

**Journal:** ESSD

**Title:** Improved estimate of global gross primary production for reproducing its long-term variation, 1982–2017

**MS No.:** essd-2019-126

**MS Type:** Data description paper

Dear Prof. Yuyu Zhou,

We appreciate the valuable opportunity to further revise our manuscript “Improved estimate of global gross primary production for reproducing its long-term variation, 1982–2017” (**MS No.:** essd-2019-126) for possible publication in ESSD.

We have carefully studied these comments and substantially revised our manuscript accordingly. Please find attached the point-by-point responses to the comments of the reviewer. Please note that the comments from the reviewers are in **bold** followed by our responses in regular text. The changes in our manuscript are underlined with red.

We believe the quality of the manuscript can now meet the high standard of ESSD and deeply appreciate your consideration of our manuscript.

Sincerely,

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## Response to Reviewer:

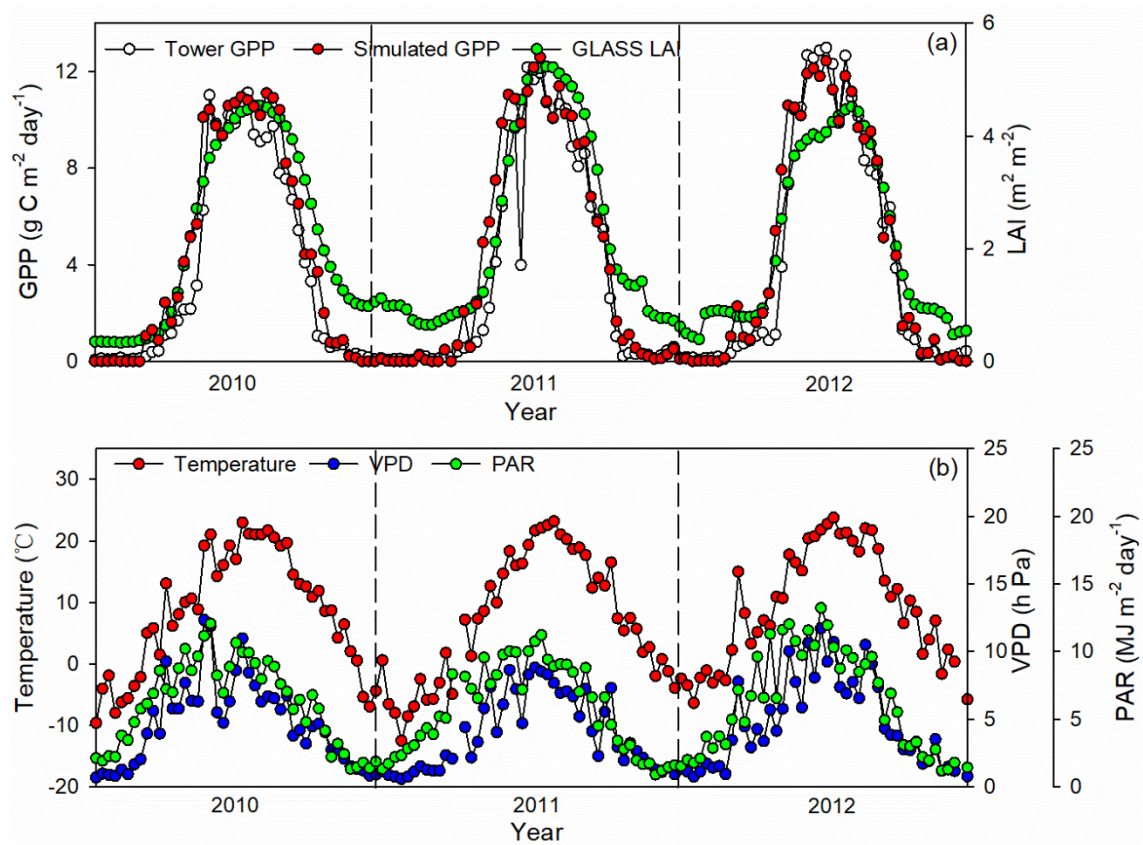
The paper entitled by ‘Improved estimate of global gross primary production for reproducing its longterm variation, 1982-2017’ proposed a new light use efficiency model (rEC-LUE) by integrating part of the Faquhar’s model (a process model for leaf-scale photosynthesis of C3 plant) into EC-LUE model, which showed an obvious improvement of the simulation performance. Based on the simulations of this new model, the authors produced a new global GPP data very correlated to FLUXCOM GPP on annual scale. According to some of the Reviewer #2’s comments, the authors also have provided adequate figures and revisions. However, I have the same opinion as the second one of Reviewer #2, which the authors didn’t accept:

In addition, it is necessary to closely look into a few specific sites (through selection of flux sites considering climate and other environmental factors) using time-series data. I guess the proposed model would work well for some sites, and not for others. Additional discussion should follow regarding that. When simply looking at Figure 8, the proposed model appears to somewhat underestimate GPP compared to the existing ones, which implies more validations are required on spatially and temporally detailed domains.

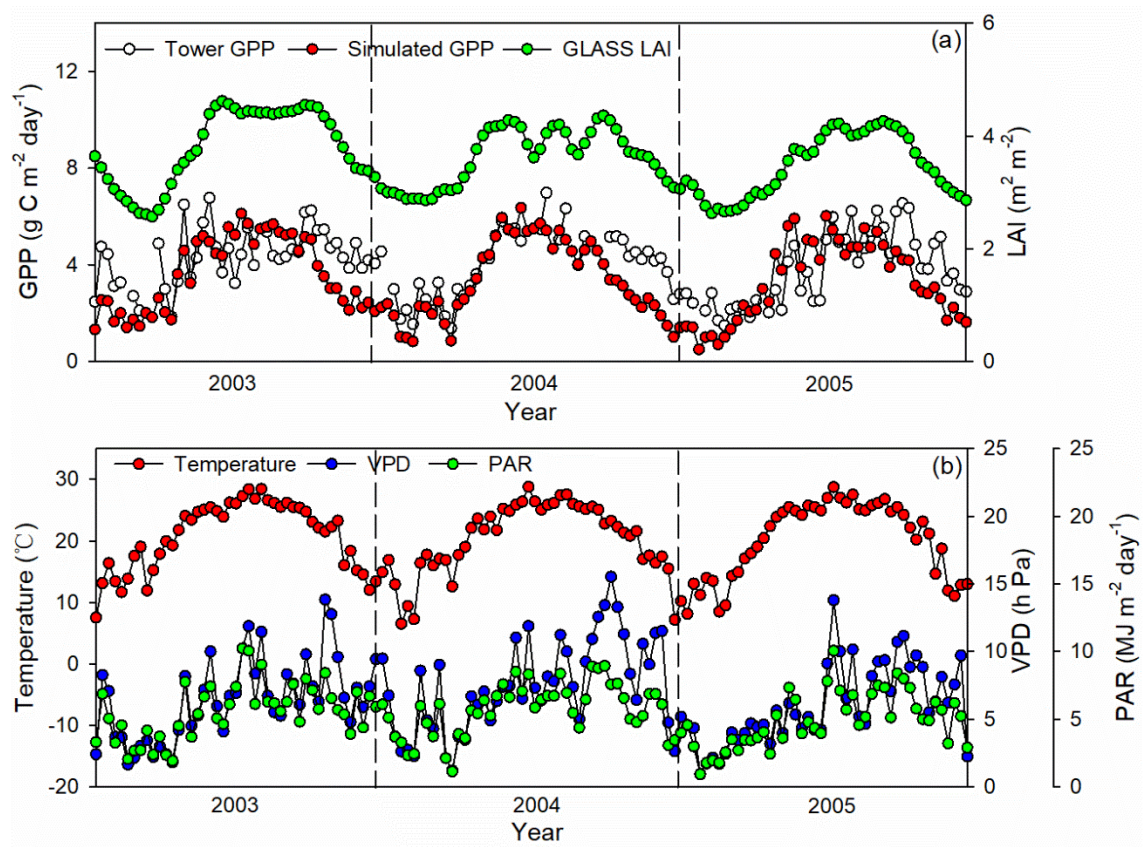
It’s very necessary that the authors could provide the time series result for some specific sites (can be one low- $r^2$  site, one median - $r^2$  site and one high- $r^2$  site) and give some analysis about the bias between models against observations.

Thanks for your comments. In our study, the model performed well at most sites except several tropical forest sites (i.e., MY-PSO, BR-Sa1 and BR-Sa3) without pronounced seasonal variations (Fig. 2). Therefore, we selected three sites with low  $R^2$  (Br-Sa3), median  $R^2$  (CN-Din), and high  $R^2$  (US-UMB) to illustrate the time-series changes of observed/simulated GPP and environmental factors (Figs S1-S3). The new analyses have been added into the revised manuscript:

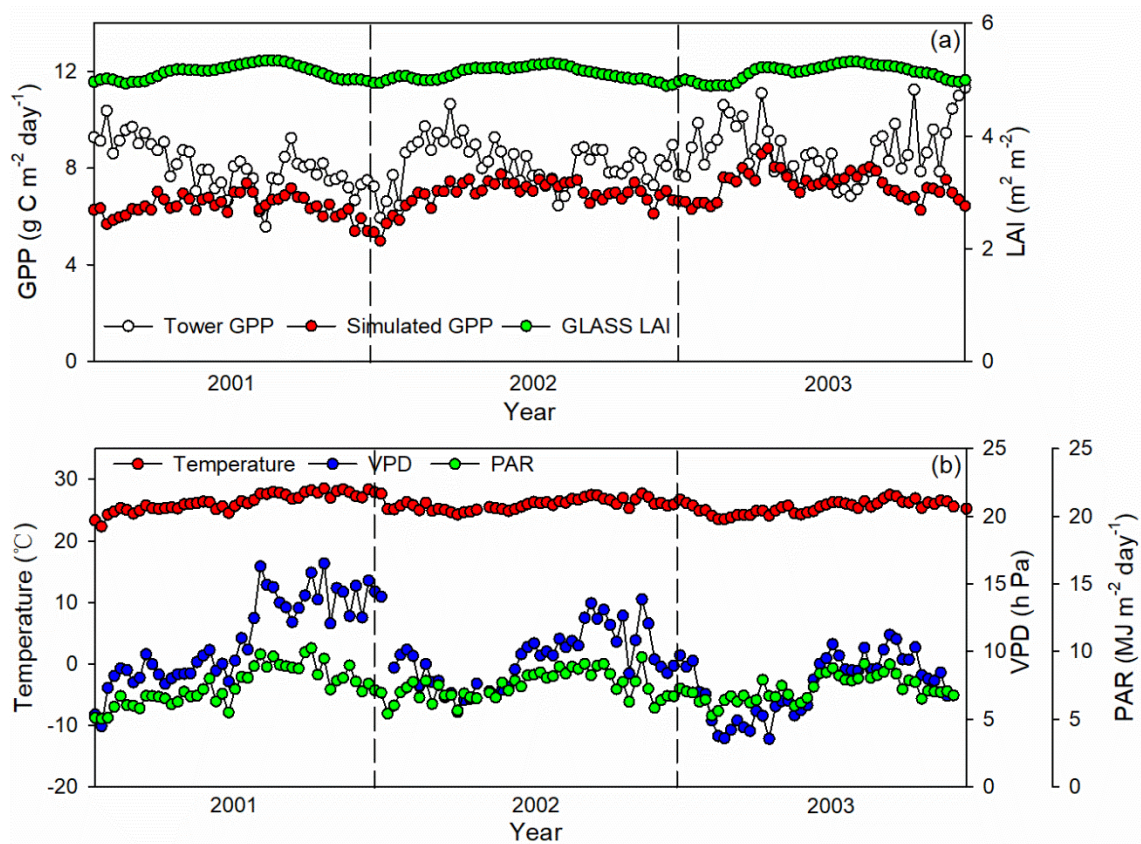
“Furthermore, we selected three sites with high  $R^2$  (US-UMB; DBF;  $R^2 = 0.93$ ), median  $R^2$  (CN-Din; EBF;  $R^2 = 0.71$ ), and low  $R^2$  (Br-Sa3; EBF;  $R^2 = 0.39$ ) to illustrate the time-series changes of observed/simulated GPP, LAI, and environmental factors (i.e., air temperature, VPD, and PAR) (Figs S1-S3). At US-UMB site, the model captured the GPP variations well all the year round with no obvious bias (Fig. S1). At CN-Din site, the model generally performed well except the underestimation at the end of the year (November-December) with decreased LAI. While, at Br-Sa3 site, the model could not capture the variations of GPP for the vegetation greenness and environmental factors varying slightly during the year (Fig. S3).” (Lines 271-277 in the revised manuscript)



**Figure S1: Time-series changes of (a) tower estimated GPP, model simulated GPP, and LAI, and (b) environmental factors (i.e., air temperature, VPD, and PAR) for low R<sup>2</sup> site in Fig. 2, taking US-UMB site as an example.**



**Figure S2: Time-series changes of (a) tower estimated GPP, model simulated GPP, and LAI, and (b) environmental factors (i.e., air temperature, VPD, and PAR) for low  $R^2$  site in Fig. 2, taking CN-Din site as an example.**



**Figure S3: Time-series changes of (a) tower estimated GPP, model simulated GPP, and LAI, and (b) environmental factors (i.e., air temperature, VPD, and PAR) for low R<sup>2</sup> site in Fig. 2, taking Br-Sa3 site as an example.**

Furthermore, I think the authors should also address following questions in their paper:

To build this model:

1. The authors were using part of Farquhar's model (Equation 5) to represent the CO<sub>2</sub> effect on GPP and embedding it into the EC-LUE model directly. However, this part of Farquhar's model only includes the limitation from the electron-transport efficiency (or J<sub>max</sub>) which is a super much simplification. The explanation should be given on why the limitation from Rubisco carboxylation rate (or V<sub>cmax</sub>) and non-linear effect of J<sub>max</sub> which also include the CO<sub>2</sub> effect were ignored.

$$C_s = \frac{C_i - \phi}{C_i + 2\phi} \quad (5)$$

In this study, we considered the atmosphere CO<sub>2</sub> fertilization effect on vegetation production, and used the equations (i.e. Eqs. 5-11) of Farquhar's model to simulate leaf internal CO<sub>2</sub> concentration. These equations are not for simulating photosynthesis rate limited by electron-transport efficiency (i.e. J<sub>max</sub>). As you know, EC-LUE model is based on the principle of light use efficiency model, which is different to Farquhar's model.

