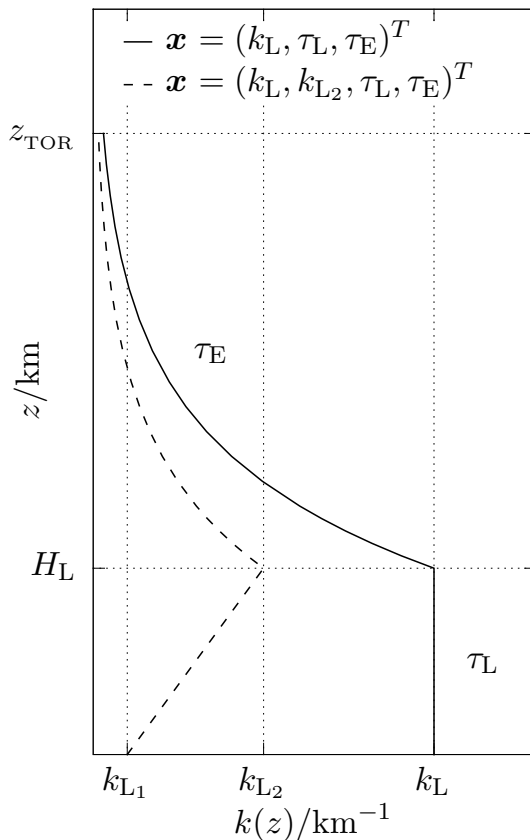
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**Fig. 1.** Parametrisation of the aerosol extinction profile below retrieval height  $z_{\text{TOR}}$ : linear from ground to height  $H_L$ , exponential from  $H_L$  to  $z_{\text{TOR}}$  and continuous at  $z = H_L$ .  $\tau_L$  and  $\tau_E$  are the partial optical depths between ground and  $H_L$  and  $H_L$  and  $z_{\text{TOR}}$ , respectively. The three-parameter case with  $k_{L_1} = k_{L_2}$  represents a well-mixed layer.

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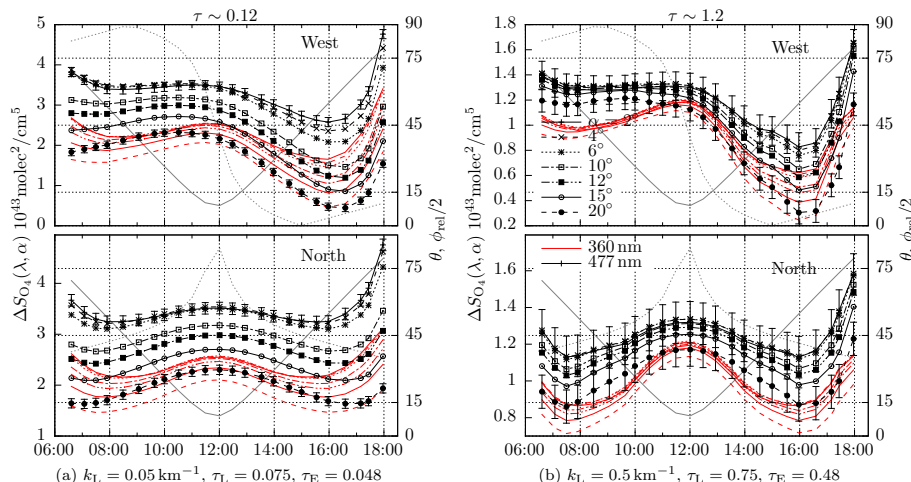
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**Fig. 2.** Differential slant column densities of  $O_4$  for the model experiment specified in Table 1 and a well-mixed ground layer of height  $H_L = 1.5\text{ km}$  with low aerosol load in (a), high aerosol load in (b). Upper plots for instrument pointing to west, lower for north. Layer parameters, see Fig. 1, refer to  $\lambda_{\text{fm}} = 477\text{ nm}$ . The exponential part of the profile has a scale height  $\zeta = 1\text{ km}$ . Solar zenith and relative solar azimuth angles  $\theta$  and  $\phi_{\text{rel}}$  referring to Shanghai ( $31^\circ 12' \text{ N}$ ,  $121^\circ 30' \text{ E}$ ) on the 5 June 2010 are here and in following figures indicated by the solid and dashed line, respectively.

