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1 *LAI, fAPAR and fCover* CYCLOPES global products derived
2 from VEGETATION. Part 1: Principles of the algorithm

3
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13
14 ***Abstract***

15
16 This article describes the algorithmic principles used to generate *LAI, fAPAR* and *fCover*
17 estimates from VEGETATION observations. These biophysical variables are produced
18 globally at 10 days temporal sampling interval under lat-lon projection at 1/112° spatial
19 resolution. After a brief description of the VEGETATION sensors, radiometric calibration
20 process, based on vicarious desertic targets is first presented. The cloud screening algorithm
21 was then fine tuned using a global network of cloudiness observations. Atmospheric
22 correction is then achieved using the SMAC code with inputs coming from meteorological
23 values of pressure, ozone and water vapour. Aerosol optical thickness is derived from MODIS
24 climatology assuming continental aerosol type. The Roujean BRDF model is then adjusted for
25 red, near infrared and short wave infrared bands used to the remaining cloud free observations
26 collected over a time window of ± 15 days. Outliers due to possible cloud contamination or
27 residual atmospheric correction are iteratively eliminated and prior information is used to get
28 more robust estimates of the three BRDF kernel coefficients. Nadir viewing top of canopy
29 reflectance in the three bands is input to the biophysical algorithm to compute the products at
30 10 days sampling interval. This algorithm is based on training neural networks over
31 SAIL+PROSPECT radiative transfer model simulations for each biophysical variable.

1 Details on the way the training data base was generated and the neural network designed and
2 calibrated are presented. Finally, theoretical performances are discussed. Validation over
3 ground measurement data sets and inter-comparison with other similar biophysical products
4 are presented and discussed in a companion paper. The CYCLOPES products and associated
5 detailed documentation are available at <http://postel.mediasfrance.org>.

6 ***Key words: LAI, fAPAR, fCover,***

1 Introduction

2 Many applications require an exhaustive monitoring of surface characteristics including
3 biogeochemical cycle and climate modelling, resource evaluation (water, agriculture or forest
4 production). Surface process models need a large number of inputs and parameters that depend on
5 plant functional types. These relatively complex models need to be validated, calibrated or ‘corrected’
6 through assimilation of a number of measured variables. Among those, canopy structure
7 characteristics, *LAI* (leaf area index) and *fAPAR* (fraction of photo-synthetically active radiation
8 absorbed by the canopy) are key variables that are both used in surface process models and retrieved
9 from remote sensing observations in the reflective solar domain. Other applications such as change
10 detection in land cover due either to hazards, climatic variation or changes in land use (deforestation,
11 afforestation, reforestation, changes in species or cultural practices), require a relatively high frequency
12 monitoring of the Earth surface. Variables such as the previous ones in addition to *fCover* (vegetation
13 cover fraction) retrieved from remote sensing observations, may offer a promising alternative to
14 vegetation indices that are currently used although they did not relate simply to canopy biophysical
15 variables.

16
17 For almost 25 years, the Earth surface has been monitored at medium to coarse resolution (not better
18 than 1km² pixels) thanks to the series of NOAA/AVHRR sensors. More recently, new sensors have
19 been launched with improved spectral (e.g. SEAWIFS/VEGETATION/MODIS/MERIS) or directional
20 (e.g. POLDER/MISR) sampling, and higher spatial resolution (e.g. MERIS/MODIS). However, space
21 agencies were used to provide only low level products such as top of atmosphere reflectances,
22 sometimes top of canopy reflectances and vegetation indices, but no true biophysical variables such as
23 *LAI*, *fAPAR* or *fCover*. Biophysical products derived from the observations of a selection of these
24 sensors have just been recently developed, and made available to the scientific community. Modellers
25 may now download these products from service centres and plug them directly into their applications.
26 This is the basis of the success of MODIS products.

27
28 The Global Monitoring of the Environment and Security programme was initiated few years ago by
29 Europe to better coordinate the use of satellite Earth observations. This programme is now considered
30 as the European component of GEOSS (Global Earth Observation System of Systems) as defined
31 GEO (Group of Earth Observation). Within GMES, Europe supported projects aiming at the
32 development of biophysical products from currently available sensors. The CYCLOPES project was
33 aiming at the development and actual production of global fields of *LAI*, *fAPAR* and *fCover* for the
34 1998-2003 period. These products were evaluated by associated users for applications focusing on
35 carbon cycle, climate modelling and change detection.

1
2 Although the CYCLOPES project was ultimately targeting to develop products from the fusion of an
3 ensemble of satellites, one of the first tasks achieved was to develop specific products for each series
4 of sensors. This paper present the results obtained for the VEGETATION sensor. After briefly
5 presenting the characteristics of VEGETATION observations, and the way it is radiometrically
6 calibrated, several steps required to derive the biophysical products from the low level products
7 available at VITO (<http://www.vgt.vito.be>) will be described. It includes cloud screening,
8 atmospheric correction, BRDF normalization, temporal compositing, and the derivation of the
9 biophysical variables and some theoretical validation elements. A companion paper will present the
10 actual validation of these products with some inter-comparison with MODIS Collection 4 (Weiss et
11 al., 2007).