**2. In the introduction, the authors only talked about on the environmental factors they considered in their model, but didn't look into the other factors, for example: soil moisture, which is considered much more important than VPD in many other light use efficiency models (e.g. VPM, CASA, Horn's model and even the original EC-LUE model). The state of art on choosing the environmental factors should be investigated.**

Thanks for your deep thoughts. Indeed, soil moisture is a very important factor to impact vegetation production. Although, there are a few satellite-based soil moisture products available currently, such as SMAP, and AMSR-E soil moisture products. However, there are large data gaps across the spatial and temporal scales. In addition, many process-based terrestrial ecosystem models represent the effects of reduced soil moisture on canopy carbon assimilation using an empirical drought factor commonly referred as b factor (Cox et al., 1998). The b-factor approach has been shown to overestimate plant responses to seasonal and experimentally induced drought (Ukkola et al., 2016; Eller et al., 2018). The b-factor has a large impact on the modelled global carbon budget, suppressing 30–40% of the annual gross primary productivity (GPP) in large areas of arid and semiarid ecosystems (Trugman et al., 2018). Therefore, none of light use efficiency models used satellite-based soil moisture to indicate the water stress of vegetation production. VPM model used satellite-based normalized difference water index (NDWI; Xiao et al., 2005) to indicate moisture condition of vegetation canopy. CASA and original EC-LUE model used the ratio of ecosystem evapotranspiration to potential evapotranspiration or net radiation (Potter et al., 1993; Yuan et al., 2007).

In the revised EC-LUE model, we used VPD as the water scalar to simulate the impacts of water stress on vegetation production based on the following two reasons. First, VPD plays an important role in regulating photosynthesis. Previous numerous studies have highlighted that the leaf and canopy photosynthetic rate declines when the atmospheric VPD increases due to stomatal closure (Fletcher et al., 2007; Grossiord et al., 2020). A recent study highlighted that increases in VPD rather than changes in precipitation would be a dominant influence on vegetation productivity (Konings et al., 2017). Second, there is a substantial long-term change of VPD driven by climate warming. Rising air temperature increases the saturated vapor pressure at a rate of  $\sim 7\%/K$  according to the Clausius–Clapeyron relationship, and therefore, VPD will increase if the atmospheric water vapor content does not increase by exactly the same amount as the saturated vapor pressure. Numerous studies indicated significant changes in the VPD in both humid areas and continental areas located far from oceanic humidity (Van Wijngaarden and Vincent, 2004; Simmons et al. 2010; Pierce et al., 2013; Willett et al. 2014). Our recent study also shows the strong regulations of long-term VPD changes to global vegetation growth (Yuan et al., 2019), and thus it is necessary to integrate the VPD impacts into the light use efficiency models. In addition, we did not take more efforts to integrate the impacts of soil moisture on vegetation production, mostly because there are strong correlations of interannual variability between VPD and soil moisture. Several previous studies have reported highly consistent interannual variability of VPD and soil moisture (Zhou et al., 2019a, b).

In our manuscript, there is one paragraph to discuss the water stress in our model:

“The evaporation fraction (EF), namely the ratio of evapotranspiration (ET) to net radiation (Rn), was used to indicate the water stress on vegetation growth in the original EC-LUE model (Yuan et al., 2007; 2010). While the atmospheric VPD was used to indicate water stress to avoid the aggregated errors from ET simulations in the revised EC-LUE model. Physiologically, vegetation production is sensitive to both atmospheric VPD and soil moisture availability to roots. Several studies have reported highly consistent interannual variability of VPD and soil moisture (Zhou et al., 2019a, b). In addition,



recent studies highlighted that the increase in VPD had a larger limitation to the surface conductance and evapotranspiration than soil moisture over short time scales in many biomes (Novick et al., 2016; Sulman et al., 2016). Other studies have also suggested substantial impacts of VPD on vegetation growth (de Cárcer et al., 2018; Ding et al., 2018), forest mortality (Williams et al., 2013), and crop yields (Lobell et al., 2014). It is increasingly important to integrate the atmospheric water constraint to the carbon and water flux modeling.” (Lines 364-373 in the revised manuscript)

#### **To force the models:**

**1. Because of weather impact, the EC sites also produce very bad quality data, which could not represent the CO<sub>2</sub> exchange under real condition. To some extent, the bad quality data could be filtered out by the quality flag. To calibrate the parameters, only the good quality data should be used for all these models.**

In our manuscript, we used the quality flags of each variable to exclude the bad quality data with less than 20% of good data. We revised manuscript as follow:

“In the FLUXNET dataset, there were quality flag ranged from 0 to 1 to indicate percentage of measured and good quality gap-filled data. For each variable, we used the daily/monthly values with more than 80% of good quality data (quality flag > 0.8). We aggregated the daily values to 8-day time step. And only the 8-day measurements with more than 5-day valid values were used.” (Lines 121-124 in the revised manuscript)

**2. It’s better to keep the observed GPP and simulated GPP on the same time scale when calibrating the parameters.**

Yes, we calibrated the parameters using flux tower GPP and other variables at 8-day time scale, and we validated the model (Figs 1-2) and produced the GPP product at the same time scale (8-day). For the main objective of our manuscript is focused on the improvement of the LUE models in interannual variations and produced a long-term GPP dataset. We aggregated the 8-day simulated GPP to yearly scale to examine the ability of the LUE models in reproducing the interannual variations in GPP (Fig. 3).

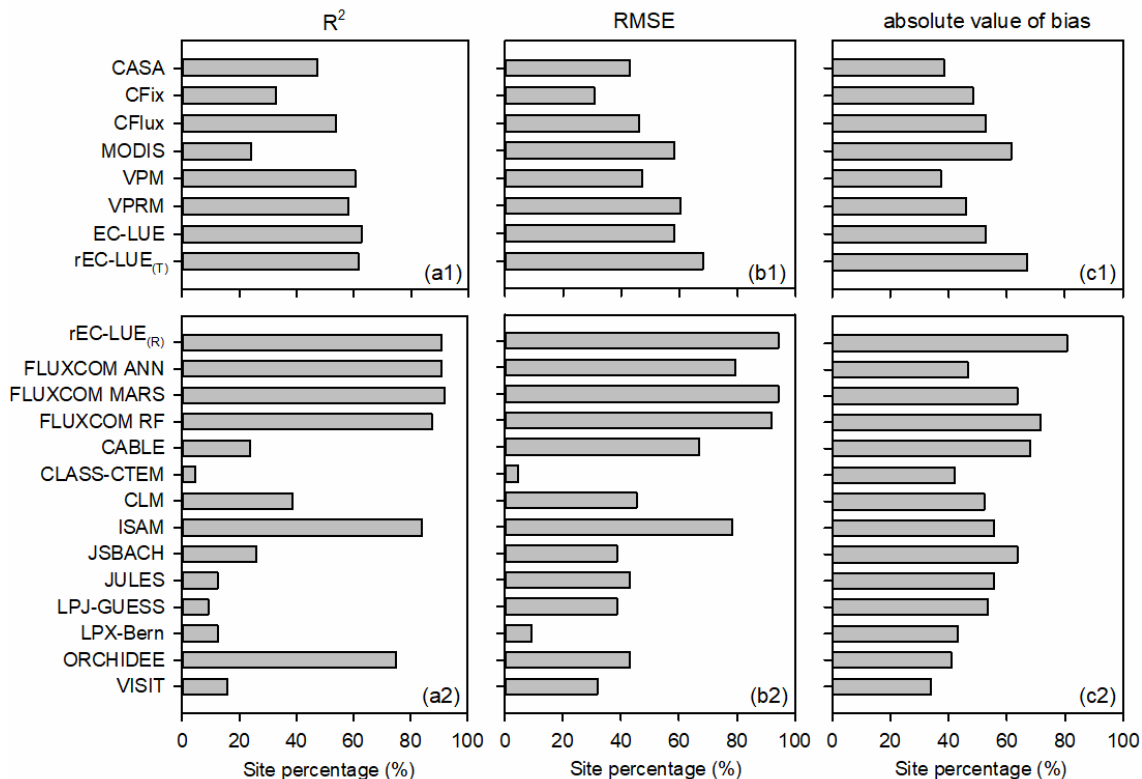
#### **To evaluate the models:**

**1. Besides giving some examples for specific sites, the site-specific R<sup>2</sup>, RMSE and bias for each model should also be given either in a plot (a scatter, histogram or boxplot) or in a table (the R<sup>2</sup>/RMSE/bias of the best model and rEC-LUE model for each site). The statistical hypothesis test between different models could help to find the best model.**

In the revised manuscript, we evaluated the performance of each model in Fig.4 at each site by comparing the simulated GPP against the tower estimated GPP at monthly step using R<sup>2</sup>, RMSE, and absolute value of bias. For each site, we compared the R<sup>2</sup>/RMSE/absolute value of bias of the individual model with the averaged value of all the models (one site has an averaged R<sup>2</sup>/RMSE/absolute value of bias). Then we added a boxplot to show which model in Fig. 4a and Fig. 4b had more sites with higher R<sup>2</sup>/lower RMSE/lower absolute value of bias than the averaged values of all the models (Fig. S4). The Results section was modified accordingly as follows:

“Additionally, we examined the model performance of the revised EC-LUE model, other LUE models, machine learning models, and processed-based models in TRENDY at monthly step by comparing against EC tower estimated GPP (Fig. 4). In comparison with seven LUE models driven by EC tower-based meteorology dataset, the overall R<sup>2</sup> of the revised EC-LUE

model was 0.71, higher than the original EC-LUE model and other LUE models ( $R^2$  ranged from 0.55 to 0.61) (Fig. 4a). For each site, we compared the  $R^2$ /RMSE/absolute value of bias of the individual model with the averaged value of all the eight LUE models (each site has an averaged  $R^2$ /RMSE/absolute value of bias) (Fig. S4a1-c1). The revised EC-LUE model had higher  $R^2$  than the mean  $R^2$  of the eight LUE models at 62% sites, which was comparable with the original EC-LUE model (63% sites) and VPM model (60% sites) (Fig. S4a1). Moreover, the revised EC-LUE model showed the lower RMSE and bias compared to mean values of all eight LUE models at 68% and 67% sites respectively, which indicated the better performance compared to the other LUE models at most sites (Fig. S4b1-c1). By using the global reanalysis meteorology data, we compared the performance of the revised EC-LUE model with three existing machine learning model products and ten process-based model products in TRENDY (Fig. 4b). The overall  $R^2$  of the revised EC-LUE model ( $R^2 = 0.57$ ) was higher than that of other models ( $R^2$  ranged from 0.02 to 0.54) (Fig. 4b). The revised EC-LUE model, FLUXCOM ANN, and FLUXCOM MARS had more sites (over 90%) with higher  $R^2$  than the mean  $R^2$  (Fig. S4a2). And the revised EC-LUE model, FLUXCOM MARS, and FLUXCOM RF showed the lower RMSE at more than 90% sites (Fig. S4b2). Compared to the other models, the revised EC-LUE model had highest site percentage (81%) with lower absolute value of bias (Fig. S4c2).” (Lines 291-306 in the revised manuscript)