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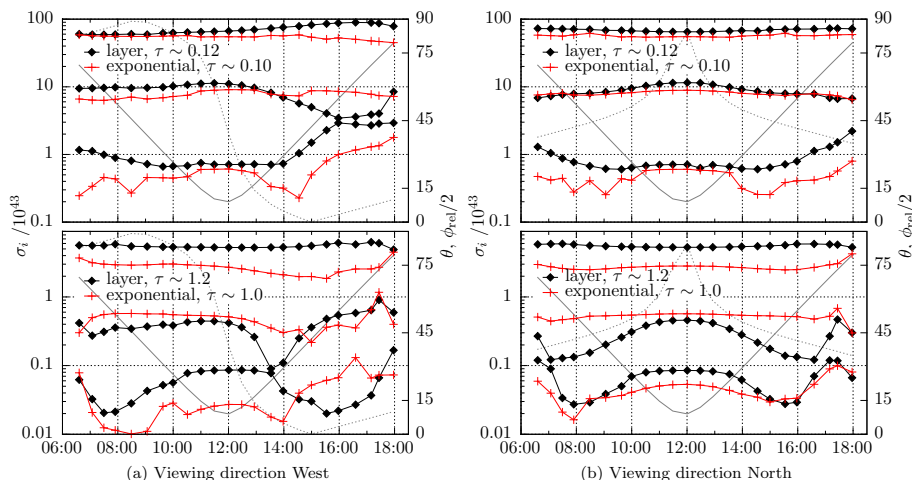
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**Fig. 3.** The three non-zero singular values  $\sigma_i$  of the Jacobian matrix for the forward model (without  $\tilde{I}$ ) using the profile parameters  $k_L$ ,  $\tau_L$  and  $\tau_E$  for the same ground layer profiles as in Fig. 2. For comparison also shown  $\sigma_i$  for exponential profiles with scale height  $\zeta = 1$  km and ground extinction coefficients  $k_L$  of  $0.1$  and  $1 \text{ km}^{-1}$ , respectively. **(a)** Viewing direction west, **(b)** north.

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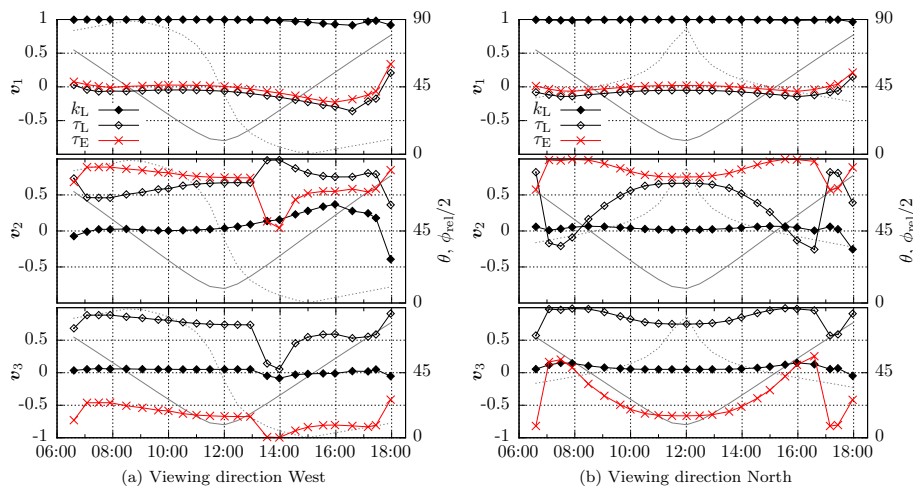
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**Fig. 4.** Singular vectors related to the singular values in Fig. 3 for the ground layer profile with  $\tau \sim 1.2$  (Fig. 2b), again for viewing directions west **(a)** and north **(b)**.

