Contrasting impacts of humidity on the ozonolysis of monoterpenes: insights into the multi-generation chemical mechanism

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9 Abstract. Secondary organic aerosol (SOA) formed from the ozonolysis of biogenic monoterpenes is a 10 major source of atmospheric organic aerosol. It has been previously found that relative humidity (RH) 11 can influence the SOA formation from some monoterpenes, yet most studies only observed the increase 12 or decrease in SOA yield without further explanations of molecular-level mechanisms. In this study, we 13 chose two structurally different monoterpenes (limonene with an endocyclic double bond and an 14 exocyclic double bond, Δ^3 -carene with only an endocyclic double bond) to investigate the effect of RH 15 in a set of oxidation flow reactor experiments. We find contrasting impacts of RH on the SOA formation: 16 limonene SOA yield increases by ~100% as RH increases, while there is a slight decrease in Δ^3 -carene SOA yield. By analyzing SOA chemical composition and reaction mechanisms, the enhancement in 17 18 limonene SOA yield can be attributed to the water-influenced reactions after ozone attack on the 19 exocyclic double bond of limonene, which leads to the increment of lower volatile organic compounds 20 under high RH condition. However, as Δ^3 -carene only has an endocyclic double bond, it cannot undergo 21 such reactions. This hypothesis is further proved by the SOA yield enhancement of β -caryophyllene, a 22 sesquiterpene that also has an exocyclic double bond. These results greatly improve our understanding 23 of how water vapor influences the ozonolysis of biogenic organic compounds and subsequent SOA 24 formation processes.

25 1 Introduction

Secondary organic aerosol (SOA), as an important type of ambient fine particulate matter (PM_{2.5}:
aerosols with aerodynamic diameter ≤ 2.5 μm) (Guo et al., 2014; Huang et al., 2014), has caused a series
of negative impacts on human health (Pye et al., 2021), air quality (Zhang et al., 2016) and global climate
(Levy et al., 2013). SOA produced from the oxidation of biogenic volatile organic compounds (BVOCs)
is a major component of SOA in heavy forest regions during summer (Sindelarova et al., 2014; Ahmadov

et al., 2012), and contributes by a large fraction (~40%-80%) to global OA budget (Cholakian et al.,
2019).

33 Monoterpenes, mostly emitted from coniferous trees, account for $\sim 11\%$ in total BVOCs 34 (Sindelarova et al., 2014; Kanakidou et al., 2005). Limonene is one of the most abundant monoterpenes, 35 with the annual emission budget of 11.4 Tg yr⁻¹ (Guenther et al., 2012). Apart from the biogenic source, 36 limonene can also be released from the indoor emission, mainly from essential oils (Ravichandran et al., 37 2018; De Matos et al., 2019; Mot et al., 2022). Limonene has an endocyclic double bond and an exocyclic 38 double bond, and is thus more reactive than other monoterpenes towards oxidants such as ozone (O_3) , 39 hydroxyl radical (OH), and nitrate radical (NO₃) (Chen and Hopke, 2010; Atkinson and Arey, 2003). Δ³-40 carene is another kind of monoterpene that dominates the monoterpene emission from Scots pine (Bäck 41 et al., 2012). Different from limonene, Δ^3 -carene contains only one endocyclic double bond, which is 42 similar to most other monoterpenes.

Ozonolysis is an important reaction pathway for limonene and Δ^3 -carene. Although reactions with 43 44 OH and NO₃ are faster than that with O₃ for both two monoterpenes (Atkinson, 1991; Khamaganov and 45 Hites, 2001; Chen et al., 2015; Shaw et al., 2018), the atmospheric concentration of the latter 46 monoterpene is much higher than that of the former (Sbai and Farida, 2019). The contributions of O₃-47 reactions with limonene and Δ^3 -carene to tropospheric degradation are 47% and 24%, respectively, in 48 the daytime (Ziemann and Atkinson, 2012). In pristine areas where NO₃ concentration is very low, 49 ozonolysis is also the dominant fate for limonene and Δ^3 -carene in the nighttime. In addition, it has been 50 previously found that the ozonolysis of monoterpenes can produce more extremely low volatility 51 products than OH-initiated oxidation, which contributes by a large fraction to the SOA production 52 (Jokinen et al., 2015). For either limonene or Δ^3 -carene, the first step for ozonolysis is attacking on the 53 endocyclic double bond to form two types of stabilized Criegee intermediates (sCI) with low energy (Fig. 54 S1) (Drozd and Donahue, 2011; Chen et al., 2019). The sCI will then trigger a series of chemical reactions, 55 like isomerization, decomposition and addition reactions. Correspondingly, the major components in Δ^3 -56 carene SOA are caric acid, OH-caronic acid, and caronic acid (Ma et al., 2009; Thomsen et al., 2021), 57 while the major components from limonene SOA are limonaldehyde, keto-limonon aldehyde, limononic 58 acid and keto-limononic acid (Pathak et al., 2012; Wang and Wang, 2021).

59 Water is ubiquitous in the atmosphere and can affect the formation mechanism of SOA and its

60 relevant physical and chemical properties (Sun et al., 2013). A number of field measurements have shown 61 that the average molecular weight of the water/organic phase and activity coefficient of condensed 62 organics would be changed due to the change of relative humidity (RH) (Seinfeld et al., 2001; Li et al., 63 2020). In addition, several laboratory studies have demonstrated that RH can influence the ozonolysis of 64 monoterpenes in different ways. Most of those studies have reported either an inhibitory effect or a 65 negligible effect of high RH on the particle formation (Bonn and Moortgat, 2002; Fick et al., 2002; Zhao 66 et al., 2021; Ye et al., 2018). Nevertheless, few other studies found that high RH can promote SOA 67 formation from the ozonolysis of limonene (Yu et al., 2011; Gong et al., 2018; Xu et al., 2021), but the 68 reason of this promotion effect remains unclear.

69 To fully examine the effects of water on SOA formation from the ozonolysis of monoterpenes, 70 especially the related chemical processes, we used an oxidation flow reactor (OFR) to investigate the 71 ozonolysis of limonene and Δ^3 -carene under different RH conditions in this study. An ultra-high 72 performance liquid chromatography with a quadrupole time-of-flight mass spectrometer (UPLC-Q-TOF-73 MS) was deployed to analyze the molecular chemical composition of the SOA, which provided insights 74 into the physical and chemical processes influenced by the water content. With these state-of-the-art 75 techniques, we proposed mechanisms that may explain the inhibitory or enhancing RH effects on SOA 76 formation for different monoterpenes.

77 **2 Experimental methods**

78 **2.1 Oxidation flow reactor experiments**

79 A series of dark ozonolysis experiments of limonene and Δ^3 -carene were conducted in a custom-80 made oxidation flow reactor (OFR). The OFR is a 602 mm long stainless cylinder with a volume of 2.5 81 L (Fig. S2) (Liu et al., 2019; Liu et al., 2014). A zero-air generator (XHZ2000B, Xianhe, China) was 82 used to generate dry clean air as the carrier gas for the OFR. As shown in Fig. S2, there are four gas paths 83 upstream of the OFR: the first path is the precursor gas channel through which monoterpenes are injected 84 via a syringe pump (ISPLab 01, Shenchen, China); the second path is for the flow of 300 sccm dry zero 85 air passing through a mercury lamp ($\lambda = 185$ nm) to generate O₃; the third path is connected to a water 86 bubbler to generate wet air; the fourth path is the extra dry zero air entering the OFR. The RH in the OFR 87 was controlled by adjusting the ratio of the wet and dry zero air flows. A water recycle system was

88 equipped to keep the temperature (T) around at 298 K. The total flow was 0.9 L min⁻¹, resulting in an 89 average residence time of 167 s. The RH and T in the OFR were monitored by a T/RH Sensor (HM40, 90 VAISALA, Finland). The concentration of ozone and the consumption of the precursor gas were 91 measured with an ozone monitor (Model 106L, 2B Technologies, USA) and a gas chromatography with 92 flame ionization detector (GC-FID 7890B, Agilent Technologies, USA), respectively. The GC was 93 equipped with a DB-624 column (30 m \times 0.32 mm, 1.8 μ m film thickness) whose temperature was set 94 to ramp from 100 °C to 180 °C at a rate of 20 °C min⁻¹, and then held at 180 °C for 2 min. Before each 95 experiment, O₃ was introduced into the OFR to clean it until the background aerosol mass concentration 96 reached $< 1 \ \mu g \ m^{-3}$.

