

Supplementary Information for

Formation drivers and photochemical effects of ClNO₂ in a coastal city of Southeast China

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41 observation-average and case conditions.

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43 study site.

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45 ClNO₂ case (Unit: ppb).

46 **Table S3.** Summary of ClNO₂ peak concentrations at different types of sites in China and other countries (Unit:
47 ppb).

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53 **Text S1.** Detailed information of this observation site and instruments.

54 The observation site is on the top of the teaching building (over 70 meters) of Institute of Urban
55 Environment, Chinese Academy of Sciences in Xiamen, surrounded by school and residential buildings,
56 Xinglin Bay, and several major transportation roads (Fig.S1). The observation site, called as the Atmospheric
57 Observation Supersite, is equipped with complete measurement instruments to observe trace gases, aerosol
58 compositions, and meteorological parameters. The continuous gas analyzers (Thermo Fisher Scientific, USA)
59 were employed to measure O₃ (TEI 49*i*), CO (TEI 48*i*), SO₂ (TEI 43*i*), and NO_x (TEI 42*i*). VOC species were
60 detected by a gas chromatography system equipped with a mass spectrometer and flame ionization detector
61 (GC-MS/FID, TH-300B, Wuhan, China). PM_{2.5} mass concentrations and its inorganic compositions (NO₃⁻,
62 SO₄²⁻, NH₄⁺, and Cl⁻) were monitored using the tapered element oscillating microbalance (TEOM1405,
63 Thermo Scientific Corp., MA, USA) and the Monitor for AeRosols and Gases in ambient Air (MARGA; ADI
64 2080, Applikon Analytical B.V., the Netherlands), respectively. The particle surface area concentrations (*S_a*)
65 were obtained from the ambient particle number size distribution detecting by the Scanning Mobility Particle
66 Sizer (SMPS, TSI Inc.) and Aerodynamic Particle Size Spectrometer (APS). Meteorological factors, including
67 air temperature (T), relative humidity (RH), atmospheric pressure (P), ultraviolet radiation (UV), wind speed
68 (WS), and wind direction (WD) were measured by the weather station with a sonic anemometer (150WX,
69 Airmar, USA). The data of boundary layer height (BLH) was gotten from the European Centre for Medium-
70 Range Weather Forecasts (ECMWF) ERA5 hourly reanalysis dataset. Photolysis frequency (including *J*(O¹D),
71 *J*(NO₂), *J*(HCHO), *J*(HONO), *J*(NO₃), and *J*(H₂O₂)) were detected by a photolysis spectrometer (PFS-100,
72 Focused Photonics Inc., Hangzhou, China). The HCHO analyzer (FMS-100, Focused Photonics Inc.,
73 Hangzhou, China) was used to observe the concentrations of HCHO.

74 The concentrations of ClNO₂ and N₂O₅ were measured by an iodide-adduct Chemical Ionization-
75 Atmospheric Pressure Interface-Long Time of Flight (Aerodyne Research Inc, USA and ToFwerk AG,
76 Switzerland) mass spectrometer (I⁻-ToF-CIMS). The ambient air is drawn into the sampling chamber through
77 a perfluoroalkoxy (PFA) pipeline, approximately 2 meters in length with an inner diameter of 1/4 inch, at a
78 flow rate of 10 standard liters per minute (SLPM). Subsamples of approximately 2 SLPM are then introduced
79 into the CIMS setting. Within this setup, methyl iodine gas (CH₃I) emitted from the heated CH₃I permeation
80 tube (VICI) undergoes ionization as it passes through a soft X-ray ionization source (ToFwerk AG, P-type),
81 carried by an ultra-high purity nitrogen gas (N₂) flow at 99.999% purity and a rate of 2.7 SLPM. Reagent ions
82 (I⁻ and I(H₂O)⁻) selectively participate in the ion-molecule reaction (IMR) chamber with the target gas,

83 producing iodide clusters in the IMR chamber. These clusters are then quantified using the Γ^- -ToF-CIMS
84 instrument. The background signals of the CIMS instrument were determined by introducing dry N_2 into the
85 inlet for a duration of 20 minutes. To prevent the buildup of particulate matter on the inlet tubing, we replaced
86 the tubing weekly and flushed it with deionized water, followed by a 20-minute stream of N_2 for drying.
87 Additionally, to reduce interference from the sampling inlet, we utilized a pump with a flow rate of 10 SLPM
88 to draw the ambient sample.