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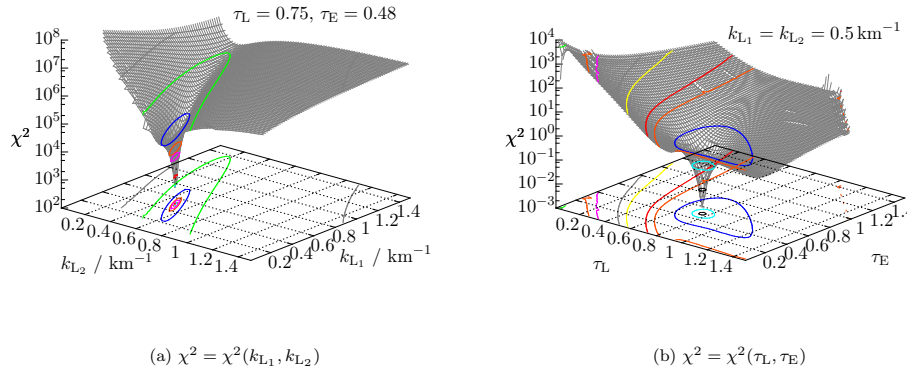
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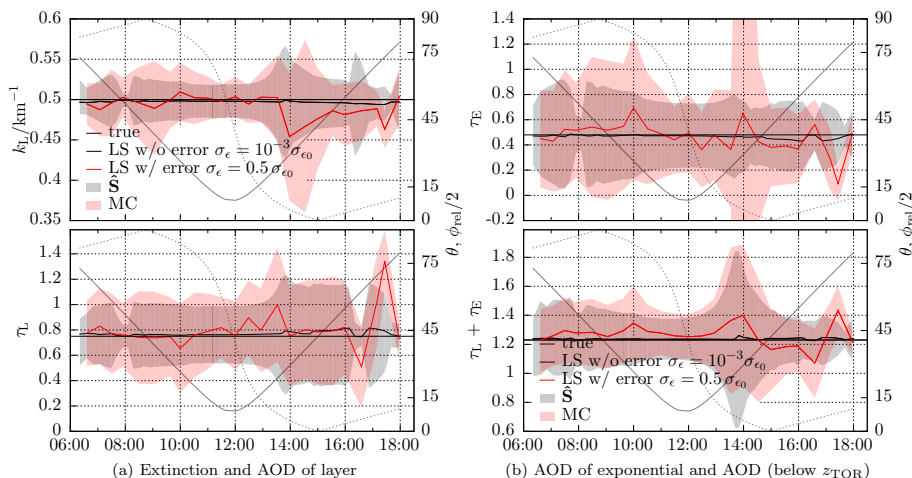




**Fig. 5.** Error-free least squares object function (residual)  $\chi^2(\mathbf{x})$ , Eq. (7), for the well-mixed ground layer with  $\tau \sim 1.2$  (Fig. 2b) represented by the four parameters  $k_{L_1}$ ,  $k_{L_2}$ ,  $\tau_L$  and  $\tau_E$ , Eq. (15). **(a)** For constant  $\tau_L$ ,  $\tau_E$ , **(b)** for constant  $k_{L_1}$ ,  $k_{L_2}$  at the minimum, respectively. The viewing direction is west and  $\theta \sim 10^\circ$ ,  $\phi_{\text{rel}} \sim 170^\circ$ .

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**Fig. 6.** Least squares retrieval (without  $\tilde{I}$ ) of  $x = (k_L, \tau_L, \tau_E)^T$  for the well-mixed ground layer profile with true parameters  $k_L = 0.5 \text{ km}^{-1}$ ,  $\tau_L = 0.75$  and  $\tau_E = 0.48$  (see Fig. 2b) and viewing direction west. The retrieval without errors added to the DSCDs assumes a precision of less than  $10^{-2}\%$ . The retrieval with random noise of about 5% added to the SCDs is the mean value of 100 Monte Carlo (MC) runs, its standard deviation given by the red shaded area. The grey shaded area indicates the corresponding 1- $\sigma$  variation calculated from  $\hat{\mathbf{S}}$ , Eq. (23) (around the error-free retrieval).

