

# 1 Impact of crop type on the GHG emissions of a rewetted cultivated 2 peatland

3 Kristiina Lång\*, Henri Honkanen, Jaakko Heikkinen, Sanna Saarnio, Tuula Larmola, Hanna Kekkonen  
4 Natural Resources Institute Finland, Latokartanonkaari 9, FI-00790 Helsinki, Finland

5 *Correspondence to:* Kristiina Lång (kristiina.lang@luke.fi)

6

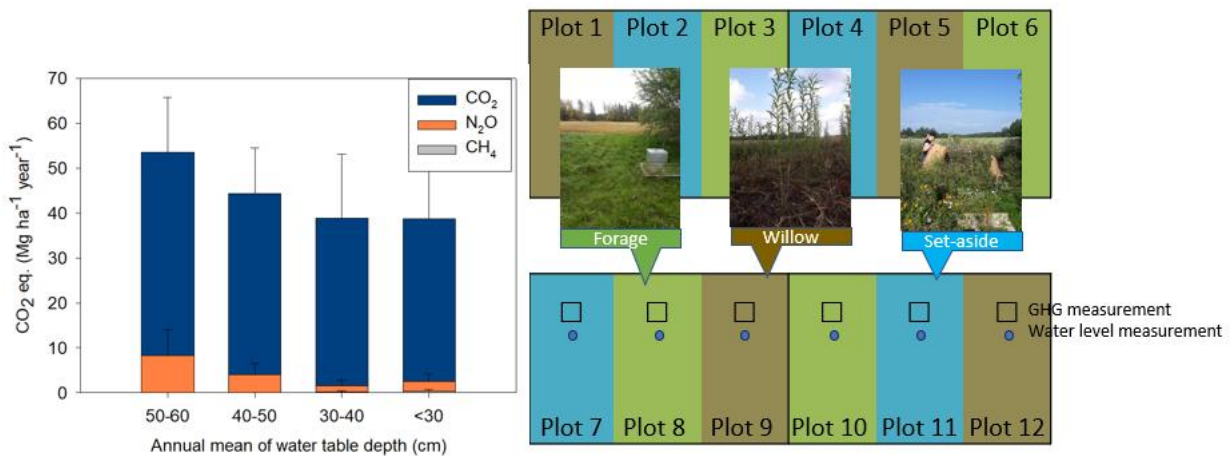
7 **Abstract.** Raising the water table is an effective way to abate greenhouse gas emissions from cultivated peat soils. We  
8 experimented a gradual water table rise at a highly degraded agricultural peat soil site with plots of willow, forage and mixed  
9 vegetation (set-aside) in southern Finland. We measured the emissions of carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>) and nitrous  
10 oxide (N<sub>2</sub>O) for four years. The mean annual ground water table depth was about 80, 40, 40 and 30 cm in 2019-2022,  
11 respectively. The results indicated that a 10 cm raise in the water table depth was able to slow down annual CO<sub>2</sub> emissions  
12 from soil respiration by 0.87 Mg CO<sub>2</sub>-C ha<sup>-1</sup>. CH<sub>4</sub> fluxes changed from uptake to emissions with a raise in the water table  
13 depth, and the maximum mean annual emission rate was 11 kg CH<sub>4</sub>-C ha<sup>-1</sup>. Nitrous oxide emissions ranged from 2 to 33 kg  
14 N<sub>2</sub>O-N ha<sup>-1</sup> year; they were high from bare soil in the beginning of the experiment but decreased towards the end of the  
15 experiment. Short rotation cropping of willow reached net sequestration of carbon before harvest, but all treatments and years  
16 showed net loss of carbon based on the net ecosystem carbon balance. Overall, the short rotation coppice of willow had the  
17 most favourable carbon and greenhouse gas balance over the years (10 Mg CO<sub>2</sub> equivalent on the average over four years).  
18 The total greenhouse gas balance of the forage and set-aside treatments did not go under 27 Mg CO<sub>2</sub> equivalent ha<sup>-1</sup> year<sup>-1</sup>  
19 highlighting the challenge in curbing peat decomposition in highly degraded cultivated peatlands.

20

21 Keywords: peatland, greenhouse gas, ground water table, paludiculture, land use

22

23 Graphical abstract



24

25

26

27

## 28 **1. Introduction**

29 Cultivated peatlands are a major source of greenhouse gas (GHG) emissions globally (Strack et al., 2022). Conventional  
30 cultivation requires lowering the water table depth (WTD) which makes all peat above the drainage depth prone to microbial  
31 decomposition. Intensive management together with the high carbon and nitrogen content of peat makes agricultural peat  
32 soils the highest CO<sub>2</sub> and N<sub>2</sub>O emitters per unit area compared to any other land use types on peat soils (Maljanen et al.,  
33 2010). Their GHG emissions currently diminish the net carbon sink of peat-rich countries significantly which can also be  
34 turned to an advance: the climate change mitigation potential of drained peatlands is high (Humpenöder et al., 2020; Leifeld  
35 and Menichetti, 2018) and cost per mitigated unit of CO<sub>2</sub> equivalent low (Lehtonen et al., 2022).

36 WTD is the major controller of GHG fluxes from peat soils (Evans et al., 2021; Wilson et al., 2016). A global meta-analysis  
37 on water table manipulation studies showed that WTD explained most of the variation in GHG emissions but e.g. climate  
38 zone had some influence as well (Huang et al., 2021). Rewetting has been found to diminish the release of CO<sub>2</sub> and N<sub>2</sub>O from  
39 decomposition but the switch from aerobic to anaerobic decomposition may change the ecosystem from a sink to source of  
40 CH<sub>4</sub>. However, the average increase in CH<sub>4</sub> emissions usually does not compromise the net GHG mitigation potential (Bianchi  
41 et al., 2021; Guenther et al., 2020; Mander et al., 2023) but data is needed to understand the factors regulating CH<sub>4</sub> emissions  
42 that can sometimes be high after rewetting (Nielsen et al., 2023).

43 Paludiculture, i.e. crop production in wet conditions on peat soils, is a GHG mitigation method that allows for slowing down  
44 peat decomposition while still maintaining agricultural income from peatlands for the landowner (Tanneberger et al., 2022).  
45 It is an opportunity for the agribusiness to improve the overall sustainability (Freeman et al., 2022; Liu et al., 2023) and it  
46 produces clearly more societal benefits regarding ecosystem services than conventional management (Liu et al., 2023). As  
47 regards to GHG mitigation, the raise in WTD reduces carbon losses from peat decomposition but export of carbon in the  
48 harvest impairs the carbon balance of the system (Beetz et al., 2013). Emissions of N<sub>2</sub>O are generally found to be low in  
49 paludiculture (Bianchi et al., 2021) but they can remain high if fertilisers are applied (Bockermann et al., 2024). Emissions  
50 of CH<sub>4</sub> are affected by the crop type, harvest management and N fertilisation (Boonman et al., 2023) but they can be efficiently  
51 reduced by leaving an oxidised, non-waterlogged, layer on the peat surface to facilitate microbial oxidation of CH<sub>4</sub> (Kandel  
52 et al., 2020). Solutions for paludiculture implementation are e.g. forage and willow that can be produced in wet conditions  
53 because their roots improve the bearing capacity of the peat and thus ease machine work in wet conditions. Compared to  
54 restoration to natural conditions, paludiculture leads to compromises, as both ecosystem services and economic productivity  
55 are expected to be maintained, and it is not well known how these two aspects are best harmonised in practice. Set-aside is  
56 often not a planned management option, but wet fields drift to non-productive use when the drainage system degrades, and  
57 there are limited data on the GHG balance of such fields.

58 We established an experimental site with forage, willow and set-aside treatments in wet management on highly decomposed  
59 cultivated peat soil in southern Finland in 2019-2023. As the target WTD of -20 cm below the surface was reached only  
60 periodically, we cannot call the site a paludiculture site, but the results can be used to discuss the effects and practical issues  
61 during the transition period to paludiculture. Our research questions were 1) What is the carbon and GHG balance of a  
62 moderately rewetted drained peatland, 2) How much does harvesting reduce the potential to improve the carbon balance and  
63 3) Do CH<sub>4</sub> emissions compromise the GHG mitigation in wet management?

## 66 **2. Materials and methods**

### 67 **2.1. The site and management**

68 The site was located in southern Finland (60.22 °N, 24.78 °E, 110 m a.s.l.) and it has been in cultivation at least since the 19<sup>th</sup>  
69 century. The field has been in a crop rotation with cereals and grass during the latest decades. The climate is boreal humid

70 with long term (1991–2021) annual mean temperature of 5.2 °C and precipitation of 621 mm (Jokinen et al. 2021). The sum  
 71 of annual global radiation is 3358 MJ m<sup>-2</sup> and total sunshine duration 1699 hours. Typically, the soil is frozen and has a snow  
 72 cover from December to March-April. The field was a highly decomposed fen with peat depth ranging from 0.8 to over 2 m.  
 73 Organic carbon content was 25% and pH 5.5 in the surface layer (0–20 cm) (Table 1). The original subsurface drainage  
 74 system with tile drains was replaced by modern plastic pipes surrounded by gravel in the 1960s. The distance between the  
 75 pipes was 18 m until 1979 when it was changed to 9 m. The drainage depth was 60–80 cm, and a control well was installed  
 76 prior to the experiment to restrict water outflow and raise the ground water table. The adjustable tube inside the well was set  
 77 to a position letting the water out when the water table reached 20 cm depth below the soil surface.

78  
 79 **Table 1:** Soil properties ( $\pm$ standard deviation) in the 0-20 cm layer in 2021

Variable	Value
Decomposition status (von Post)	8 (7-9)
Bulk density (g cm <sup>-3</sup> )	0.39 $\pm$ 0.05
Porosity (%)	0.80 $\pm$ 0.02
Ash (%)	42 $\pm$ 3.8
pH	5.4 $\pm$ 0.09
C (g kg <sup>-1</sup> )	286 $\pm$ 24.6
N (g kg <sup>-1</sup> )	15.2 $\pm$ 1.24
Tot P (g kg <sup>-1</sup> )	0.97 $\pm$ 0.08
Soluble P (g kg <sup>-1</sup> )	0.01 $\pm$ 0.001
K (g kg <sup>-1</sup> )	0.17 $\pm$ 0.03
Mn (g kg <sup>-1</sup> )	0.15 $\pm$ 0.02
S (g kg <sup>-1</sup> )	2.01 $\pm$ 0.13
Al (g kg <sup>-1</sup> )	1.41 $\pm$ 0.12
Fe (g kg <sup>-1</sup> )	5.92 $\pm$ 0.63

80  
 81 The site was established in 2018 and it consists of twelve experimental plots (9  $\times$  6 m) in four blocks (see the graphical  
 82 abstract). Four replicate plots with either grass mixture for forage (sown with *Poa trivialis* and *Festuca pratensis*, replanted  
 83 in 2019 and 2021 with *Phleum pratense*, *Festuca pratensis*, *Lolium multiflorum* and *Poa pratensis*), bog bilberry (*Vaccinium*  
 84 *uliginosum*; aka bog blueberry or bog whortleberry) or willow variety Klara (hybrid of *Salix schwerinii* Amgunkaja  $\times$  *Salix*  
 85 *viminialis* Ivar) were randomly assigned within the four blocks. The grass was seeded and bilberry seedlings and willow  
 86 saplings planted in June 2018 (Table S1). The bilberry did not grow roots, and those plots were left to develop to set-aside  
 87 during the following years, thus we named this treatment as “set-aside”. The number of species in all twelve plots was  
 88 determined once in the summer 2021.