90 **Text S2. The calibrations of $ClNO_2$ and N_2O_5 and uncertainty analysis.**

91 In our study, the calibrations of $ClNO_2$ and N_2O_5 were based on previous established methods (Thaler et
92 al., 2011; Wang et al., 2016; Wang et al., 2022). A nitrogen (N_2) flow at a rate of $50 \text{ mL} \cdot \text{min}^{-1}$, containing 6
93 ppm of Cl_2 , was directed over a slurry composed of sodium nitrite ($NaNO_2$) and sodium chloride ($NaCl$). This
94 slurry facilitated the production of $ClNO_2$, with $NaCl$ added to minimize NO_2 formation as a secondary
95 product. Subsequently, the resulting mixture containing $ClNO_2$ was conditioned to a specified RH and then
96 sampled using the CIMS instrument. To quantify the concentrations of $ClNO_2$, the mixed flow was directly
97 fed into a cavity attenuated phase shift spectroscopy (CAPS) instrument to measure the baseline levels of NO_2 ,
98 subsequently, it was passed through a thermal dissociation tube heated to $380 \text{ }^\circ\text{C}$, causing $ClNO_2$ to decompose
99 into NO_2 , determined by the CAPS instrument. The differences in measured NO_2 concentrations between with
100 and without thermal dissociation corresponded to the concentrations of $ClNO_2$. For the calibration of N_2O_5 ,
101 O_3 was generated by passing approximately 30 sccm of ultrapure zero air through a mercury lamp (UVP). O_3
102 then reacted with a 30 sccm flow rate of NO_2 to produce NO_3 , which subsequently reacted with NO_2 to yield
103 a flow of N_2O_5 . This N_2O_5 -enriched flow was utilized to calibrate the CIMS measurements of N_2O_5 . By
104 adjusting the RH, a mixed flow containing stable N_2O_5 was introduced into the CIMS instrument, allowing
105 for the acquisition of a normalized humidity-dependent curve for N_2O_5 . Although the absolute concentrations
106 of the N_2O_5 source were not directly quantified due to the absence of an N_2O_5 -specific detector, the N_2O_5 -
107 enriched flow was passed through a supersaturated $NaCl$ solution assuming 100% conversion efficiency from
108 N_2O_5 to $ClNO_2$. The dependences of $ClNO_2$ and N_2O_5 sensitivity on RH are shown in Fig. S2, indicating that
109 the sensitivities of $ClNO_2$ and N_2O_5 depended on the variations of RH values. The sensitivities of $ClNO_2$ and
110 N_2O_5 were 0.055 ± 0.018 and $0.11 \pm 0.063 \text{ ncps} \cdot \text{ppb}^{-1}$, respectively. The detection limit (3σ) of $ClNO_2$ and
111 N_2O_5 was 1.3 and 0.61 ppt, respectively.

Text S3. The model configuration of machine learning.

The gradient boosted tree implemented from XGBoost was chosen as machine learning methods. It has been optimized to run in distributed computing environment and can handle a large amount of input data. Compared to neural networks, the results of gradient boosting tree models are more interpretable, enabling them to link the results with the recognizable chemical features. The XGBoost algorithm is a tree-based machine learning model known for its excellent performance in speed and accuracy. It can aggregate weak learners into a strong model, enhancing single generalization ability and robustness, thereby improving prediction accuracy.

In this study, ClNO₂ concentrations are as dependent variable, and trace gases (SO₂, CO, NO₂, NO, O₃, and N₂O₅), PM_{2.5}, inorganic compositions (NO₃⁻, SO₄²⁻, NH₄⁺, and Cl⁻), and meteorological parameters (T, RH, UV, WS, WD, and BLH) are as the argument. In the XGBoost model, 70% of the observed data are used as the training set, and the rest 30% data are used as the testing set. Five cross-validation is employed to adjust hyperparameters to stabilize the predictive ability of the model. The adjusted hyperparameters include maximum depth, learning rate, gamma, minimum child weight, and estimate. R², mean square error (MSE), and root mean square error (RMSE) are used to assess the model performance. The Shapely additive explanations (SHAP) model is an interpreter package designed to investigate the contributions of each feature to the model predictions. Its approach involves an additive explanatory model which treats all features as contributors, a concept inspired by cooperative game theory. For each predicted sample, the SHAP model provides a Shapley value, which is the sum of the values assigned to each feature.