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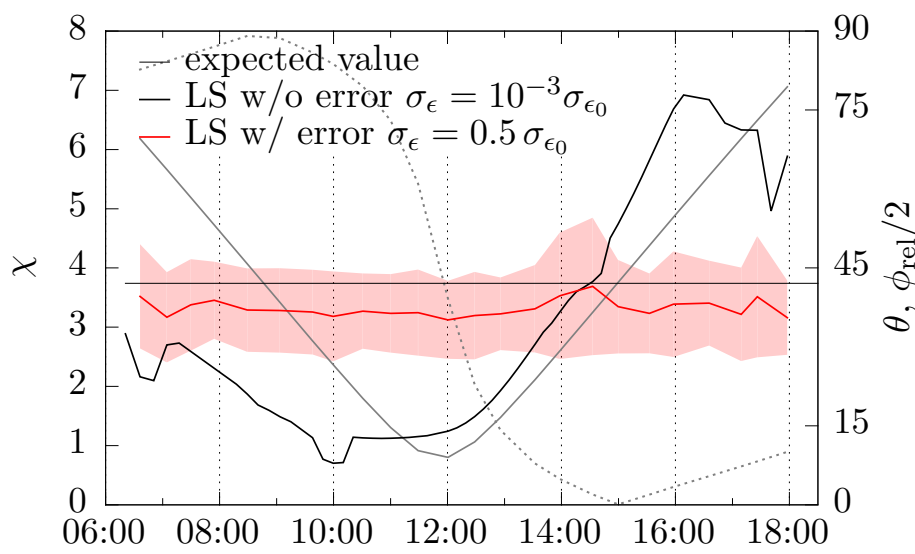
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**Fig. 7.** Residual for the retrieval shown in Fig. 6.

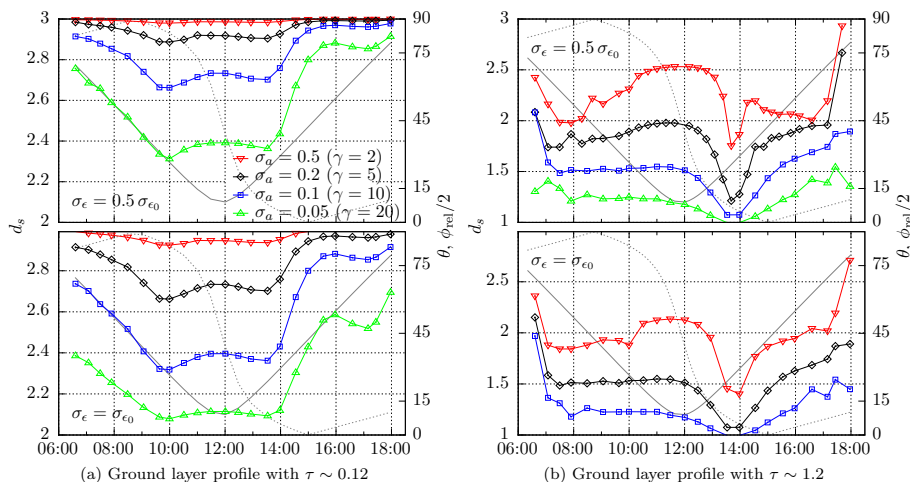
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**Fig. 8.** Degrees of freedom of the signal  $d_s = \text{tr}(\mathbf{A})$  for the regularised least squares retrieval, Eq. (11), of  $\mathbf{x} = (k_L, \tau_L, \tau_E)^T$  for the ground layer profiles with **(a)**  $\tau \sim 0.12$  and **(b)**  $\tau \sim 1.2$  (see Fig. 2) for different  $\sigma_\epsilon$  levels of regularisation and levels of measurement noise (viewing direction west).

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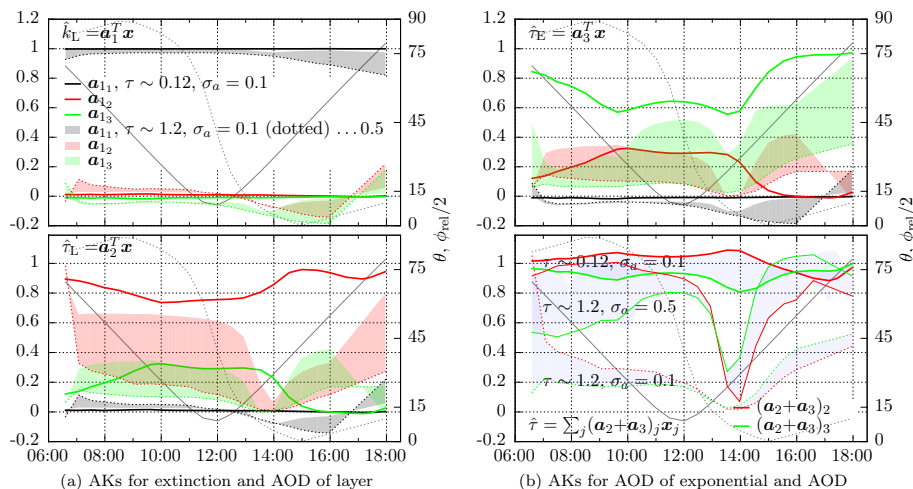
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**Fig. 9.** (a, b) Top: row vectors  $\mathbf{a}_i$  of the averaging kernel (AK) matrix for the regularised least squares retrieval of  $\mathbf{x} = (k_L, \tau_L, \tau_E)^T$  as in Fig. 8. For the ground layer profile with  $\tau \sim 0.12$  (bold lines) only the regularisation parameter  $\sigma_a = 0.1$  is shown. For  $\tau \sim 1.2$  the range of parameters  $\sigma_a = 0.1 \dots 0.5$  is indicated by the shaded areas. (b) Bottom: illustrates how well the AOD  $\tau = \tau_L + \tau_E$  can in principle be reconstructed. Ideally, the contributions from  $\tau_L$  (red) and  $\tau_E$  (green) should be equally one.