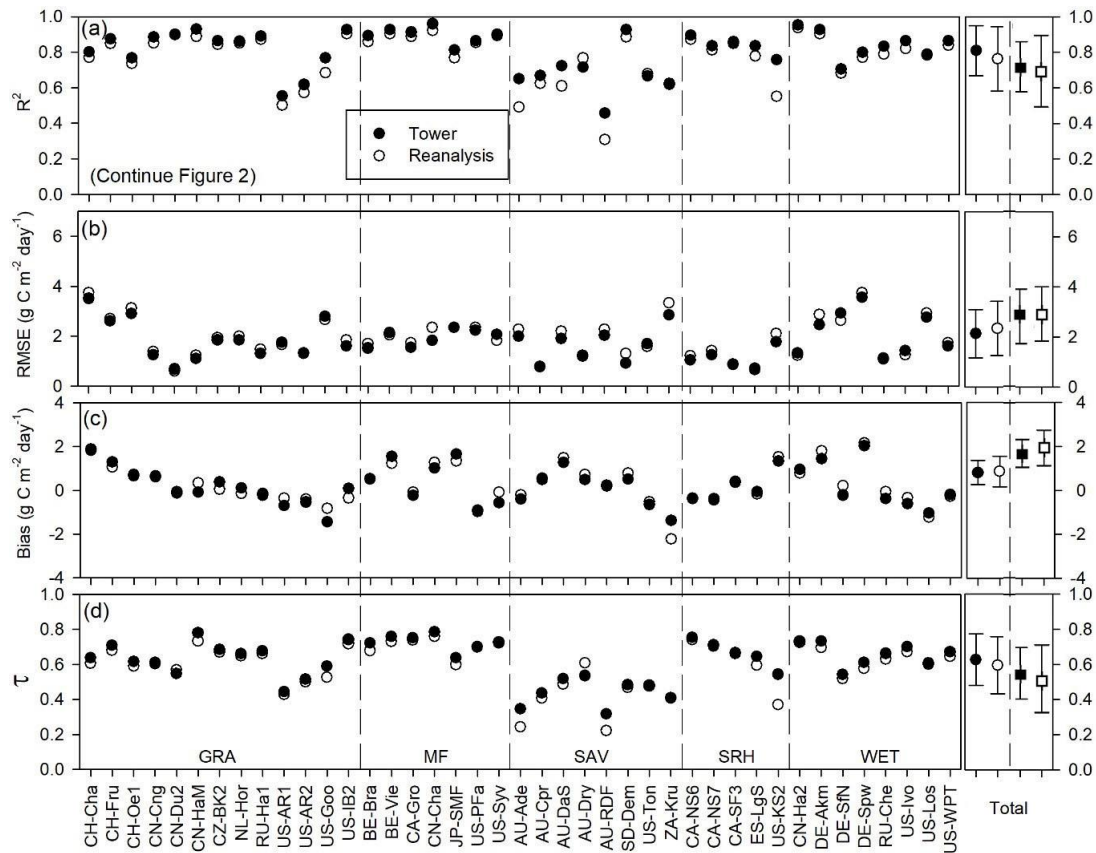
**Figure S4: Percentage of sites for each model in Fig.4 where (a1, a2)  $R^2$  of individual model > averaged  $R^2$  of all models, (b1, b2) RMSE of individual model < averaged RMSE of all models, and (c1, c2) RMSE of individual model < averaged**



**absolute value of bias of all models. All the monthly GPP simulations in (a1)-(c1) derived from tower-derived meteorology data, and all the monthly GPP simulations in (a2)-(c2) derived from global reanalysis meteorology data.**

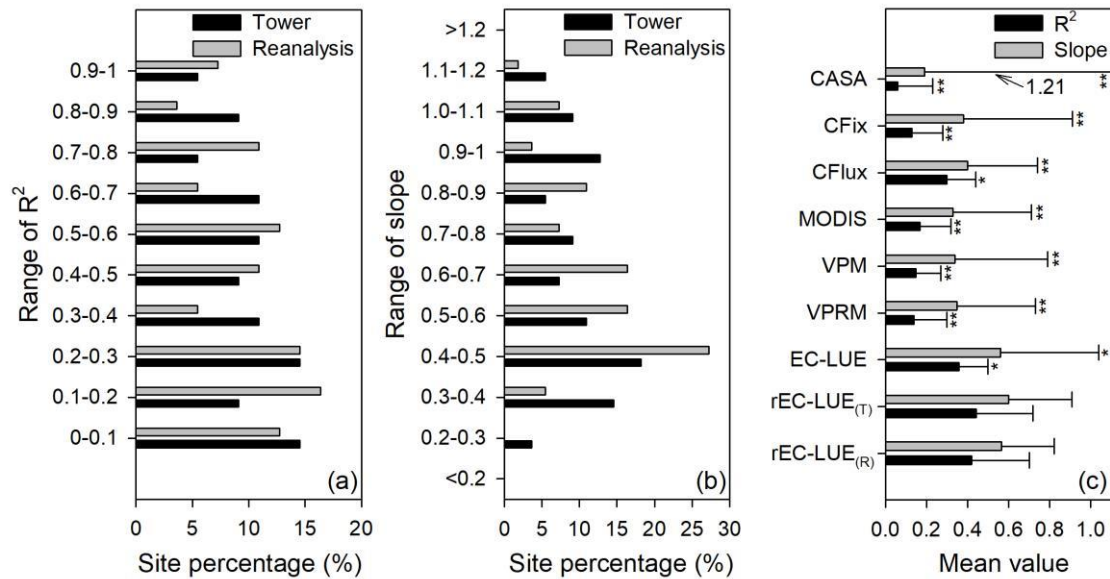
Besides the model comparison at monthly step, we compared the overall accuracy of the revised EC-LUE model and the original EC-LUE model in Fig. 2 using four metrics ( $R^2$ /RMSE/bias/seasonal index  $\tau$ ). The result indicated the performance of the revised EC-LUE model was better than the original EC-LUE model. Yuan et al. (2014) compared the performance of seven models (i.e., EC-LUE, VPRM, VPM, MODIS, CFlux, CFix, and CASA), and the EC-LUE model and CFlux performed better than other models. Therefore, we can conclude the result that the revised EC-LUE model was in high accuracy when compared with other EC-LUE models. We updated Fig. 2 and the results:

**“On average, the revised EC-LUE model showed higher  $R^2$  and  $\tau$ , and lower bias and RMSE than the original EC-LUE model for both (Fig. 2).”** (Lines 270-271)



**Figure 2: Comparisons of 8-day mean GPP between the model simulated GPP and tower estimated GPP. Solid and open dots indicate the GPP simulations of the revised EC-LUE model derived from tower-derived meteorology data and meteorological reanalysis dataset, respectively; solid and open squares indicate the GPP simulations of the original EC-LUE model derived from tower-derived meteorology data and meteorological reanalysis dataset, respectively.**

**2. Figure 3: the highest slope is still under 0.5, does it mean the GPP is underestimated for most sites?**



**Figure 3: Site percentage of (a) correlation coefficients ( $R^2$ ), and (b) regression slopes between the model simulated and tower-based interannual variabilities in GPP. (c) Averaged values (error bars represent the standard deviations) of  $R^2$  and slope for various LUE models. rEC-LUE<sub>(T)</sub> and rEC-LUE<sub>(R)</sub> indicate the revised EC-LUE models derived from tower-derived meteorology data and meteorological reanalysis dataset. The  $R^2$  and slopes of the other seven LUE models (i.e., EC-LUE, VPRM, VPM, MODIS, CFlux, CFix, and CASA) in the figure were obtained from the study by Yuan et al. (2014). \*\* and \* indicate a significant difference in statistic variables ( $R^2$  and slope) between the rEC-LUE<sub>(T)</sub> and other LUE models (i.e., rEC-LUE<sub>(T)</sub> and other seven LUE models) at  $p$ -value < 0.01 and  $p$ -value < 0.05, respectively.**

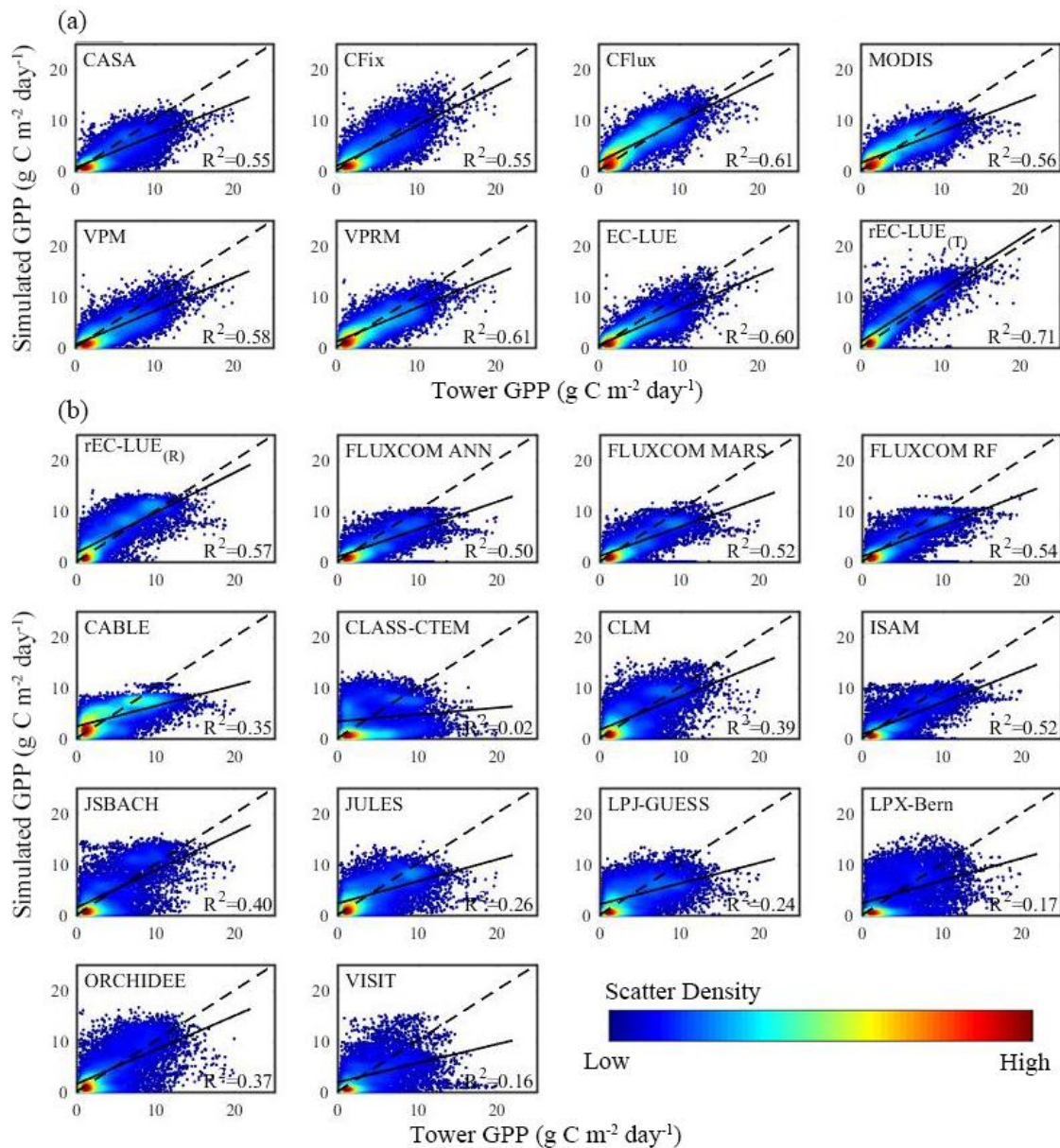
Yes, it is indeed. Although the revised EC-LUE model has improved the performance for reproducing interannual variability of GPP, there are still uncertainties in simulating interannual variability of GPP. As we discussed in the manuscript, we did not integrate the regulation of soil nitrogen content on vegetation production. Atmospheric nitrogen deposition has exhibited a large increasing trend in the past few decades because of the excessive fossil fuel combustion in the industrial and transportation sectors and the abuse of nitrogenous fertilizer in the agricultural practice (Galloway et al., 2004). However, leaf water, starch, lignin, and cellulose overlap with the absorption characters of nitrogen in the shortwave infrared bands, making it difficult to retrieve the nitrogen content (Kokaly and Clark, 1999). What's more, canopy structures, background, and illumination/viewing geometry can further decrease the capacity to detect leaf nitrogen (Yoder and Pettigrew-Crosby, 1995; Knyazikhin et al., 2013). Therefore, the future studies need strengthen the algorithm development of nitrogen impacts on vegetation production, and we discuss this issue in the manuscript (Line 426-440 in the revised manuscript).