12 **2 VEGETATION sensors characteristics and radiometric** 13 **calibration**

14 **2.1 VEGETATION sensors characteristics**

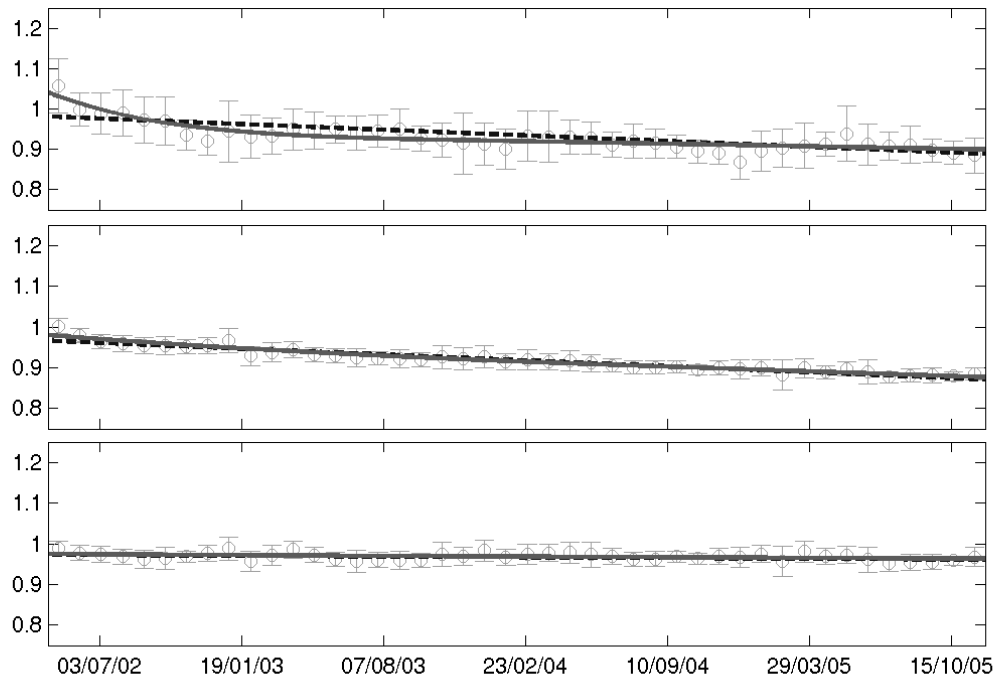
15 Europe has launched the first VEGETATION sensor in 1998 aboard SPOT4. A second instrument was
16 launched in 2002 aboard SPOT5 to ensure the continuity of observations. The two instruments are
17 identical and provide global observations of the surface from a sun-synchronous orbit at 822 km
18 altitude, with an inclination of 96.7°, a period of 26 days and an equatorial crossing time of 10:30.
19 Because of the large swath (101°, equivalent to 2200 km), about 90% of the equatorial areas are
20 imaged each day, the remaining 10% being imaged the next day. For latitudes higher than 35° (North
21 and South), all regions are acquired at least once a day. The instrumental concept relies on a linear
22 array of 1728 CCD detectors providing a spatial resolution around 1.15 km with minimum variations
23 for off-nadir pixel size thanks to the telecentric design of the optics. Four spectral bands are available:
24 B0 (450 nm, $\Delta\lambda=40$ nm); B2 (645 nm, $\Delta\lambda=70$ nm); B3 (835 nm, $\Delta\lambda=110$ nm); SWIR (1165 nm,
25 $\Delta\lambda=170$ nm). The stability of the platform, the accurate knowledge on its position and attitude and
26 post processing of the images allow to achieve a multi-temporal registration accuracy around 200 m
27 (rms). (Sylvander et al., 2003). The system and the corresponding products are described with more
28 details in (Henry, 1999) and (Maisongrande et al., 2004).

29 The P products kindly copied from the archive stored at the VITO processing and archiving centre in
30 Mol (Belgium) were used within this CYCLOPES project. P products are extracts of a segment along
31 a single orbit. P products are geometrically and radiometrically corrected and correspond thus to top of
32 atmosphere reflectance. These products are delivered under the lat-lon projection at 1/112° spatial
33 resolution.

1 **2.2 Radiometric calibration**

2 CYCLOPES project was aimed at processing data from many sensors with consistent algorithms. The
3 first step consists in obtaining a common calibration reference, thus giving priority to a cross
4 calibration of sensors compared to an absolute calibration. Calibration over desert sites is well suited
5 for this purpose because it provides large extent homogeneous areas to minimise misregistration, a
6 temporal stability, low variability of the overlying atmosphere including low cloudiness, and reduced
7 directional effects. Twenty areas, each larger than 100x100km² in North Africa and Saudi Arabia
8 corresponding to these criteria were selected (Cosnefroy et al., 1996). The POLDER instruments
9 offering a very good characterisation of the site Bi-Directional Reflectance Factor (BRDF) in 8
10 spectral bands (443nm to 910nm) were used as the reference sensors. Each site observed by
11 VEGETATION may be also observed by POLDER under very similar directional configuration.
12 POLDER in-flight calibration has been thoroughly studied (Hagolle et al., 1999) and its accuracy is
13 estimated better than 5% (3 sigma). POLDER 2 instrument was thus used here to cross-calibrate other
14 sensors including VEGETATION. More details on the desert cross calibration method can be found in
15 (Cabot et al., 2000; Hagolle et al., 1999).. However, for VEGETATION SWIR band, it is not possible
16 to use POLDER as a reference since it has no spectral band in the SWIR domain. In this case, the
17 reference sensor is VEGETATION 1 (Hagolle et al., 2004). Continuous monitoring of these desartic
18 sites allows updating the original radiometric calibration coefficients by adjusting an empirical model
19 accounting for the degradation of performances of the sensor (Figure 1). Note that a non linear model
20 based on the sum of an exponential model describing the more rapid evolution of the optical system
21 just after launch particularly important in the blue band, and a linear model describing the slow
22 degradation of the system.

1



3

4 **Figure 1.** Time course of the calibration coefficient obtained using 20 desert sites (Cabot et al., 2000; Hagolle et
 5 al., 1999) for VEGETATION 2 using POLDER 2 as a reference: top: Blue band (B0); centre: red band (B2);
 6 bottom: near-infrared band (B3). The calibration coefficient corresponds here to the correction to apply to the
 7 pre-launch calibration of VEGETATION2. Symbols are monthly averages with standard deviation, and curves
 8 are the best fit of a linear model (dashed line) or the non linear model (plain line).

9 **3 Principles of the algorithm**

10 Top of atmosphere calibrated reflectance values were transformed into top of canopy reflectance
 11 values computed in a standard geometrical configuration and for evenly spaced dates. These top of
 12 canopy directionally normalized reflectance values are used as inputs to the biophysical algorithm to
 13 get LAI, fAPAR and fCover estimates. Four steps are therefore necessary to get the biophysical
 14 products. They include chronologically cloud screening, atmospheric correction, BRDF normalization
 15 and temporal compositing and finally biophysical algorithm. These four steps are described hereafter.

16 **3.1 Cloud screening**

17 Cloud screening for VEGETATION as defined by Kempeneers et al. (2000) uses a constant threshold
 18 on blue band (B0) based on Top Of Atmosphere reflectance ($R^{\text{TOC}}(\text{B0})$). In this domain, clouds are
 19 relatively bright because of the importance of scattering, while surfaces are generally dark because of
 20 strong absorption features of many vegetation and soil constituents. Such threshold provides good
 21 results globally, regardless of both surface type and viewing geometry. However, performances are
 22 limited over bright surfaces. Hagolle et al. (2005) developed an alternative cloud screening method
 23 based on surface reflectance computed using the SMAC code (Rahman and Dedieu, 1994) and

1 corrected only from Rayleigh scattering and gaseous absorption. This allows accounting for first order
2 effects in the observation geometry. Because threshold values in B0 may depend on land cover, the 1
3 km spatial resolution GLC2000 global land cover classification was used to provide this piece of
4 information (Bartholomé and Belward, 2005). Specific thresholds were adjusted empirically for
5 barren, sparse herbaceous, artificial surfaces and associated areas, mosaic cropland/shrub cover, and
6 shrub cover classes were tuned: the pixel is flagged as cloudy if ($R^{\text{TOC}}(\text{B0}) > 0.24$) and
7 ($\text{NDVI}^{\text{TOC}} < 0.05$). For the other classes, the algorithm of Hagolle et al. (2005) is applied. To
8 separate snow from clouds a snow detection algorithm was finally implemented based on spectral
9 variation of snow (Dozier, 1989). Thresholds associated to each of the four bands were empirically
10 tuned (Berthelot, 2004). Note that observations partly saturated in B0 and B2 are taken into account
11 for snow detection. Note also that cloud detection is not active over salt lakes.