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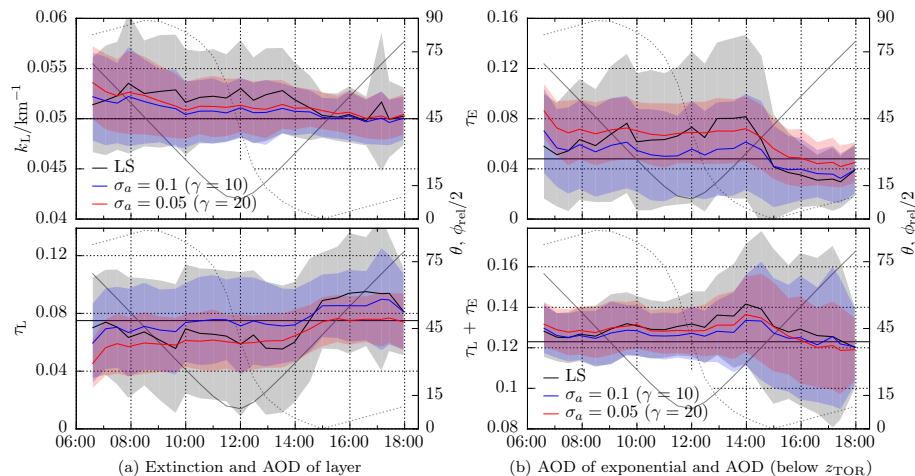
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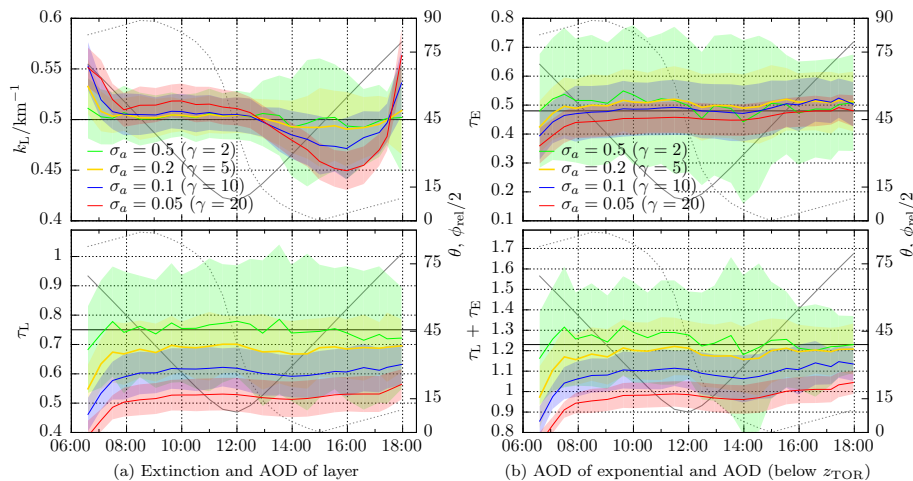




**Fig. 10.** Retrieval of  $x = (k_L, \tau_L, \tau_E)^T$  for the ground layer profile with  $\tau \sim 0.12$  (Fig. 2a), viewing direction west) for the unregularised (LS) and regularised problem and a random measurement error with  $\sigma_e = \sigma_{e_0}$  (Table 1). Lines indicate averages of 100 samples and shaded areas the standard deviations. Compare to Fig. 3a top.