97 The experimental conditions are shown in Table 1. In these OFR experiments, the precursor 98 (limonene or Δ^3 -carene) concentration was set to ~320-340 ppb. A high O₃ concentration of ~6 ppm was 99 used to realize an equivalent aging time of 0.41 day in the real atmosphere, assuming an average ambient 100 O₃ concentration of 28 ppb (Sbai and Farida, 2019) (see Section S1 for the calculation). Under such 101 conditions, most of the precursors were consumed, since the residence time was almost five and three 102 times of the half-life for limonene and Δ^3 -carene, respectively. Correspondingly, the O₃ consumption for 103 limonene and $\Delta 3$ -carene were ~250 ppb and ~100 ppb, respectively. A series of RH conditions ranging 104 from dry (1-2%) to 60% with a step of $\sim 10\%$ were used to investigate the effects of water content on 105 SOA production and composition (see Table 1). All materials used in the experiments have been 106 described in Section S2.

107 **Table 1.** Experimental conditions and results.

Exp.	[Precursor] (ppb)	[O]3 (ppm)	T (K)	RH (%)	N(13.8-723.4 nm) ^a (cm ⁻³)	M _(13.8-723.4 nm) ^b (μg m ⁻³)	D _(mean) c (nm)	SOA yield (%)
_				liı	monene			
1	321±39	5.7	298	1–2	6.9×10 ⁵	980.9	138.2	62.9
2	321±39	6.0	298	10±2	1.3×10^{6}	1377.5	126.8	88.4
3	321±39	5.9	298	20±2	9.0×10 ⁵	1573.3	150.9	90.2
4	321±39	5.9	298	30±2	1.4×10^{6}	1573.3	128.9	100.9
5	321±39	6.0	298	40±2	1.7×10^{6}	2051.4	130. 7	131.6

6	321±39	5.5	298	50±2	1.5×10^{6}	1962.7	137.8	125.9
7	321±39	5.5	298	60±2	1.5×10^{6}	2211.1	139.0	141.8
				Δ^3 -	carene			
8	341±28	6.1	298	1–2	9.5×10 ⁴	346.0	195.8	19.4
9	341±28	6.4	298	10±2	1.4×10 ⁵	300.3	163.4	16.8
10	341±28	6.4	298	20±2	9.4×10 ⁴	244.9	176.9	13.7
11	341±28	6.0	298	30±2	5.9×10^{4}	241.2	205.1	13.5
12	341±28	6.3	298	40±2	4.6×10 ⁴	205.8	203.2	11.5
13	341±28	6.3	298	50±2	6.8×10^{4}	196.7	180.7	11.0
14	341±28	6.3	298	60±2	5.6×10 ⁴	198.5	190.2	11.1

111 **2.2 SOA particle analysis**

112 **2.2.1 SOA yield**

113 The SOA particle size distribution was measured with a scanning mobility particle sizer (SMPS), which consists of a differential mobility analyzer (DMA) (model 3082, TSI Inc., USA) and a 114 condensation particle counter (CPC) (model 3776, TSI Inc., USA). The samples were measured by SMPS 115 116 every 5 minutes with a sampling flow and a sheath flow of 0.3 L min⁻¹ and 3 L min⁻¹, respectively. The 117 SOA mass concentration was calculated from the volume concentration measured with SMPS and the aerosol density, which was estimated to be 1.25 cm⁻³ for limonene- and 1.09 g cm⁻³ for Δ^3 -carene-SOA 118 119 (Thomsen et al., 2021; Watne et al., 2017). 120 The SOA yield (Y) for individual organic gas can be calculated as: $Y = \frac{\Delta M}{\Delta HC}$ 121 122 Where ΔM is the total mass concentration of SOA, ΔHC is the mass concentration of reacted precursor 123 (Ng et al., 2007; Odum et al., 1996).

124 2.2.2 Ultra-high performance liquid chromatography quadrupole time-of-flight mass spectrometry
 125 analysis

126 An ultra-high performance liquid chromatography (UPLC, UltiMate 3000, Thermo Scientific)

127 coupled with a quadrupole time-of-flight mass spectrometry (Q-TOFMS, Bruker Impact HD) was used 128 to analyze the molecular-level chemical composition of SOA. First, the SOA particles were collected on 129 the PTFE filters (47 mm diameter, 0.22 μ m pore size, Jinteng, China). Next, these filters were dissolved 130 and extracted by 5 mL methanol for two times. Extracts were then filtered through PTFE syringe filters 131 (0.22 μ m pore size), and were concentrated to near dryness by nitrogen-blowing. At last, the samples 132 were redissolved in a 200 μ L solution with 0.1% (v/v) formic acid in 50:50 methanol/ultrapure water 133 mixture.

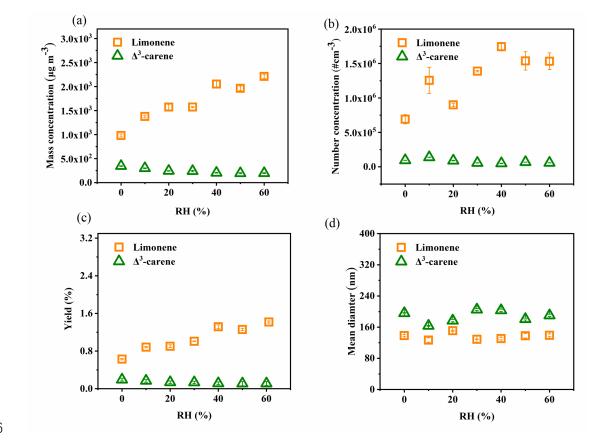
134 The parameters of LC-MS were set as follows: capillary voltage 4000 V, nebulizer pressure 0.4 bar, 135 dry heater temperature 200°C, end plate voltage -500 V, and flow of dry gas 4 L min⁻¹. A C₁₈ column 136 (100 Å, 3 mm particle size, 2.1 mm×50 mm, Waters, USA) was used with a column temperature of 35°C. The mobile phase was 0.1 % formic acid in methanol (A) and 0.1 % formic acid in ultra-high purity 137 138 water (B) with a flow of 200 µL min⁻¹. The injection volume was 5 µL. The MS was operated in negative 139 ion mode, and the detection molecular weight range was from m/z 50 to 1500. The temperature ramp 140 program was: 0-3min with 0%-3% phase B, 3-25min with 3%-50% phase B, 25-43min with 50%-90% 141 phase B, 43-48 min with 90%-3% phase B, 48-60min with 3% phase B.

142 **3 Results and discussion**

143 **3.1 SOA production under different RH conditions**

144 SOA formation of a representative experiment is shown in Fig. S3. It is found that the formed SOA 145 are mainly in the size range of 60-200 nm, and the number concentration and mass concentration are 146 relatively stable during the course of the OFR experiment. SOA formation from limonene and Δ^3 -carene 147 in terms of particle number concentration, particle mass concentration, and SOA yield as a function of 148 RH are illustrated in Fig. 1a-c. We find that all the above-mentioned 3 parameters of limonene-SOA 149 increase with the increasing RH. The increment of particle mass concentration and SOA yield from the 150 ozonolysis of limonene is ~100% higher at wet (60% RH) than at dry conditions. In contrast, SOA 151 formation from Δ^3 -carene is suppressed by ~40% under high RH. The distinct effects of RH on SOA 152 formation from the ozonolysis of limonene and Δ^3 -carene found in this study agree with most previous 153 studies (Yu et al., 2011; Jonsson et al., 2006b; Bonn et al., 2002; Gong and Chen, 2021; Li et al., 2019b). 154 As shown in Table 2, Yu et al. (2011) reported a positive correlation between SOA production and RH 155 for the ozonolysis of limonene in the chamber experiments without OH scavenger. Their experimental 156 condition is similar to that in our study regarding the absence of OH scavenger and, thus, similar results 157 were observed. However, in the presence of OH scavenger, results are quite different. Jonsson et al. (2006) 158 observed a similar enhancement effect of high RH on SOA production with 2-butanol as the OH scavenger, while Bonn et al. (2002) found a negligible or suppressive effect with cyclohexane as the OH 159 160 scavenger. It should be noted that the OH scavenger not only has the ability to scavenge OH but also 161 produces additional products which may influence the reactions of target precursors. For example, there 162 is no difference between 2-butanol and cyclohexane in the scavenging ability of OH radical, though 2-163 butanol will produce more HO₂ than cyclohexane and, consequently, $R \cdot$ will react with HO₂ to produce 164 more hydroxyl acids and hydroxyl per-acid products, most of which have low volatility and, thus high 165 partitioning into the particle phase. According to previous studies, the influence of different OH 166 scavengers can vary (Jonsson et al., 2008). This may explain the different findings with and without OH 167 scavenger for limonene-SOA. With regard to Δ^3 -carene, similar results are found in the absence of OH 168 scavenger, namely, high RH has negligible or slightly suppressive effect on SOA production (Bonn et al., 169 2002; Fick et al., 2002). Same as limonene, the presence of OH scavenger and its different chemical 170 nature can explain the different results found previously (Jonsson et al., 2006a; Bonn et al., 2002).