89  
 90 **2.2 Ancillary measurements**

91 Biomass growth of willow was monitored by cutting three willow individuals from each plot for determining the above-  
 92 ground biomass each June. The leaves, and stem + branches were separated and weighed to determine the fresh biomass. The  
 93 woody biomass was cut in 10 cm pieces and dried at 65 °C for two weeks. The root biomass around one of the monitored  
 94 plants per plot was determined by taking 50  $\times$  80  $\times$  20 cm peat samples from three layers: 0–20, 20–40 and 40–60 cm once  
 95 per year. Visible large (>2 mm) and fine roots were manually separated from the peat, dried and weighed. For determining  
 96 fine roots, the peat samples were mixed, and a 1 kg subsample was taken. Annual growth in stem, stool and coarse roots was  
 97 calculated by subtracting the value from the previous year. Annual turnover rate of fine roots was assumed to be three times  
 98 the biomass of fine roots as in Pacaldo et al. (2014). For example, biomass increment in 2019 was calculated with the  
 99 following equation:

100  
 101 
$$\text{Annual Growth} = F_{19} + (S_{20} - S_{19}) + (St_{20} - St_{19}) + (Cr_{20} - Cr_{19}) + 3 * Fr_{19} \quad (1)$$

103 , where F19 is foliage in 2019, S19 and S20 are stems in 2019 and 2020, St19 and St20 are stools in 2019 and 2020, Cr19  
104 and Cr20 are coarse roots in 2019 and 2020 and Fr19 is fine roots in 2019. Subsamples were taken for determining the C  
105 content of the dried biomass in 2019 and 2020, and the mean values were used for the following years. The yield per hectare  
106 was estimated to be the weight of 25,000 individuals, based on 80 cm × 50 cm spacing.

107 Soil temperature was measured first at the depth of 10 cm (but at the depth of 5 cm from May 2020 on to achieve better  
108 response of CO<sub>2</sub> to air temperature) in each treatment with Elcolog sensors (Elcoplast Oy, Tampere, Finland). The sampling  
109 rate was one hour in summer and 2.5 hours in winter. The air temperature, precipitation and radiation data were taken from  
110 the Jokioinen weather station of Finnish Meteorological Institute (FMI 2024, CC BY 4.0) located about 10 km from the site.  
111 Continuous photosynthetically active radiation (PAR) data was produced with global radiation data from FMI and corrected  
112 using the ratio of 2.04 for global radiation and the PAR (Meek et al., 1984).

113 WTD was measured from monitoring pipes at the corners of the site at the time of the opaque chamber measurements until  
114 2021 when monitoring pipes were installed also in the centre of each plot. During summers 2021 and 2022, there were also  
115 HOBO Water Level data loggers (Onset, Bourne, United States) in each plot for continuous water level monitoring with a  
116 sampling rate of one hour. In winter when the loggers were not used, WTD was measured manually from monitoring pipes  
117 when the water was not frozen.

118 Leaf area index (LAI) was measured at the same time with the transparent chamber measurements with a portable LAI meter  
119 (SunScan; Delta-T Devices Ltd, Cambridge, United Kingdom). LAI values > 3 were set to 3 as they were assumed to not  
120 affect photosynthesis due to saturation of the reflectance (Aparicio et al., 2000). When harvesting the grass plots, the previous  
121 measured LAI value was extrapolated to the moment just before harvesting, after which the LAI value was set to 1 as  
122 measured. In 2022, we measured green canopy cover with the Canopeo app (Patrignani and Ochsner, 2015) instead of LAI.  
123 Based on our experiences, and due to the operation and physical design of the LAI device, it did not provide as comprehensive  
124 picture of the biomass inside the gas measurement collar as Canopeo. Vegetation index has been found to be faster to measure  
125 and less dependent on the ambient light conditions than the light interception method (Shepherd et al. 2018). LAI was indexed  
126 by dividing by maximum value 3 and green canopy cover by 100 (values from 0% to 100%) so that the generated vegetation  
127 index range was 0–1. The vegetation index was set to 0 from the end of November until mid-April when the snow and frost  
128 covered the ground or no green vegetation was present.

129 Soil samples for analysing the soil properties were taken first in October 2018 and another sampling was conducted in June  
130 2021 with additional analysis. As there were no significant changes in the soil properties between these samplings, we present  
131 only the results of the second soil sampling in Table 1. The samples were taken from the 0-20 cm layer using a soil corer with  
132 a diameter of 3 cm. Approx. 20 subsamples were pooled to make composite sample that was air-dried and sieved (2 mm) for  
133 the chemical analyses. Soil core samples for dry bulk density and porosity (diameter 5 cm) were taken from the surface layer  
134 (0–17.5 cm) of each plot in Oct 2020 using the Kopec corer, and the samples were dried at 37 °C for a week. Soil acidity was  
135 determined using the ISO 10390 method. Nutrient content was analysed as described in Vuorinen and Mäkitie (1955). Soil  
136 carbon and nitrogen were determined using the dry combustion method (Leco TruMac CN, LECO corporation, MI).

137

### 138 **2.3. GHG measurements**

139 Dark respiration of the plants together with soil respiration (ecosystem respiration) and fluxes of N<sub>2</sub>O and CH<sub>4</sub> were measured  
140 using opaque chambers biweekly or once per month in the winter between 3/2019 – 3/2023. In each plot, a 60 cm × 60 cm  
141 steel collar was installed at the depth of 10–15 cm. The location of the collars was one metre from the short edge of the plot  
142 and three metres from the edges of the adjacent plots. An aluminium chamber (height 40 cm) mounted at the top of the collar  
143 was sealed with water in the groove of the upper edge of the collar. In the winter, NaCl was added to the water to avoid ice  
144 formation. The clear aluminium surface reflected effectively light and kept the temperature change moderate inside the  
145 chamber. The measurements were done during the daytime between 10 am and 2 pm approximately every two weeks in

146 summertime, and monthly in the winter. The chambers were closed for 30 minutes, and four 20 ml gas samples were taken  
147 with a 60-ml plastic syringe to pre-evacuated vials (Exetainer, Labco Limited, UK) in 10-minute intervals starting  
148 immediately after closing. Prior to sampling, the syringe was pumped five times to mix the air in the chamber. The samples  
149 were analysed with a gas chromatograph (Agilent 7890 Agilent Technologies, Inc., Wilmington, DE, USA) equipped with  
150 flame ionizer and electron capture detectors, and a nickel catalyst for converting CO<sub>2</sub> to CH<sub>4</sub>. The gas chromatograph had a  
151 2 ml sample loop and a backflush system for separating water from the sample and flushing the precolumn between the runs.  
152 The precolumn and analytical columns consisted of 1.8 and 3 m long steel columns, respectively, and were packed with  
153 80/100 mesh Hayesep Q (Supelco Inc., Bellefonte, PA, USA). Nitrogen was used as the carrier gas and a standard gas mixture  
154 of known concentration of CO<sub>2</sub>, N<sub>2</sub>O and CH<sub>4</sub> was used for a calibration curve with seven concentration points. An  
155 autosampler (222 XL Liquid handler, Gilson Medical Electronics, France) fed the samples to the loop of the gas  
156 chromatograph.

157 Net ecosystem exchange (NEE) including photosynthesis and respiration of the soil and plants was measured approximately  
158 every two weeks during the growing season using a transparent chamber (60 × 60 × 60 cm) made of polycarbonate plexiglass  
159 (1 mm, light transmission 95%). The chamber was equipped with a Vaisala GMP-343 probe for CO<sub>2</sub> measurement and a  
160 temperature and humidity sensor (Vaisala Oy, Vantaa, Finland) and two fans for mixing the air during the measurement. PAR  
161 was measured with LI-190 quantum PAR sensor (LI-COR, Lincoln, Nebraska, USA) inside the chamber. Four measurements  
162 with different amount of entering light were taken from each plot on each measurement day in order to cover a large range of  
163 light conditions and facilitate the gap filling by modelling. One or two layers of a white fabric shroud and one blackout curtain  
164 were used to acquire measurement results in different light conditions (approximately 100%, 50%, 25%, 0% of ambient  
165 radiation). The measurement with 0% radiation gave the estimate for ecosystem respiration (ER). The measurements were  
166 done in the same collars as the opaque chamber measurements. Each measurement took one minute with a five second  
167 sampling rate, or two minutes in early or late growing season when the change in flux was minor. The chamber was flushed  
168 after each measurement to reconstitute ambient CO<sub>2</sub> and air humidity contents. After closing the chamber, a lag time of 10  
169 seconds was applied to exclude the time when the flux was not yet stabilised. Clear sky conditions were preferred to avoid  
170 problems related to changing cloud cover and to achieve the widest possible range of available light. The temperature change  
171 inside the chamber was less than 1.5 degrees which was also used as a criterion for data filtering.

172 The change of CO<sub>2</sub> concentration during the chamber enclosure was assumed to be linear. The measurement results of CO<sub>2</sub>  
173 as parts per million (ppm) unit were converted to g m<sup>-2</sup> h<sup>-1</sup> by the ideal gas law using measured temperature inside the chamber.  
174 If the flux was not yet stabilized at the beginning (first 4 datapoints) of the measurement, outliers were defined with Matlab  
175 isoutlier command resulting removal of 210 of total 23066 datapoints in 1564 flux measurements.

176 If the snow cover was thicker than 20 cm, a concentration snow gradient method as in Maljanen et al. (2003) was used to  
177 determine the GHG fluxes. A probe made of a steel pipe (Ø 3 mm), with a three-way valve and a plastic syringe, was used to  
178 sample 15 ml of air just above the snow cover, in the bottom of the snow cover and at every 10 cm in between in three  
179 replicate locations per plot. The gas was stored in the pre-evacuated vials and the concentrations were determined gas  
180 chromatographically.

181 Measurements for bare soil respiration were made in unvegetated subplots in 7/2019–12/2022. For willow, the large 60 × 60  
182 cm frames were used but for forage and set-aside we installed one sheet metal air ventilation pipe 27 cm in diameter and 30  
183 cm in length to the depth of 5–10 cm next to the opaque chamber collars in the 8 plots of grass and set-aside. All green  
184 vegetation within the chamber area was removed and root growth was limited by cutting around the chamber occasionally  
185 with a knife. For the measurements, the cylinders were closed with a cover equipped with a CO<sub>2</sub> sensor (GMP-343; Vaisala  
186 Oyj, Vantaa, Finland) and a small fan. One measurement lasted for one minute with a five second sampling rate.  
187 Measurements were taken about once in a week or two, more frequently in summer than in winter. In winters 2021 – 2022

188 and 2022 – 2023 this method was not used due to too high snow depth but measurements with the snow gradient method were  
189 utilized (Maljanen et al., 2003).