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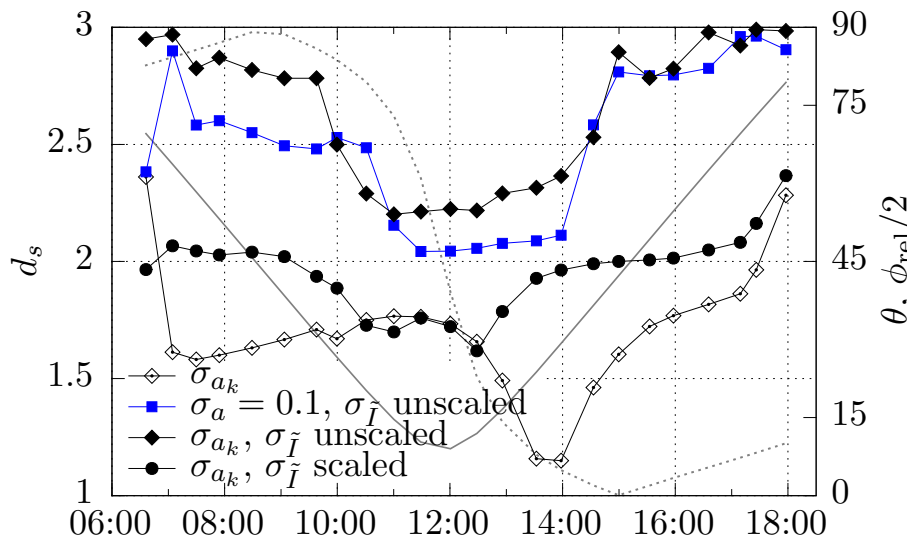




**Fig. 11.** As Fig. 10 for the ground layer profile with  $\tau \sim 1.2$ . See also Figs. 3a bottom and 6.

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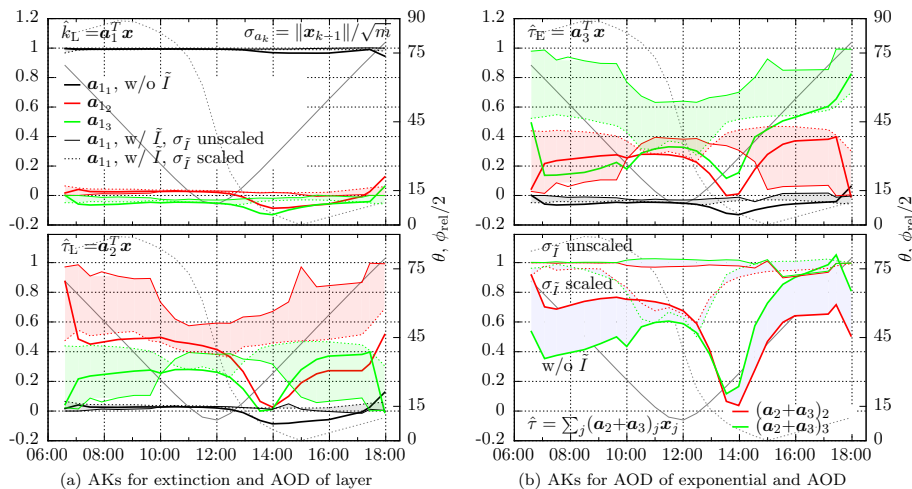
**Fig. 12.** As Fig. 8, here for the ground layer profile with  $\tau \sim 1.2$  and choice of the regularisation parameter as in Eq. (19). Shown are cases without  $\tilde{I}$  (open symbols), with  $\tilde{I}$  and fix  $\sigma_{\tilde{I}}$  (squares and diamonds) and scaled  $\sigma_{\tilde{I}}$  according to Eq. (20) (filled circles).

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**Fig. 13.** As Fig. 9, here for the ground layer profile with  $\tau \sim 1.2$  and choice of the regularisation parameter as in Eq. (19). Shown are the cases without  $\tilde{I}$  (bold lines), with  $\tilde{I}$  and fix  $\sigma_{\tilde{I}}$  (thin lines) and scaled  $\sigma_{\tilde{I}}$  according to Eq. (20) (dashed lines).