171 The enhancement in limonene-SOA production under high RH can be due to several reasons from 172 either physical or chemical processes. First, the hygroscopic growth of the particles (i.e., absorption of 173 water content) can lead to higher mass concentration under higher RH, but the enhancement should be 174 at most ~30% as the growth factor (GF, the ratio of wet and dry diameter: Dwet/Ddry) of limonene-SOA is 175 \leq 1.1 (Varutbangkul et al., 2006). However, we do not observe an obvious change in the mean diameter 176 when comparing dry and wet conditions (Fig. 1d). In addition, hygroscopic growth should also occur for 177 Δ^3 -carene SOA, but no obvious enhancement in particle mass is observed (Fig. 1a). Therefore, it is 178 suggested that physical processes regarding hygroscopic growth play a minor role in the enhancement in 179 limonene-SOA under high RH. As a consequence, we believe that chemical processes are likely the 180 reason of the enhancement in limonene-SOA under high RH. Water can influence chemical processes in 181 the gas phase or in the particle phase. Particle-phase reactions can promote the growth of small particles 182 and, thus, mainly lead to larger particle sizes. As the observed SOA enhancement is mainly from high 183 number concentration particles rather than the large size particles (Fig. 1b and 1d), it is likely that the 184 water-participated gas-phase reactions are the most possible reasons for the limonene-SOA enhancement.



185 The reaction mechanism is analyzed below based on the mass spectra information on the SOA.

186

187 **Figure 1.** The effect of RH on the SOA formation: (a) number concentration, (b) mass concentration, (c)

188 SOA yield, (d) mean diameter.

189

190				Table 2. Com	parison with	previous st	Table 2. Comparison with previous studies on the effect of RH.	f RH.			
	Precursor	O ₃ concentration		щ			SOA Mass				
Precursor	concentration	(qdd)	Reactor	5	T (K)	RH (%)	Concentration	SOA Yield (%)	\mathbf{M}^{a}	Nb	Reference
	(dqq)			scavenger			(µg/m³)				
	000	1000	flow			0.02 and	NT MC	NT MC	100 JUC	ى ب	
	1000		reactor	cyclonexane	7±067	32.5	- IMI'NI			1	BOIII et al. (2002)
		100±5				18±2,		$7.0 \pm 0.7;$			
	320		chamber	N.M.°	296±2	50±3	24; 58; 120	$17.4\pm 1.3;$	+ ^d (7 times)	+ ^d (8 times)	Yu et al. (2011)
						and 82±2		$53.4{\pm}1.9$			
	15 1 20	430.9	flow			30 C /	2.7-10.5 and 62-	6.8-26.4 and	р -	- -	
limonene	US DUB CI		reactor	2-DULANOI	298±0.4	CQ-7 >	229	77.4-285.7	₽ +	8 +	Jonsson et al. (2000)
		endocyclic					endocyclic	endocyclic	exocyclic		
	CIROCYCIIC	(270) and	flow		000	10 50	(~11) and	(~ 7.4) and	(+ ^d) and	NT MG	Gong and Chen
	(C 21) allu	exocyclic (12200)	reactor	Z-DUIA1101	067	00-01	exocyclic (22-	exocyclic (23.8-	endocyclic	TAT.VI	(2021)
							51)	55.3)	(-c)		
	1085	900±10	flow	euvu	208	<i>د</i> ع-د	150-200-210	MN	р+	ပျ	[i et a] (2010)
	C001		reactor		0/7	70-0	1.70, 200, 210	TATINT	-	I	LI U al. (2017)

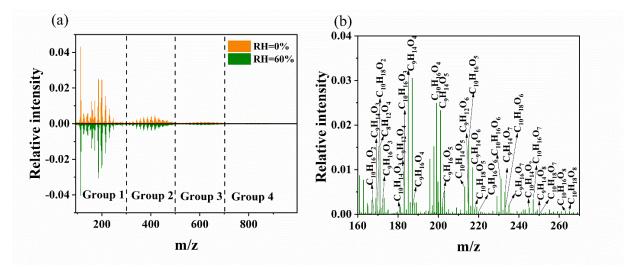
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	371+30	5786±203	flow	enon	208	0-60	080 0_711 1	67 0-141 8	() times)	+ ^d (3 times)	this study
	10-170		reactor		0		1.1177-0.000	0.111-0.70			
	000	1000	flow		C - 20C	0.02 and	N.M ^c	N.M°	7 - 3 0	e	
	1000		reactor	cyclonexane	77067	32.5			no ellect)	Bonn et al. (2002)
		2300	flow	100000			0.78-3.8 and	2.1-10.1and	-	- -	
~	14.2 and 29.4		reactor	2-Dultanol	298±0.4	CQ-7 >	15.3-94;	19.8-116.7	a ┿	; +	Jonsson et al. (2000)
∆°-carene	1111	$900{\pm}10$	flow		000	<i>C7 C</i>	75. 80. 00		٥	٥	
	1111		reactor	попе	067	70-0	10; 00; 20	IN.NI	1	1	LI EI al. (2019)
	0C - 17 C	6257±140	flow		000	02.0	216 0 108 E	1 11 7 01	ں ا		
	041±∠0		reactor	попе	067	00-0	0.040.0-1.20.0	19.4-11.1	l		uns stuay
191	^a M means the	^a M means the change trend total particle mass concentration. ^b N means total particle number concentration. ^c N.M. means not mentioned. ^d Positive sign (+) means the	particle mass o	concentration. ^b]	N means tota	l particle nui	mber concentration.	° N.M. means not	mentioned. ^d Pc	sitive sign (+) n	neans the

mass or number concentration increases with RH. ^e Negative sign (-) means the mass or number concentration decreases with RH.

193 **3.2 Molecular analysis of SOA particles**

194 The UPLC/ESI-Q-TOF-MS was used to examine the SOA molecular composition under high and 195 low RH conditions. As shown in Fig. 2a, the mass spectra of limonene-SOA are divided into four groups: 196 monomeric group (<m/z 300), dimeric group (m/z 300-500), trimeric group (m/z 500-700), and 197 tetrameric group (m/z 700-1000), corresponding to products containing one, two, three, and four 198 oxygenated limonene units, respectively (Bateman et al., 2009). Most of the SOA molecules are 199 monomers (>60%) (Fig. 2b) and dimers (~25%), while trimers and tetramers contribute to very small 200 fractions (<10% and ~3%) (Table S1). Correspondingly, the distribution of Δ^3 -carene-SOA can be 201 divided into four groups (Fig. S4), comparable to that of limonene-SOA. Most of the SOA molecules are 202 monomers (\sim 70%) and dimers (\sim 25%), while trimers and tetramers contribute to smaller proportions (\sim 2%) 203 and <1%, respectively) (Table S2). Although the SOA mass concentration increases by $\sim 100\%$ under 204 high RH condition, the relative intensities of MS peaks do not significantly change with varying RH 205 conditions. In other words, we did not observe an obvious change in the overall MS patterns, and the 206 fractions of the four groups only slightly differed under different RH conditions, e.g., the fraction of 207 monomers was 62% under dry condition and 66% under wet conditions. However, if we take a closer 208 look, the intensities and contributions of specific peaks are quite different with varying RH. For example, 209 the relative intensity of $C_{10}H_{16}O_2$, a possible first-generation product (Gong et al., 2018), decreases by 210 \sim 20% with increasing RH from dry to 60% (Table S3). This is likely due to the multi-generation reactions 211 influenced by water vapor concentration, as discussed below with the proposed reaction mechanism of 212 limonene ozonolysis.



213

214 Figure 2. UPLC/ (-) ESI-Q-TOF-MS mass spectra of SOA from limonene ozonolysis. (a) MS under

215 high and low RH conditions; (b) the identification of monomers under high RH condition.