**3. Figure 4: most of the GPP are close to 0. Please check the data quality, if data is good, could try data only in growing season.**

Thanks for your constructive comment. Yes, we excluded the measurements with poor data quality using quality flag

(refer comment #1 in “To force the models”) and non-growing season. For it is difficult to define growing season globally, we used 0°C as the temperature threshold of growing season. We updated the results:

“Additionally, we examined the model performance of the revised EC-LUE model, other LUE models, machine learning models, and processed-based models in TRENDY at monthly step by comparing against EC tower estimated GPP (Fig. 4). In comparison with seven LUE models driven by EC tower-based meteorology dataset, the overall  $R^2$  of the revised EC-LUE model was 0.71, higher than the original EC-LUE model and other LUE models ( $R^2$  ranged from 0.55 to 0.61) (Fig. 4a). For each site, we compared the  $R^2$ /RMSE/absolute value of bias of the individual model with the averaged value of all the eight LUE models (each site has an averaged  $R^2$ /RMSE/absolute value of bias) (Fig. S4a1-c1). The revised EC-LUE model had higher  $R^2$  than the mean  $R^2$  of the eight LUE models at 62% sites, which was comparable with the original EC-LUE model (63% sites) and VPM model (60% sites) (Fig. S4a1). Moreover, the revised EC-LUE model showed the lower RMSE and bias compared to mean values of all eight LUE models at 68% and 67% sites respectively, which indicated the better performance compared to the other LUE models at most sites (Fig. S4b1-c1). By using the global reanalysis meteorology data, we compared the performance of the revised EC-LUE model with three existing machine learning model products and ten process-based model products in TRENDY (Fig. 4b). The overall  $R^2$  of the revised EC-LUE model ( $R^2 = 0.57$ ) was higher than that of other models ( $R^2$  ranged from 0.02 to 0.54) (Fig. 4b).” (Lines 291-303 in the revised manuscript)



**Figure 4: Comparisons between estimated GPP based on EC measurements and GPP simulations in growing season (defined as temperature larger than 0°C) by the various models (including LUE models, machine learning models, and process-based models in TRENDY) at monthly scale. The comparison of GPP simulations were simulated using (a) tower-derived meteorology data and (b) global reanalysis meteorology data, respectively (see method 2.4). rEC-LUE<sub>(T)</sub> and rEC-LUE<sub>(R)</sub> indicate the simulations of the revised EC-LUE model derived from tower-derived meteorology data and global reanalysis meteorology data, respectively.**

#### 4. Could you give a reason why the $r^2$ of FLUXCOM GPP is lower than their paper (Jung et al., 2017)?

Jung et al. (2017) assessed the consistency of global GPP anomaly patterns between FLUXCOM and TRENDY datasets, and didn't validate the FLUXCOM GPP products against the tower-based GPP. The reviewer may talk about the paper of Tramontana et al. (2016), which developed four machine learning models (FLUXCOM ANN, MARS, RF, and KRR) and evaluated the performance of FLUXCOM products based on the measurements of eddy covariance towers.

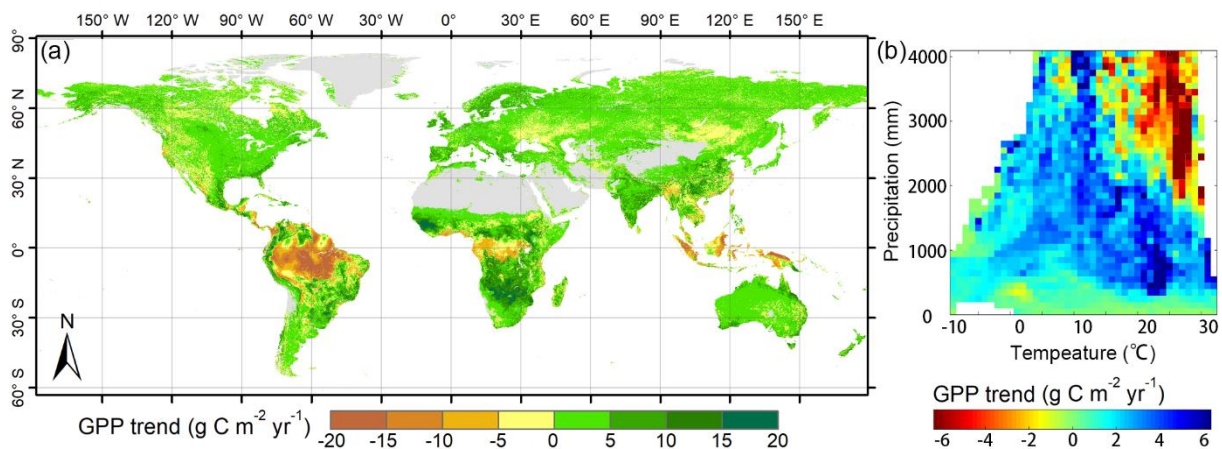
Our validation result is different with that of Tramontana et al (2016), and there are three major causes as followings:

(1) Tramontana et al (2016) and this study used difference forcing data when conducted the model validation, which may be the most important cause for different validation results. Tramontana et al. (2016) developed and validated FLUXCOM ANN, MARS, RF, and KRR models based on the in situ measured meteorology data. On contrary, this study used global FLUXCOM products driven by the reanalysis meteorological datasets. The uncertainties of global reanalysis meteorological datasets may largely decrease the accuracy of global FLUXCOM products.

(2) Tramontana et al (2016) developed four machine learning models (i.e., RF, MARS, ANN, and KRR), and validated the accuracy based on the median estimates of four methods against the measurements of eddy covariance towers. However, there are only the products of three methods (ANN, MARS, and RF) available for global analyses. In our manuscript, we directly validated three products individually, and did not use median values of three products, which may be another cause for lower  $R^2$  of FLUXCOM GPP. In addition, KRR product is not available, and thus we cannot evaluate the contribution of KRR methods to validation of Tramontana et al (2016).

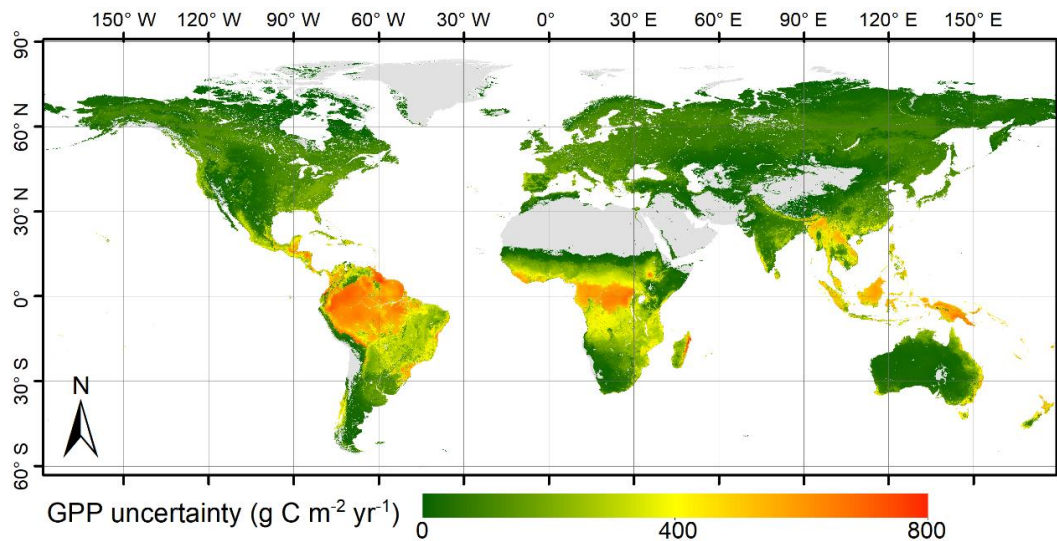
(3) Tramontana et al. (2016) developed and evaluated the four machine learning models using the La Thuile dataset, while we validated the monthly FLUXCOM GPP products using the FLUXNET2015 dataset. The two datasets have differences in site coverage, study period, and data processing methods, which probably induce some differences.

#### 5. Figure 6: it seems the negative GPP trend happen in the region where the uncertainty is relatively big. This should be discussed.



**Figure 6: Spatial pattern of global GPP trend simulated by the revised EC-LUE models during 1982–2017: (a) trend of annual GPP, (b) trend of annual GPP at different temperature and precipitation gradients.**





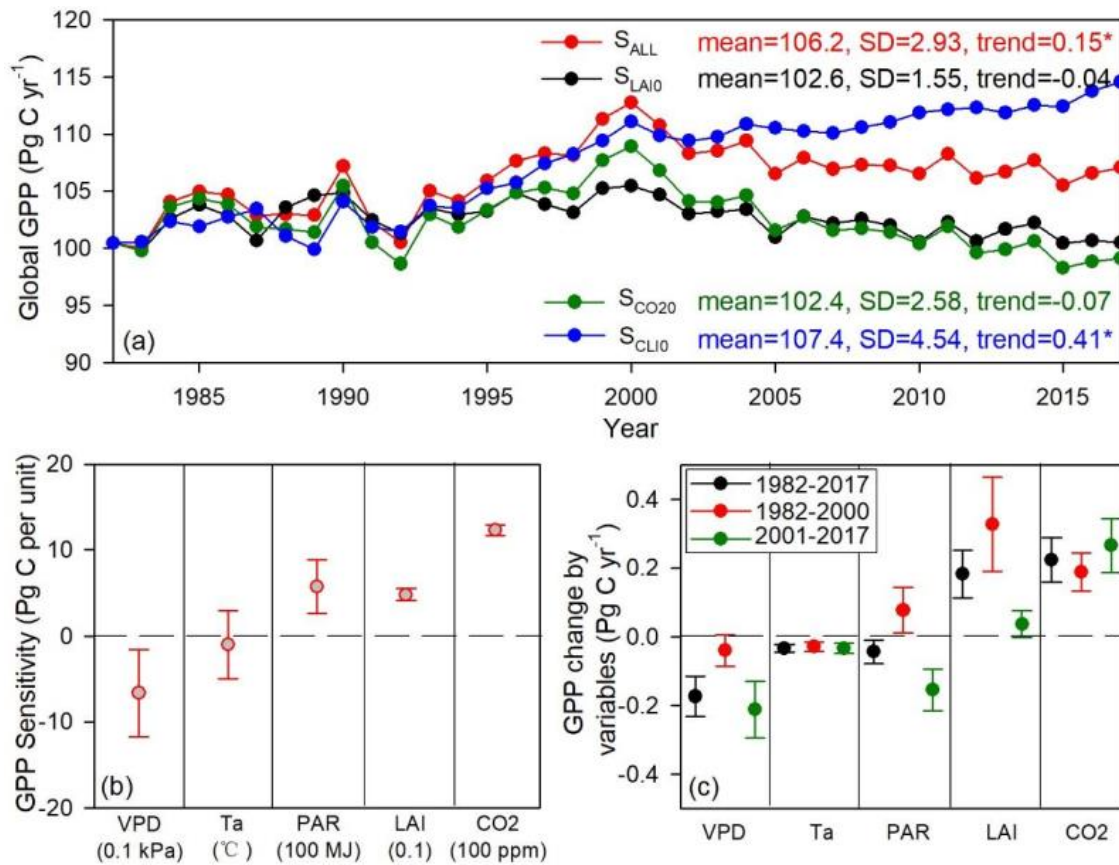
**Figure 7: Spatial pattern of the uncertainty in global GPP simulated by the revised EC-LUE model.**

The GPP uncertainties were high in tropical forests (Fig. 7). It may be introduced by cloudiness contamination and the quality of optical satellite data. Coincidentally, a decrease of GPP was found in these areas in our study. Similarly, many studies reported a decrease of tropic carbon by using satellite-based optical and microwave data (Samanta et al., 2010; Fan et al., 2019; Liu et al., 2015). We discussed these issues in our manuscript.

“The uncertainties of our GPP dataset were low in high and middle latitude areas but high in tropical areas (Fig. 7). This is consistent with the validations at site level that the revised EC-LUE model showed the lowest accuracy over the tropical evergreen broadleaf forest sites (Fig. 2). Similarly, other satellite-based models exhibited a large uncertainty in the GPP simulations over tropical forest areas (Ryu et al., 2011; Yuan et al., 2014). For example, MODIS GPP product (MOD17) underestimated GPP at high productivity sites over the tropical evergreen forests (de Almeida et al., 2018). Regarding the quality of satellite data, a high cloud cover exists over tropical regions, introducing large uncertainties to FAPAR/LAI and other vegetation indices (e.g., NDVI and EVI). As suggested by de Almeida et al. (2018), the lack of reliable MOD15 FAPAR data during January to April as a result of the cloudiness contamination could have substantially affected the seasonality of GPP estimates. Besides, the quality of satellite data can even affect the evaluation of the interannual variations in GPP. Using MODIS EVI data, Saleska et al. (2007) reported a large-scale green up in the Amazon evergreen forests during the drought of 2005. However, an opposite conclusion was drawn when the cloud-contaminated data were excluded from the analysis (Samanta et al., 2010). In our study, a significant decrease of GPP was found in the Amazon evergreen forests, which may be resulted from the sharp increase in VPD after the late 1990s (Yuan et al., 2019). Studies using optical satellite data can be influenced by the cloudiness contamination. Recently studies using cloud free satellite-based microwave data also reported a carbon loss in tropic forest (Liu et al., 2015; Fan et al., 2019).” (Lines 406-420 in the revised manuscript)

**6. Figure 8: it’s incredible that the GPP sensitivity to VPD is so much higher than to air temperature. Can you compare this with the sensitivity derived from other LUE model or GPP products?**





**Figure 8: Long-term changes in global GPP and the environmental regulations: (a) Global summed GPP derived from the four experimental simulations in section 2.5, (b) GPP sensitivity to climate variables (i.e., VPD, Ta, and PAR), LAI, and atmospheric CO<sub>2</sub>, (c) contributions of climate variables (i.e., VPD, Ta, and PAR), LAI, and atmospheric CO<sub>2</sub> to GPP changes over 1982–2017, 1982–2000, and 2001–2017. \* indicates the significant level at  $p$ -value<0.05.**

In our manuscript, global GPP decreased by  $6.67 \pm 5.04$  Pg C with an increased VPD of 0.1 kPa and global GPP slightly decreased by about 0.5 Pg C with an increased temperature of 1°C. GPP sensitivity to VPD (0.1 kPa) was so much higher than to air temperature (1°C). Increased temperature has a positive effect on vegetation productivity at high latitude and a negative effect on vegetation productivity at low latitude. The influence of increased temperature on vegetation growth will partly counteract globally. While the increased VPD suppresses vegetation growth at both low and high latitude. Therefore, at global scale, the influence of VPD can be higher than temperature. Similar result can be obtained from other models. Yuan et al. (2019) quantified the impacts of VPD and temperature on long-term changes of GPP using two satellite-based models (EC-LUE and MODIS model). For the EC-LUE model, global GPP decreased by  $13.82 \pm 3.12$  Pg C with an increased VPD of 0.1 kPa and global GPP increased by about 6 Pg C with an increased temperature of 1°C. For the MODIS model, global GPP decreased by  $18.29 \pm 3.65$  Pg C with an increased VPD of 0.1 kPa and global GPP increased by about 1 Pg C with an increased temperature of 1°C.