12
13 Cloud screening was validated at global scale during April 1999 by comparison to quasi-simultaneous
14 cloud fraction observations observed over 11500 synoptic nebulosity ground stations provided by
15 Météo-France (Berthelot et al., 2004). Comparison with CYCLOPES cloud mask shows good
16 performances over Europe, with however some over detection above snow, and under detection in
17 tropical regions, due to the difficulty to identify semi transparent clouds.

18 Once cloudy pixels are detected, cloud shadows are first projected over the surface taking into account
19 view and sun directions, assuming that clouds are at a 5 km altitude. The cloud mask is spatially
20 extended over a 2 x 2 pixels radius to prevent small contamination of clear pixels in the cloud vicinity.
21 Pixels are then declared 'suspect' over a 3 x 3 radius centred on the cloud detected pixels. Note that
22 cloudy pixels are not spatially extended over snowy pixels. Each valid pixel (not cloudy, not
23 shadowed, not saturated or not out of orbit) may thus be declared either clear (all the bands valid),
24 suspect (in the vicinity of cloudy/shadowed pixels), or snowy (with all the bands valid or with just B0
25 and/or B2 bands invalid).

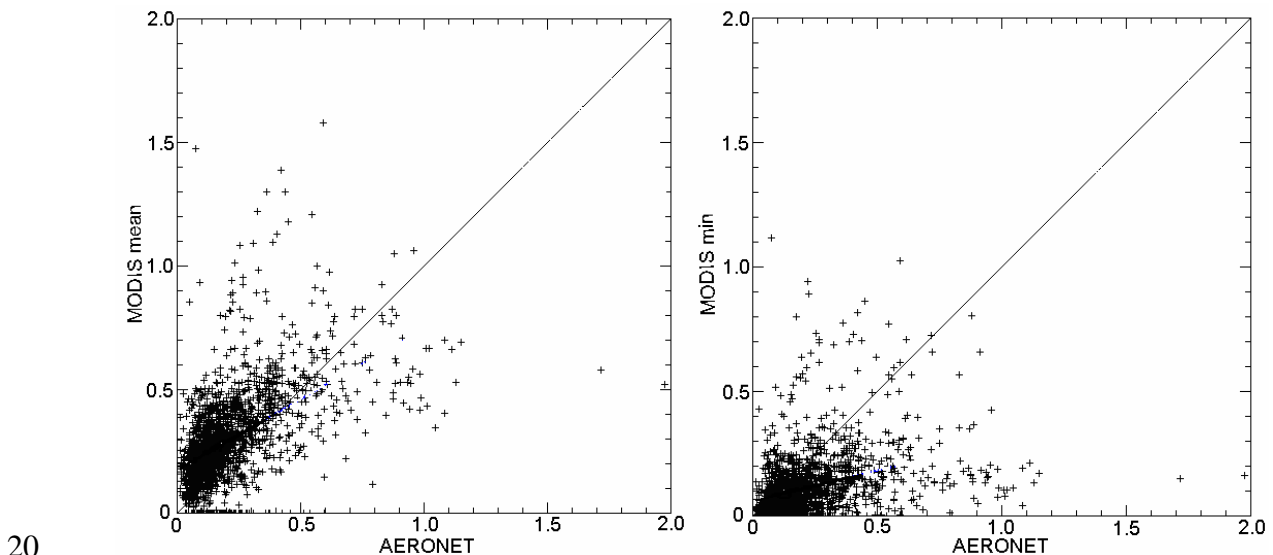
26 **3.2 Atmospheric correction**

27 Atmospheric effects are corrected using the SMAC code (Rahman and Dedieu, 1994) for the four
28 VEGETATION bands. Input atmospheric characteristics are derived from currently available
29 products: NCEP Meteorological data for pressure and water vapour, and TOMS/TOVS for ozone
30 (<ftp://oceans.gsfc.nasa.gov>). With the availability of MODIS aerosol products (Kaufman et al., 1997),
31 the possibility to achieve better atmospheric correction based on monthly mean values (MOD08_M3)
32 was investigated. The climatology was built over the period covering March 2000 up to August 2003,
33 representing 42 months. Performances of the corresponding monthly 1° x 1° spatial resolution MODIS
34 aerosol products were evaluated by comparison with AERONET measurements (Holben et al., 2001)

1 over an ensemble of 250 sites distributed globally. The monthly average of aerosol optical thickness at
2 550nm (AOT_{550}) was computed from Level 2.0 AERONET products.

3

4 Results show significant overestimation of MODIS monthly $1^\circ \times 1^\circ$ spatial average as compared to
5 cloud free AERONET measurements averaged over the month (Figure 2, left). This confirms previous
6 observations achieved over daily aerosol products (Béal et al., 2006; Levy et al., 2005). Inspection of
7 the yearly dynamics of mean, maximum and minimum MODIS monthly AOT_{550} values shows that the
8 minimum of the monthly values was in closer agreement with AERONET monthly AOT_{550} time
9 course. Figure 2 (right) shows that while significant scattering is observed between MODIS and
10 AERONET values, no systematic biases is observed. Sensitivity analysis of atmospheric correction
11 performances achieved by comparing VEGETATION 1 and 2 top of canopy reflectance Hagolle et al.
12 (2005) showed improvements as compared to the mean monthly values or to the simple latitudinal
13 gradient proposed by Berthelot and Dedieu (1997). The minimum of the monthly values was thus
14 proposed to be used for atmospheric correction. However, MODIS aerosol products present missing
15 values and unexpected large AOT_{550} values ($AOT_{550} > 0.6$) that were removed, mostly corresponding to
16 transition between deserts and vegetated areas. The gaps were thus filled with the simple latitudinal
17 gradient proposed by Berthelot and Dedieu (1997). A spatial smoothing procedure is finally applied to
18 minimize the possible discontinuities between $1^\circ \times 1^\circ$ pixels and to reach the original VEGETATION
19 spatial resolution.



21 **Figure 2.** Comparison between MODIS monthly mean (left) and minimum (right) values averaged over $1^\circ \times 1^\circ$
22 areas and the monthly mean values of aerosol optical thickness at 550 nm measured from AERONET sun
23 photometers for clear sky conditions.

24 **3.3 BRDF normalization and temporal compositing**

25 The objective of the directional normalization is to transform a set of reflectance factor measurements
26 taken at irregularly spaced dates under varying view and solar angles into a regular time series of

1 parameters describing the BRDF. The principles behind this algorithm have been already presented in
 2 Hagolle et al. (2005). However, additional improvements have been implemented including a better
 3 weighing of observations, prior constraints on LAI and the reduction in the minimum number of
 4 required observations as well as accounting for snow cover. These improvements of the original
 5 algorithm are presented hereafter.