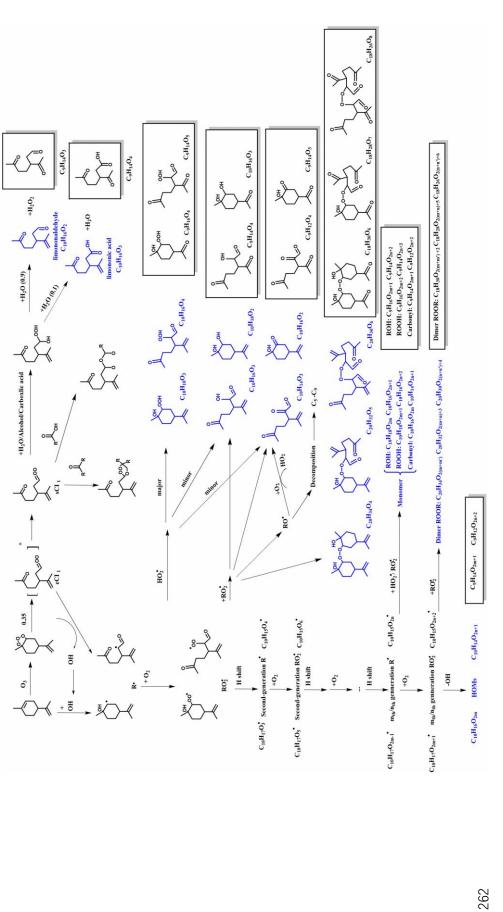
216 The proposed reaction mechanism of limonene ozonolysis is shown in Fig. 3 and Fig. 4. The initial 217 step in the reaction of O_3 with limonene is the attack of the endocyclic double bond to form eCI_1 and 218 eCI_2 (with branching ratios of 0.35 and 0.65, respectively). In the context of eCI_1 , several complex 219 reactions occur, with the most dominant reaction being the generation of hydroxyl radicals (OH) and a 220 reaction pathway known as sCI₁. The sCI₁ pathway can proceed through three distinct reactions, as 221 depicted in Fig. 3. The first pathway is the reaction with H₂O, alcohol or carboxylic acid to form a 222 carboxylic acid species with hydroxyl, which would subsequently lose a molecule of water to form 223 limononaldehyde or lose a molecule of hydrogen peroxide to form limononic acid (Grosjean et al., 1992; 224 Li et al., 2019b). The second and third pathways involve reactions of sCI₁ with carboxylic acids and 225 carbonyls, respectively, leading to the formation of anhydrides and secondary ozonides. Additionally, the 226 generated OH radicals can react with limonene, giving rise to another alkyl radical, C₁₀H₁₇O. These 227 alkyl radicals react with O_2 and form peroxy radicals (RO₂·). The atmospheric fate of produced RO₂· in 228 the absence of NO_x includes the reaction with RO₂ or HO₂ (Atkinson and Arey, 2003) and the 229 unimolecular H shift. The RO_2 + HO_2 route mainly form hydroperoxide (ROOH), and the minor fraction 230 is to form alcohols and carbonyls (Atkinson and Arey, 2003). The products of bimolecular reactions 231 between RO2. and RO2. are alcohols, carbonyls, alkoxy radicals, peroxides and ROOR dimers (Hammes 232 et al., 2019; Peng et al., 2019). The H shift of RO₂ can form second-generation R and trigger a main 233 generation channel of highly oxidized molecules (HOMs), i.e., R. would go through a process of repeated 234 oxygen addition and hydrogen-atom shift to form HOMs with high O/C ratios of > 0.7-0.8 (Molteni et 235 al., 2018; Bianchi et al., 2019).

In addition to the eCI_1 route, the eCI_2 pathway is also responsible for the generation of various products (Fig. 4). Since the reaction of the hydroxyl radical (OH) attacking limonene is already depicted in Fig. 3, our main emphasis in Fig. 4 is on the pathways involved in the generation of SCI. First, sCI_2 reacts with H₂O and decomposes to limononaldehyde and H₂O₂. Additionally, sCI_2 could experience an O₂ addition, ·OH loss and isomerization to produce two types of RO₂·, which can undergo the similar reactions as the RO₂· formed from the sCI_1 route, and the major products are also shown in Fig. 4.

242 Since limonene and Δ^3 -carene both have an endocyclic double bond, the similar reactions as

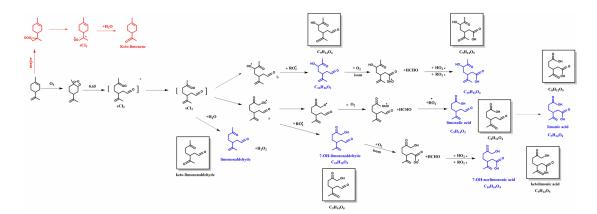
243 mentioned above can occur for the ozonolysis of Δ^3 -carene (Fig. S5), and most corresponding formula 244 in Fig. S5 could be identified in Table S4. However, the reactivity of limonene towards O₃ is expected to 245 be higher owing to its exocyclic double bond. As shown in Fig. 4, the attack of O₃ to the exocyclic double 246 bond mainly leads to sCI₃ (highlighted in red) with the unpaired electrons outside the ring (Leungsakul 247 et al., 2005). sCI₃ can react with H₂O to form a carbonyl called keto-limonene. It should be noted that this reaction can occur not only for limonene, but also for all the products that retain the exocyclic double 248 249 bond. As a result, the compounds that are colored in blue in Fig. 3 and Fig. 4 can undergo further reactions 250 to generate products with an additional carbonyl (see the boxes in Fig. 3 and Fig. 4). Furthermore, their 251 molecular formula shown in Table S5 have been identified using the Q-TOF-MS. This mechanism can 252 well explain the decrease in the relative intensity of $C_{10}H_{16}O_2$ from high RH to low RH and the increase 253 in the relative intensity of C₉H₁₄O₃ from low RH to high RH (Table S3).

254 In such progress, we cannot rule out the possibility that relative humidity (RH) may influence the 255 generation of other free radicals (Ma et al., 2009), thereby impacting the formation of secondary organic 256 aerosols (SOA), such as, OH-radical reactions (Bonn et al., 2002; Fick et al., 2002). However, Molar OH 257 radical yields were reported as 0.65±0.10 (Hantschke et al., 2021), 0.86±0.11 (Aschmann et al., 2002) 258 and 0.56 to 0.59 (Wang et al., 2019) for Δ^3 -carene, while for limonene, the reported yields were 0.67±0.10 259 (Aschmann et al., 2002) and 0.76±0.06 (Herrmann et al., 2010). It seems that the OH radicals produced 260 from limonene and Δ^3 -carene are quite similar within the range of uncertainties. Therefore, the increased 261 ozone consumption by limonene is primarily attributed to the presence of its exocyclic double bond.





264 TOF-MS.



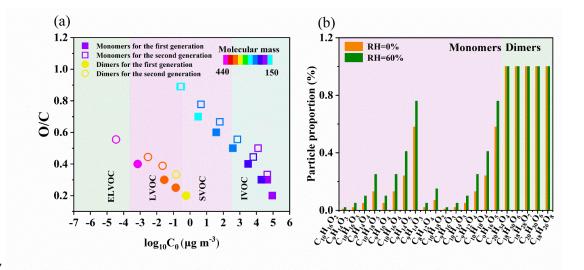
265

Figure 4. Proposed formation mechanisms for SOA formation from eCI₂ and exocyclic double bond
oxidation under high RH. The compounds in blue and in boxes are identified using UPLC/ (-) ESI-QTOF-MS.

269

270 **3.3 Processes leading to the increase or decrease in SOA formation**

271 Based on the results and mechanisms shown above, we present evidence that high humidity 272 enhances limonene-SOA formation. First, the presence of water vapor enhances the formation of 273 carbonyls from the reaction of exocyclic double bond, and the oligomerization of these carbonyls 274 generates more dimers including hemiacetal (or acetal) formation and aldol condensation (Zhang et al., 275 2022; Kroll et al., 2005; Jang et al., 2003). As shown in Table S6, 54 out of the total 187 dimers were 276 exclusively observed for limonene under high humidity conditions, contributing to a corresponding intensity of ~19%. These dimers can be classified as low-volatile organic compounds (LVOC; $3 \times 10^{-4} <$ 277 $C_0 < 0.3 \ \mu g \ m^{-3}$) and extremely low-volatile organic compounds (ELVOCs; $C_0 < 3 \times 10^{-4} \ \mu g \ m^{-3}$) (Fig. 278 279 5a), and thus promote the nucleation and new particle formation in different ways. This finding is similar 280 to that from a previous study showing that high RH can promote dimer formation from the ozonolysis of 281 α -pinene (Kristensen et al., 2014). Second, we find that high RH can also promote the formation of 282 HOMs, although the mechanism remains unclear. As shown in Table S3, many HOMs proposed from the 283 mechanism are detected under high RH condition but not detected under low RH condition, including both monomers and dimers. Many HOMs have low volatilities and, thus, can also promote new particle 284 285 formation. Overall, the promoted dimer and HOM formation greatly enhance the new particle number 286 concentration under high RH condition (Fig. 6).



287

Figure 5. (a) Distribution of the limonene-SOA in the two-dimensional volatility basis set (2D-VBS)
space. (b) Partitioning coefficients of limonene monomers and dimers under low and high RH conditions.