190

#### 191 **2.4. Flux modelling**

192 Gross photosynthesis (GP) can be determined as the difference of NEE and ER. Instantaneous GP was estimated for each  
193 measurement occasion by (equation 2),

194

$$195 \quad GP = NEE - ER \quad (2)$$

196

197 , where the full darkened transparent chamber measurement result (ER) is subtracted from the light-dependent flux (NEE)  
198 measured during the same day. Thus, we follow the sign convention with positive ER and negative GP values.

199 The gaps in GP and ER data between the measurement occasions were predicted using hourly timeseries of the ancillary data.

200 Hourly time points for vegetation index, WTD and soil temperature and were acquired from the measured values by linear  
201 interpolation. Gaps in soil temperature were filled with the modified soil temperature model (Zheng et al., 1993) using the air  
202 temperature. Air temperature and PAR were assumed to be the same for all plots, whereas we used plot specific vegetation  
203 index and the soil temperature from the certain treatment. Hourly ER and GP were modelled using nonlinear regression  
204 (fitnlm function in MATLAB) for all 8 plots in forage and set-aside treatments. Empirical models were used for ER as in  
205 Lohila et al. (2003) and for GP as in Kandel et al. (2013). Instead of the phytomass indices used in the above publications,  
206 we used vegetation index formed according to the LAI and Canopeo measurements (index from 0 to 1) to describe the stage  
207 of the crop growth.

208 We used the following equation first defined by (Long and Hällgren, 1993) for GP to estimate empirical coefficients ( $A_{max}$   
209 and  $k$ ):

210

$$211 \quad GP = \frac{A_{Max} * PAR}{k + PAR} * VI * T_{Scale} \quad (3)$$

212

213 , where PAR is the measured photosynthetically active radiation ( $\mu\text{mol m}^{-2} \text{s}^{-1}$ ), VI is the vegetation index,  $A_{max}$  is the  
214 asymptotic maximum ( $\text{g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ), and  $k$  is a half-saturation value ( $\mu\text{mol m}^{-2} \text{s}^{-1}$ ).  $T_{Scale}$  represents the temperature  
215 sensitivity of photosynthesis and follows the equation presented by (Raich et al., 1991):

216

$$217 \quad T_{Scale} = \frac{(T - T_{min})(T - T_{max})}{(T - T_{min})(T - T_{max}) - (T - T_{opt})^2} \quad (4)$$

218

219 , where  $T$  is the measured temperature, photosynthetically active minimum temperature  $T_{min}$  is  $-2 \text{ }^\circ\text{C}$ , maximum  $T_{max}$  is  $40 \text{ }^\circ\text{C}$   
220 and the optimum is  $20 \text{ }^\circ\text{C}$  as in (Kandel et al., 2013).

221 ER was estimated using data from the opaque and fully darkened transparent chambers. The empirical coefficients ( $R0_s$ ,  $R0_p$ ,  
222  $E0_s$  and  $b$ ) were estimated with a nonlinear regression model similarly as in the case of GP. Annual fluxes were computed as  
223 sum of the hourly fluxes with a trapezoidal method (trapz function in Matlab 2019b). ER consists of autotrophic ( $R_{auto}$ ), i.e.,  
224 plant respiration and heterotrophic ( $R_{hetero}$ ), i.e., soil respiration (LLOYD and TAYLOR, 1994) with extension of WTD as in  
225 (Karki et al., 2014):

226

$$227 \quad ER = R_{hetero} + R_{auto} \quad (5)$$

228

$$228 \quad R_{hetero} = R0_s * \exp\left(E0_s \left(\frac{1}{56.02} - \frac{1}{T_{soil} + 46.02}\right)\right) + b * WTD \quad (6)$$

$$R_{auto} = VI * R_{0p} * \exp \left( b_d \left( \frac{1}{10+273} - \frac{1}{T_{air}+273} \right) \right) \quad (7)$$

230  
 231 , where  $T_{soil}$  is the measured soil temperature, VI is the vegetation index,  $T_{air}$  is the measured air temperature,  $R_{0s}$  is soil  
 232 respiration at the reference temperature 10 °C ( $\text{g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ),  $R_{0p}$  is plant respiration at the reference temperature at 10 °C  
 233 ( $\text{g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ),  $b$  is the effect of WTD,  $E_0s$  is ecosystem sensitivity and  $b_d$  is the temperature dependence of dark respiration  
 234 set to 5000 as in (Lohila et al., 2003). Bare soil respiration was estimated like ER but using only equation 6. The estimated  
 235 parameters  $R_{0s}$ ,  $E_s$  and  $b$  and model R2 are shown in Table S3.

236  
 237  
 238

## 239 **2.5. Data processing and analysis**

240 For the transparent chamber measurements, the criterion  $R^2 > 0.9$  for the fitted linear assumption of flux measurements would  
 241 exclude a large amount of data, especially with a small change in  $\text{CO}_2$ , leading to a biased dataset. Therefore, we decided to  
 242 add the criterion  $S_{xy} < 2.3 \text{ g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$  for dataset like in (Kutzbach et al., 2007) ( $S_{xy}$  is the standard deviation of the residuals  
 243 and  $2.3 \text{ g m}^{-2} \text{ h}^{-1}$  is the 95% percentile of measurements). This procedure resulted in the removal of 59 values out of total  
 244 1467 measurements. In the modelling phase, fitted values were examined, and outliers were removed to avoid distortion.  
 245 Outliers were defined as observations with an absolute value of standardised residuals greater than three. In 2019, 3 out of  
 246 260 GP values and 3 of 243 ER values were removed. In 2020, none of 200 GP values and 2 of 230 ER values were removed.  
 247 2 out of 365 GP values and 4 of 247 ER values were removed in 2021 and 12 out of 583 GP values and 2 of 323 ER values  
 248 were removed in 2022. The model's estimated parameters  $A_{max}$ ,  $k$  of GP,  $R_{0s}$ ,  $R_{0p}$ ,  $E_s$  and  $b$  of ER and model correlations are  
 249 shown in Table S2. The measured versus model predicted values of GP and ER are shown by treatments and years in Fig. S1.  
 250 For bare soil respiration measurements in set-aside and forage, the same criteria were used as for transparent chamber ( $R^2 >$   
 251  $0.9$  and  $S_{xy} < 95\%$ ) leading to a removal of 12 values of total 601. In bare soil measurements in midsummer 2022, there  
 252 occurred 24 flux measurements (9 values in one plot, 0–5 in others) of total 147 values which were unexplained high (3–18  
 253  $\text{g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ ). Values were of the same magnitude as values measured immediately after ploughing in (Honkanen et al.,  
 254 2023). We decided to remove these values as outliers to avoid model distortion. Soil respiration of willow was defined with  
 255 opaque chamber method and such outliers did not occur in these measurements. In modelling phase, outliers defined as  
 256 observations with an absolute value of standardised residuals greater than three, were removed resulting removal of 13  
 257 measurements of total 984 measurements (including all plots).

258 A linear regression model was fitted to calculate gas concentrations and the ideal gas law was used to solve the flux rate for  
 259 every enclosure of the opaque chambers. Nonlinear responses of  $\text{CO}_2$  indicated a leaking chamber or other problem in the  
 260 measurement and thus, if the  $R^2$  of  $\text{CO}_2$  was less than 0.9, also the results of  $\text{CH}_4$  and  $\text{N}_2\text{O}$  were discarded. In addition, sudden  
 261 variations in  $\text{CH}_4$  fluxes due to ebullition were filtered by selecting only flux rates with the intercept between 1.5 and 2.4  
 262 ppm. These criteria resulted to 176, 117 and 118 discarded values out of 1044 in the case of  $\text{CH}_4$ ,  $\text{CO}_2$  and  $\text{N}_2\text{O}$ , respectively.  
 263 All data cleaning and processing was done with Matlab (The Math Works, Inc., MATLAB, version 2019b).

264

## 265 **2.6. GHG balance**

266 The annual net ecosystem carbon balance was constructed as the sum of the hourly values of NEE and yield data for each  
 267 year in the case of forage and set-aside treatments. Modelling was used to fill the gaps between the measurement occasions  
 268 to create a continuous series of hourly values. For willow, annual estimates of carbon loss in soil respiration were available  
 269 from the chamber measurements from the unvegetated frames but as the willows were too high for the chambers their net  
 270 production had to be estimated based on biomass accumulation during four years (Pacaldo et al., 2014). The presented net  
 271 ecosystem carbon balance of willow is thus the sum of average annual  $\text{CO}_2\text{-C}$  from soil respiration and average annual amount

272 of carbon bound in the biomass during the four first cultivation years. The cumulative annual fluxes of CH<sub>4</sub> and N<sub>2</sub>O for each  
 273 management practice were calculated by interpolating the emissions between consecutive sampling days. Global warming  
 274 potentials 27 and 273 were used for CH<sub>4</sub> and N<sub>2</sub>O, respectively, to convert the results to CO<sub>2</sub> equivalents for the total GHG  
 275 balance (Forster et al. 2021).

276

## 277 2.7. Statistical analyses

278 Linear mixed models were used to find variables explaining variation in the gas fluxes. Crop, year, WTD and all their  
 279 interactions were denoted as fixed effects. Block and block × year were assumed to be independent and normally distributed  
 280 random effects. The most suitable covariance structure was chosen using Akaike's Information Criterion (AIC). The models  
 281 were fitted using the residual maximum likelihood (REML) method and degrees of freedom were estimated using the  
 282 Kenward-Roger method. The residuals were plotted against the fitted values and the normality of the residuals were checked  
 283 using boxplots. The data was log--transformed when needed to normalise the distributions. The method of Tukey-Kramer  
 284 was used for all pairwise comparisons of means with a significance level of 0.05. After the first model run with all relevant  
 285 variables the non-significant variables were removed one by one to find the most relevant effects. All statistical analyses were  
 286 performed using the SAS Enterprise Guide v7.1 (SAS Institute Inc., Cary, NC, USA).

287

## 288 3. Results

### 289 3.1 Climate and site variables

290 Annual mean temperature was 6.9, 6.0, 5.8 and 5.8 °C and annual precipitation 750, 600, 660, and 546 mm in 2019 – 2022,  
 291 respectively. Number of days with a snow-cover on the soil within each modelling year (April to March) was 13, 81, 108 and  
 292 118, respectively. The annual mean temperature during the study years was higher than the long-term average of 5.2°C in  
 293 1991–2020 (Jokinen et al., 2021). Two study years exhibited lower and two higher annual precipitation as compared to the  
 294 long-term mean of 621 mm. The WTD showed an increasing trend in time and high within-year variation (Fig. 1). The average  
 295 WTD was -54, -41, -39 and -27 cm in 2019–2022, respectively. WTD varied from -89 to -4, -77 to 2, -120 to 1.4 and -100 to  
 296 1.8 cm in 2019 – 2022, respectively.

297 The forage yields were 6.3±0.9, 8.9±0.7, 11±0.8 and 9.4±0.9 Mg DM ha<sup>-1</sup> in 2019-2022, respectively. There were two harvests  
 298 in 2020 and three in the other years. The plots were dominated by *Phleum pratense* and *Festuca pratensis* in 2021. The dry  
 299 mass yields of willow were 30±14 and 73±28 Mg DM ha<sup>-1</sup> in the harvests of February 2021 and 2023. Most of the C  
 300 accumulation occurred in the stem (59%), followed by stool (25%) and roots (9%) and foliage (7%) (Table 2). Vegetation of  
 301 the set-aside plots in 2021 was dominated by wild plants belonging to Families *Asteraceae*, *Cichoriaceae* and  
 302 *Caryophyllaceae*. Bog bilberry covered one percentage or less on each of the four replicate plots. The set-aside vegetation  
 303 had the highest species diversity, 19 vascular plants compared to 12 at willow plots and 9 at forage plots, the two latter  
 304 including crop plants.