Text S4. The box model configuration and output.

The observation-based model (OBM) coupled with the Master Chemical Mechanism (MCM) version 3.3.1 was utilized to assess the impacts of ClNO₂ on photochemically atmospheric oxidation. As delineated in earlier studies (Xue et al., 2015; Tham et al., 2016; Xia et al., 2021; Peng et al., 2021; Peng et al., 2022), established chlorine chemistry mechanisms have been integrated. The impacts of dilution mixing were included for all species by introducing a dilution factor, defined as a function of the variation of the planetary boundary layer (PBL) height. In our study, ClNO₂, N₂O₅, VOCs, HCHO, HONO, CO, O₃, NO, NO₂, SO₂, along with meteorological factors as observation constrained were input into the box model at an hourly resolution. We performed the OBM for 5 days to initialize the unconstrained compounds and radicals before starting the simulation. We focused on elucidating the influence of ClNO₂ photolysis on the formation of RO_x radicals and O₃, the AOC. Generally, the reactions of HO₂ + NO and RO₂ + NO are the major O₃ production

145 pathways (Eq.1), and the O₃ loss pathways (Eq.2) include NO₂ + OH/RO₂, O₃ photolysis, O₃ + OH/HO₂
 146 radicals, O₃/NO₃+ VOCs. The O₃ production rate minus the O₃ loss rate was used to calculate the net O₃
 147 production rate (Eq.3).

$$148 \quad P(O_3) = k_1[HO_2][NO] + \sum(k_2[RO_2][NO]) \quad (1)$$

$$149 \quad L(O_3) = k_3[O_1D][H_2O] + k_4[O_3][OH] + k_5[O_3][HO_2] + k_6[NO_2][OH] + \sum(k_7[O_3][VOCs]) +$$

$$150 \quad 2 \sum(k_8[NO_3][VOCs]) \quad (2)$$

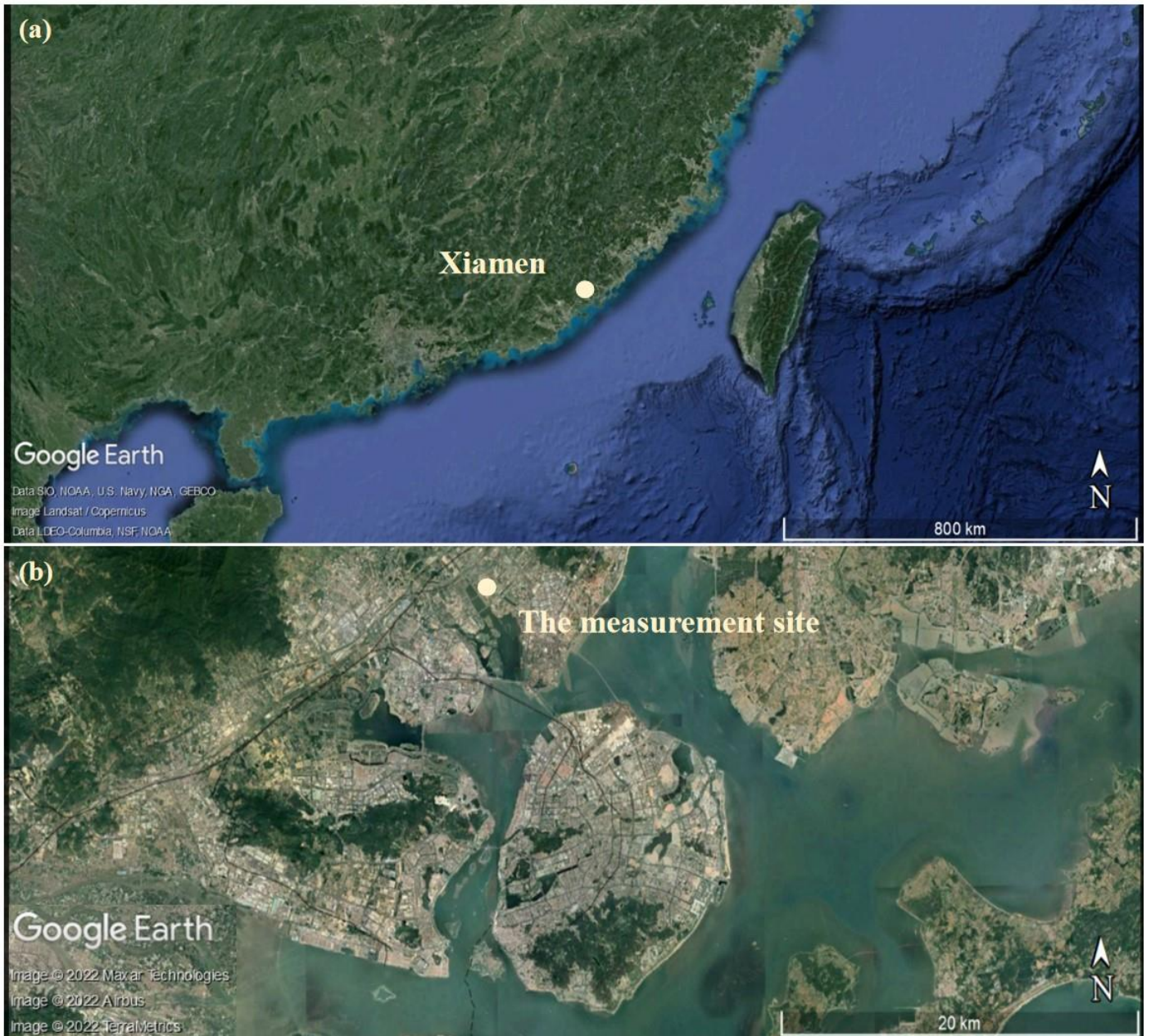
$$151 \quad P_{net}(O_3) = P(O_3) - L(O_3) \quad (3)$$