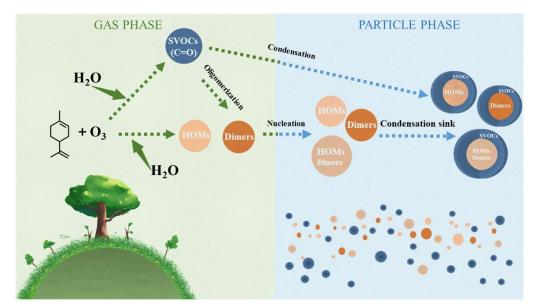




Figure 6. Schematic diagram of the possible mechanisms for the enhancement of limonene-SOA.

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High particle number concentration generally provides more surface areas for semi-volatile organic compounds (SVOCs; $0.3 < C_0 < 300 \ \mu g \ m^{-3}$) to condense on, which results in higher condensation sink (CS). In the OFR, the fates of SVOCs include condensing on aerosol, getting lost on the wall, and reacting with OH radicals to form functionalization and/or fragmentation products (Palm et al., 2016; Li et al., 2019a). The promoted condensation by higher CS leads to a higher fraction of SVOCs getting into the particle phase rather than getting lost on the wall or becoming smaller fragments staying in the gas phase,

300 and thus promoting SOA formation (Li et al., 2019a). Furthermore, the transformation from C-C double 301 bond to carbonyl shown in Fig. 3 and Fig. 4 decreases the volatility of molecules, which can largely 302 influence the gas-particle partitioning of the monomeric compounds (Fig. 5b). For example, the C₀ values 303 of $C_{10}H_{16}O_2$ and $C_{10}H_{16}O_3$ are 90701 and 19968 µg m⁻³, corresponding to partitioning coefficients of 304 0.01 and 0.05, respectively (Fig. 5b and Table S3), with an SOA mass concentration of $\sim 1000 \ \mu g \ m^{-3}$ 305 under dry condition. When they are converted to carbonyls $C_9H_{14}O_3$ and $C_9H_{14}O_4$, the values of C_0 306 become 45556 and 6479 µg m⁻³, corresponding to partitioning coefficients of 0.02 and 0.13, respectively 307 (Fig. 5b and Table S3), with the same SOA loading. This enhancement in partitioning coefficient can 308 largely promote the condensation of SVOCs and, thus, enhance the SOA mass concentration. In addition, 309 the enhanced SOA formation can further influence the equilibrium, e.g., the partitioning coefficient of $C_{10}H_{16}O_3$ increases from 0.05 to 0.10 when SOA mass concentration increases from ~1000 µg m⁻³ under 310 311 dry condition to $\sim 2000 \ \mu g \ m^{-3}$ under wet condition (Fig. 5b and Table S3). The distribution of saturation 312 vapor pressure for monomers and dimers identified by MS has also been shown in Fig. 5a. As can be 313 seen from this figure, around 50% monomers are categorized as SVOCs, thus having the large fraction 314 in the particle phase when converting from dry to wet conditions. Overall, the different fate and 315 partitioning of SVOCs largely enhance the amount of SVOCs in the particle phase (Fig. 6).

316 Concluding the analysis above, high humidity promotes the SOA formation from the ozonolysis of 317 limonene in two steps: nucleation of new particles and condensation of SVOCs on them (Fig. 6). These 318 two steps are closely related to the multi-generation reactions of the exocyclic C=C bond, which are 319 unlikely to happen for the ozonolysis of Δ^3 -carene. Interestingly, Gong and Chen (2021) have found that 320 high RH can inhibit the SOA formation from the first-generation oxidation of limonene ozonolysis, but 321 enhance the SOA formation from the second-generation oxidation (Gong and Chen, 2021), their results 322 agree well with the results and analysis shown here. In contrast, Li et al. (2019b) found negligible change 323 in dimers and HOMs in limonene-O₃ system when changing RH from 0 to 60%. The discrepancy is 324 mainly attributed to the different experimental conditions. The ozone exposure in this study is ~ 18 times 325 higher than in Li et al. (2019b), while the limonene concentration in this study is only \sim 30% of that in 326 their study. These two conditions both favor the multi-generation reactions occurred at the exocyclic 327 double bond of limonene and its products. Thus, we believe this leads to the different results regarding 328 the formation of HOMs and dimers.

329 Regarding Δ^3 -carene, the mechanisms and processes are almost opposite to those of limonene. First, 330 water vapor reacts with sCl₁ or sCl₂ to promote the formation of α -hydroxyalkyl-hydroperoxides (Fig. 331 S5). Their subsequent products without second ozonolysis of exocyclic double bond have higher 332 volatility, and may most likely prevail in the gas phase. In addition, it has been found that α -hydroxyalkyl 333 hydroperoxides preferentially decompose into aldehydes and H₂O₂ (Kumar et al., 2014; Chen et al., 2016), 334 i.e., 3-caronaldehyde for Δ^3 -carene, which has higher volatility than the products from other reaction 335 pathways. Correspondingly, the number and relative intensity of HOMs and dimers detected under high 336 RH conditions are both lower than those under low RH conditions (Table S7). Furthermore, out of a total 337 of 178 dimers, 63 dimers were exclusively identified under low RH conditions (Table S6). As a result, 338 high RH shows an inhibitory effect on the SOA formation from Δ^3 -carene ozonolysis.

339 To investigate the multi-generation reactions of limonene under low-concentration conditions, we 340 conducted low-concentration limonene ozonolysis experiments, and the results are shown in Fig. S6. In 341 these experiments, the limonene and O₃ concentrations were 20.5 ppb and 5.7 ppm, respectively. 342 According to the experimental results, the number concentration of SOA formed from limonene 343 ozonolysis increased by approximately 1.4 times under high RH, which is similar to the increase observed 344 under high-loading conditions. The mass concentration increased by approximately 1.3 times at a 345 precursor concentration of 20.5 ppb. The relatively small increase in mass concentration compared to the 346 high-concentration conditions may be attributed to the less pronounced distribution of SVOCs at low 347 mass concentrations. This result indicates that the enhancement effect on limonene SOA by high RH is 348 still valid for low precursor concentrations.

349 To further confirm the assumption that water-influenced multi-generation reactions of the exocyclic 350 double bond enhance the SOA formation, we conducted two comparative analyses: firstly, we examined 351 the ozonolysis of the endocyclic double bond in limonene, leaving the exocyclic double bond unreacted. 352 This was done by applying a low O_3 concentration (~67 ppb), since the reaction of O_3 with endocyclic 353 double bond is \sim 30 times faster than the reaction of O₃ with exocyclic double bond (Shu and Atkinson, 354 1994). Interestingly, when limonene was oxidized at only the endocyclic double bond, we observed a 355 slight decrease in both the number and mass concentrations as the RH increased (Fig. S7). This is similar 356 to the results obtained for Δ^3 -carene, which contains only one endocyclic double bond. Secondly, we 357 compared the ozonolysis of structurally similar β -caryophyllene, which has an exocyclic C-C double bond that can undergo further reactions (Fig. S8). As expected, we observe a large enhancement in SOA formation under high RH condition (Table S8 and Fig. S9). This implies that monoterpenes, sesquiterpenes, and other BVOCs with two unsaturation double bonds may follow similar reaction mechanisms during ozonolysis, and thus have a RH dependency in SOA production.

362 4 Conclusions

363 In this study, the effect of humidity on SOA production from the ozonolysis of two monoterpenes 364 (limonene and Δ^3 -carene) was investigated with an OFR. Contrasting impacts of RH on the SOA 365 formation were observed: limonene-SOA yield increases by ~100% when RH changes from ~1% to 366 ~60%, while Δ^3 -carene-SOA yield slightly decreases. By analyzing the chemical composition of SOA 367 with ESI-Q-TOF-MS, we find that the multi-generation reactions of the exocyclic C-C double bond are 368 likely the driving force of the enhancement in limonene-SOA. The presence of water promotes the 369 formation of carbonyls from the reaction of exocyclic double bond, and further favors the formation of 370 dimers and HOMs. This leads to promoted new particle formation and subsequent condensation of 371 SVOCs. These reactions also lower the volatilities of the SVOCs, and further promote the gas-particle partitioning. Moreover, this hypothesis is proved by a similar behavior of the ozonolysis of β -372 373 caryophyllene (sesquiterpene with an exocyclic double bond) in SOA enhancement under high RH 374 condition. The results in this study suggest that multi-generation reactions play an important role in SOA 375 formation from the ozonolysis of BVOCs, which are significantly influenced by humidity. This impact is largely dependent on the molecular structure of the SOA precursors (e.g., with or without the exocyclic 376 377 double bond), thus highlighting the importance to consider the molecular structure of monoterpenes in 378 modeling and field studies of biogenic SOA.

379

380 **Data availability.** Experimental data are available upon request to the corresponding authors.

381 **Supplement.** The supplement related to this article is available online.