305

306 **Table 2:** Four-year cumulative carbon balance of willow (±standard deviation). Negative sign indicates sequestered carbon  
 307 and positive sign released carbon to the atmosphere.

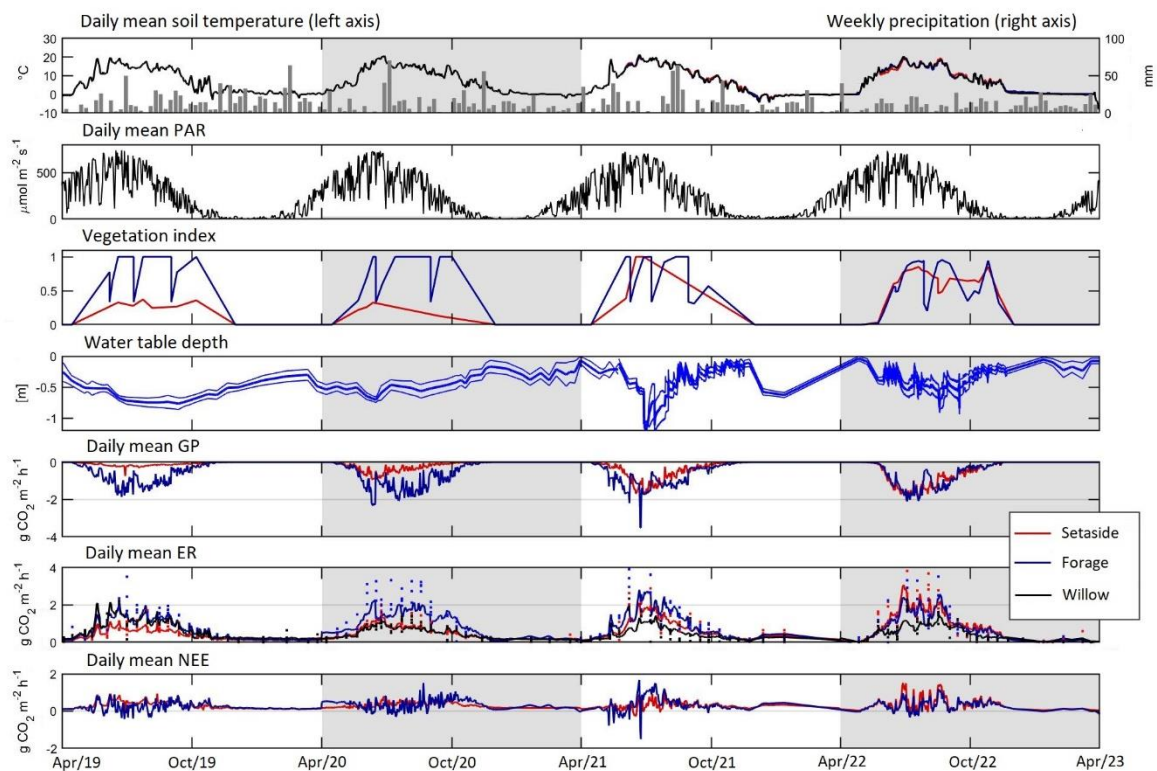
Component	Mg C ha <sup>-1</sup> 4 yrs <sup>-1</sup>	% of total
Stem (harvested)	-50.7±14.3	59
Foliage	-6.2±0.8	7
Aboveground stool	-12.6±3.2	15
Underground stool	-8.5±3.8	10
Coarse roots	-3.9±1.5	4
Fine roots	-4.6±0.6	5
Total sequestered carbon	-86.5±19.5	



Soil respiration	43.5±2.7
Net ecosystem exchange	-43.1±21.1
Net ecosystem carbon balance	7.6±7.7

308

309



310

311

312

313

314

**Figure 1:** Daily mean of soil temperature and precipitation, photosynthetically active radiation (PAR), vegetation index, water table depth (site mean±std), gross photosynthesis, measured (dots) and model predicted (line) ecosystem respiration (soil respiration for willow) and net ecosystem exchange. Annual modelling periods (Apr - Mar) are marked with light grey or white background.

315

316

317

### 3.2. Carbon balance

318

319

320

321

322

323

324

325

326

327

328

Model predicted maximum hourly GP was -0.7, -3.2, -4.3 and -4.8 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> in the set-aside in 2019 – 2020, and -3.9, -5.9, -4.3 and -4.8 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> in the forage plots in 2019–2022, respectively (Fig. S1). Maximum measured GP value was -1.1, -2.4, -3.4 and -4.5 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> for set-aside and -3.4, -6.2, -4.8 and -4.2 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> for forage in 2019–2020, respectively. Annual values of GP varied from -9.3 to -12 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in the forage and from -1.5 to -10 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in the set-aside treatment (Table 3). The variables initially included in the analysis were annual mean WTD, crop type and year as main effects and all their interactions. Finally, only crop, year and their interaction were left in the model (Table S4). Inclusion of WTD did not result in a meaningful estimate as both the productivity of especially the set-aside established in 2019 and the WTD increased with years, and the observed change in GP did not in reality increase with WTD but with time. Forage and set-aside treatments differed significantly (p<0.001), there was an increasing trend in productivity from 2019–2022 (p=0.0009) and the differences between the crop types were highest in 2019–2021 (p<0.001).

330 **Table 3:** The estimated annual sums ( $\pm$ standard deviation) of gross photosynthesis (GP), ecosystem respiration (ER), net  
 331 ecosystem exchange (NEE), carbon exported in the harvested yield, net ecosystem carbon balance (NECB), N<sub>2</sub>O and CH<sub>4</sub>  
 332 effluxes and the total emissions (global warming potential of one hundred years; GWP-100) with either NEE or NECB  
 333 representing CO<sub>2</sub> emissions in the forage and set-aside plots, and selected data for willow. Significant differences ( $p < 0.05$ )  
 334 between treatments within a year are denoted with different letters ( $n=4$ ).

Year	Variable and unit	Forage	Set-aside	Willow <sup>a</sup>
2019	GP Mg CO <sub>2</sub> -C ha <sup>-1</sup>	-9.31 $\pm$ 0.75a	-1.48 $\pm$ 0.32b	
	ER Mg CO <sub>2</sub> -C ha <sup>-1</sup>	14.4 $\pm$ 2.17a	8.25 $\pm$ 2.40b	
	NEE Mg CO <sub>2</sub> -C ha <sup>-1</sup>	5.08 $\pm$ 1.80	6.77 $\pm$ 2.41	
	C in yield Mg C ha <sup>-1</sup>	3.17 $\pm$ 0.49	0	
	NECB Mg C ha <sup>-1</sup>	8.25 $\pm$ 2.13	6.77 $\pm$ 2.41	
	Soil respiration Mg CO <sub>2</sub> -C ha <sup>-1</sup>	12.8 $\pm$ 4.99	11.4 $\pm$ 1.82	14.8 $\pm$ 0.76
	N <sub>2</sub> O-N kg ha <sup>-1</sup>	11.9 $\pm$ 7.60a	32.6 $\pm$ 12.1b	17.4 $\pm$ 10.3
	CH <sub>4</sub> -C kg ha <sup>-1</sup>	-0.28 $\pm$ 0.75	-1.00 $\pm$ 0.73	-1.64 $\pm$ 0.26
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEE) <sup>b</sup>	23.7 $\pm$ 5.67	38.8 $\pm$ 19.8	
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NECB) <sup>c</sup>	35.3 $\pm$ 6.86	38.8 $\pm$ 19.8	
2020	GP Mg CO <sub>2</sub> -C ha <sup>-1</sup>	-11.7 $\pm$ 1.17a	-3.57 $\pm$ 0.46b	
	ER Mg CO <sub>2</sub> -C ha <sup>-1</sup>	19.3 $\pm$ 1.85a	10.6 $\pm$ 1.49b	
	NEE Mg CO <sub>2</sub> -C ha <sup>-1</sup>	7.64 $\pm$ 1.75	7.05 $\pm$ 1.12	
	C in yield Mg C ha <sup>-1</sup>	3.35 $\pm$ 1.28	0	
	NECB Mg C ha <sup>-1</sup>	11.0 $\pm$ 2.02	7.05 $\pm$ 1.12	
	Soil respiration Mg CO <sub>2</sub> -C ha <sup>-1</sup>	9.09 $\pm$ 4.86	10.6 $\pm$ 0.97	10.0 $\pm$ 0.79
	N <sub>2</sub> O-N kg ha <sup>-1</sup>	6.26 $\pm$ 3.39	6.59 $\pm$ 2.97	4.61 $\pm$ 2.99
	CH <sub>4</sub> -C kg ha <sup>-1</sup>	-0.36 $\pm$ 0.40	-1.01 $\pm$ 0.56	-1.13 $\pm$ 0.33
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEE)	30.4 $\pm$ 7.64	28.7 $\pm$ 3.73	
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEBC)	42.6 $\pm$ 8.69	28.7 $\pm$ 3.73	
2021	GP Mg CO <sub>2</sub> -C ha <sup>-1</sup>	-9.46 $\pm$ 1.20a	-6.34 $\pm$ 0.68b	
	ER Mg CO <sub>2</sub> -C ha <sup>-1</sup>	17.4 $\pm$ 1.40a	13.5 $\pm$ 1.82b	
	NEE Mg CO <sub>2</sub> -C ha <sup>-1</sup>	7.95 $\pm$ 2.16	7.12 $\pm$ 2.16	
	C in yield Mg C ha <sup>-1</sup>	5.54 $\pm$ 0.46	0	
	NECB Mg C ha <sup>-1</sup>	13.5 $\pm$ 1.88a	7.12 $\pm$ 2.16b	
	Soil respiration Mg CO <sub>2</sub> -C ha <sup>-1</sup>	7.82 $\pm$ 2.30	11.5 $\pm$ 2.44	8.99 $\pm$ 2.70
	N <sub>2</sub> O-N kg ha <sup>-1</sup>	6.49 $\pm$ 3.88a	2.18 $\pm$ 0.24b	5.75 $\pm$ 6.25
	CH <sub>4</sub> -C kg ha <sup>-1</sup>	7.92 $\pm$ 12.7	0.58 $\pm$ 1.87	3.89 $\pm$ 6.12
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEE)	32.2 $\pm$ 8.16	27.1 $\pm$ 8.03	
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEBC)	52.2 $\pm$ 6.93	27.1 $\pm$ 8.03	
2022	GP Mg CO <sub>2</sub> -C ha <sup>-1</sup>	-9.34 $\pm$ 2.13	-9.65 $\pm$ 2.08	
	ER Mg CO <sub>2</sub> -C ha <sup>-1</sup>	14.4 $\pm$ 3.29	16.5 $\pm$ 3.21	
	NEE Mg CO <sub>2</sub> -C ha <sup>-1</sup>	5.10 $\pm$ 1.15	6.82 $\pm$ 1.15	
	C in yield Mg C ha <sup>-1</sup>	4.72 $\pm$ 0.50	0	
	NECB Mg C ha <sup>-1</sup>	5.46 $\pm$ 6.37	6.82 $\pm$ 1.15	
	Soil respiration Mg CO <sub>2</sub> -C ha <sup>-1</sup>	8.40 $\pm$ 1.48	15.0 $\pm$ 5.32	9.70 $\pm$ 2.3
	N <sub>2</sub> O-N kg ha <sup>-1</sup>	9.54 $\pm$ 4.49a	3.07 $\pm$ 0.86b	1.69 $\pm$ 1.10b
	CH <sub>4</sub> -C kg ha <sup>-1</sup>	7.74 $\pm$ 0.59	11.3 $\pm$ 7.72	10.9 $\pm$ 12.9
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEE)	22.9 $\pm$ 3.34	26.6 $\pm$ 4.61	
	GWP <sub>100</sub> Mg CO <sub>2</sub> eq ha <sup>-1</sup> (NEBC)	43.3 $\pm$ 3.18	26.6 $\pm$ 4.61	