152 Where, k_i stands for the rate constant.

153 The AOC is calculated by the sum of the rates of CH₄, CO, and VOCs oxidized by atmospheric oxidants
 154 (O₃, OH, Cl, and NO₃ radicals) (Xue et al., 2015; Yi et al., 2023), used by Eq. (4).

$$155 \quad AOC = \sum_i k_{Y_i}[Y_i][X] \quad (4)$$

156 Where, $[Y_i]$ is the concentrations of reduced species (VOCs, CO, and CH₄), $[X]$ is the concentrations of
 157 oxidants (O₃, OH, Cl, and NO₃ radicals), and k_{Y_i} represents the reaction rate constant of Y_i and X .



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159 **Fig. S1** Measurement locations. (a) Xiamen City in the southeast of China. (b) Location of the measurement
160 site in Xiamen. (The topographic image is provided by © Google Earth.)

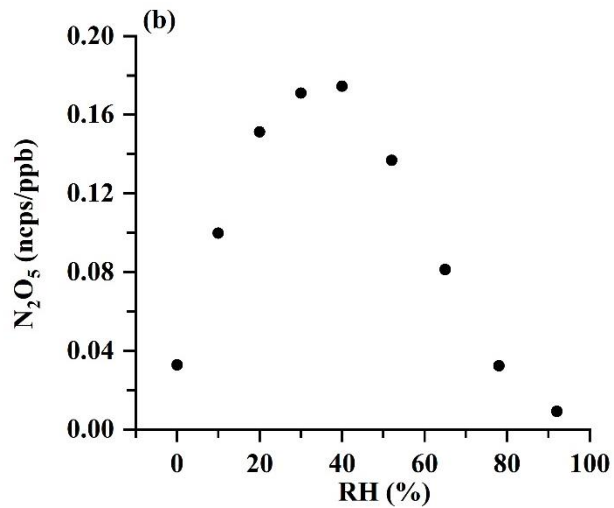
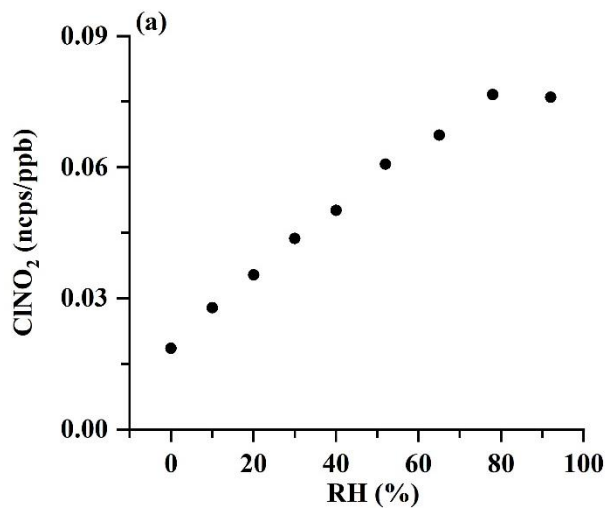
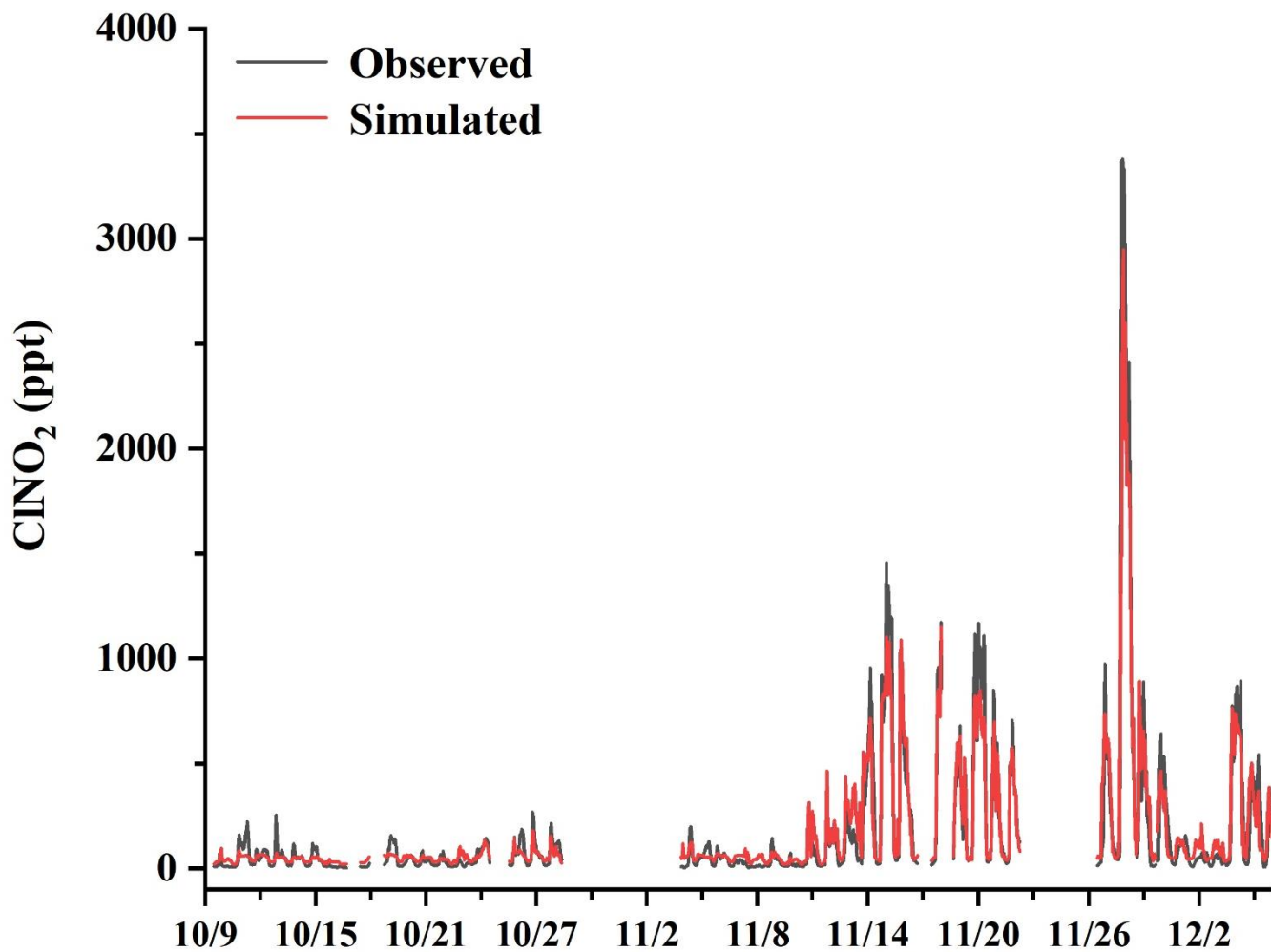
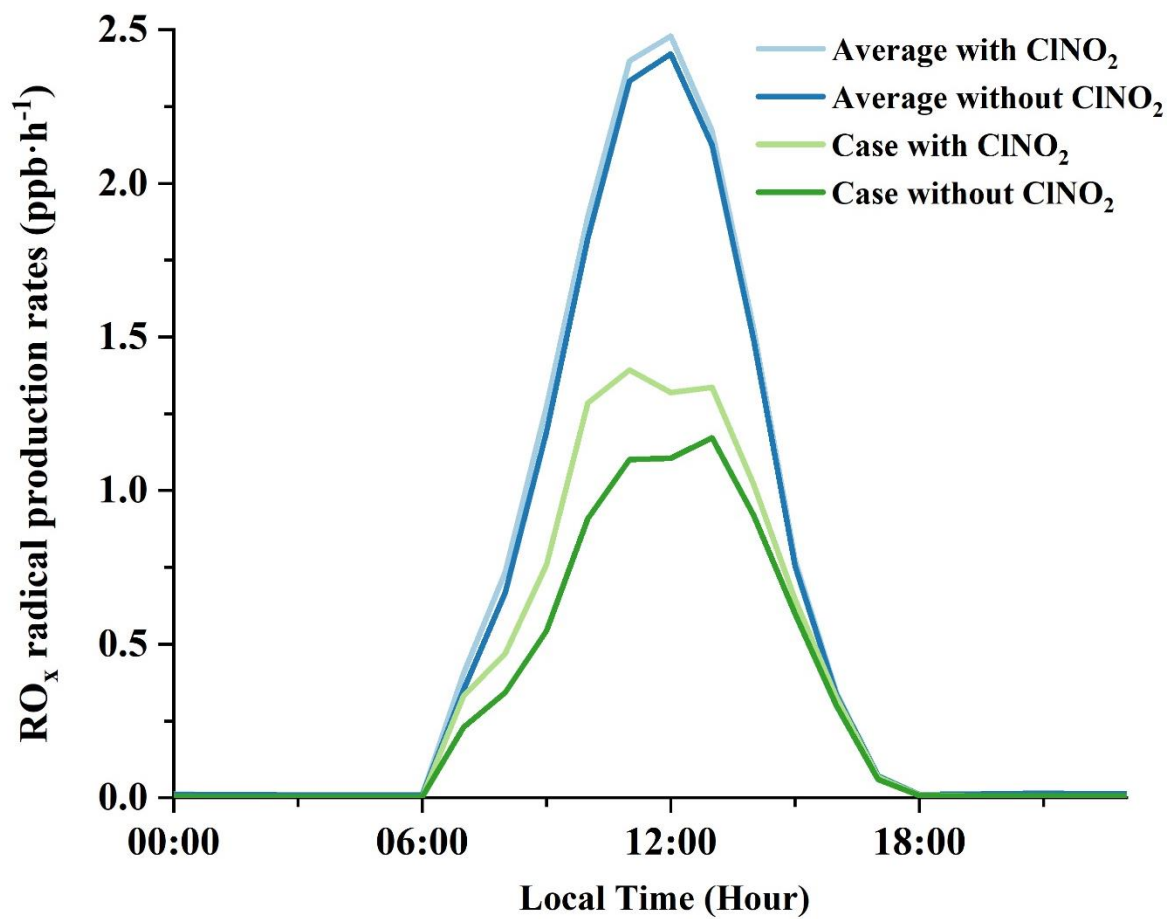