6 **3.3.1 Reflectance Model**

7 Roujean's model linear BRDF model (Roujean et al., 1992) was used for processing top of canopy
 8 reflectance data

9 $R(\theta_v, \theta_s, \phi)$: It is a where reflectance depends on:

$$10 \quad R(\theta_v, \theta_s, \phi) = k_0 + k_1 f_1(\theta_v, \theta_s, \phi) + k_2 f_2(\theta_v, \theta_s, \phi) \quad (1)$$

12 θ_v and θ_s are the view and sun zenith angles and ϕ the relative azimuth angle between the two
 13 directions and k_i are coefficients to be tuned for each band. Functions f_1 and f_2 are fixed angular kernels
 14 representing the angular distribution related to "geometric" and "volumetric" surface scattering
 15 processes. The technique used to estimate the k_i coefficients is described in detail in Geiger et al.
 16 (Geiger et al., 2006) and is based on a customized linear regression. The main principles are
 17 summarized below.

18 **3.3.2 Weighting reflectance observations**

19 A weight factor w depending on time, view and sun directions is associated to each reflectance
 20 observation collected in the temporal compositing window, and used in the linear regression:

$$21 \quad w = w_\theta(\theta_v, \theta_s) \cdot w_t(t) \quad (2)$$

22 The angular component of the weight, $w_\theta(\theta_v, \theta_s)$ is assumed to be inversely proportional to the
 23 corresponding uncertainties σ_R . Atmospheric correction uncertainties contribute the more to the non-
 24 correlated (random) part of the TOC-reflectance errors and was therefore set proportional to the air-
 25 mass $\eta(\theta_{vj}, \theta_{sj})$ and uncertainties for nadir viewing and illumination geometry, $\sigma_{R(0,0,0)}$.

$$26 \quad w_\theta(\theta_v, \theta_s) = \frac{1}{\sigma_R} = \frac{1}{\eta(\theta_v, \theta_s) \sigma_{R(0,0,0)}} = \frac{2}{\left(\frac{1}{\cos \theta_v} + \frac{1}{\cos \theta_s} \right) \sigma_{R(0,0,0)}} \quad (3)$$

27 As described in (Geiger et al., 2006), the value of $\sigma_{R(0,0,0)}$ can be approximated as a linear function of
 28 reflectance R : $\sigma_{R(0,0,0)} = c_1 + c_2 R$ with coefficients c_1 and c_2 given in Table 1.

29

Bands	c_1	c_2
Blue	0.009	0.14
Red	0.005	0.05
NIR	0.003	0.03
SWIR	0.005	0.03

1 **Table 1.** Values of c_1 and c_2 coefficients to approximate the error $\sigma_{R(0,0,0)}$ for the four VEGETATION bands.

2 The temporal component of the weight $w_i(t)$ was designed to put more emphasis on data collected
3 close to the centre t_0 of the temporal window of width τ (in days) where all observations are collected
4 to adjust the BRDF model. A gaussian function centred on t_0 and truncated at $\pm\tau/2$ is used to compute
5 $w_i(t)$. The standard deviation of the gaussian was selected so that $w_i(t \pm \tau/2) = 0.5$ to provide
6 enough weight at the extremities of the window. A compromise was found between the resulting
7 number of cloud-free observations typically available and the user requirements on the temporal
8 resolution of the products. A value of $\tau=30$ days was chosen, corresponding thus to the temporal
9 resolution. Note that pixels flagged as ‘suspect’ are used in the compositing period.

10 3.3.3 Constraints on k_i coefficients

11 Prior values of k_1 and k_2 coefficients were computed as the average of the estimated values derived by
12 adjusting the BRDF model without constraints over a representative set of VEGETATION data. These
13 values (Table 2) were then introduced as a priori constraints in the model inversion process. Taking
14 into account this prior information of the form $k_i = k_{i\text{ap}} \pm \sigma_{\text{ap}}[k_i]$ considerably reduces the
15 uncertainties in k_i estimates when the directional sampling achieved with the available observations is
16 limited. It also increases the efficiency of the outlier elimination scheme described below. Note that a
17 minimum of two valid observations in the compositing window was required.

18

Bands	$k_{1\text{ap}}$	$\sigma_{\text{ap}}[k_1]$	$k_{2\text{ap}}$	$\sigma_{\text{ap}}[k_2]$
Blue (B0)	0.00	0.07	0.08	0.29
Red (B2)	0.02	0.05	0.17	0.30
NIR (B3)	0.04	0.07	0.67	0.34
SWIR	0.05	0.06	0.41	0.28

19 **Table 2:** A priori information on k_1 and k_2 coefficients of Roujean’s BRDF model for the four bands
20 of VEGETATION instrument bands.

21 3.3.4 Iterative elimination of outliers

22 Undetected clouds can introduce non-Gaussian outliers in the error structure of reflectance and
23 decrease the performances of the inversion process. An iterative scheme for eliminating unreliable
24 observations as described in Hagolle et al. (2005) was therefore implemented. Detection of outliers is
25 carried out by means of the reflectance factor values in the blue channel which is the most affected by

1 residual clouds. Those outlier dates are then eliminated in all bands. The decision criterion for
 2 elimination is based on the “relative error”, σ_{rel} defined as:

$$3 \quad \sigma_{rel} = \sqrt{\frac{\sum_{t=1}^n (\hat{R}(t) - R(t))^2}{\sum_{t=1}^n (\bar{R} - R(t))^2}} \quad (4)$$

4 Where $\hat{R}(t)$ is the estimated reflectance for observation date t , and \bar{R} is the reflectance value
 5 averaged over the n available observations considered within the temporal window. If σ_{rel} computed
 6 after the first iteration exceeds the threshold value $\sigma_{thr} = 0.25$ then all observations exceeding the
 7 model fit by more than $\sqrt{\frac{1}{n} \sum_{t=1}^n (\hat{R}(t) - R(t))^2}$ (1 standard deviation) are eliminated.

8 Otherwise and if $\sigma_{rel} > 0.5 \sigma_{thr}$, all observations deviating from the model fit by more than
 9 $2\sqrt{\frac{1}{n} \sum_{t=1}^n (\hat{R}(t) - R(t))^2}$ (2 standard deviation) are eliminated. The model fit is then repeated and the relative
 10 error σ_{rel} is recalculated. The procedure is iterated up to four times as long as $\sigma_{rel} > 0.5 \sigma_{thr}$ and the
 11 number of eliminated observations does not exceed one third of the initially available ones. Note that
 12 temporal and angular weighing is applied only for the last iteration to save computation time. Note
 13 that the threshold values were empirically derived and additional tests have shown that the results
 14 were not too sensitive to their precise values which explains the simple numbers used (0.25 and 0.5).

15 **3.3.5 Snow pixels**

16 Reflectance properties of land surfaces can be drastically changed over small time scales by snow fall
 17 and melt. To prevent instabilities in the inversion due to reflectances with inconsistent values, it has to
 18 be decided whether the window is considered mainly with or without snow. If less than half of the
 19 total number of reflectance observations during a reduced period of ten days centred on t_0 is flagged
 20 as snow-covered, then all observations over the full compositing period flagged as snowy are
 21 eliminated and the inversion is performed only with the snow-free observations. Conversely, when the
 22 majority of observations are flagged as ‘snow’, the snow-free reflectances are eliminated and the
 23 BRDF model adjustment is applied only over snow observations. If the number of snow-free equals
 24 that of snow-covered observations, the full composition period is used for the decision criterion.