335 <sup>a</sup>All components of the carbon balance are not available for willow, see chapter 2.6, <sup>b</sup>With NEE representing CO<sub>2</sub>, <sup>c</sup>With  
 336 NEBC representing CO<sub>2</sub>

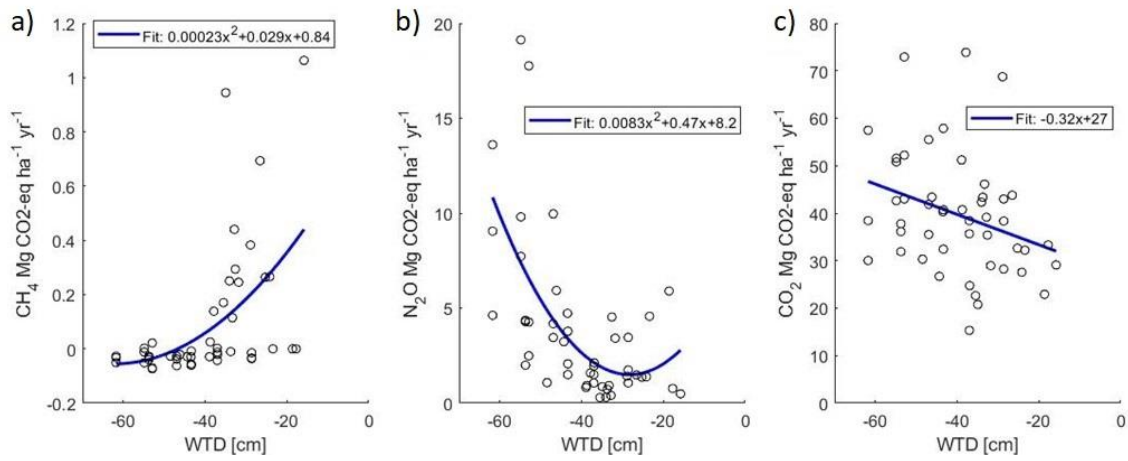
338 Modelled maximum hourly ER was 2.4, 2.3, 3.0 and 4.7 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> in the set-aside and 3.5, 3.4, 4.8 and 4.5 g CO<sub>2</sub>-C m<sup>-2</sup>  
339 h<sup>-1</sup> in the forage plots (Fig. S1). Measured maximum ER with the opaque chamber method was 1.7, 2.0, 2.9 and 4.6 g CO<sub>2</sub>-C  
340 m<sup>-2</sup> h<sup>-1</sup> for the set-aside and 3.5, 4.2, 4.0 and 4.8 g CO<sub>2</sub>-C m<sup>-2</sup> h<sup>-1</sup> for forage. Annual ER varied from 14 to 19 Mg CO<sub>2</sub>-C ha<sup>-1</sup>  
341 yr<sup>-1</sup> in the forage and from 8 to 17 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in set-aside treatment (Table 3). Variables initially included in the  
342 analysis were annual mean WTD, crop type and year as main effects as well as their interactions. WTD did not well explain  
343 the variation in the annual ER estimate, likely for the same reason as for GP as plant respiration is related to the biomass of  
344 the vegetation, which increased during the experimental years. The best model was based on the crop type and year as main  
345 effects, and their interactions (Table S4). In this model, crop type explained ER well, ER increased in time, and ER increased  
346 between years more in forage than set-aside.

347 Hourly model-predicted NEE varied from -2.9 to 3.7 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> in the set-aside and from -4.6 to 3.7 g CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> in the  
348 forage treatment (results not shown). There were 29, 28, 16 and 47 days annually with negative daily NEE in the forage plots  
349 during the study years, respectively, and fewer such days (1, 0, 9 and 6) in the set-aside plots in 2019–2022 (Fig. 1). The  
350 cumulative annual balance ranged from 5.1 to 8.0 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in the forage and from 6.8 to 7.1 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in  
351 the set-aside treatment (Table 3) and the treatments did not differ statistically. The net ecosystem carbon balance (NECB)  
352 that accounts the amount of carbon exported in the harvested yield varied from 5.5 to 13.5 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in the forage  
353 treatment and were equal to NEE in the set-aside treatment (Table 3). The NECB values differed statistically between the  
354 forage and set-aside treatments across all years (p>0.001) and inclusion of additional effects in the analysis did not improve  
355 the model (Table S4).

356 Annual sum of respiration varied from 8 to 15 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> in the different treatments and years (Table 3). The  
357 proportion of soil respiration of the total ecosystem respiration varied from 45 to 90% in the forage plots and from 85 to 100%  
358 in the set-aside plots in 2019–2022. In the set-aside plots, estimated annual bare soil respiration exceeded the estimated ER  
359 in all plots in 2019, two plots in 2020 and one plots in 2021 and 2022 and those values were not used in the above calculation,  
360 and thus it is assumed that total respiration constituted only of soil respiration in 2019. Annual cumulative soil respiration  
361 was explained by WTD (Fig. 2; p=0.053) and crop type (p=0.033) so that forage and set-aside treatments were significantly  
362 different in the whole dataset and respiration increased in the order forage<willow<set-aside (Fig. S4; Table S4). Plots of the  
363 bare soil respiration in relation to WTD and temperature show that there is a clear trend of decreasing respiration with raising  
364 WTD (Fig. S2). Three individual curves indicate a contrasting trend, but these three estimations are based on a small number  
365 of measurement results. Based on all annual estimates of soil respiration, a 0.1 m raise in WTD reduces respiration by 0.87  
366 Mg CO<sub>2</sub>-C ha<sup>-1</sup> yr<sup>-1</sup>.

367 The cumulative total amount of C in the above and below ground willow biomass was 86.5 Mg C ha<sup>-1</sup> during the four study  
368 years (Table 2). About 40% of the carbon in the biomass was left at the site after harvest, and soil respiration amounted to  
369 43.5 Mg ha<sup>-1</sup>, leading to a strongly negative cumulative NEE of -43 Mg ha<sup>-1</sup>. Carbon export in the harvest changed the net  
370 balance to net loss of 7.6 Mg, corresponding to an average annual CO<sub>2</sub> rate of 7 Mg of CO<sub>2</sub>.

371



372

373

374

375

376

377

378

**Fig. 2.** Plot-wise mean annual fluxes of CH<sub>4</sub> (a), N<sub>2</sub>O (b) and soil respiration (c) (CO<sub>2</sub> eq.) as related to the mean annual WTD.

379

380

381

382

383

384

385

386

387

### 3.3. CH<sub>4</sub> fluxes

Hourly fluxes of CH<sub>4</sub> varied between -50 and 30 μg CH<sub>4</sub>-C m<sup>-2</sup> h<sup>-1</sup> during the first half of the experimental period (Fig. 3). In conjunction with the raise in the WTD the values the hourly fluxes increased and varied between -40 and 900 μg m<sup>-2</sup> h<sup>-1</sup> during the latter half of the period. The annual flux of CH<sub>4</sub> varied from -1.6 to 11 kg CH<sub>4</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> with an increasing trend towards the end of the measurement period (Table 3; Fig. S3). When the mean annual WTD was below 40 cm the soil was mainly consuming CH<sub>4</sub>, but the consumption tended to change to emissions as the WTD raised (Fig. 2). Variation in the annual cumulative fluxes of each plot was explained by the WTD (p=0.015) but not by crop (Table S4). The increasing trend between years 2019–2022 was also shown in the mixed model analysis as year had a significant effect (p=0.0003) and the effect of WTD decreased with time (years).

388

389

390

391

392

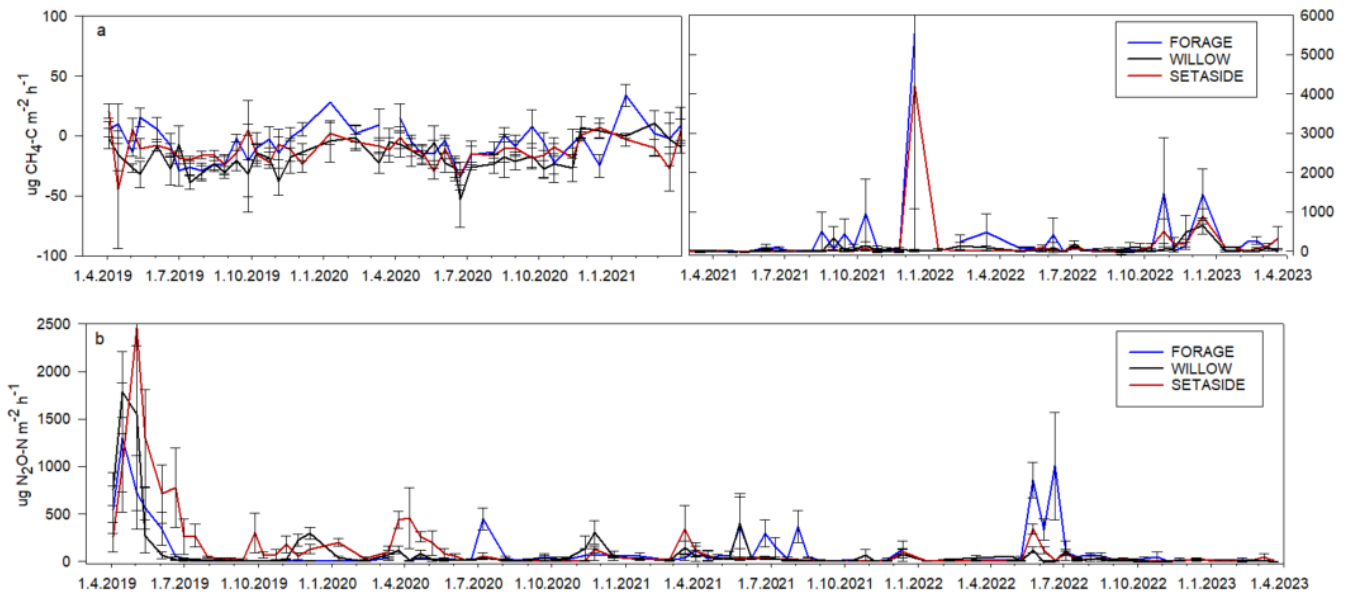
393

394

395

### 3.4. N<sub>2</sub>O fluxes

Hourly fluxes of N<sub>2</sub>O varied between -3 and 2500 μg N<sub>2</sub>O-N m<sup>-2</sup> h<sup>-1</sup> during the four years with the highest emissions during the first four months (Fig. 3). Annual fluxes of N<sub>2</sub>O varied from 1.7 to 33 kg N<sub>2</sub>O-N ha<sup>-1</sup> yr<sup>-1</sup> (Table 3). The emissions declined in time (Fig. S3) especially in the case of set-aside and willow whereas those of forage did not show such a trend (Table 3). WTD explained variation in N<sub>2</sub>O fluxes well (p=0.015) (Fig. 2; Table S4). There were some interactions of year and crop, but the crop type did not affect N<sub>2</sub>O emissions systematically between years. Annual N<sub>2</sub>O fluxes of the forage and willow treatments differed in the whole timeseries (p=0.026).