Fig. S2 The dependences of ClNO₂ and N₂O₅ sensitivity on relative humidity.



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190 **Fig. S3** Comparison of observed ClNO₂ concentrations and simulated concentrations of ClNO₂ by the
191 XGBoost model.



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 201 **Fig. S4** RO_x (OH + HO₂ + RO₂) radicals production rates induced by ClNO₂ photolysis under the observation-
 202 average conditions and the high ClNO₂ case.
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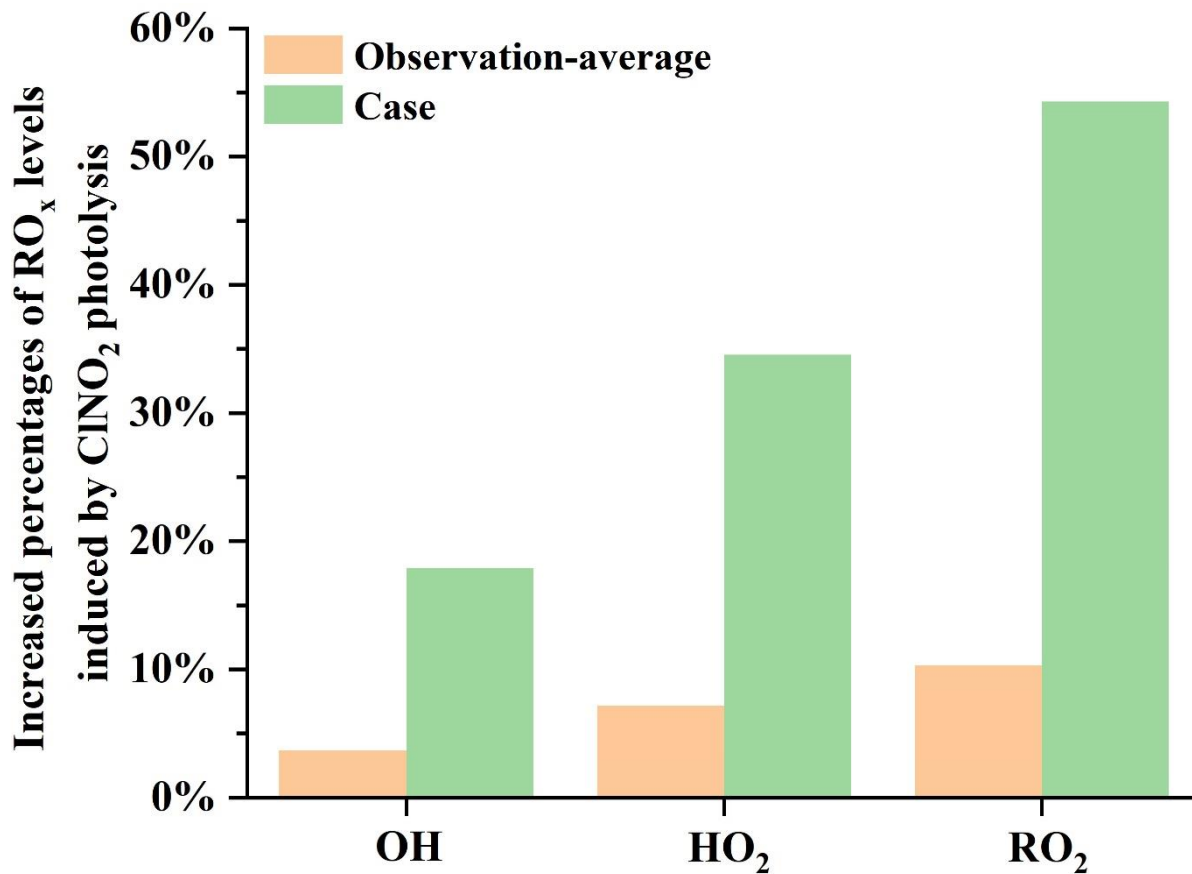


Fig. S5 Increased percentages of RO_x (OH, HO₂, RO₂) radicals induced by ClNO₂ photolysis under the observation-average conditions and the high ClNO₂ case.

217 **Table S1.** Measurement techniques, time resolutions, and detection limit of observation instruments at our
 218 study site.

Parameters	Techniques	Time resolutions	Limit of detection
ClNO ₂ and N ₂ O ₅	I ⁻ -ToF-CIMS	1 min	~ 1 ppt
CH ₄	NCMS6300	15 min	21 ppb
VOCs	GC-MS/FID	1 hour	0.02-0.30 ppb
HCHO	Hantzsch fluorimetry	1 s	0.05 ppb
<i>J</i> (O ¹ D), <i>J</i> (NO ₂), <i>J</i> (HONO), <i>J</i> (NO ₃), <i>J</i> (HCHO), and <i>J</i> (H ₂ O ₂)	Photolysis spectrometer	8 s	a
CO	Infrared absorption	1 min	40.00 ppb
SO ₂	Pulsed UV fluorescence	1 min	0.50 ppb
O ₃	UV photometry	1 min	1.00 ppb
NO	Chemiluminescence	1 min	0.50 ppb
NO ₂	Chemiluminescence	1 min	0.50 ppb

219 a. Process-specific, 5 orders of magnitude lower than maximum at noon.