25 **3.4 Biophysical algorithms**

26 The last three steps allow getting relatively clean top of canopy reflectance values in a standard
 27 geometrical configuration (nadir viewing, sun zenith angle corresponding to the median value in the
 28 compositing period) at regular time steps. These will be used as input to the biophysical algorithm to

1 estimate the corresponding *LAI*, *fAPAR* and *fCover* biophysical variables. Before describing the
2 biophysical algorithms, a clear definition of CYCLOPES biophysical products must be given:

- 3 • ***LAI***: the CYCLOPES products correspond to some effective *LAI* where only clumping at the
4 landscape level is accounted for: pixels are considered made of patches of different surface
5 types. This alteration of the main definition (half the developed area of green vegetation
6 elements per unit horizontal soil) is due to the radiative transfer models and the embedded
7 assumptions on canopy architecture considered as explained later.
- 8 • ***fAPAR***: The *fAPAR* CYCLOPES products correspond to a black-sky value (no diffuse
9 illumination) at 10:00 local solar time. It is a fair approximation of the daily integrated value
10 for non cloudy days (Baret et al., 2004).
- 11 • ***fCover*** refers only to the green vegetation elements.

12 The biophysical algorithm is based on radiative transfer model inversion. A neural network approach
13 was selected for many reasons: neural networks are known to be computationally very efficient, which
14 is very important for operational applications over a long time series of global data. A recent review
15 on canopy characteristics retrieval methods (Baret and Buis, 2007) showed that neural networks when
16 trained over radiative transfer model simulations provides good estimation performances because of
17 their efficient interpolation capacity (Leshno et al., 1993). This model inversion technique has been
18 applied with success to remote sensing of land surfaces since about a decade (Abuelgasim et al., 1998;
19 Atkinson and Tatnall, 1997; Baret et al., 1995; Danson et al., 2003; Fang and Liang, 2005; Gong et al.,
20 1999; Kimes et al., 1998; Smith, 1992, 1993; Weiss et al., 2002). Recently, several authors developed
21 operational biophysical products for medium resolution sensors, with top of canopy reflectance as
22 inputs: (Lacaze, 2005) for POLDER, (Bacour et al., 2006) for MERIS.

23
24 Algorithm calibration consists in generating a learning database from which the network architecture
25 and the synaptic coefficients and bias are tuned. The radiative transfer model used, distribution of the
26 input variables will first be described. Calibration of the network as well as management of extreme
27 cases will then be presented. Note that our algorithm is designed to be applied to any surface type
28 conversely to what is implemented for MODIS (Myneni et al., 2002) or GLOBCARBON (Plummer et
29 al., 2005) products for which the algorithm is tuned for each biome type.

30 **3.4.1 Generating the learning data base**

31 Radiative transfer models were run to simulate the actual VEGETATION observations. The SAIL
32 model (Verhoef, 1984, 1985) with the hot-spot correction implemented by (Kuusk, 1991) was used. It
33 assumes the canopy as a turbid medium for which leaves are randomly distributed in space. Canopy
34 structure is characterized by *LAI*, the average leaf angle inclination (ALA) assuming an ellipsoidal
35 distribution (Campbell, 1990) and the hot-spot parameter. To account for clumping at the landscape
36 scale, each pixel was supposed to be made of a fraction $\nu Cover$ of pure vegetation and $(1-\nu Cover)$ of

1 pure bare soil. The extensive biophysical variables (LAI , $fAPAR$, $fCover$) and reflectances were thus
2 computed at the pixel level according to this assumption.

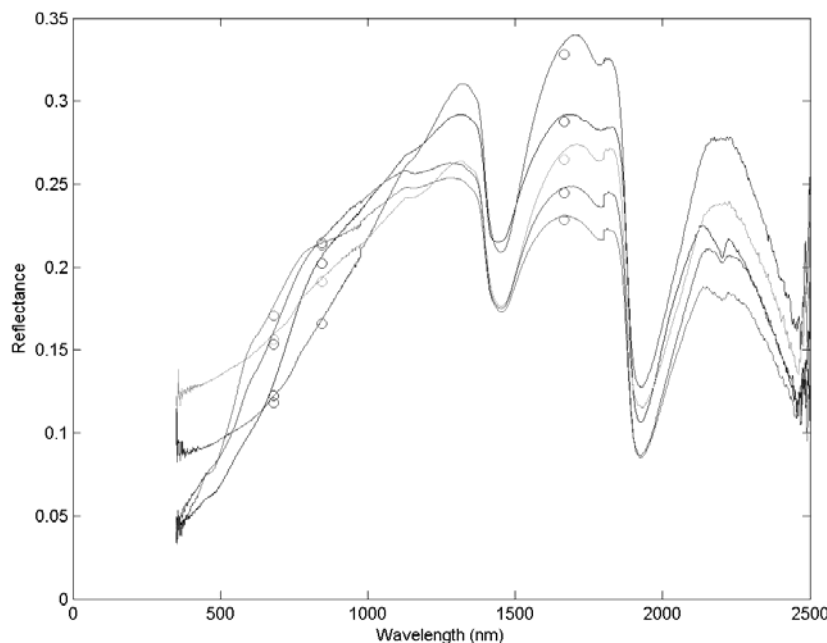
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4 Leaf optical properties are simulated with the PROSPECT model (Jacquemoud and Baret, 1990) using
5 specific absorption coefficients calibrated by (Fourty and Baret, 1997). It requires 5 input variables:
6 the mesophyll structure parameter (N), chlorophyll (Cab), dry matter (Cdm), water (Cw) and brown
7 pigment (Cbp) contents. Water content was tied to the dry matter content assuming that green leaves
8 have a relative water content (H) varying within a relatively small range (Table 3).

9

10 Reflectance values of soils were simulated using 5 typical soil reflectance spectra multiplied by a
11 brightness coefficient B_s . These 5 reflectance spectra (Figure 3) were selected to represent with a good
12 accuracy within VEGETATION bands (RMSE<0.01) the reflectance spectra of a large soil database
13 where soil types, roughness, moisture and observational geometry vary over a large range (Liu et al.,
14 2003). The distribution of the brightness coefficient observed over this data base was used in the
15 generation of the training database. Note that here again selection of the soil type and brightness
16 coefficient is independent on any prior knowledge of the surface type.

17



18

19 **Figure 3.** Reflectance of the 5 soils used to represent the possible range of spectral shapes.