396

397

398

**Fig. 3.** Fluxes of CH<sub>4</sub> (a) and N<sub>2</sub>O (b) in 2019-2023. The error bars denote standard error. Note the different scale in the y-axis of (a) for the latter half of the period.

399

400

401

### 3.5. Global warming potential

402

403

404

405

406

407

408

## 4. Discussion

409

410

411

412

413

414

415

416

417

418

419

420

421

422

423

424

425

NEE values of 5–8 Mg C ha<sup>-1</sup> per year in the forage plots were of the same magnitude as values reported for grass cultivation in northern Europe (Maljanen et al., 2010). They were, however, 6–10 times higher than NEE reported for year 2002 in a nearby field (Lohila et al., 2004), highlighting the spatial and temporal variation in soil emissions. During the first two years of the experiment, the CH<sub>4</sub> fluxes of the forage plots were negative indicating net consumption of CH<sub>4</sub> by the soil microorganism. The CH<sub>4</sub> oxidation rates were generally higher than average values reported from Nordic cultivated peat soils which have shown net positive values for grass fields (Maljanen et al., 2007, 2010). There was a change from negative fluxes of CH<sub>4</sub> to relatively high emissions after the annual mean water table rose above -40 cm during the two latter years of the experiment. However, compared to rewetted agricultural sites in the temperate zone, the values of ca. 8 kg CH<sub>4</sub>-C per hectare were clearly lower than the average of 180 kg CH<sub>4</sub>-C ha<sup>-1</sup> yr<sup>-1</sup> found in temperate paludiculture-like grassland ecosystems (Bianchi et al., 2021). The N<sub>2</sub>O emissions ranging from 6 to 12 kg N per hectare annually were typical for northern European grass fields on organic soils as they were within the 95% confidence interval of the reported values from temperate and boreal regions (Hiraishi et al., 2014). After the high emission peak in the beginning of the experiment there were only short-term peaks after fertilisation. One of them was especially high and long-lasting and likely induced by heavy rainfall after a long dry period coinciding with fertilisation in May-June 2022. It is typical that high peaks after fertilisation occur when fertilisation is followed by rainfall (Dobbie et al., 1999), and fertiliser-induced peaks may be totally absent if there is no coinciding rainfall (Beetz et al., 2013). The set-aside plots with slowly evolving vegetation had clearly lower GP than the forage plots during the three first years. However, also the ER was lower in the set-aside, and the resulting NEE was of the

426 same magnitude in both treatments. Because there was no biomass export from the set-aside the NECB was lower than in the  
427 forage treatment in most years. The modelled NEE values were about double compared to long-term abandoned croplands in  
428 the Nordic countries (Maljanen et al., 2010) but in our study the plots did not represent similar ecosystems as they were  
429 “abandoned” only for a short period. N<sub>2</sub>O fluxes of the set-aside plots were extremely high in 2019 compared to results of  
430 previous studies on cultivated peat soils in Nordic countries (Maljanen et al., 2010). As the “set-aside” was fertilised and  
431 unsuccessfully planted with bog bilberry, the high emissions were likely due to abundant free mineral nitrogen in the absence  
432 of plant nutrient uptake. As the berry plants did not thrive, the soil was bare for a long period and the N<sub>2</sub>O emissions remained  
433 higher than in the other treatments throughout the summer. Such conditions were prevailing also in a similar bare fallow  
434 treatment at a nearby site in 2000-2002, yielding average N<sub>2</sub>O emissions of 25 kg N ha<sup>-1</sup> yr<sup>-1</sup> (Regina et al., 2004). During the  
435 second year, the emissions lowered but as the plots were fertilised also in 2020, they still exhibited as high emissions as the  
436 forage plots. During the last two years the N<sub>2</sub>O emissions were at a notably low level which likely resulted from ceasing of  
437 fertilisation and a slightly higher WTD leading to less peat being exposed to aerobic conditions. Raising the WTD has been  
438 found to diminish N<sub>2</sub>O emissions in several studies (van Beek et al., 2010; Leppelt et al., 2014).

439 Willow grew well at this site and the mean annual yields were in the higher end of the range 4-16 Mg ha<sup>-1</sup> estimated for  
440 northern climate conditions (Viherä-Aarnio et al., 2022). Carbon lost in soil respiration was lower than the amount sequestered  
441 in the willow biomass in all years except in 2019, leading to highly negative NEE during the whole rotation. However, the  
442 amount of carbon exported in harvest exceeded the NEE and the yielding NECB indicated net loss of carbon to the  
443 atmosphere. Although the average annual NECB calculated from the four-year carbon balance (1.9 Mg C ha<sup>-1</sup> yr<sup>-1</sup>) was low  
444 compared to the forage or set-aside treatments, it still indicated a climate-warming end result in short-rotation cropping of  
445 willow on peat soil. It is possible to achieve a net positive NEBC in willow cultivation on mineral soils (Harris et al., 2017;  
446 Morrison et al., 2019) but in peatlands the high rate of soil respiration inevitably reduces this potential (Kasimir et al., 2018).

447 The target WTD was not reached for most of the time likely because there was unexcepted lateral water outflow from the  
448 site. Our strategy of raising the WTD at a limited area within a field parcel was thus not successful. The larger the area where  
449 the water outflow is restricted, the better the result likely is, and catchment level water management planning is often  
450 recommended for the best results (Mitsch and Wilson, 1996; Pasquet et al., 2015).

451 The annual emissions can be compared to a well-drained cereal site (oats; mean WTD 68 cm) on the same field (Honkanen  
452 et al., 2023), as the distance between these two experiments was just about 20 m. In 2020, when similar measurements were  
453 conducted in both experiments, the total GHG balance (GWP100 with harvest) was 29 Mg in the set-aside and 43 Mg in the  
454 forage treatment while it was 39 Mg CO<sub>2</sub>-eq. in the conventionally managed cereal plots. As the comparable number for  
455 willow was 10 Mg in the willow treatment, it can be argued that the set-aside and willow cultivation with a moderate raise of  
456 WTD were better management options than cultivation of annual crops with a typical drainage depth. It was also clear that  
457 willow had the best GHG balance of these three management options, which is in agreement with findings of grassland and  
458 willow cropping in southern England (Harris et al., 2017). However, the total emissions were still relatively high suggesting  
459 that this kind of moderately wet management is not an efficient climate mitigation measure. This was also shown by the  
460 modelling results of (Kasimir et al., 2018) concluding that fully rewetted peatland had the most favourable carbon balance  
461 and less emissions from soil in a comparison of four different peatland management scenarios. However, management  
462 decisions, like cutting height also play a role in determining the final carbon balance in short-rotation cropping (Berhongaray  
463 et al., 2017).

464 Set-aside is a relevant management option to study because many cultivated peat fields end up as uncultivated plots when  
465 their drainage system degrades, and the landowner finds them too wet for cultivation. The annual total emissions were lower  
466 in the set-aside plots compared to the forage in 2020 and 2021, and also in 2022 if the carbon exported in harvest is taken into  
467 account. However, they were not especially low as compared to cultivated peat soils in general. Thus, leaving cultivated peat  
468 soils uncultivated without active rewetting is not desirable form of land management as these sites drift out from food

469 production, but the GHG emissions can remain high. A recent Swedish study also found that setting aside did not reduce  
470 GHG emissions from a drained peat soil (Keck et al., 2024)

471 Our set-aside plots were actually intended to be vegetated by bog bilberry, a native mire plant that could become a novel  
472 antioxidant-rich ingredient for food (Lätti et al., 2010) or pharmaceutical (Esposito et al., 2019) industries. However, we soon  
473 noticed that the seedlings did not grow roots indicating that formerly agriculturally cultivated peat was not a suitable substrate  
474 for this plant. As the nutrient content of the topsoil did not show large deviations from the reported ranges supporting the  
475 growth of bog bilberry (Jacquemart, 1996), it is likely that the pH of 5.4 at our site was too high. Bog bilberry is usually found  
476 in soils with pH below 5. However, in recent trials it has been successfully grown on Chinese farmlands with pH 5–6 but low  
477 pH improved the growth also there (Duan et al., 2022).

478 There are usually high uncertainties in the GHG measurements, and this is especially true regarding the combination of  
479 methods chosen for the willow treatment. The carbon balance of willow was determined using a combination of the pool-  
480 based and flux-based methods, which can differ by several magnitudes (Berhongaray et al., 2017). The most reliable method  
481 for measuring the carbon balance of willow stands is likely the eddy covariance method, which is not feasible in experiments  
482 with small plots. Part of the uncertainty also arises from the simplicity of the models. For example, soil respiration was  
483 modelled only based on soil temperature and WTD, although it can be affected also e.g. by changes in microbial community  
484 composition or activity (Yang et al. 2022) and soil moisture which does not always well follow changes in WTD (Smith et  
485 al. 2018). Estimating vegetation cover using measured LAI is also problematic, as it reflects weakly the amount of active  
486 chlorophyll (Delegido et al., 2015; Gregersen et al., 2013). It is especially difficult to assess active vegetation at the beginning  
487 and end of the growing season. However, the influence on the annual balance is minor due to low temperature and radiation  
488 at that time. With the Canopeo application, the models were significantly better as it was possible to determine the green leaf  
489 area better than with the previously used LAI measurement with the SunScan instrument. The measurement results of PAR  
490 values feature uncertainties due to abrupt changes in cloudiness or fogging and dirt on the plexiglass. Due to technical  
491 problems, FMI data and another PAR sensor was used to fill the gaps in the PAR measurements especially in 2021. The  
492 plexiglass surfaces were kept as clean as possible, fogging was kept low by using a short measurement time, and clear sky  
493 conditions were preferred which should reduce the uncertainty occurred in measurements. Model predicted soil temperature  
494 in gap filling may cause some error, but the filled gaps were not long, and the error was mostly diurnal with low significance  
495 for the annual balances. Regarding biweekly N<sub>2</sub>O and CH<sub>4</sub> measurements, there is a high risk of missing short-term peaks,  
496 for example due to freeze-thaw cycles (Lammirato et al., 2021). Also, if the measurements hit peaks, the emissions may be  
497 overestimated due to interpolation of the gaps in the data particularly during times with infrequent measurements.