**7. Some letters in this paper look weird: 'v', 'x', 'z' and '%' are extremely small comparing other letters, please check the format before upload.**

Thank you, we checked these formats.

## References

1. Copernicus Climate Change Service (C3S): ERA5: Fifth generation of ECMWF atmospheric reanalyses of the global climate, C3S Climate Data Store, available at: <https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era5>, 2017.
2. Cox PM, Huntingford C, Harding RJ.: A canopy conductance and photosynthesis model for use in a GCM land surface scheme. *J. Hydrol.*, 212–213, 79–94, 1998.
3. de Almeida, C.T., Delgado, R.C., Galvao, L.S., de Oliveira Cruz e Aragao, L.E., Concepcion Ramos, M.: Improvements of the MODIS Gross Primary Productivity model based on a comprehensive uncertainty assessment over the Brazilian Amazonia, *ISPRS J. Photogramm. Remote Sens.*, 145, 268–283, doi:10.1016/j.isprsjprs.2018.07.016, 2018.
4. Doelling, D. R., Loeb, N. G., Keyes, D. F., Nordeen, M. L., Morstad, D., Nguyen, C., Wielicki, B. A., Young, D. F., and Sun, M.: Geostationary Enhanced Temporal Interpolation for CERES Flux Products, *J. Atmos. Ocean. Tech.*, 30, 1072–1090, 2013.
5. Eller, C. B., Rowland, L., Oliveira, R. S., Bittencourt, P. R. L., Barros, F. V., da Costa, A. C. L., Meir, P., Friend, A. D., Mencuccini, M., Sitch, S., and Cox, P.: Modelling tropical forest responses to drought and El Niño with a stomatal optimization model based on xylem hydraulics. *Philos. Trans. R. Soc. London, Ser. B*, 373, 20170315, 2018.
6. Fan, L., Wigneron, J.-P., Ciais, P., Chave, J., Brandt, M., Fensholt, R., Saatchi, S. S., Bastos, A., Al-Yaari, A., Hufkens, K., Qin, Y., Xiao, X., Chen, C., Myneni, R. B., Fernandez-Moran, R., Mialon, A., Rodriguez-Fernandez, N. J., Kerr, Y., Tian, F., and Penuelas, J.: Satellite-observed pantropical carbon dynamics, *Nature Plants*, 5, 944–951, 2019.
7. Fletcher, A.L., Sinclair, T.R., Allen, L.H.: Transpiration responses to vapor pressure deficit in well watered 'slow-wilting' and commercial soybean, *Environ. Exp. Bot.*, 61, 145–151, doi:10.1016/j.envexpbot.2007.05.004, 2007.
8. Galloway, J.N., Dentener, F.J., Capone, D.G., Boyer, E.W., Howarth, R.W., Seitzinger, S.P., Asner, G.P., Cleveland, C.C., Green, P.A., Holland, E.A., Karl, D.M., Michaels, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J.: Nitrogen cycles: past, present, and future, *Biogeochemistry*, 70, 153–226, doi:10.1007/s10533-004-0370-0, 2004.
9. Grossiord, C., Buckley, T. N., Cernusak, L. A., Novick, K. A., Poulter, B., Siegwolf, R. T. W., Sperry, J. S., and McDowell, N. G.: Plant responses to rising vapor pressure deficit, *New Phytologist*, 226, 1550–1566, 2020.
10. Harris, I. C.: CRU JRA v1.1: A forcings dataset of gridded land surface blend of Climatic Research Unit (CRU) and Japanese reanalysis (JRA) data, Unit, available at: <https://catalogue.ceda.ac.uk/uuid/13f3635174794bb98cf8ac4b0ee8f4ed>, 2019.
11. Huffman, G. J., Adler, R. F., Morrissey, M. M., Bolvin, D. T., Curtis, S., Joyce, R., McGavock, B., and Susskind, J.: Global Precipitation at One-Degree Daily Resolution from Multisatellite Observations, *J. Hydrometeorol.*, 2, 36–50, 2001.
12. Kim, H.: Global Soil Wetness Project Phase 3 Atmospheric Boundary Conditions (Experiment 1), Data set, <https://doi.org/10.20783/DIAS.501>, 2017.
13. Jung, M., Reichstein, M., Schwalm, C.R., Huntingford, C., Sitch, S., Ahlstrom, A., Arneeth, A., Camps-Valls, G., Ciais, P., Friedlingstein, P., Gans, F., Ichii, K., Ain, A.K.J., Kato, E., Papale, D., Poulter, B., Raduly, B., Rodenbeck, C., Tramontana, G., Viovy, N., Wang, Y.-P., Weber, U., Zaehle, S., Zeng, N.: Compensatory water effects link yearly global land CO<sub>2</sub> sink changes to temperature, *Nature*, 541, 516–520, doi:10.1038/nature20780, 2017.
14. Knyazikhin, Y., Schull, M.A., Stenberg, P., Mottus, M., Rautiainen, M., Yang, Y., Marshak, A., Latorre Carmona, P.,

- Kaufmann, R.K., Lewis, P., Disney, M.I., Vanderbilt, V., Davis, A.B., Baret, F., Jacquemoud, S., Lyapustin, A., Myneni, R.B.: Hyperspectral remote sensing of foliar nitrogen content, *Proc. Natl. Acad. Sci. U.S.A.*, 110, E185-E192, doi:10.1073/pnas.1210196109, 2013.
15. Kokaly, R.F., Clark, R.N.: Spectroscopic determination of leaf biochemistry using band-depth analysis of absorption features and stepwise multiple linear regression, *Remote Sens. Environ.*, 67, 267-287, doi:10.1016/s0034-4257(98)00084-4, 1999.
  16. Konings, A.G., Williams, A.P., Gentine, P.: Sensitivity of grassland productivity to aridity controlled by stomatal and xylem regulation, *Nat. Geosci.*, 10, 284-+, doi:10.1038/ngeo2903, 2017.
  17. Liu, Y. Y., van Dijk, A. I. J. M., de Jeu, R. A. M., Canadell, J. G., McCabe, M. F., Evans, J. P., and Wang, G.: Recent reversal in loss of global terrestrial biomass, *Nat. Clim. Change*, 5, 470-474, 2015.
  18. Pierce, D.W., Westerling, A.L., Oyler, J.: Future humidity trends over the western United States in the CMIP5 global climate models and variable infiltration capacity hydrological modeling system, *Hydrol. Earth Syst. Sci.*, 17, 1833-1850, doi:10.5194/hess-17-1833-2013, 2013.
  19. Potter, C.S., Randerson, J.T., Field, C.B., Matson, P.A., Vitousek, P.M., Mooney, H.A., Klooster, S.A. Terrestrial ecosystem production: a process model-based on global satellite and surface data. *Glob. Biogeochem. Cycles* 7, 811–84, 1993.
  20. Rawson, H.M., Begg, J.E., Woodward, R.G.: The effect of atmospheric humidity on photosynthesis, transpiration and water use efficiency of leaves of several plant species, *Planta*, 134, 5-10, doi:10.1007/bf00390086, 1977.
  21. Ryu, Y., Baldocchi, D.D., Kobayashi, H., van Ingen, C., Li, J., Black, T.A., Beringer, J., van Gorsel, E., Knohl, A., Law, B.E., Rouspard, O.: Integration of MODIS land and atmosphere products with a coupled-process model to estimate gross primary productivity and evapotranspiration from 1 km to global scales, *Global Biogeochem. Cycles*, 2510.1029/2011gb004053, 2011.
  22. Saleska, S.R., Didan, K., Huete, A.R., da Rocha, H.R.: Amazon forests green-up during 2005 drought, *Science*, 318, 612-612, doi:10.1126/science.1146663, 2007.
  23. Samanta, A., Ganguly, S., Hashimoto, H., Devadiga, S., Vermote, E., Knyazikhin, Y., Nemani, R.R., Myneni, R.B.: Amazon forests did not green-up during the 2005 drought, *Geophys. Res. Lett.*, 3710.1029/2009gl042154, 2010.
  24. Simmons, A.J., Willett, K.M., Jones, P.D., Thorne, P.W., Dee, D.P.: Low-frequency variations in surface atmospheric humidity, temperature, and precipitation: Inferences from reanalyses and monthly gridded observational data sets, *J. Geophys. Res. D: Atmos.*, 11510.1029/2009jd012442, 2010.
  25. Tramontana, G., Jung, M., Schwalm, C. R., Ichii, K., Camps-Valls, G., Ráduly, B., Reichstein, M., Arain, M. A., Cescatti, A., Kiely, G., Merbold, L., Serrano-Ortiz, P., Sickert, S., Wolf, S., and Papale, D.: Predicting carbon dioxide and energy fluxes across global FLUXNET sites with regression algorithms, *Biogeosciences*, 13, 4291–4313, <https://doi.org/10.5194/bg-13-4291-2016>, 2016.
  26. Trugman AT, Medvigy D, Mankin JS, Anderegg WRL.: Soil moisture stress as a major driver of carbon cycle uncertainty. *Geophys. Res. Lett.*, 45, 6495–6503, 2018
  27. Ukkola AM, De Kauwe MG, Pitman AJ, Best MJ, Abramowitz G, Haverd V, Decker M, Haughton N.: Land surface models systematically overestimate the intensity, duration and magnitude of seasonal-scale evaporative droughts. *Environ.*

Res. Lett., 11,104012, 2016.

28. Van Wijngaarden, W.A., Vincent, L.A.: Trends in relative humidity in Canada from 1953–2003. In 15th Symp. on Global Change and Climate Variations, 2004.
29. Weedon, G. P., Balsamo, G., Bellouin, N., Gomes, S., Best, M. J., and Viterbo, P.: The WFDEI meteorological forcing data set: WATCH Forcing Data methodology applied to ERA-Interim reanalysis data, *Water Resour. Res.*, 50, 7505–7514, 2014.
30. Willett, K.M., Dunn, R.J.H., Thorne, P.W., Bell, S., de Podesta, M., Parker, D.E., Jones, P.D., Williams, C.N., Jr.: HadISDH land surface multi-variable humidity and temperature record for climate monitoring, *Climate of the Past*, 10, 1983-2006, doi:10.5194/cp-10-1983-2014, 2014.
31. Xiao, X.M., Zhang, Q.Y., Hollinger, D., Aber, J., Moore, B.: Modeling gross primary production of an evergreen needleleaf forest using MODIS and climate data, *Ecol. Appl.*, 15, 954-969, doi:10.1890/04-0470, 2005.
32. Yoder, B.J., Pettigrew-Crosby, R.E.: Predicting nitrogen and chlorophyll content and concentrations from reflectance spectra (400-2500 nm) at leaf and canopy scales, *Remote Sens. Environ.*, 53, 199-211, doi:10.1016/0034-4257(95)00135-n, 1995.
33. Yuan, W., Cai, W., Xia, J., Chen, J., Liu, S., Dong, W., Merbold, L., Law, B., Arain, A., Beringer, J., Bernhofer, C., Black, A., Blanken, P.D., Cescatti, A., Chen, Y., Francois, L., Gianelle, D., Janssens, I.A., Jung, M., Kato, T., Kiely, G., Liu, D., Marcolla, B., Montagnani, L., Raschi, A., Rouspard, O., Varlagin, A., Wohlfahrt, G.: Global comparison of light use efficiency models for simulating terrestrial vegetation gross primary production based on the La Thuile database, *Agric. For. Meteorol.*, 192, 108-120, doi:10.1016/j.agrformet.2014.03.007, 2014.
34. Yuan, W., Liu, S., Zhou, G., Zhou, G., Tieszen, L.L., Baldocchi, D., Bernhofer, C., Gholz, H., Goldstein, A.H., Goulden, M.L., Hollinger, D.Y., Hu, Y., Law, B.E., Stoy, P.C., Vesala, T., Wofsy, S.C., AmeriFlux, C.: Deriving a light use efficiency model from eddy covariance flux data for predicting daily gross primary production across biomes, *Agric. For. Meteorol.*, 143, 189-207, doi:10.1016/j.agrformet.2006.12.001, 2007.
35. Yuan, W., Zheng, Y., Piao, S., Ciais, P., Lombardozzi, D., Wang, Y., Ryu, Y., Chen, G., Dong, W., Hu, Z., Jain, A.K., Jiang, C., Kato, E., Li, S., Lienert, S., Liu, S., Nabel, J.E.M.S., Qin, Z., Quine, T., Sitch, S., Smith, W.K., Wang, F., Wu, C., Xiao, Z., Yang, S.: Increased atmospheric vapor pressure deficit reduces global vegetation growth, *Science Advances*, 5, eaax1396, 2019.
36. Zhou, S., Williams, A. P., Berg, A. M., Cook, B. I., Zhang, Y., Hagemann, S., Lorenz, R., Seneviratne, S. I., and Gentine, P.: Land-atmosphere feedbacks exacerbate concurrent soil drought and atmospheric aridity, *Proc. Natl. Acad. Sci. U.S.A.*, 116, 18848-18853, 2019a.
37. Zhou, S., Zhang, Y., Williams, A. P., and Gentine, P.: Projected increases in intensity, frequency, and terrestrial carbon costs of compound drought and aridity events, *Science Advances*, 5, eaau5740, 2019b.