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233 **Table S2.** The observation data used in the box model under the observation-average conditions and the high

234 ClNO₂ case (Unit: ppb).

Parameters	Average	Case	Parameters	Average	Case
T (K)	296.77	295.63	M2HEX	0.04	0.04
RH (%)	65.98	88.48	M3HEX	0.06	0.08
SO ₂	1.93	2.34	NC7H16	0.06	0.06
NO ₂	12.94	26.38	TOLUENE	1.12	1.73
NO	3.19	5.85	NC8H18	0.02	0.02
CO	461.59	769.56	PXYL	0.40	0.73
O ₃	31.07	11.43	EBENZ	0.19	0.42
C2H4	0.89	1.41	NC9H20	0.01	0.02
C2H2	0.78	1.53	STYRENE	0.04	0.09
C2H6	2.01	2.64	OXYL	0.11	0.20
C3H6	0.24	0.23	TM124B	0.02	0.03
C3H8	1.90	3.28	TM123B	0.01	0.01
IC4H10	0.74	1.30	NC11H24	0.01	0.01
BUT1ENE	0.05	0.06	NC12H26	0.02	0.03
NC4H10	1.07	1.97	HCHO	2.50	3.92
CBUT2ENE	0.02	0.03	ACR	0.06	0.06
TBUT2ENE	0.02	0.03	CH3COCH3	1.98	2.26
IC5H12	0.61	1.46	MEK	0.47	0.52
NC5H12	0.27	0.50	CH3CL	0.39	0.33
C5H8	0.02	0.02	C4H6	0.01	0.02
M22C4	0.02	0.02	IPOPOL	0.17	0.11
M23C4	0.05	0.06	MTBE	0.10	0.19
M2PE	0.05	0.08	ETHACET	1.26	2.14
M3PE	0.08	0.13	JNO ₂ (s ⁻¹)	0.002106	0.000981
HEX1ENE	0.00	0.00	ClNO ₂	0.17	0.96
NC6H14	0.09	0.15	N ₂ O ₅	0.02	0.01
BENZENE	0.18	0.32	Cl ₂	0.01	0.05
CHEX	0.02	0.02	HONO	0.48	0.77

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240 **Table S3.** Summary of ClNO₂ peak concentrations at different types of sites in China and other countries (Unit:
 241 ppb).

Observation Area	Observation Sites	ClNO ₂	References
Wangdu, China	Rural site	~3.5	1
Beijing, China	Urban site	~3.0	2
Jinan, China	Urban site	~0.8	3
Mt. Tai, China	Mountain site	~2.0	4
Changzhou, China	Suburban site	~1.3	5
Shanghai, China	Urban site	~5.7	5
Nanjing, China	Rural site	~3.7	6
Xiamen, China	Urban site	~3.4	This study
Heshan, China	Rural site	~8.1	7
Shenzhen, China	Background site	~1.7	8
Hong Kong, China	Mountain site	~4.7	9
Seoul, Korea	Urban site	~2.5	10
Hesen, Germany	Rural site	~0.8	11
London, UK	Urban site	~0.7	12
Boulder, USA	Urban site	~1.3	13
Houston, USA	Marine boundary layer	~1.2	14
Los Angeles, USA	Marine boundary layer	~3.5	15
Calgary, Canada	Urban site	~0.3	16

242 The references as follows: 1 (Tham et al., 2016), 2 (Ma et al., 2023), 3 (Wang et al., 2017a), 4 (Wang et al.,
 243 2017b), 5 (Li et al., 2023), 6 (Xia et al., 2020), 7 (Yun et al., 2018), 8 (Niu et al., 2022), 9 (Wang et al., 2016),
 244 10 (Jeong et al., 2019), 11 (Phillips et al., 2012), 12 (Bannan et al., 2015), 13 (Thornton et al., 2010), 14
 245 (Osthoff et al., 2008), 15 (Riedel et al., 2012), 16 (Mielke et al., 2011).

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