20

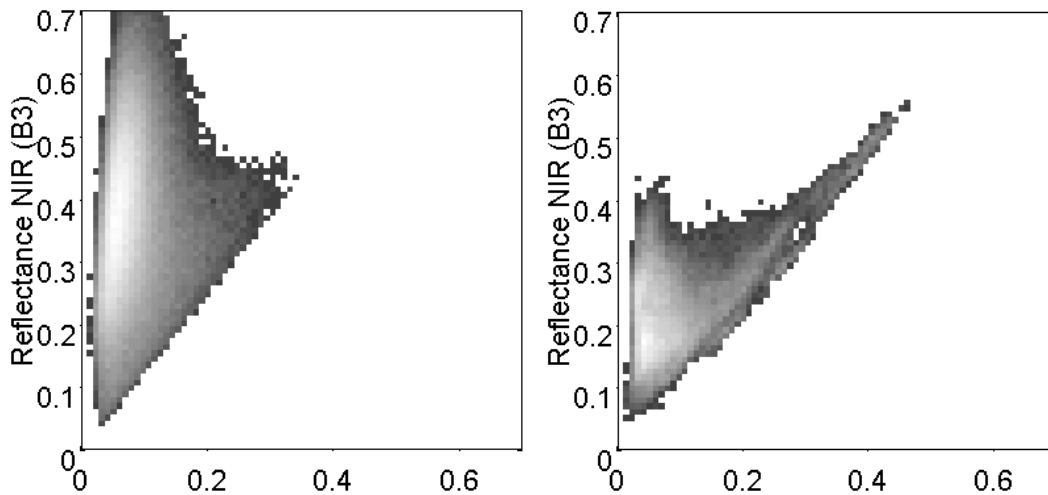
21 Table 3 shows the distribution of all the input variables. Note that for LAI, a uniform distribution law
22 was used in order to put more emphasis on the larger LAI values that are more difficult to retrieve.
23 Truncated gaussian or uniform distributions were used for the other variables, trying to describe the
24 actual distribution over the Earth. Observation geometry was derived from computing orbit dynamics
25 of VEGETATION assuming uniform distribution of sites in latitude and longitude as well as date of

1 observations (Table 3). A full orthogonal experimental plan was used to combine the 13 input
 2 variables by splitting the whole range of variation of each variable into a small number of classes. This
 3 resulted to 196608 cases simulated.

	Input variables	Unit	min	mean	max	std	law	Class
Observation geometry	Latitude	°	-60°	0°	60°	-	Uniform	1
	Longitude	°	0°	180°	360°	-	Uniform	1
	Day of year	day	1	92	183	-	Uniform	1
Canopy structure	LAI	-	0	3	6	-	Uniform	6
	ALA	°	30°	60°	80°	20°	Gauss	4
	Hot	-	0.001	0.1	1	0.3	Gauss	1
	vCover	-	0	1	1	0.2	Gauss	2
Leaf optical properties	N	-	1	1.5	2.5	1	Gauss	4
	Cab	$\mu\text{g}\cdot\text{cm}^{-2}$	30	50	90	30	Gauss	4
	Cdm	$\text{g}\cdot\text{cm}^{-2}$	0.002	0.0075	0.02	0.0075	Gauss	4
	H	-	0.65	0.75	0.85	-	Uniform	4
	Cbp	-	0	0	1.5	0.2	Gauss	4
Background	Bs	-	0.2	1	2.2	0.7	Gauss	4

4 **Table 3.** Distribution of the input variables of the radiative transfer models used to generate the training database
 5

6 For any combination of the input variables, top of canopy reflectance was computed for each
 7 wavelength and then integrated according to the spectral sensitivity of VEGETATION bands. The
 8 blue band (B0) was not used here because it was considered too much contaminated by residual
 9 atmospheric effects. A white gaussian noise of 0.04 (absolute value) was added to the simulated
 10 reflectances to account for uncertainties attached to measurements and models. It is quite difficult to
 11 quantify precisely the uncertainties attached to top of canopy reflectances used as input to the
 12 biophysical algorithm since they are induced by several possible processes including radiometric
 13 calibration, BRDF normalization, atmospheric correction, residual cloud contamination and of course
 14 the suitability of the radiative transfer models to describe photon transport in actual canopies. In
 15 addition uncertainties have some structure that is very difficult to evaluate that may however impact
 16 largely results. The uncertainties level selected for all the three bands corresponds to the order of
 17 magnitude of radiometric calibration (about 3 to 5 %), BRDF normalization (lower than 0.02
 18 absolute), and 0.01 up to 0.10 for atmospheric corrections depending on assumptions and bands (Beal
 19 2006). However, complementary sensitivity analysis was conducted showing a relatively small effect
 20 of the uncertainty level (in the range 0.00-0.05) on theoretical performances.



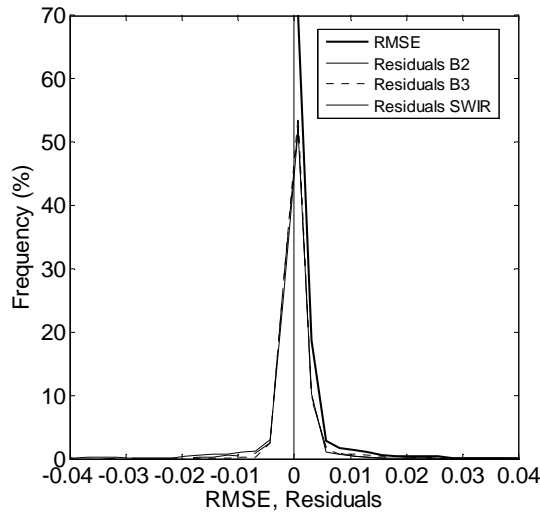
1
 2 **Figure 4.** On the left, density plots of simulated reflectance values in red (B2) and near infrared (B3) feature
 3 plan (196608 cases). On the right, the same figure with actual VEGETATION data acquired over a subset of
 4 196608 pixels randomly selected over the BELMANIP sites for years 1999-2003 and for which water bodies,
 5 cloud and snow contaminated pixels were eliminated as far as possible.

6
 7 First evaluation of the simulations was achieved by simply plotting red and near infrared reflectance
 8 values. It shows (Figure 4, left) the typical pattern as expected, with the soil line bounding the bottom
 9 of the scatter plot. The realism of the simulated reflectance in the training database was further
 10 evaluated by comparison over a database of the same number of cases representing the possible range
 11 of variation of land surface properties: 196608 pixels were randomly selected in an ensemble of sites
 12 (BELMANIP, (Baret et al., 2006)) over the 1999-2003 period. Results show that the distribution of
 13 simulated cases (Figure 4, left) differs from that of actual VEGETATION observations (Figure 4,
 14 right): more cases are concentrated for low reflectance values in the red band, and higher near infrared
 15 reflectance values were simulated as compared to actual observations. These differences are due
 16 assumptions on the distribution and co-distribution of the variables in the training data base: for
 17 example, LAI distribution over the Earth is not uniform, and higher LAI values might not be observed
 18 with a planophile leaf inclination distribution (small ALA values), although these cases were
 19 simulated in a systematic way assuming the independency between LAI and ALA. Better agreement
 20 between simulated and actual distribution of reflectance data might improve the algorithm
 21 performances. Further efforts will have to be directed towards this issue.