382 Author contributions. LD and SZ designed the experiments and SZ carried them out. SZ performed

data analysis with assistance from KL, LD, ZY, and JL. SZ and KL wrote the paper with contributions

from all co-authors.

385 **Declaration**. The authors declare that they have no conflict of interest.

19

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391 Reference

- 392 Ahmadov, R., McKeen, S. A., Robinson, A. L., Bahreini, R., Middlebrook, A. M., de Gouw, J. A.,
- 393 Meagher, J., Hsie, E. Y., Edgerton, E., Shaw, S., and Trainer, M.: A volatility basis set model for
- 394 summertime secondary organic aerosols over the eastern United States in 2006, J. Geophys. Res.-Atmos.,
- 395 117, D6301, <u>https://doi.org/10.1029/2011JD016831</u>, 2012.
- 396 Aschmann, S. M., Arey, J., and Atkinson, R.: OH radical formation from the gas-phase reactions of O3
- 397 with a series of terpenes, Atmos. Environ., 36, 4347-4355, <u>https://doi.org/10.1016/S1352-</u>
 398 2310(02)00355-2, 2002.
- 399 Atkinson, R.: Kinetics and mechanisms of the gas-phase reactions of the NO₃ radical with organic
- 400 compounds, J. Phys. Chem. Ref. Data, 20, 459-507, <u>https://doi.org/10.1063/1.555887</u>, 1991.
- 401 Atkinson, R. and Arey, J.: Gas-phase tropospheric chemistry of biogenic volatile organic compounds: a
- 402 review, Atmos. Environ., 37, 197-219, <u>https://doi.org/10.1016/S1352-2310(03)00391-1</u>, 2003.
- 403 Bäck, J., Aalto, J., Henriksson, M., Hakola, H., He, Q., and Boy, M.: Chemodiversity of a Scots pine
- 404 stand and implications for terpene air concentrations, Biogeosciences, 9, 689-702, 405 https://doi.org/10.5194/bg-9-689-2012, 2012.
- 406 Bateman, A. P., Nizkorodov, S. A., Laskin, J., and Laskin, A.: Time-resolved molecular characterization
- 407 of limonene/ozone aerosol using high-resolution electrospray ionization mass spectrometry, Phys. Chem.
- 408 Chem. Phys., 11, 7931-7942, <u>https://doi.org/10.1039/b905288g</u>, 2009.
- 409 Bianchi, F., Kurtén, T., Riva, M., Mohr, C., Rissanen, M. P., Roldin, P., Berndt, T., Crounse, J. D.,
- 410 Wennberg, P. O., Mentel, T. F., Wildt, J., Junninen, H., Jokinen, T., Kulmala, M., Worsnop, D. R.,
- 411 Thornton, J. A., Donahue, N., Kjaergaard, H. G., and Ehn, M.: Highly oxygenated organic molecules
- 412 (HOM) from gas-phase autoxidation involving peroxy radicals: a key contributor to atmospheric aerosol,
- 413 Chem. Rev., 119, 3472-3509, https://doi.org/10.1021/acs.chemrev.8b00395, 2019.

- 414 Bonn, B. and Moortgat, G. K.: New particle formation during α and β -pinene oxidation by O₃, OH and
- 415 NO₃, and the influence of water vapour: particle size distribution studies, Atmos. Chem. Phys., 2, 183-
- 416 196, <u>https://doi.org/10.5194/acp-2-183-2002</u>, 2002.
- 417 Bonn, B., Schuster, G., and Moortgat, G. K.: Influence of water vapor on the process of new particle
- 418 formation during monoterpene ozonolysis, J. Phys. Chem. A, 106, 2869-2881,
- 419 <u>https://doi.org/10.1021/jp012713p</u>, 2002.
- 420 Chen, H., Ren, Y., Cazaunau, M., Dalele, V., Hu, Y., Chen, J., and Mellouki, A.: Rate coefficients for the
- 421 reaction of ozone with 2-and 3-carene, Chem. Phys. Lett., 621, 71-77,
 422 <u>https://doi.org/10.1016/j.cplett.2014.12.056</u>, 2015.
- 423 Chen, L., Huang, Y., Xue, Y., Shen, Z., Cao, J., and Wang, W.: Mechanistic and kinetics investigations
- 424 of oligomer formation from Criegee intermediate reactions with hydroxyalkyl hydroperoxides, Atmos.
- 425 Chem. Phys., 19, 4075-4091, <u>https://doi.org/10.5194/acp-19-4075-2019</u>, 2019.
- 426 Chen, L., Wang, W., Wang, W., Liu, Y., Liu, F., Liu, N., and Wang, B.: Water-catalyzed decomposition
- 427 of the simplest Criegee intermediate CH₂OO, Theor. Chem. Acc., 135, 131, 428 https://doi.org/10.1007/s00214-016-1894-9, 2016.
- 429 Chen, X. and Hopke, P. K.: A chamber study of secondary organic aerosol formation by limonene
- 430 ozonolysis, Indoor Air, 20, 320-328, <u>https://doi.org/10.1111/j.1600-0668.2010.00656.x</u>, 2010.
- 431 Cholakian, A., Beekmann, M., Coll, I., Ciarelli, G., and Colette, A.: Biogenic secondary organic aerosol
- 432 sensitivity to organic aerosol simulation schemes in climate projections, Atmos. Chem. Phys., 19, 13209-
- 433 13226, <u>https://doi.org/10.5194/acp-19-13209-2019</u>, 2019.
- 434 de Matos, S. P., Teixeira, H. F., de Lima, Á. A. N., Veiga-Junior, V. F., and Koester, L. S.: Essential oils
- 435 and isolated terpenes in nanosystems designed for topical administration: a review, Biomolecules, 9, 138,
- 436 <u>https://doi.org/doi:10.3390/biom9040138</u>, 2019.
- 437 Drozd, G. T. and Donahue, N. M.: Pressure dependence of stabilized Criegee intermediate formation
- 438 from a sequence of alkenes, J. Phys. Chem. A, 115, 4381-4387, <u>https://doi.org/10.1021/jp2001089</u>, 2011.
- 439 Fick, J., Pommer, L., Andersson, B., and Nilsson, C.: A study of the gas-phase ozonolysis of terpenes:
- 440 the impact of radicals formed during the reaction, Atmos. Environ., 36, 3299-3308,
- 441 https://doi.org/10.1016/s1352-2310(02)00291-1, 2002.
- 442 Gong, Y. and Chen, Z.: Quantification of the role of stabilized Criegee intermediates in the formation of

- 443 aerosols in limonene ozonolysis, Atmos. Chem. Phys., 21, 813-829, <u>https://doi.org/10.5194/acp-21-813-</u>
 444 <u>2021</u>, 2021.
- 445 Gong, Y., Chen, Z., and Li, H.: The oxidation regime and SOA composition in limonene ozonolysis: roles
- 446 of different double bonds, radicals, and water, Atmos. Chem. Phys., 18, 15105-15123,
- 447 <u>https://doi.org/10.5194/acp-18-15105-2018</u>, 2018.
- 448 Grosjean, D., Williams, E. L., and Seinfeld, J. H.: Atmospheric oxidation of selected terpenes and related
- 449 carbonyls: gas-phase carbonyl products, Environ. Sci. Technol., 26, 1526-1533,
 450 https://doi.org/10.1021/es00032a005, 1992.
- 451 Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang, X.:
- 452 The model of emissions of gases and aerosols from nature version 2.1 (MEGAN2.1): an extended and
- 453 updated framework for modeling biogenic emissions, Geosci. Model Dev., 5, 1471-1492,
- 454 <u>https://doi.org/10.5194/gmd-5-1471-2012</u>, 2012.
- 455 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L.,
- 456 Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China, Proc. Natl. Acad. Sci.
- 457 U. S. A., 111, 17373-17378, https://doi.org/10.1073/pnas.1419604111, 2014.
- 458 Hammes, J., Lutz, A., Mentel, T., Faxon, C., and Hallquist, M.: Carboxylic acids from limonene oxidation
- 459 by ozone and hydroxyl radicals: insights into mechanisms derived using a FIGAERO-CIMS, Atmos.
- 460 Chem. Phys., 19, 13037-13052, <u>https://doi.org/10.5194/acp-19-13037-2019</u>, 2019.
- 461 Hantschke, L., Novelli, A., Bohn, B., Cho, C., Reimer, D., Rohrer, F., Tillmann, R., Glowania, M.,
- 462 Hofzumahaus, A., Kiendler-Scharr, A., Wahner, A., and Fuchs, H.: Atmospheric photooxidation and
- 463 ozonolysis of Δ 3-carene and 3-caronaldehyde: rate constants and product yields, Atmos. Chem. Phys.,
- 464 21, 12665-12685, 10.5194/acp-21-12665-2021, 2021.
- 465 Herrmann, F., Winterhalter, R., Moortgat, G. K., and Williams, J.: Hydroxyl radical (OH) yields from the
- 466 ozonolysis of both double bonds for five monoterpenes, Atmos. Environ., 44, 3458-3464,
- 467 <u>https://doi.org/10.1016/j.atmosenv.2010.05.011</u>, 2010.
- 468 Huang, X., Yun, H., Gong, Z., Li, X., He, L., Zhang, Y., and Hu, M.: Source apportionment and secondary
- 469 organic aerosol estimation of PM2.5 in an urban atmosphere in China, Sci. China-Earth Sci., 57, 1352-
- 470 1362, https://doi.org/10.1007/s11430-013-4686-2, 2014.
- 471 Jang, M. S., Carroll, B., Chandramouli, B., and Kamens, R. M.: Particle growth by acid-catalyzed