498

## 499 **5 Conclusions**

500 This study gave valuable insights to the practical implementation and climate mitigation potential of three management  
501 options relevant for cultivated peatlands with raised WTD: forage, willow and set-aside. The results indicate that wet  
502 management of cultivated peat soils considerably reduces the soil respiration and N<sub>2</sub>O emissions. Significant counteracting  
503 effect of increased CH<sub>4</sub> emissions are avoided as long as the WTD does not rise close to the soil surface. However, compared  
504 to full rewetting, partial rewetting remains a compromise solution to climate warming as it is likely that the peat layer will  
505 eventually be lost. It is important to develop incentives to inundate large, connected peatland areas to ensure water availability  
506 and maintenance of high enough water table for efficient control of peat decomposition.

507

508 **Authorship contributions.** KL and HK designed the experiment. JH, HH, SS and TL developed the methodology. HK and  
509 TL planned, supervised and partly conducted the field work. KL, HH and HK analysed and visualised the data. KL and HH  
510 wrote the original manuscript. All authors were involved in revising the text.

511

512 **Data availability.** The data will be available in Zenodo two years after publication.

513

514 **Competing interests.** The contact author has declared that none of the authors has any competing interests.

515

516 **Acknowledgements.** The authors are grateful to the technical staff of Natural Resources Institute Finland for skilled work in  
517 the field and laboratory.

518

519 **Financial support.** This study was part of the project SOMPA (Novel soil management practices – key for sustainable  
520 bioeconomy and climate change mitigation), funded by the Strategic Research Council at the Research Council of Finland,  
521 grant No 312912 and INSURE (Indicators for successful carbon sequestration and greenhouse gas mitigation by rewetting  
522 cultivated peat soils) subproject of the EJP Soil project that has received funding from the European Union’s Horizon 2020  
523 research and innovation programme under grant agreement No 862695.

524

525

526

527

## 528 **References**

529 Aparicio, N., Villegas, D., Casadesus, J., Araus, J. L., and Royo, C.: Spectral Vegetation Indices as Nondestructive Tools  
530 for Determining Durum Wheat Yield, *Agron. J.*, 92, 83–91, <https://doi.org/10.2134/agronj2000.92183x>, 2000.

531 van Beek, C. L., Pleijter, M., Jacobs, C. M. J., Velthof, G. L., van Groenigen, J. W., and Kuikman, P. J.: Emissions of N<sub>2</sub>O  
532 from fertilized and grazed grassland on organic soil in relation to groundwater level, *Nutr. Cycl. Agroecosystems*, 86, 331–  
533 340, <https://doi.org/10.1007/s10705-009-9295-2>, 2010.

534 Beetz, S., Liebersbach, H., Glatzel, S., Jurasinski, G., Buczko, U., and Hoepfer, H.: Effects of land use intensity on the full  
535 greenhouse gas balance in an Atlantic peat bog, *Biogeosciences*, 10, 1067–1082, <https://doi.org/10.5194/bg-10-1067-2013>,  
536 2013.

537 Berhongaray, G., Verlinden, M. S., Broeckx, L. S., Janssens, I. A., and Ceulemans, R.: Soil carbon and belowground carbon  
538 balance of a short-rotation coppice: assessments from three different approaches, *GCB Bioenergy*, 9, 299–313,  
539 <https://doi.org/10.1111/gcbb.12369>, 2017.

540 Bianchi, A., Larmola, T., Kekkonen, H., Saarnio, S., and Lang, K.: Review of Greenhouse Gas Emissions from Rewetted  
541 Agricultural Soils, *WETLANDS*, 41, 108, <https://doi.org/10.1007/s13157-021-01507-5>, 2021.

542 Bockermann, C., Eickenscheidt, T., and Droesler, M.: Adaptation of fen peatlands to climate change: rewetting and  
543 management shift can reduce greenhouse gas emissions and offset climate warming effects, *BIOGEOCHEMISTRY*,  
544 <https://doi.org/10.1007/s10533-023-01113-z>, 2024.

545 Boonman, C. C. F., Heuts, T. S., Vroom, R. J. E., Geurts, J. J. M., and Fritz, C.: Wetland plant development overrides  
546 nitrogen effects on initial methane emissions after peat rewetting, *Aquat. Bot.*, 184, 103598,  
547 <https://doi.org/10.1016/j.aquabot.2022.103598>, 2023.

548 Delegido, J., Verrelst, J., Rivera, J. P., Ruiz-Verdú, A., and Moreno, J.: Brown and *green* LAI mapping through spectral  
549 indices, *Int. J. Appl. Earth Obs. Geoinformation*, 35, 350–358, <https://doi.org/10.1016/j.jag.2014.10.001>, 2015.

550 Dobbie, K. E., McTaggart, I. P., and Smith, K. A.: Nitrous oxide emissions from intensive agricultural systems: Variations  
551 between crops and seasons, key driving variables, and mean emission factors, *J. Geophys. Res. Atmospheres*, 104, 26891–  
552 26899, <https://doi.org/10.1029/1999JD900378>, 1999.

553 Duan, Y., Guo, B., Zhang, L., Li, J., Li, S., Zhao, W., Yang, G., Zhou, S., Zhou, C., Song, P., Li, P., Fang, L., Hou, S., Shi,  
554 D., Zhao, H., and Guo, P.: Interactive climate-soil forces shape the spatial distribution of foliar N:P stoichiometry in  
555 *Vaccinium uliginosum* planted in agroforests of Northeast China, *Front. Ecol. Evol.*, 10,  
556 <https://doi.org/10.3389/fevo.2022.1065680>, 2022.



- 557 Eposito, D., Overall, J., and Grace, M. H.: Alaskan Berry Extracts Promote Dermal Wound Repair Through Modulation of  
558 Bioenergetics and Integrin Signaling, *Front. Pharmacol.*, 10, <https://doi.org/10.3389/fphar.2019.01058>, 2019.
- 559 Evans, C. D., Peacock, M., Baird, A. J., Artz, R. R. E., Burden, A., Callaghan, N., Chapman, P. J., Cooper, H. M., Coyle,  
560 M., Craig, E., Cumming, A., Dixon, S., Gauci, V., Grayson, R. P., Helfter, C., Heppell, C. M., Holden, J., Jones, D. L.,  
561 Kaduk, J., Levy, P., Matthews, R., McNamara, N. P., Misselbrook, T., Oakley, S., Page, S. E., Rayment, M., Ridley, L. M.,  
562 Stanley, K. M., Williamson, J. L., Worrall, F., and Morrison, R.: Overriding water table control on managed peatland  
563 greenhouse gas emissions, *Nature*, <https://doi.org/10.1038/s41586-021-03523-1>, 2021.
- 564 FMI 2024. Finnish Meteorological Institute open data on weather monitoring. Available at:  
565 <https://opendata.fmi.fi/wfs?request=GetCapabilities>
- 566 Freeman, B. W. J., Evans, C. D., Musarika, S., Morrison, R., Newman, T. R., Page, S. E., Wiggs, G. F. S., Bell, N. G. A.,  
567 Styles, D., Wen, Y., Chadwick, D. R., and Jones, D. L.: Responsible agriculture must adapt to the wetland character of mid-  
568 latitude peatlands, *Glob. Change Biol.*, 28, 3795–3811, <https://doi.org/10.1111/gcb.16152>, 2022.
- 569 Gregersen, P. L., Culetic, A., Boschian, L., and Krupinska, K.: Plant senescence and crop productivity, *Plant Mol. Biol.*, 82,  
570 603–622, <https://doi.org/10.1007/s11103-013-0013-8>, 2013. Guenther, A., Barthelmes, A., Huth, V., Joosten, H., Jurasinski,  
571 G., Koebisch, F., and Couwenberg, J.: Prompt rewetting of drained peatlands reduces climate warming despite methane  
572 emissions, *Nat. Commun.*, 11, 1644–1644, <https://doi.org/10.1038/s41467-020-15499-z>, 2020.
- 573 Harris, Z. M., Alberti, G., Viger, M., Jenkins, J. R., Rowe, R., McNamara, N. P., and Taylor, G.: Land-use change to  
574 bioenergy: grassland to short rotation coppice willow has an improved carbon balance, *GCB Bioenergy*, 9, 469–484,  
575 <https://doi.org/10.1111/gcbb.12347>, 2017.
- 576 Hiraishi, T., Krug, T., Tanabe, K., Srivastava, N., Baasansuren, J., Fukuda, M., and Troxler, T. G.: Supplement to the 2006  
577 IPCC guidelines for national greenhouse gas inventories: Wetlands, IPCC, Switzerland, 2014.
- 578 Honkanen, H., Kekkonen, H., Heikkinen, J., Kaseva, J., and Lång, K.: Minor effects of no-till treatment on GHG emissions  
579 of boreal cultivated peat soil, *Biogeochemistry*, <https://doi.org/10.1007/s10533-023-01097-w>, 2023.
- 580 Huang, Y., Ciais, P., Luo, Y., Zhu, D., Wang, Y., Qiu, C., Goll, D. S., Guenet, B., Makowski, D., De Graaf, I., Leifeld, J.,  
581 Kwon, M. J., Hu, J., and Qu, L.: Tradeoff of CO<sub>2</sub> and CH<sub>4</sub> emissions from global peatlands under water-table drawdown,  
582 *Nat. Clim. Change*, 11, 618–622, <https://doi.org/10.1038/s41558-021-01059-w>, 2021.
- 583 Humpenöder, F., Karstens, K., Lotze-Campen, H., Leifeld, J., Menichetti, L., Barthelmes, A., and Popp, A.: Peatland  
584 protection and restoration are key for climate change mitigation, *Environ. Res. Lett.*, 15, 104093,  
585 <https://doi.org/10.1088/1748-9326/abae2a>, 2020.
- 586 Jacquemart, A.-L.: *Vaccinium Uliginosum* L., *J. Ecol.*, 84, 771–785, <https://doi.org/10.2307/2261339>, 1996.
- 587 Jokinen, P., Pirinen, P., Kaukoranta, J.-P., Kangas, A., Alenius, P., Eriksson, P., Johansson, M., and Wilkman, S.: Tilastoja  
588 Suomen ilmastosta ja merestä 1991-2020, Ilmatieteen laitos, 2021.
- 589 Kandel, T. P., Elsgaard, L., and Laerke, P. E.: Measurement and modelling of CO<sub>2</sub> flux from a drained fen peatland  
590 cultivated with reed canary grass and spring barley, *Glob. Change Biol. Bioenergy*, 5, 548–561,  
591 <https://doi.org/10.1111/gcbb.12020>, 2013.
- 592 Kandel, T. P., Karki, S., Elsgaard, L., Labouriau, R., and Laerke, P. E.: Methane fluxes from a rewetted agricultural fen  
593 during two initial years of paludiculture, *Sci. Total Environ.*, 713, 136670–136670,  
594 <https://doi.org/10.1016/j.scitotenv.2020.136670>, 2020.
- 595 Karki, S., Elsgaard, L., Audet, J., and Laerke, P. E.: Mitigation of greenhouse gas emissions from reed canary grass in  
596 paludiculture: effect of groundwater level, *Plant Soil*, 383, 217–230, <https://doi.org/10.1007/s11104-014-2164-z>, 2014.
- 597 Kasimir, Å., He, H., Coria, J., and Nordén, A.: Land use of drained peatlands: Greenhouse gas fluxes, plant production, and  
598 economics, *Glob. Change Biol.*, 24, 3302–3316, <https://doi.org/10.1111/gcb.13931>, 2018.
- 599 Keck, H., Meurer, K. H. E., Jordan, S., Kätterer, T., Hadden, D., and Grelle, A.: Setting-aside cropland did not reduce  
600 greenhouse gas emissions from a drained peat soil in Sweden, *Front. Environ. Sci.*, 12,  
601 <https://doi.org/10.3389/fenvs.2024.1386134>, 2024.