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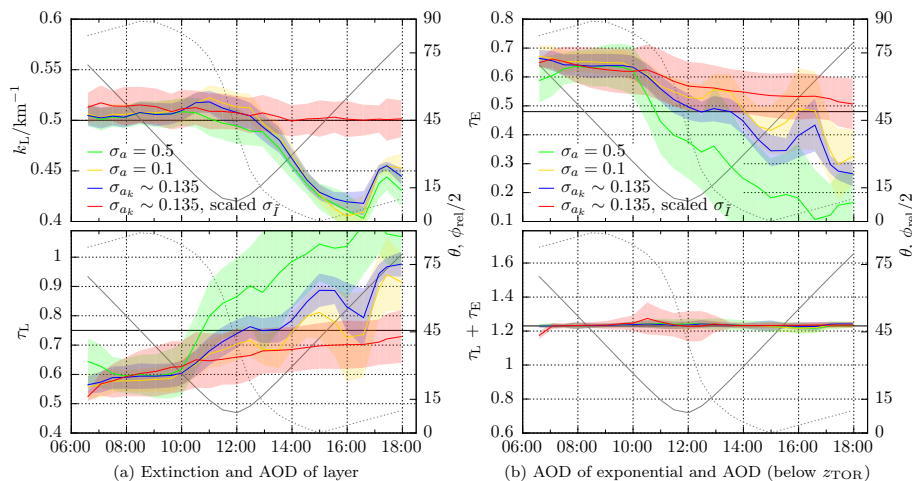
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**Fig. 14.** As Fig. 11, here with relative intensities  $\tilde{I}$ , for different choices of the regularisation parameter (fix and acc. to Eq. 19) and  $\sigma_j$  scaled acc. to Eq. (20) in one case (red line).

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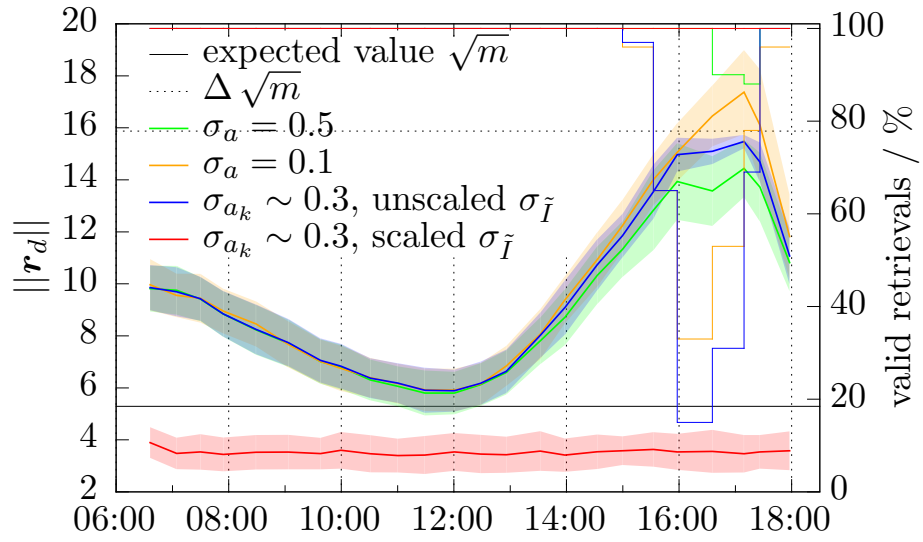
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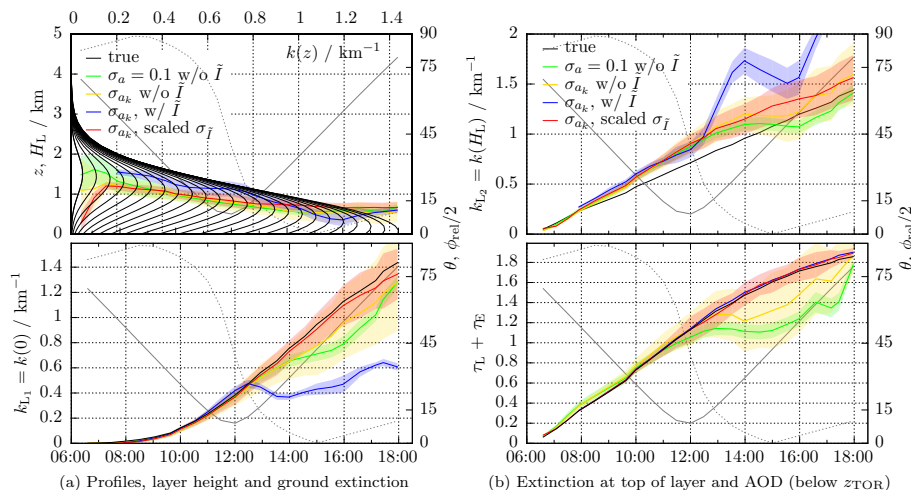


**Fig. 15.** Data residuals as in Fig. 7, here for the regularised cases of Fig. 14. Also shown the percentage of valid retrievals according to Eq. (17).