- 472 heterogeneous reactions of organic carbonyls on preexisting aerosols, Environ. Sci. Technol., 37, 3828-
- 473 3837, 10.1021/es021005u, 2003.
- 474 Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V.-M., Junninen, H., Paasonen, P., Stratmann, F.,
- 475 Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and Sipilä, M.: Production of
- 476 extremely low volatile organic compounds from biogenic emissions: Measured yields and atmospheric
- 477 implications, Proc. Natl. Acad. Sci. U. S. A., 112, 7123-7128,
- 478 <u>https://doi.org/doi:10.1073/pnas.1423977112</u>, 2015.
- 479 Jonsson, A. M., Hallquist, M., and Ljungstrom, E.: Impact of humidity on the ozone initiated oxidation
- 480 of limonene, Δ^3 -carene, and α -pinene, Environ. Sci. Technol., 40, 188-194, 481 https://doi.org/10.1021/es051163w, 2006a.
- 482 Jonsson, A. M., Hallquist, M., and Ljungstrom, E.: Impact of humidity on the ozone initiated oxidation
- 483 of limonene, Delta(3)-carene, and alpha-pinene, Environ. Sci. Technol., 40, 188-194,
 484 <u>https://doi.org/10.1021/es051163w</u>, 2006b.
- 485 Jonsson, A. M., Hallquist, M., and Ljungstrom, E.: Influence of OH scavenger on the water effect on
- 486 secondary organic aerosol formation from ozonolysis of limonene, Δ³-carene, and α-pinene, Environ. Sci.
 487 Technol., 42, 5938-5944, https://doi.org/10.1021/es702508y, 2008.
- 488 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen,
- 489 R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J.,
- 490 Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and
- 491 Wilson, J.: Organic aerosol and global climate modelling: a review, Atmos. Chem. Phys., 5, 1053-1123,
- 492 <u>https://doi.org/10.5194/acp-5-1053-2005</u>, 2005.
- 493 Khamaganov, V. G. and Hites, R. A.: Rate constants for the gas-phase reactions of ozone with isoprene,
- 494 α and β -pinene, and limonene as a function of temperature, J. Phys. Chem. A, 105, 815-822,
- 495 <u>https://doi.org/10.1021/jp002730z</u>, 2001.
- 496 Kristensen, K., Cui, T., Zhang, H., Gold, A., Glasius, M., and Surratt, J. D.: Dimers in α-pinene secondary
- 497 organic aerosol: effect of hydroxyl radical, ozone, relative humidity and aerosol acidity, Atmos. Chem.
- 498 Phys., 14, 4201-4218, https://doi.org/10.5194/acp-14-4201-2014, 2014.
- 499 Kroll, J. H., Ng, N. L., Murphy, S. M., Varutbangkul, V., Flagan, R. C., and Seinfeld, J. H.: Chamber
- 500 studies of secondary organic aerosol growth by reactive uptake of simple carbonyl compounds, J.

- 501 Geophys. Res.-Atmos., 110, 10.1029/2005JD006004, 2005.
- 502 Kumar, M., Busch, D. H., Subramaniam, B., and Thompson, W. H.: Role of tunable acid catalysis in
- 503 decomposition of α-Hydroxyalkyl hydroperoxides and mechanistic implications for tropospheric
- 504 chemistry, J. Phys. Chem. A, 118, 9701-9711, https://doi.org/10.1021/jp505100x, 2014.
- 505 Leungsakul, S., Jaoui, M., and Kamens, R. M.: Kinetic Mechanism for Predicting Secondary Organic
- 506 Aerosol Formation from the Reaction of d-Limonene with Ozone, Environ. Sci. Technol., 39, 9583-9594,
- 507 https://doi.org/10.1021/es0492687, 2005.
- 508 Levy, H., II, Horowitz, L. W., Schwarzkopf, M. D., Ming, Y., Golaz, J.-C., Naik, V., and Ramaswamy,
- 509 V.: The roles of aerosol direct and indirect effects in past and future climate change, J. Geophys. Res.510 Atmos., 118, 4521-4532, https://doi.org/10.1002/jgrd.50192, 2013.
- 511 Li, J. Y., Zhang, H. W., Ying, Q., Wu, Z. J., Zhang, Y. L., Wang, X. M., Li, X. H., Sun, Y. L., Hu, M.,
- 512 Zhang, Y. H., and Hu, J. L.: Impacts of water partitioning and polarity of organic compounds on
- secondary organic aerosol over eastern China, Atmos. Chem. Phys., 20, 7291-7306,
 https://doi.org/10.5194/acp-20-7291-2020, 2020.
- 515 Li, K., Liggio, J., Lee, P., Han, C., Liu, Q., and Li, S.-M.: Secondary organic aerosol formation from α-
- 516 pinene, alkanes, and oil-sands-related precursors in a new oxidation flow reactor, Atmos. Chem. Phys.,
- 517 19, 9715-9731, <u>https://doi.org/10.5194/acp-19-9715-2019</u>, 2019a.
- 518 Li, X., Chee, S., Hao, J., Abbatt, J. P. D., Jiang, J., and Smith, J. N.: Relative humidity effect on the
- 519 formation of highly oxidized molecules and new particles during monoterpene oxidation, Atmos. Chem.
- 520 Phys., 19, 1555-1570, https://doi.org/10.5194/acp-19-1555-2019, 2019b.
- 521 Liu, Q., Liggio, J., Breznan, D., Thomson, E. M., Kumarathasan, P., Vincent, R., Li, K., and Li, S.-M.:
- 522 Oxidative and Toxicological Evolution of Engineered Nanoparticles with Atmospherically Relevant
- 523 Coatings, Environ. Sci. Technol., 53, 3058-3066, 10.1021/acs.est.8b06879, 2019.
- 524 Liu, Y., Liggio, J., Harner, T., Jantunen, L., Shoeib, M., and Li, S.-M.: Heterogeneous OH Initiated
- 525 Oxidation: A Possible Explanation for the Persistence of Organophosphate Flame Retardants in Air,
- 526 Environ. Sci. Technol., 48, 1041-1048, 10.1021/es404515k, 2014.
- 527 Ma, Y., Porter, R. A., Chappell, D., Russell, A. T., and Marston, G.: Mechanisms for the formation of
- 528 organic acids in the gas-phase ozonolysis of 3-carene, Phys. Chem. Chem. Phys., 11, 4184-4197,
- 529 <u>https://doi.org/10.1039/b818750a</u>, 2009.