# Improved estimate of global gross primary production for reproducing its long-term variation, 1982-2017

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**Abstract.** Satellite-based models have been widely used to simulate vegetation gross primary production (GPP) at the site, regional, or global scales in recent years. However, accurately reproducing the interannual variations in GPP remains a major  
20 challenge, and the long-term changes in GPP remain highly uncertain. In this study, we generated a long-term global GPP dataset at 0.05° latitude by 0.05° longitude and 8-day interval by revising a light use efficiency model (i.e. EC-LUE model). In the revised EC-LUE model, we integrated the regulations of several major environmental variables: atmospheric CO<sub>2</sub> concentration, radiation components, and atmospheric vapor pressure deficit (VPD). These environmental variables showed substantial long-term changes, which could greatly impact the global vegetation productivity. Eddy covariance (EC)  
25 measurements at 95 towers from the FLUXNET2015 dataset, covering nine major ecosystem types around the globe, were used to calibrate and validate the model. In general, the revised EC-LUE model could effectively reproduce the spatial, seasonal, and annual variations in the tower estimated GPP at most sites. The revised EC-LUE model could explain 71% of the spatial variations in annual GPP over 95 sites. At more than 95% of the sites, the correlation coefficients ( $R^2$ ) of seasonal changes between tower estimated and model simulated GPP are larger than 0.5. Particularly, the revised EC-LUE model improved the  
30 model performance in reproducing the interannual variations in GPP, and the averaged  $R^2$  between annual mean tower estimated and model simulated GPP is 0.62 over all 55 sites with observations longer than 5-years, which is significantly higher than those of original EC-LUE model ( $R^2 = 0.36$ ) and other LUE models ( $R^2$  ranged from 0.06 to 0.30 with an average value of 0.16). At the global scale, GPP derived from light use efficiency models, machine learning models, and processes-based biophysical models exist substantial differences in magnitude and interannual variations. The revised EC-LUE model  
35 quantified the mean global GPP from 1982 to 2017 as  $106.2 \pm 2.9$  Pg C yr<sup>-1</sup> with the trend 0.15 Pg C yr<sup>-1</sup>. Sensitivity analysis



indicated that GPP simulated by the revised EC-LUE model was sensitive to atmospheric CO<sub>2</sub> concentration, VPD, and radiation. Over the period of 1982–2017, the CO<sub>2</sub> fertilization effect on the global GPP ( $0.22 \pm 0.07 \text{ Pg C yr}^{-1}$ ) could be partly offset by increased VPD ( $-0.17 \pm 0.06 \text{ Pg C yr}^{-1}$ ). The long-term changes in the environmental variables could be well reflected in global GPP. Overall, the revised EC-LUE model is able to provide a reliable long-term estimate of global GPP. The GPP dataset is available at <https://doi.org/10.6084/m9.figshare.8942336> (Zheng et al., 2019).

## 1 Introduction

Vegetation gross primary production (GPP) is the largest carbon flux component within terrestrial ecosystems and plays an essential role in regulating the global carbon cycle (Canadell et al., 2007; Zhao et al., 2010). As a primary variable of the terrestrial ecosystem cycle, GPP estimates will substantially determine other variables of the carbon cycle (Yuan et al., 2011).  
45 Satellite-based GPP models have been developed based on the light use efficiency (LUE) principle (Monteith, 1972; Potter et al., 1993; Running et al., 2004; Xiao et al., 2005; Yuan et al., 2007). Thus far, LUE models have been a major tool for investigating the spatio-temporal changes in GPP and the environmental dominantes, either independently or by combining with other ecosystem models (Keenan et al., 2016; Smith et al., 2016).

However, current LUE models exhibit poor performance in reproducing the interannual variations in GPP. A previous study  
50 indicated that seven LUE models could only explain 6–36% of the interannual variations in GPP at 51 eddy covariance (EC) towers (Yuan et al., 2014). Similarly, a model comparison showed that none of the examined 16 processes-based biophysical models or the 3 remote sensing products (BESS, MODIS C5, and MODIS C5.1) could consistently reproduce the observed interannual variations in GPP at 11 forest sites in North America (Keenan et al., 2012). Seven LUE models simulated the long-term trends in global GPP varied from  $-0.15$  to  $1.09 \text{ Pg C yr}^{-1}$  over the period 2000–2010 (Cai et al., 2014). An important  
55 reason for the poor performance in modeling the interannual variability is that the effect of environmental regulations on vegetation production is not completely integrated into the LUE models (Stocker et al., 2019). In particular, the long-term changes in several environmental variables are very important for accurately simulating the GPP series at the decadal scale.

Several environmental variables should be included in GPP models. Firstly, as we all know the rising atmospheric CO<sub>2</sub> concentration in the past few decades substantially stimulated global vegetation growth (Zhu et al., 2016; Liu et al., 2017).  
60 Field experiments using greenhouses or open-top chambers showed that an increase of approximately 300 ppm in CO<sub>2</sub> concentration can increase the photosynthesis of C3 plants on the order of 60% (Norby et al., 1999). Free-air CO<sub>2</sub> enrichment (FACE) experiments generally confirmed the enhancement in net primary production (NPP) with the rising CO<sub>2</sub> concentration (Ainsworth and Long, 2005). For example, four FACE experiments indicated that the forest NPP consistently increased at the median of  $23 \pm 2\%$  when the ambient CO<sub>2</sub> concentration was elevated to approximately 550 ppm (Norby et al., 2005).  
65 According to observations, the atmospheric CO<sub>2</sub> concentration has risen by approximately 20% from 340 ppm (1982) to 410 ppm (2018) (<https://www.esrl.noaa.gov/>). However, the effects of CO<sub>2</sub> fertilization on GPP have not been integrated in most current satellite-based LUE models.

Secondly, solar radiation, or more specifically the photosynthetic active radiation (PAR) substantially influences the vegetation production of the terrestrial ecosystem (Alton et al., 2007; Kanniah et al., 2012; Krupkova et al., 2017). Study indicated that the solar radiation incident at the earth surface underwent significant decadal variations (Wild et al., 2005). A comprehensive analysis based on the datasets of worldwide distributed sites indicated significant decreases in solar radiation (2% per decade) from the late 1950s to 1990 in the regions of Asia, Europe, North America, and Africa (Gilgen et al., 1998). A later assessment by Wild et al. (2005) showed that the radiation increased at widespread locations since the mid-1980s.

However, not only the total amount of solar radiation or PAR incident at the earth surface but more importantly, their partitioning into direct and diffuse radiations, impact the vegetation productivity (Urban et al., 2007; Kanniah et al., 2012). Increased proportion of diffuse radiation enhances vegetation photosynthesis, because a higher blue/red light ratio within the diffuse radiation may lead to higher light use efficiency (Gu et al., 2002; Alton et al., 2007). For example, the sharply increased diffuse radiation induced by the 1991 Mount Pinatubo eruption enhanced the noontime vegetation productivity of a deciduous forest in the next 2 years (Gu et al. 2003). Besides volcanic aerosols, clouds could also reduce the total and direct radiation, while increase the proportion of diffuse radiation. Yuan et al. (2010) found that the higher LUE at European forests than North America was because of the higher ratio of cloudy days in Europe. Yuan et al. (2014) further proved that the significantly underestimated GPP during cloudy days by six LUE models was because the effects of diffuse radiation on LUE were neglected in these models.

Thirdly, atmospheric vapor pressure deficit (VPD) is another factor that should be included in GPP models. As an important driver of atmospheric water demand for plants, VPD influences terrestrial ecosystem function and photosynthesis (Rawson et al., 1977; Yuan, et al., 2019). Rising air temperature increases the saturated vapor pressure at a rate of  $\sim 7\%/K$  according to the Clausius–Clapeyron relationship, and therefore, VPD will increase if the atmospheric water vapor content does not increase by exactly the same amount as the saturated vapor pressure. Numerous studies indicated significant changes in the relative humidity (ratio of actual water vapor pressure to saturated water vapor pressure) in both humid areas and continental areas located far from oceanic humidity (Van Wijngaarden and Vincent, 2004; Pierce et al., 2013). In particular, the global averaged land surface relative humidity decreased sharply after the late 1990s (Simmons et al. 2010; Willett et al. 2014); [and the global averaged land surface VPD increased sharply after the late 1990s \(Yuan, et al., 2019\)](#). The leaf and canopy photosynthetic rate declines when the atmospheric VPD increases due to stomatal closure (Fletcher et al., 2007). A recent study highlighted that increases in VPD rather than changes in precipitation would be a dominant influence on vegetation productivity (Konings et al., 2017). However, currently the influence of long-term VPD variations is not well expressed in many LUE models.

We have developed a LUE model, namely the EC-LUE model, by integrating remote sensing data and eddy covariance data to simulate daily GPP (Yuan et al., 2007; 2010). The model has been evaluated using the observations at EC towers located in Europe, North America, China, and East Asia, covering various ecosystem types (Yuan et al., 2007; 2010; Li et al., 2013). In this study, we revised the EC-LUE model by integrating the impacts of several environmental variables (i.e., atmospheric CO<sub>2</sub> concentration, radiation components, and atmospheric VPD) across a long-term temporal scale. Firstly, we evaluated the effectiveness of the revised EC-LUE model in determining the spatial, seasonal, and interannual variations in GPP from

multiple eddy covariance sites. Secondly, a global GPP dataset at 0.05° spatial resolution was generated based on the optimized model. Finally, we analyzed the contributions of the aforementioned environmental variables to the global GPP and discussed the spatial and interannual variations in GPP from different datasets.

## 105 2 Data and Methods

### 2.1 Data from the eddy covariance towers

The FLUXNET2015 dataset (<http://www.fluxdata.org>) includes over 200 variables of carbon fluxes, energy fluxes, and meteorological variables collected and processed at sites by the FLUXNET community. In our study, ninety-five EC sites in FLUXNET2015 dataset were utilized to optimize the parameters and evaluate the performance of the revised EC-LUE model, including nine major terrestrial ecosystem vegetation types (Table 1): evergreen broadleaf forests (EBF), evergreen needleleaf forests (ENF), deciduous broadleaf forests (DBF), mixed forests (MF), grasslands (GRA), savannas (SAV), shrubland (SHR), wetlands (WET), and croplands (CRO). More information about the characteristics of these sites can be referred to the FLUXNET website. For each site, the daily GPP, PAR, air temperature ( $T_a$ ), and VPD were used in our study. The GPP variable (GPP\_NT\_VUT\_REF) used in this study was estimated from night-time partitioning method. The corresponding net ecosystem exchange (NEE) was generated using variable friction velocity (USTAR) threshold for each year (VUT), in which 40 versions of NEE were created by using different percentiles of USTAR thresholds. The model efficiency between each version and the others 39 versions were calculated to test their similarities and the reference (REF) NEE was selected as the one with higher model efficiency sum (the most similar to the others 39). The daily meteorological 120 variables were gap-filled or downscaled from the ERA-interim reanalysis dataset in both space and time (Vuichard and Papale, 2015). The gap-filled technique of the carbon flux measurements and meteorological variables is the marginal distribution sampling (MDS) method described in Reichstein et al. (2005). [In the FLUXNET dataset, there were quality flag ranged from 0 to 1 to indicate percentage of measured and good quality gap-filled data.](#) For each variable, we [used the daily/monthly values with more than 80% of good quality data \(quality flag > 0.8\).](#) We aggregated the daily values to 8-day time step. ~~Only~~[And only](#) the 8-day measurements with more than 5-day valid values were used.