- 602 Kutzbach, L., Schneider, J., Sachs, T., Giebels, M., Nykanen, H., Shurpali, N. J., Martikainen, P. J., Alm, J., and Wilmking,  
603 M.: CO<sub>2</sub> flux determination by closed-chamber methods can be seriously biased by inappropriate application of linear  
604 regression, *Biogeosciences*, 4, 1005–1025, <https://doi.org/10.5194/bg-4-1005-2007>, 2007.
- 605 Lammirato, C., Wallman, M., Weslien, P., Klemetsson, L., and Rütting, T.: Measuring frequency and accuracy of annual  
606 nitrous oxide emission estimates, *Agric. For. Meteorol.*, 310, 108624, <https://doi.org/10.1016/j.agrformet.2021.108624>,  
607 2021.
- 608 Lätti, A. K., Jaakola, L., Riihinen, K. R., and Kainulainen, P. S.: Anthocyanin and Flavonol Variation in Bog Bilberries  
609 (*Vaccinium uliginosum* L.) in Finland, *J. Agric. Food Chem.*, 58, 427–433, <https://doi.org/10.1021/jf903033m>, 2010.
- 610 Lehtonen, H., Huan-Niemi, E., and Niemi, J.: The transition of agriculture to low carbon pathways with regional  
611 distributive impacts, *Environ. Innov. Soc. Transit.*, 44, 1–13, <https://doi.org/10.1016/j.eist.2022.05.002>, 2022.
- 612 Leifeld, J. and Menichetti, L.: The underappreciated potential of peatlands in global climate change mitigation strategies,  
613 *Nat. Commun.*, 9, 1071–1071, <https://doi.org/10.1038/s41467-018-03406-6>, 2018.
- 614 Leppelt, T., Dechow, R., Gebbert, S., Freibauer, A., Lohila, A., Augustin, J., Droesler, M., Fiedler, S., Glatzel, S., Hoepfer,  
615 H., Jaerveoja, J., Laerke, P. E., Maljanen, M., Mander, U., Maekiranta, P., Minkkinen, K., Ojanen, P., Regina, K., and  
616 Stromgren, M.: Nitrous oxide emission budgets and land-use-driven hotspots for organic soils in Europe, *Biogeosciences*,  
617 11, 6595–6612, <https://doi.org/10.5194/bg-11-6595-2014>, 2014.
- 618 Liu, W., Fritz, C., Van Belle, J., and Nonhebel, S.: Production in peatlands: Comparing ecosystem services of different land  
619 use options following conventional farming, *Sci. Total Environ.*, 875, 162534,  
620 <https://doi.org/10.1016/j.scitotenv.2023.162534>, 2023.
- 621 LLOYD, J. and TAYLOR, J.: On the Temperature-Dependence of Soil Respiration, *Funct. Ecol.*, 8, 315–323,  
622 <https://doi.org/10.2307/2389824>, 1994.
- 623 Lohila, A., Aurela, M., Regina, K., and Laurila, T.: Soil and total ecosystem respiration in agricultural fields: effect of soil  
624 and crop type, *Plant Soil*, 251, 303–317, <https://doi.org/10.1023/A:1023004205844>, 2003.
- 625 Lohila, A., Aurela, M., Tuovinen, J., and Laurila, T.: Annual CO<sub>2</sub> exchange of a peat field growing spring barley or  
626 perennial forage grass, *J. Geophys. Res.-Atmospheres*, 109, D18116–D18116, <https://doi.org/10.1029/2004JD004715>,  
627 2004.
- 628 Long, S. P. and Hällgren, J.-E.: Measurement of CO<sub>2</sub> assimilation by plants in the field and the laboratory, in:  
629 *Photosynthesis and Production in a Changing Environment: A field and laboratory manual*, edited by: Hall, D. O., Scurlock,  
630 J. M. O., Bolhàr-Nordenkamp, H. R., Leegood, R. C., and Long, S. P., Springer Netherlands, Dordrecht, 129–167,  
631 [https://doi.org/10.1007/978-94-011-1566-7\\_9](https://doi.org/10.1007/978-94-011-1566-7_9), 1993.
- 632 Maljanen, M., Liikanen, A., Silvola, J., and Martikainen, P.: Measuring N<sub>2</sub>O emissions from organic soils by closed  
633 chamber or soil/snow N<sub>2</sub>O gradient methods, *Eur. J. Soil Sci.*, 54, 625–631, <https://doi.org/10.1046/j.1365-2389.2003.00531.x>, 2003.
- 635 Maljanen, M., Hytonen, J., Makiranta, P., Alm, J., Minkkinen, K., Laine, J., and Martikainen, P. J.: Greenhouse gas  
636 emissions from cultivated and abandoned organic croplands in Finland, *Boreal Environ. Res.*, 12, 133–140, 2007.
- 637 Maljanen, M., Sigurdsson, B. D., Guomundsson, J., Oskarsson, H., Huttunen, J. T., and Martikainen, P. J.: Greenhouse gas  
638 balances of managed peatlands in the Nordic countries - present knowledge and gaps, *Biogeosciences*, 7, 2711–2738,  
639 <https://doi.org/10.5194/bg-7-2711-2010>, 2010.
- 640 Mander, U., Espenberg, M., Melling, L., and Kull, A.: Peatland restoration pathways to mitigate greenhouse gas emissions  
641 and retain peat carbon, *BIOGEOCHEMISTRY*, <https://doi.org/10.1007/s10533-023-01103-1>, 2023.
- 642 Meek, D. W., Hatfield, J. L., Howell, T. A., Idso, S. B., and Reginato, R. J.: A Generalized Relationship between  
643 Photosynthetically Active Radiation and Solar Radiation<sup>1</sup>, *Agron. J.*, 76, 939–945,  
644 <https://doi.org/10.2134/agronj1984.00021962007600060018x>, 1984.
- 645 Mitsch, W. J. and Wilson, R. F.: Improving the Success of Wetland Creation and Restoration with Know-How, Time, and  
646 Self-Design, *Ecol. Appl.*, 6, 77–83, <https://doi.org/10.2307/2269554>, 1996.
- 647 Morrison, R., Rowe, R. L., Cooper, H. M., and McNamara, N. P.: Multi-year carbon budget of a mature commercial short  
648 rotation coppice willow plantation, *GCB Bioenergy*, 11, 895–909, <https://doi.org/10.1111/gcbb.12608>, 2019.

- 649 Nielsen, C. K., Elsgaard, L., Jørgensen, U., and Lærke, P. E.: Soil greenhouse gas emissions from drained and rewetted  
650 agricultural bare peat mesocosms are linked to geochemistry, *Sci. Total Environ.*, 896, 165083,  
651 <https://doi.org/10.1016/j.scitotenv.2023.165083>, 2023.
- 652 Pacaldo, R. S., Volk, T. A., and Briggs, R. D.: Carbon Sequestration in Fine Roots and Foliage Biomass Offsets Soil CO<sub>2</sub>  
653 Effluxes along a 19-year Chronosequence of Shrub Willow (*Salix x dasyclados*) Biomass Crops, *BioEnergy Res.*, 7, 769–  
654 776, <https://doi.org/10.1007/s12155-014-9416-x>, 2014.
- 655 Pasquet, S., Pellerin, S., and Poulin, M.: Three decades of vegetation changes in peatlands isolated in an agricultural  
656 landscape, *Appl. Veg. Sci.*, 18, 220–229, <https://doi.org/10.1111/avsc.12142>, 2015.
- 657 Patrignani, A. and Ochsner, T. E.: Canopeo: A Powerful New Tool for Measuring Fractional Green Canopy Cover, *Agron.*  
658 *J.*, 107, 2312–2320, <https://doi.org/10.2134/agronj15.0150>, 2015.
- 659 Raich, J. W., Rastetter, E. B., Melillo, J. M., Kicklighter, D. W., Steudler, P. A., Peterson, B. J., Grace, A. L., Moore III, B.,  
660 and Vorosmarty, C. J.: Potential Net Primary Productivity in South America: Application of a Global Model, *Ecol. Appl.*, 1,  
661 399–429, <https://doi.org/10.2307/1941899>, 1991.
- 662 Regina, K., Syvasalo, E., Hannukkala, A., and Esala, M.: Fluxes of N<sub>2</sub>O from farmed peat soils in Finland, *Eur. J. Soil Sci.*,  
663 55, 591–599, <https://doi.org/10.1111/j.1365-2389.2004.00622.x>, 2004.
- 664 Shepherd, M.J., Lindsey, L.E. and Lindsey, A.J. Soybean Canopy Cover Measured with Canopeo Compared with Light  
665 Interception. *Agricultural & Environmental Letters* 3: 180031. doi:10.2134/aerl2018.06.0031, 2018.
- 666 Strack, M., Davidson, S. J., Hirano, T., and Dunn, C.: The Potential of Peatlands as Nature-Based Climate Solutions, *Curr.*  
667 *Clim. Change Rep.*, 8, 71–82, <https://doi.org/10.1007/s40641-022-00183-9>, 2022.
- 668 Tanneberger, F., Birr, F., Couwenberg, J., Kaiser, M., Luthardt, V., Nерger, M., Pfister, S., Oppermann, R., Zeitz, J., Beyer,  
669 C., van der Linden, S., Wichtmann, W., and Närmann, F.: Saving soil carbon, greenhouse gas emissions, biodiversity and  
670 the economy: paludiculture as sustainable land use option in German fen peatlands, *Reg. Environ. Change*, 22, 69,  
671 <https://doi.org/10.1007/s10113-022-01900-8>, 2022.
- 672 Viherä-Aarnio, A., Jyske, T., and Beuker, E.: Pajut biokiertoaloudessa - Materiaaleja, arvoaineita, ympäristöhyötyjä :  
673 Synteesiraportti, Luonnonvarakeskus, 2022.
- 674 Vuorinen, J. and Mäkitie, O.: The method of soil testing in use in Finland, *Maatalouskoelaitoksen maatumkimusosasto*,  
675 Helsinki, 44 pp., 1955.
- 676 Wilson, D., Blain, D., Couwenberg, J., Evans, C., Murdiyarso, D., Page, S., Renou-Wilson, F., Rieley, J., Strack, M., and  
677 Tuittila, E.: Greenhouse gas emission factors associated with rewetting of organic soils, *Mires Peat*, 14, 1–28, 2016.
- 678 Zheng, D., Hunt, E., and Running, S.: A daily soil temperature model based on air temperature and precipitation for  
679 continental applications, *Clim. Res.*, 2, 183–191, <https://doi.org/10.3354/cr002183>, 1993.
- 680