22
 23 However, more importantly is the realism with which reflectance data are simulated as compared with
 24 actual observations. The mismatch, δ , between each VEGETATION measurements and the closest
 25 simulated reflectance was computed according to equation [5] over the ensemble Ω (196608 cases) of
 26 measured (ρ_b) values for the three VEGETATION bands (b) considered here.

$$\delta = \min_{\Omega} \left(\sqrt{\frac{\sum_{b=1}^3 (\rho_b - \hat{\rho}_b)^2}{3}} \right) \quad (5)$$

2 This was repeated for all the VEGETATION measurements collected previously (196608 pixels).
 3 Figure 5 shows that most observations may be matched very closely with a simulated case. Note that
 4 the assumed uncertainties level (0.04 RMSE value) is larger than the RMSE value associated to the
 5 reflectance mismatch. This is mainly explained by the fact that model uncertainties are only
 6 marginally described by the mismatch test that concentrates on reflectance values independently on
 7 the input canopy characteristics. Distribution of the residuals for each of the three bands confirms the
 8 consistency between simulated and observed reflectance values.



9
 10 **Figure 5.** Distribution of the mismatch values (δ) as computed by equation [5] over a database made of 2000
 11 VEGETATION observations.

12
 13 Finally, $fAPAR$ values for θ_s at 10:00 solar time and $fCover$ were also computed by running the
 14 radiative transfer code in the forward direction with the same input variables as those used for
 15 reflectance simulations.

16 3.4.2 The networks

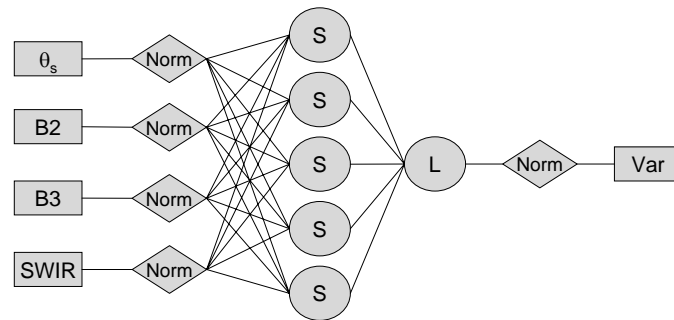
17 Inputs of the network are made of the median value of the sun zenith angle observed during the
 18 compositing period and top of canopy nadir reflectance in the three VEGETATION bands B2, B3 and
 19 SWIR. Outputs are the biophysical variables: LAI, $fAPAR$ and $fCover$. The learning data base made
 20 of pairs of inputs and outputs was split in three parts by randomly selecting cases: half of the cases
 21 were used to train the network, a quarter to test the hyper-specialization during the training process.
 22 The last quarter was used for theoretical validation. Input and output (X) data were first normalized

1 (X_{Norm}) according to: $X_{Norm} = 2 * (X - X_{Min}) / (X_{Max} - X_{Min}) - 1$ where X_{Max} and X_{Min} are respectively the
2 maximum and minimum values observed in the data base for variable X .

3

4 For each biophysical variable, a dedicated back-propagation artificial neural network (Rummelhart et
5 al., 1986) was trained. After several tests, this solution was preferred to that consisting in a single
6 network with three outputs as proposed by (Bacour et al., 2006). An optimal architecture was selected
7 for each biophysical variable which converged towards the one depicted by Figure 6. For each
8 variable, five networks were trained with different initial solutions, and the one providing the best
9 performances (over the validation data set) was selected. The Levenberg-Marquardt minimization
10 algorithm was used in the learning process because of its convergence performances (Ngia and
11 Sjoberg, 2000). Note that 26 synaptic and 6 biases coefficients needed to be adjusted, which
12 corresponded to about 3000 training data for one coefficient to tune which was far above the minimum
13 value advised by (Harrel, 2001).

14



15

16 **Figure 6.** Architecture of the neural network used to estimate LAI , $fAPAR$ and $fCover$ (here labelled Var) from
17 VEGETATION normalized top of canopy reflectance in bands B2, B3 and SWIR. Sun zenith angle (θ_s)
18 corresponding to the median value during the compositing period is also used as input. 'Norm' corresponds to
19 the normalization of the inputs and outputs. Neurones with 'S' and 'L' mean respectively tangent-sigmoid and
20 linear transfer functions.

21 3.4.3 Theoretical performances

22 The theoretical performances were evaluated over the validation data set made of 25% (37850 cases)
23 of the simulated cases extracted from the database. It shows gradually improved performances (Figure
24 7) from LAI (RMSE=1.10; RRMSE=37%), $fAPAR$ (RMSE=0.08; RRMSE=12%) and $fCover$
25 (RMSE=0.06; RRMSE=10%). The poorest performances observed for LAI are mainly due to reduced
26 sensitivity of reflectances to LAI for the larger LAI values. Estimated LAI values always keep lower
27 than $LAI=5.5$, which is probably due to the regularization of the inversion process achieved with the
28 distribution of LAI values in the training data base as noticed by (Baret and Buis, 2007).

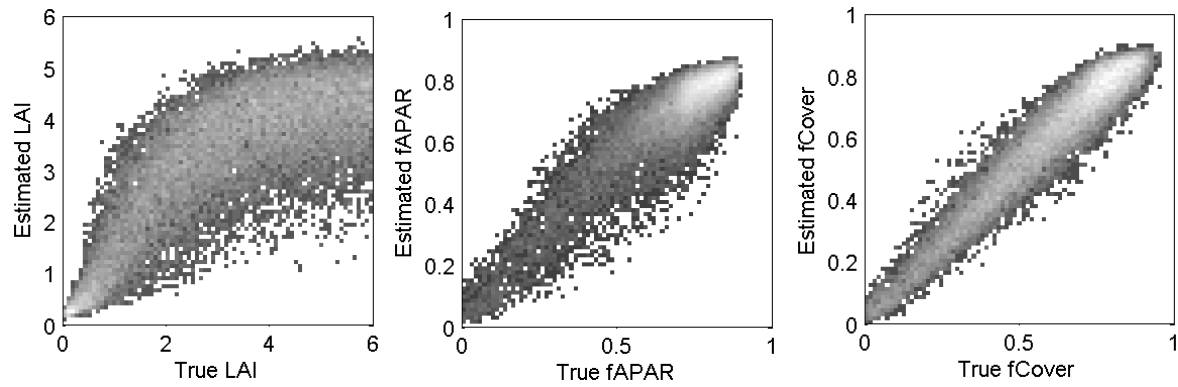


Figure 7. Density plots showing the theoretical performances evaluated over the validation data set.

A theoretical error model was adjusted based on the residuals computed over the validation data set (Figure 8). It shows an increase of the uncertainties up to *LAI* values close to 3.0, and then a stabilization showing however an underestimation of the actual *LAI* values. This behaviour results from the non linearity of the *LAI*-reflectance relationship which saturates for the larger *LAI* values, i.e. reflectance is little sensitivite to *LAI* variation. In the mean time, some other variables such as canopy structure (*ALA*, *hot*) and leaf characteristics (C_{ab} in the red band, C_w in the SWIR band, leaf mesophyll structure parameter, *N*) still impact reflectance values. The effect of these variables may be confounded with *LAI* effects, increasing the uncertainties attached to *LAI* retrievals. However, some regularization due to the distribution of *LAI* values in the training database tends to make the solution closer to the mode of *LAI* distribution, resulting in an underestimation of *LAI* for the larger *LAI* values.