125 <<Table 1>>

### 2.2 Data at the global scale

The global scale datasets used in this study are shown in Table 2. The meteorological reanalysis dataset was derived from the second Modern-Era Retrospective analysis for Research and Applications (MERRA-2) dataset. It was produced by NASA's Global Modeling and Assimilation Office that used an upgraded version of the GEOS-5 (Rienecker et al., 2011). It has been validated carefully using surface meteorological datasets and enhanced assimilation system to reduce the uncertainty in various meteorological variables globally. In our study, we obtained the daily mean air temperature ( $T_a$ , °C), mean dew point

temperature ( $T_d$ , °C), total direct PAR ( $PAR_{dr}$ , MJ m<sup>-2</sup> d<sup>-1</sup>), and total diffuse PAR ( $PAR_{df}$ , MJ m<sup>-2</sup> d<sup>-1</sup>) at 0.625° in longitude by 0.5° in latitude from 1982 to 2017. VPD was calculated from air temperature and dew point temperature:

$$SVP = 6.112 \times e^{\frac{17.67T_a}{T_a+243.5}} \quad (1)$$

$$135 \quad RH = e^{\frac{17.625T_d}{T_d+243.04} - \frac{17.625T_a}{T_a+243.04}} \quad (2)$$

$$VPD = SVP \times (1 - RH) \quad (3)$$

where SVP is the saturated vapor pressure (k Pa), and RH is the relative humidity. We aggregated the daily variables (air temperature, VPD, direct PAR, and diffuse PAR) to 8-day interval temporal resolution. And these variables were resampled to the spatial resolution of 0.05° latitude by 0.05° longitude using the bilinear interpolation method.

140 The 8-day Global LAnd Surface Satellite-leaf area index (GLASS LAI) dataset at 0.05° latitude by 0.05° longitude was adopted to indicate vegetation growth from 1982 to 2017. It was produced using the general regression neural networks (GRNNs) trained with the fused MOD15 LAI and CYCLOPES LAI and the preprocessed MODIS/AVHRR reflectance data over the BELMANIP sites (Xiao et al., 2016). Products validation and comparison showed that the GLASS LAI product was spatially complete and temporally continuous with lower uncertainty (Xu et al., 2018).

145 Additionally, the MCD12Q1 product with IGBP classification scheme was used as land cover map. The ISLSCP II C4 Vegetation Percentage map was used to separate the C3 and C4 crop. The NOAA's Earth System Research Laboratory (ESRL) CO<sub>2</sub> concentration dataset was used to express the CO<sub>2</sub> fertilization effect.

<<Table 2>>

### 2.3 The revised EC-LUE model

150 The terrestrial vegetation GPP can be expressed as follows in the revised EC-LUE model:

$$GPP = (\varepsilon_{msu} \times APAR_{su} + \varepsilon_{msh} \times APAR_{sh}) \times C_s \times \min(T_s, W_s) \quad (4)$$

where  $\varepsilon_{msu}$  is the maximum LUE of sunlit leaves;  $APAR_{su}$  is the PAR absorbed by sunlit leaves;  $\varepsilon_{msh}$  is the maximum LUE of shaded leaves;  $APAR_{sh}$  is the PAR absorbed by shaded leaves;  $C_s$ ,  $T_s$ , and  $W_s$  represent the downward regulation scalars of atmospheric CO<sub>2</sub> concentration ([CO<sub>2</sub>]), air temperature, and VPD on LUE ranging from 0 to 1;  $min$  represents the minimum value.

155 The effect of atmospheric CO<sub>2</sub> concentration on GPP is determined by the following equations (Farquhar et al., 1980; Collatz et al., 1991):

$$C_s = \frac{C_i - \varphi}{C_i + 2\varphi} \quad (5)$$

$$C_i = C_a \times \chi \quad (6)$$

160 where  $\varphi$  is the CO<sub>2</sub> compensation point in the absence of dark respiration (ppm);  $C_i$  is the leaf internal CO<sub>2</sub> concentration;  $C_a$  is the atmospheric CO<sub>2</sub> concentration;  $\chi$  is the ratio of leaf internal to atmospheric CO<sub>2</sub> concentration which can be estimated as follows (Prentice et al., 2014; Keenan et al., 2016):

$$\chi = \frac{\varepsilon}{\varepsilon + \sqrt{VPD}} \quad (7)$$

$$\varepsilon = \sqrt{\frac{356.51K}{1.6\eta^*}} \quad (8)$$

165 where  $\varepsilon$  is a parameter related to the ‘carbon cost of water’, which means the sensitivity of VPD to  $\chi$ ;  $K$  is the Michaelis–Menten coefficient of Rubisco;  $\eta^*$  is the viscosity of water relative to its value at 25 °C (Korson et al., 1969).

$$K = K_c \left(1 + \frac{P_o}{K_o}\right) \quad (9)$$

where  $P_o$  is the partial pressure of  $O_2$ ;  $K_c$  and  $K_o$  are the Michaelis–Menten constants for  $CO_2$  and  $O_2$  (Keenan et al., 2016):

$$K_c = 39.97 \times e^{\frac{79.43 \times (T_a - 298.15)}{298.15 \times R \times T_a}} \quad (10)$$

$$170 \quad K_o = 27480 \times e^{\frac{36.38 \times (T_a - 298.15)}{298.15 \times R \times T_a}} \quad (11)$$

where  $T_a$  is air temperature (unit: K);  $R$  is the molar gas constant (8.314 J mol<sup>-1</sup> K<sup>-1</sup>).

$T_s$  and  $W_s$  can be expressed as follows:

$$T_s = \frac{(T_a - T_{min}) \times (T_a - T_{max})}{(T_a - T_{min}) \times (T_a - T_{max}) - (T_a - T_{opt}) \times (T_a - T_{opt})} \quad (12)$$

$$W_s = \frac{VPD_o}{VPD_o + VPD} \quad (13)$$

175 where  $T_{min}$ ,  $T_{opt}$ , and  $T_{max}$  are the minimum, optimum, and maximum temperatures for vegetation photosynthesis, respectively (Yuan et al., 2007);  $VPD_o$  is the half-saturation coefficient of the VPD constraint equation (k Pa).

$APAR_{su}$  and  $APAR_{sh}$  can be expressed as follows (Chen et al., 1999):

$$APAR_{su} = \left( PAR_{dir} \times \frac{\cos(\beta)}{\cos(\bar{\theta})} + \frac{PAR_{dif} - PAR_{dif,u}}{LAI} + C \right) \times LAI_{su} \quad (14)$$

$$APAR_{sh} = \left( \frac{PAR_{dif} - PAR_{dif,u}}{LAI} + C \right) \times LAI_{sh} \quad (15)$$

$$180 \quad PAR_{dif,u} = PAR_{dif} \times \exp\left(\frac{-0.5 \times \Omega \times LAI}{\cos(\bar{\theta})}\right) \quad (16)$$

where  $PAR_{dir}$  is the direct PAR;  $PAR_{dif}$  is the diffuse PAR;  $PAR_{dif,u}$  is the diffuse PAR under the canopy;  $C$  represents the multiple scattering effects of direct radiation;  $\Omega$  is the clumping index, which is set according to vegetation types (Tang et al., 2007);  $\theta$  is the solar zenith angle;  $\beta$  is the mean leaf–sun angle, which is set to 60°;  $\bar{\theta}$  is the representative zenith angle for diffuse radiation transmission and can be expressed by LAI (Chen et al., 1999):

$$185 \quad \cos(\bar{\theta}) = 0.537 + 0.025 \times LAI \quad (17)$$

The LAIs of shaded leaves ( $LAI_{sh}$ ) and sunlit leaves ( $LAI_{su}$ ) in Eqs. (14) and (15) are computed following Chen et al (1999):

$$LAI_{su} = 2 \times \cos(\theta) \times \left( 1 - e^{-0.5 \times \Omega \times \frac{LAI}{\cos(\bar{\theta})}} \right) \quad (18)$$

$$LAI_{sh} = LAI - LAI_{su} \quad (19)$$

## 2.4 Model calibration and validation

190 Cross-validation method was used to calibrate and validate the revised EC-LUE model. Fifty percent of the sites were randomly selected to calibrate model parameters for each vegetation type, and the remaining 50% of the sites were used to validate the model. This parameterization process was repeated until all possible combinations of 50% sites were achieved for each vegetation type. The nonlinear regression procedure (Proc NLIN) in the Statistical Analysis System (SAS, SAS Institute Inc., Cary, NC, USA) was applied to optimize the model parameters ( $\epsilon_{msu}$ ,  $\epsilon_{msh}$ ,  $\phi$ , and  $VPD_0$ ) using 8-day estimated GPP based on  
195 EC measurements. The mean GPP simulations of 8-day from all validation runs only were used to model validation. At global scale, mean calibrated parameter values (Table 3) were used to produce GPP dataset at  $0.05^\circ \times 0.05^\circ$  spatial resolution and 8-day temporal resolution over 1982-2017. In order to investigate the uncertainties of the global GPP dataset, 10,000 sets of optimized parameters were randomly selected to simulate global GPP by assuming a normal distribution of these parameters (Table 3). The uncertainty of global GPP simulations was determined by the mean absolute deviation (MAD) of all the 10,000  
200 simulations (Khair et al., 2017).

<<Table 3>>

Three metrics, the coefficient of determination ( $R^2$ ), RMSE, and bias (the difference between observations and simulations) were adopted to evaluate the performance of the revised EC-LUE model. Additionally, Kendall's coefficient of rank correlation  $\tau$  (Kanji, 1999) was used to quantify the agreement of seasonal changes between the simulated and tower estimated GPP. The  
205 Kendall coefficient measured the tendency coherence between predicted and observed GPP by comparing the ranks assigned to successive pairs. If  $GPP_{sim,j} - GPP_{sim,i}$  and  $GPP_{obs,j} - GPP_{obs,i}$  have the same sign (positive or negative), the pair would be concordant, or discordant. A time-series data with  $n$  observations, the Kendall's coefficient of rank correlation  $\tau$  can be expressed:

$$\tau = \frac{C-D}{n(n-1)/2} \quad (20)$$

210 where  $n(n-1)/2$  is the total combination of pairs, C is the number of concordant pairs, and D is the number of discordant pairs. The Kendall's coefficient ranged from -1 (C = 0) to 1 (D = 0). The Kendall's coefficient is much closer to 1, which means a stronger positive relationship between the seasonal patterns of the simulated and tower estimated GPP.

In addition, we compared the model performance of the revised EC-LUE model with seven light use efficiency models, three machine learning methods and ten process-based ecosystem models. The participated light use efficiency models include  
215 CASA (Potter et al., 1993), CFlux (Turner et al., 2006; King et al., 2011), CFix (Veroustraete et al., 2002), MODIS (Running et al., 2004), VPM (Xiao et al., 2005), VPRM (Mahadevan et al., 2008), and EC-LUE (Yuan et al., 2007). We calibrated the model parameters of all seven light use efficiency models based on the eddy covariance measurements using the same parameterization method as the revised EC-LUE model (see the above method), and then compared the GPP simulations of seven LUE models driven by EC tower-based meteorology data against the estimated GPP based on EC measurements. For  
220 the comparison with machine learning models and process-based ecosystem models, we collected their global monthly GPP products released by FLUXCOM (Jung et al., 2017) and TRENDY program (version 5) (Le Quéré et al., 2016), respectively.



FLUXCOM program uses the Artificial Neural Network method (FLUXCOM ANN), the Multivariate Adaptive Regression Splines method (FLUXCOM MARS), and the Random Forest method (FLUXCOM RF); and TRENDY program includes the CSIRO Atmosphere and Biosphere Land Exchange (CABLE) (Zhang et al., 2013), the coupled Canadian Land Surface Scheme and Canadian Terrestrial Ecosystem Model (CLASS-CTEM) (Melton and Arora, 2016), the Community Land Model (CLM) (Oleson et al., 2013), the Integrated Science Assessment Model (ISAM) (Jain et al., 2013), the land component of the Max Planck Institute Earth System Model (JSBACH) (Reick et al., 2013), the Joint UK Land Environment Simulator (JULES) (Clark et al., 2011), the Lund-Postdam-Jena General Ecosystem Simulator (LPJ-GUESS) (Smith et al., 2014), the Land surface Processes and eXchanges (LPX-Bern) (Stocker et al., 2014), the ORganizing Carbon and Hydrology In Dynamic Ecosystems (ORCHIDEE) (Krinner et al., 2005), and the Vegetation Integrated Simulator for Trace Gases (VISIT) (Kato et al., 2013). The monthly GPP simulations at all investigated EC sites were derived from their global products, and equally we obtained the monthly GPP simulations of the revised EC-LUE model from its global dataset driven by MERRA-2 reanalysis dataset.