- 530 Molteni, U., Bianchi, F., Klein, F., El Haddad, I., Frege, C., Rossi, M. J., Dommen, J., and Baltensperger,
- 531 U.: Formation of highly oxygenated organic molecules from aromatic compounds, Atmos. Chem. Phys.,
- 532 18, 1909-1921, https://doi.org/10.5194/acp-18-1909-2018, 2018.
- 533 Mot, M.-D., Gavrilas, S., Lupitu, A. I., Moisa, C., Chambre, D., Tit, D. M., Bogdan, M. A., Bodescu, A.-
- 534 M., Copolovici, L., Copolovici, D. M., and Bungau, S. G.: Salvia officinalis L. essential oil:
- 535 characterization, antioxidant properties, and the effects of aromatherapy in adult patients, Antioxidants,
- 536 11, 808, https://doi.org/10.3390/antiox11050808, 2022.
- 537 Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Secondary
- organic aerosol formation from m-xylene, toluene, and benzene, Atmos. Chem. Phys., 7, 3909-3922,
 https://doi.org/10.5194/acp-7-3909-2007, 2007.
- 540 Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle
- partitioning and secondary organic aerosol yields, Environ. Sci. Technol., 30, 2580-2585,
 https://doi.org/10.1021/es950943+, 1996.
- 543 Palm, B. B., Campuzano-Jost, P., Ortega, A. M., Day, D. A., Kaser, L., Jud, W., Karl, T., Hansel, A.,
- 544 Hunter, J. F., Cross, E. S., Kroll, J. H., Peng, Z., Brune, W. H., and Jimenez, J. L.: In situ secondary
- 545 organic aerosol formation from ambient pine forest air using an oxidation flow reactor, Atmos. Chem.
- 546 Phys., 16, 2943-2970, https://doi.org/10.5194/acp-16-2943-2016, 2016.
- 547 Pathak, R. K., Salo, K., Emanuelsson, E. U., Cai, C., Lutz, A., Hallquist, A. M., and Hallquist, M.:
- 548 Influence of ozone and radical chemistry on limonene organic aerosol production and thermal
- 549 characteristics, Environ. Sci. Technol., 46, 11660-11669, <u>https://doi.org/10.1021/es301750r</u>, 2012.
- 550 Peng, Z., Lee-Taylor, J., Orlando, J. J., Tyndall, G. S., and Jimenez, J. L.: Organic peroxy radical
- 551 chemistry in oxidation flow reactors and environmental chambers and their atmospheric relevance,
- 552 Atmos. Chem. Phys., 19, 813-834, https://doi.org/10.5194/acp-19-813-2019, 2019.
- 553 Pye, H. O. T., Ward-Caviness, C. K., Murphy, B. N., Appel, K. W., and Seltzer, K. M.: Secondary organic
- aerosol association with cardiorespiratory disease mortality in the United States, Nat. Commun., 12,
- 555 https://doi.org/10.1038/s41467-021-27484-1, 2021.
- 556 Ravichandran, C., Badgujar, P. C., Gundev, P., and Upadhyay, A.: Review of toxicological assessment of
- 557 d-limonene, a food and cosmetics additive, Food Chem. Toxicol., 120, 668-680,
- 558 <u>https://doi.org/10.1016/j.fct.2018.07.052</u>, 2018.

- 559 Sbai, S. E. and Farida, B.: Photochemical aging and secondary organic aerosols generated from limonene
- 560 in an oxidation flow reactor, Environ. Sci. Pollut. Res., 26, 18411-18420, https://doi.org/10.1007/s11356-
- 561 <u>019-05012-5</u>, 2019.
- 562 Seinfeld, J. H., Erdakos, G. B., Asher, W. E., and Pankow, J. F.: Modeling the formation of secondary
- 563 organic aerosol (SOA). 2. The predicted effects of relative humidity on aerosol formation in the α-Pinene-,
- 564 β-Pinene-, sabinene-, Δ^3 -Carene-, and cyclohexene-ozone systems, Environ. Sci. Technol., 35, 1806-
- 565 1817, <u>https://doi.org/10.1021/es001765</u>+, 2001.
- 566 Shaw, J. T., Lidster, R. T., Cryer, D. R., Ramirez, N., Whiting, F. C., Boustead, G. A., Whalley, L. K.,
- 567 Ingham, T., Rickard, A. R., Dunmore, R. E., Heard, D. E., Lewis, A. C., Carpenter, L. J., Hamilton, J. F.,
- 568 and Dillon, T. J.: A self-consistent, multivariate method for the determination of gas-phase rate
- 569 coefficients, applied to reactions of atmospheric VOCs and the hydroxyl radical, Atmos. Chem. Phys.,
- 570 18, 4039-4054, https://doi.org/10.5194/acp-18-4039-2018, 2018.
- 571 Shu, Y. G. and Atkinson, R.: RATE CONSTANTS FOR THE GAS-PHASE REACTIONS OF O-3 WITH
- 572 A SERIES OF TERPENES AND OH RADICAL FORMATION FROM THE 0-3 REACTIONS WITH
- 573 SESQUITERPENES AT 296+/-2-K, INTERNATIONAL JOURNAL OF CHEMICAL KINETICS, 26,
- 574 1193-1205, 10.1002/kin.550261207, 1994.
- 575 Sindelarova, K., Granier, C., Bouarar, I., Guenther, A., Tilmes, S., Stavrakou, T., Müller, J. F., Kuhn, U.,
- 576 Stefani, P., and Knorr, W.: Global data set of biogenic VOC emissions calculated by the MEGAN model
- 577 over the last 30 years, Atmos. Chem. Phys., 14, 9317-9341, <u>https://doi.org/10.5194/acp-14-9317-2014</u>,
- 578 2014.
- 579 Sun, Y., Wang, Z., Fu, P., Jiang, Q., Yang, T., Li, J., and Ge, X.: The impact of relative humidity on
- aerosol composition and evolution processes during wintertime in Beijing, China, Atmos. Environ., 77,
- 581 927-934, <u>https://doi.org/10.1016/j.atmosenv.2013.06.019</u>, 2013.
- 582 Thomsen, D., Elm, J., Rosati, B., Skonager, J. T., Bilde, M., and Glasius, M.: Large discrepancy in the
- 583 formation of secondary organic aerosols from structurally similar monoterpenes, ACS Earth Space
- 584 Chem., 5, 632-644, <u>https://doi.org/10.1021/acsearthspacechem.0c00332</u>, 2021.
- 585 Varutbangkul, V., Brechtel, F. J., Bahreini, R., Ng, N. L., Keywood, M. D., Kroll, J. H., Flagan, R. C.,
- 586 Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Hygroscopicity of secondary organic aerosols formed by
- 587 oxidation of cycloalkenes, monoterpenes, sesquiterpenes, and related compounds, Atmos. Chem. Phys.,

- 588 6, 2367-2388, <u>https://doi.org/10.5194/acp-6-2367-2006</u>, 2006.
- 589 Wang, L., Liu, Y., and Wang, L.: Ozonolysis of 3-carene in the atmosphere. Formation mechanism of
- 590 hydroxyl radical and secondary ozonides, Phys. Chem. Chem. Phys., 21, 8081-8091,
- 591 10.1039/c8cp07195k, 2019.
- 592 Wang, L. Y. and Wang, L. M.: The oxidation mechanism of gas-phase ozonolysis of limonene in the
- 593 atmosphere, Phys. Chem. Chem. Phys., 23, 9294-9303, <u>https://doi.org/10.1039/d0cp05803c</u>, 2021.
- 594 Watne, A. K., Westerlund, J., Hallquist, A. M., Brune, W. H., and Hallquist, M.: Ozone and OH-induced
- 595 oxidation of monoterpenes: Changes in the thermal properties of secondary organic aerosol (SOA), J.
- 596 Aerosol Sci, 114, 31-41, <u>https://doi.org/10.1016/j.jaerosci.2017.08.011</u>, 2017.
- 597 Xu, L., Tsona, N. T., and Du, L.: Relative humidity changes the role of SO₂ in biogenic secondary organic
- 598 aerosol formation, J. Phys. Chem. Lett., 12, 7365-7372, <u>https://doi.org/10.1021/acs.jpclett.1c01550</u>,
- 599 2021.
- 600 Ye, J., Abbatt, J. P. D., and Chan, A. W. H.: Novel pathway of SO₂ oxidation in the atmosphere: reactions
- with monoterpene ozonolysis intermediates and secondary organic aerosol, Atmos. Chem. Phys., 18,
 5549-5565, https://doi.org/10.5194/acp-18-5549-2018, 2018.
- 603 Yu, K. P., Lin, C. C., Yang, S. C., and Zhao, P.: Enhancement effect of relative humidity on the formation
- and regional respiratory deposition of secondary organic aerosol, J. Hazard. Mater., 191, 94-102,
- 605 <u>https://doi.org/10.1016/j.jhazmat.2011.04.042</u>, 2011.
- 606 Zhang, H., Wang, S., Hao, J., Wang, X., Wang, S., Chai, F., and Li, M.: Air pollution and control action
- 607 in Beijing, J. Clean Prod., 112, 1519-1527, <u>https://doi.org/10.1016/j.jclepro.2015.04.092</u>, 2016.
- 608 Zhang, Y., He, L., Sun, X., Ventura, O. N., and Herrmann, H.: Theoretical Investigation on the
- 609 Oligomerization of Methylglyoxal and Glyoxal in Aqueous Atmospheric Aerosol Particles, ACS Earth
- 610 Space Chem., 6, 1031-1043, 10.1021/acsearthspacechem.1c00422, 2022.
- 611 Zhao, R. R., Zhang, Q. X., Xu, X. Z., Zhao, W. X., Yu, H., Wang, W. J., Zhang, Y. M., and Zhang, W. J.:
- 612 Effect of experimental conditions on secondary organic aerosol formation in an oxidation flow reactor,
- 613 Atmos. Pollut. Res., 12, 392-400, https://doi.org/10.1016/j.apr.2021.01.011, 2021.
- Eita Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol
- 615 formation, Chem. Soc. Rev., 41, 6582-6605, <u>https://doi.org/10.1039/c2cs35122f</u>, 2012.

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