For *fAPAR* and *fCover*, uncertainties are relatively stable over the whole dynamic range.

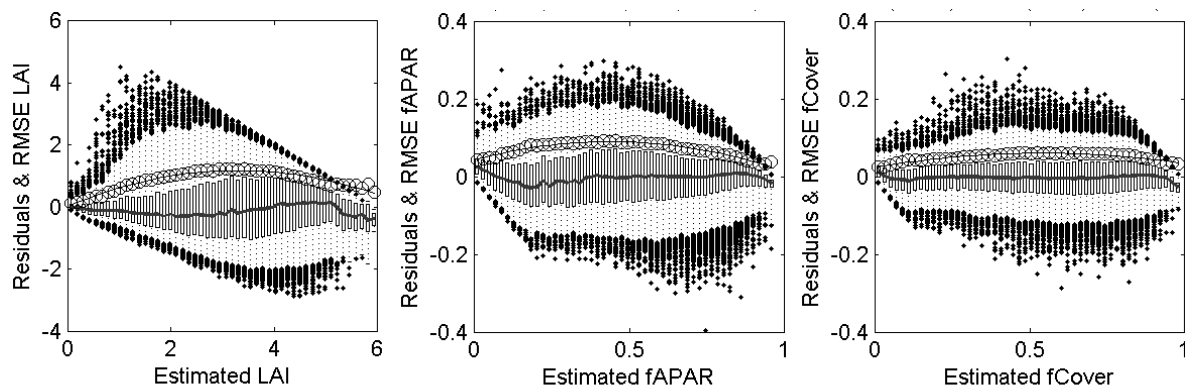


Figure 8. Distribution of the residuals (bars correspond to 25%, dots to individual values below 1% or over 99%). The solid line corresponds to the median value. The circles represent the RMSE values and thin solid line the corresponding uncertainty model.

3.4.4 Management of extreme cases

Minimum and maximum values in the training data base are used to declare ‘out of range’ the cases for which reflectance values in one band is outside this definition domain:

1
2 If $X < X_{min}$ or $X > X_{max}$, then X is declared 'out of range'

3
4 Table 4 provides the values used to declare the inputs 'out of range'. Management of estimated values
5 (Y) falling outside the definition range of the considered biophysical variable was achieved similarly
6 as previously, except that a tolerance value (Y_{tol}) is used to keep slightly negative values by resetting
7 them to Y_{min} :

8
9 If $Y_{tol} < Y < Y_{min}$, then Y is set to $Y = Y_{min}$

10 If $Y < Y_{tol}$ or $Y > Y_{max}$, then Y is declared 'out of range'

11
12 The use of the tolerance value was introduced to account for the uncertainties of the algorithm for the
13 lower amount of vegetation possibly due to approximations in the description of soil reflectance
14 properties. Table 4 provides the values of Y_{tol} , Y_{min} and Y_{max} for the three output biophysical products.
15 Note that LAI , $fAPAR$ and $fCover$ are flagged as 'invalid' over snow and set to 0.00 over lakes.

16

Inputs	$B2$	$B3$	$B4$	Outputs	LAI	$fAPAR$	$fCover$
X_{min}	0.000	0.000	0.000	Y_{min}	0.00	0.00	0.00
X_{max}	0.612	0.916	0.835	Y_{max}	6.00	1.00	1.00
				Y_{tol}	-0.20	-0.10	-0.10

17 **Table 4.** On the left, X_{min} and X_{max} reflectance values used to declare the inputs 'out of range' in each of the three
18 bands. On the right, Y_{min} , Y_{max} and Y_{tol} values used for LAI , $fAPAR$ and $fCover$ to declare the outputs 'out of
19 range'.

20 4 Conclusion

21 CYCLOPES processing chains have been developed under a flexible and modular environment based
22 on an open architecture (C++ code, CORBA integration, Java interfaces, PostgreSQL database back-
23 end). Using object-oriented components eases the integration of new algorithms while providing a
24 reliable production framework. The actual processing is operated by 8 linux workstations (Intel 32
25 Architecture). The processing chains have been carefully implemented and allow producing about one
26 year of global data within 27 days. The first processing steps, particularly those regarding cloud
27 masking and atmospheric correction takes about 18 days, while BRDF normalization and compositing
28 require about 8 days. The last step corresponding to the biophysical algorithm starting from the
29 normalized reflectances to generate LAI , $fAPAR$ and $fCover$ products requires only 1day processing
30 per year of global data. This allows relatively fast reprocessing of this last step. This is achieved
31 thanks to the efficiency of the neural network approach that offers great potentials for this type of
32 application.

1 CYCLOPES products derived from VEGETATION sensors are available freely at the following web
2 site: <http://postel.mediasfrance.org/>. The period 1998-2003 is currently available, and the series will
3 soon be extended up to 2006 making 9 years of continuous and consistent observations. Products are
4 organized by tiles of 10° zenith and 10° azimuth. One year of global data including the original *P*
5 products, intermediate products and the final *LAI*, *fAPAR* and *fCover* products representing 1.1 Tb.

6
7 The current version 3 of CYCLOPES products have been validated by comparison to ground
8 measurements and inter-comparison to other products. The corresponding results are presented in a
9 companion paper (Weiss et al., 2007). Further, complementary validation activities were conducted
10 (Garrigues et al., 2007), showing the inter-comparison with additional products as well as
11 investigating the spatial consistency of the products.

12
13 The first steps of the processing seem relatively mature although some tuning may still improve the
14 performances. Cloud masking could be improved by enhanced adjustment of the threshold values for
15 each surface type. Additionally, distinction between snow, bare surfaces, water bodies and clouds may
16 be also useful for the biophysical algorithms. Atmospheric correction could be improved by getting a
17 better climatology of aerosol optical thickness and types. However, departure from the climatology
18 due to events such as fires or sand storms, could only be corrected by developing autonomous
19 correction methods providing that enough information is carried in the radiometric signal. The
20 compositing and BRDF normalization method could also be improved with emphasis on gap filling to
21 get continuous fields of variables as desired by most users. Kalman filtering as proposed by (Samain et
22 al., 2007) may offer a very attractive solution. Nevertheless, largest improvements may come from the
23 biophysical algorithm itself. The turbid medium canopy radiative transfer model does not describe leaf
24 clumping, except at the landscape level through the *vCover* variable. Alternative models accounting
25 for clumping at the tree and shoot levels could probably improve the results and allow retrieving
26 higher values of *LAI*. However, it is not yet clear whether the measured radiometric signal carries
27 enough and specific information on these additional canopy structure features. A prior classification
28 may be used to get the required information to account for leaf clumping. Major enhancement of the
29 algorithm is expected to come from a better design of the training data base. This should incorporate
30 more realistic description of measurements and model uncertainties, particularly regarding the
31 structure of these uncertainties, as well as distribution of the variables more representative of the
32 actual ones. Further work is currently conducted in this direction, combining actual satellite
33 observations and radiative transfer models to generate a better training data base.

34
35
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