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**Fig. 16.** Retrieval of  $\mathbf{x} = (k_{L_1}, k_{L_2}, \tau_L, \tau_E)^T$  for the extinction profiles varying in time as shown in (a) top and different choices of the fit quantities, the regularisation parameter and weights for  $\tilde{I}$ . (a) Top: “layer height”  $H_L = 2\tau_L / (k_{L_1} + k_{L_2})$ , bottom: ground extinction  $k_{L_1}$ . (b) Top: extinction  $k_{L_2}$  at height  $H_L$ , bottom: AOD  $\tau = \tau_L + \tau_E$ . Lines show averages of random errors, shaded areas standard deviations as in previous figures. The maximum number (100) of valid retrievals is reached except for the first three profiles retrieved with  $\tilde{I}$  included explicitly into the fit.

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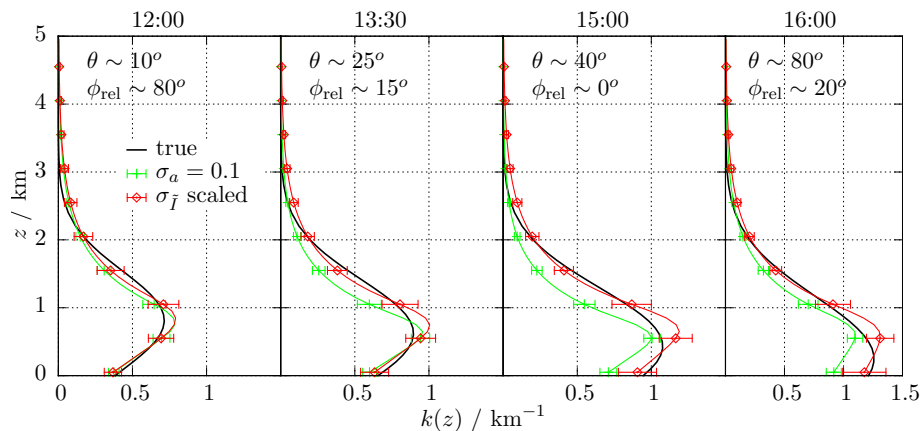


Fig. 17. Example for profiles retrieved as in Fig. 16 for  $\phi_{\text{rel}} < 90^\circ$  (same colours).

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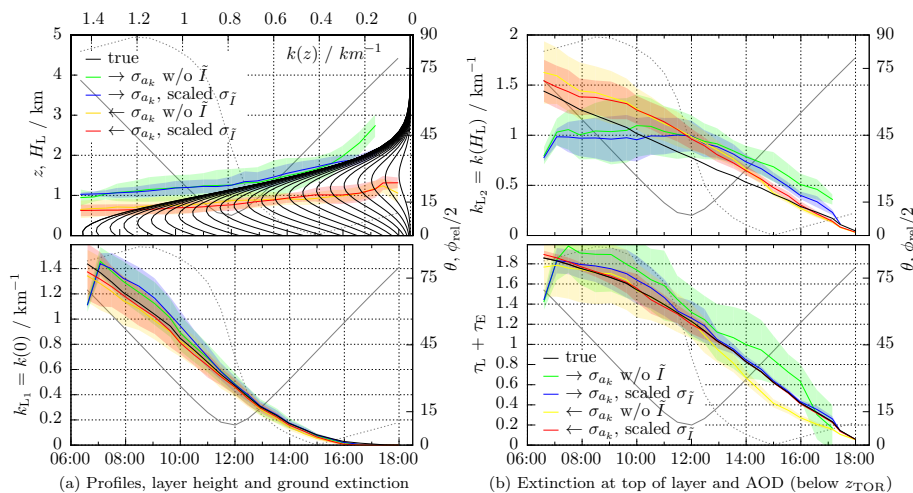
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**Fig. 18.** As Fig. 16, but here with order of profiles reversed in time for retrieval starting with the first ( $\rightarrow$ ) or last ( $\leftarrow$ ) set of DSCDs (and RIs), respectively.

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