## 2.5 Environmental contributions to long-term changes in GPP

To evaluate the contribution of the major environmental variables to GPP, including the atmospheric CO<sub>2</sub> concentration ([CO<sub>2</sub>]), climate, and satellite-based LAI, two types of experimental simulations were performed. The first simulation experiment (S<sub>ALL</sub>) was a normal model run, with all the environmental drivers changing over time. In the second type of simulation experiments (S<sub>CLI0</sub>, S<sub>LAI0</sub>, and S<sub>CO20</sub>), two driving factors could be varied with time while maintaining the third constant at an initial baseline level. For example, the S<sub>CLI0</sub> simulation experiment allowed the LAI and atmospheric [CO<sub>2</sub>] to vary with time while the climate variables were kept constant at 1982 values. The S<sub>LAI0</sub> (S<sub>CO20</sub>) simulation experiments kept LAI (atmospheric [CO<sub>2</sub>]) constant at 1982 values and varied the other two variables.

Considering the differences between the simulation results of the first type (S<sub>ALL</sub>) and the second type (S<sub>CO20</sub> and S<sub>LAI0</sub>) of experiments, the GPP sensitivities to atmospheric [CO<sub>2</sub>] ( $\beta_{CO2}$ ) and LAI ( $\beta_{LAI}$ ) were estimated as follows:

$$\Delta GPP_{(S_{ALL}-S_{CO20})i} = \beta_{CO2} \times \Delta CO2_{(S_{ALL}-S_{CO20})i} + \varepsilon \quad (21)$$

$$\Delta GPP_{(S_{ALL}-S_{LAI0})i} = \beta_{LAI} \times \Delta LAI_{(S_{ALL}-S_{LAI0})i} + \varepsilon \quad (22)$$

where  $\Delta GPP_i$ ,  $\Delta CO2_i$ , and  $\Delta LAI_i$  denote the differences in the GPP simulations, atmospheric [CO<sub>2</sub>], and LAI between the two model experiments from 1982 to 2017, and  $\varepsilon$  is the stochastic error term.

The GPP sensitivities to the three climate variables: air temperature ( $\beta_{Ta}$ ), VPD ( $\beta_{VPD}$ ), and PAR ( $\beta_{PAR}$ ) were calculated using a multiple regression approach:

$$\Delta GPP_{(S_{ALL}-S_{CLI0})i} = \beta_{Ta} \times \Delta Ta_{(S_{ALL}-S_{CLI0})i} + \beta_{VPD} \times \Delta VPD_{(S_{ALL}-S_{CLI0})i} + \beta_{PAR} \times \Delta PAR_{(S_{ALL}-S_{CLI0})i} + \varepsilon \quad (23)$$

where  $\Delta Ta_i$ ,  $\Delta VPD_i$ , and  $\Delta PAR_i$  denote the differences in Ta, VPD, and PAR time series between the two model experiments (S<sub>ALL</sub> and S<sub>CLI0</sub>), respectively. The regression coefficient  $\beta$  was estimated using the maximum likelihood analysis.

### 3 Results

#### 3.1 Model performance

255 In general, the revised EC-LUE model could effectively reproduce the spatial, seasonal, and annual variations in the tower estimated GPP at most sites (Figs. 1–3). The revised EC-LUE model explained 71% and 64% of the spatial variations in GPP across all the validation sites by using the tower-derived meteorology data and the meteorological reanalysis dataset, respectively (Fig. 1).

<<Figure 1>>

260 The revised EC-LUE model also showed a good performance in reproducing the seasonal variations in GPP at most EC sites (Fig. 2). In this study, we compared the modeled and tower GPP at 8-day step for each site to examine the model capacity in reproducing the temporal variations of GPP. In terms of GPP simulations driven by tower-derived meteorology data, the coefficients of determination ( $R^2$ ) varied from 0.26 at MY-PSO site to 0.96 at DK-Sor site, with most of them being statistically significant ( $p$ -value  $< 0.05$ ) (Fig. 2a), and the mean  $R^2$  was 0.81 over all investigated sites. The low  $R^2$  values ( $< 0.4$ ) were found  
265 at three tropical forest sites (i.e., MY-PSO, BR-Sa1 and BR-Sa3). The averaged Kendall's correlation coefficient ( $\tau$ ) was 0.63 over all sites, indicating a strong seasonal coherence between simulated and tower-estimated GPP (Fig. 2d). Similarly,  $\tau$  at tropical forest sites were generally lower than other sites. According to the RMSE and absolute bias, the revised EC-LUE model performed very well at most sites. The averaged RMSE and absolute bias over all the sites were 2.13 and 0.81 g C m<sup>-2</sup> d<sup>-1</sup>, respectively (Fig. 2b–c). In addition, there was no obvious difference between the seasonal GPP performances using the  
270 tower-derived meteorology data and the meteorological reanalysis dataset (Fig. 2). On average, the revised EC-LUE model showed higher  $R^2$  and  $\tau$ , and lower bias and RMSE than the original EC-LUE model (Fig. 2). Furthermore, we selected three sites with high  $R^2$  (US-UMB; DBF;  $R^2 = 0.93$ ), median  $R^2$  (CN-Din; EBF;  $R^2 = 0.71$ ), and low  $R^2$  (Br-Sa3; EBF;  $R^2 = 0.39$ ) to illustrate the time-series changes of observed/simulated GPP, LAI, and environmental factors (i.e., air temperature, VPD, and PAR) (Figs S1-S3). At US-UMB site, the model captured the GPP variations well all the year round with no obvious bias  
275 (Fig. S1). At CN-Din site, the model generally performed well except the underestimation at the end of the year (November-December) with decreased LAI (Fig. S2). While, at Br-Sa3 site, the model could not capture the variations of GPP for the vegetation greenness and environmental factors varying slightly during the year (Fig. S3).

<< Figure 2>>

The ability of the LUE models to reproduce the interannual variations in GPP was investigated at 55 EC towers with  
280 observations greater than 5-years (Table 1; Fig. 3). We examined the relations between the mean annual GPP simulations and observations at each site and used the coefficient correlation ( $R^2$ ) and slope of the regression relationship to investigate the model capability in simulating the interannual variations in GPP. The result showed that the revised EC-LUE model could effectively determine the interannual variations in GPP (Fig. 3). Approximately 42% and 40% of the sites showed higher  $R^2$  values ( $> 0.5$ ) by using the tower-derived meteorology data and the meteorological reanalysis dataset (Fig. 3a). The averaged  
285  $R^2$  for the revised EC-LUE model was 0.44 by using the tower-derived meteorology data, which was significantly higher than

the original EC-LUE model ( $R^2 = 0.36$ ) and other LUE models ( $R^2$  ranged from 0.06 to 0.30 with an average value of 0.16) (Fig. 3c). The averaged  $R^2$  for the revised EC-LUE model was 0.42 by using the meteorological reanalysis dataset. The averaged slopes of the revised EC-LUE model were 0.60 and 0.57 by using the tower-derived meteorology data and the meteorological reanalysis dataset (Fig. 3c).

290 <<Figure 3>>

Additionally, we examined the model performance of the revised EC-LUE model ~~at monthly step by comparing with~~ other LUE models, machine learning models, and process-based models in TRENDY ~~(Fig. 4)~~ at monthly step by comparing against EC tower estimated GPP (Fig. 4). In comparison with seven LUE models driven by EC tower-based meteorology dataset, the overall  $R^2$  of the revised EC-LUE model was 0.7671, higher than the original EC-LUE model and other LUE models ( $R^2$  ranged from 0.6555 to 0.7161) (Fig. 4a). For each site, we compared the  $R^2$ /RMSE/absolute value of bias of the individual model with the averaged value of all the eight LUE models (each site has an averaged  $R^2$ /RMSE/absolute value of bias) (Fig. S4a1-c1). The revised EC-LUE model had higher  $R^2$  than the mean  $R^2$  of the eight LUE models at 62% sites, which was comparable with the original EC-LUE model (63% sites) and VPM model (60% sites) (Fig. S4a1). Moreover, the revised EC-LUE model showed the lower RMSE and bias compared to mean values of all eight LUE models at 68% and 67% sites respectively, which indicated the better performance compared to the other LUE models at most sites (Fig. S4b1-c1). By using the global reanalysis meteorology data, we compared the performance of the revised EC-LUE model with three existing machine learning model products and ten process-based model products in TRENDY (Fig. 4b). The overall  $R^2$  of the revised EC-LUE model ( $R^2 = 0.6657$ ) was higher than that of other models ( $R^2$  ranged from 0.0602 to 0.6154) (Fig. 4b). The revised EC-LUE model, FLUXCOM ANN, and FLUXCOM MARS had more sites (over 90%) with higher  $R^2$  than the mean  $R^2$  (Fig. S4a2). And the revised EC-LUE model, FLUXCOM MARS, and FLUXCOM RF showed the lower RMSE at more than 90% sites (Fig. S4b2). Compared to the other models, the revised EC-LUE model had highest site percentage (81%) with lower absolute value of bias (Fig. S4c2).

300 <<Figure 4>>

### 3.2 Spatial-temporal patterns of global GPP

310 A global GPP dataset at  $0.05^\circ$  latitude by  $0.05^\circ$  longitude and 8-day interval was generated from 1982 to 2017 based on the revised EC-LUE model. The global GPP was  $106.2 \pm 2.9$  Pg C yr<sup>-1</sup> across the vegetated area averaged from 1982 to 2017. The GPP was high over the tropical forest areas, such as Amazon and Southeast Asia, where the moisture and temperature conditions are sufficient for photosynthesis (Fig. 5a). The GPP decreased with the decreasing gradients of temperature and precipitation (Fig. 5b). The moderate GPP was found in temperate and subhumid regions; and the lowest GPP was located in arid or cold regions, where either precipitation or temperature is limited (Fig. 5b).

315 <<Figure 5>>

Long-term trend of GPP over the period of 1982–2017 was determined using a linear regression analysis (Fig. 6). In general, the revised EC-LUE model showed an increased trend in the annual mean GPP from 1982 to 2017. Approximately 69.5% of

the vegetated areas, mainly located in temperate and humid regions, showed increased trends. The spatial pattern of the GPP trend along with the temperature and precipitation gradients was substantially heterogeneous (Fig. 6b). The decreased GPP was found in the tropic regions, especially in the Amazon forest (Fig. 6a). The extremely cold or arid areas exhibited less variations in GPP (Fig. 6b).

<<Figure 6>>

In addition, this study used the MAD of 10,000 simulations to quantify the uncertainty of estimated GPP globally (see methods). Over the globe, the mean uncertainty of estimated GPP by the revised EC-LUE model is 19.33 Pg C yr<sup>-1</sup>. The GPP uncertainties were found to be low in high and middle latitudes, but relatively high in tropical forests (about 600 g C m<sup>-2</sup> yr<sup>-1</sup>) (Fig. 7).

<<Figure 7>>

### 3.3 Contributions of environmental variables to GPP

To quantify the contributions of the environmental variables to long-term changes in GPP, we explored the sensitivity of global summed GPP to climate variables (i.e., VPD, Ta, and PAR), LAI, and atmospheric CO<sub>2</sub> (Fig. 8). The global summed GPP generated from different experimental simulations (section 2.5) exhibited differently in terms of the annual mean value, trend, and standard deviation (Fig. 8a). The normal simulated GPP ( $S_{ALL}$  GPP, all the environmental drivers changing over time) significantly increased at the rate of 0.15 Pg C yr<sup>-1</sup>, while the increasing rate of  $S_{CLI0}$  GPP (climate variables were kept constant at 1982 values) was even greater (0.41 Pg C yr<sup>-1</sup>). On the contrary, the  $S_{LAI0}$  GPP (LAI was kept constant at 1982 values) and the  $S_{CO20}$  GPP (atmospheric [CO<sub>2</sub>] was kept constant at 1982 values) showed insignificantly decreasing trend at the rate of -0.04 Pg C yr<sup>-1</sup> and -0.07 Pg C yr<sup>-1</sup> (Fig. 8a). The GPP sensitivity analysis showed that the global GPP decreased by 6.67 ± 5.04 Pg C with a 0.1 kPa increase in VPD, which was comparable to the increase in GPP with 0.1 unit greening of LAI (i.e.,  $\beta_{LAI} = 4.78 \pm 0.72$  Pg C 0.1 unit<sup>-1</sup>) or 100 MJ increase in PAR (i.e.,  $\beta_{PAR} = 5.73 \pm 3.22$  Pg C 100 MJ<sup>-1</sup>) (Fig. 8b). The global GPP increased by 12.31 ± 0.61 Pg C with a 100 ppm<sup>-1</sup> rise of atmospheric [CO<sub>2</sub>] (i.e.,  $\beta_{CO2} = 12.31 \pm 0.61$  Pg C 100 ppm<sup>-1</sup>). Over the period of 1982–2017, the increased VPD resulted in the global GPP decreases of -0.17 ± 0.06 Pg C yr<sup>-1</sup>, which could partly counteract the fertilization effect of CO<sub>2</sub> (0.22 ± 0.07 Pg C yr<sup>-1</sup>). The global GPP showed a decreased trend after 2001 due to the joint effect of increased VPD and decreased PAR (Fig. 8c). While the increased trend of GPP before 2000 was affected by the rising atmospheric [CO<sub>2</sub>], greening of LAI, and increased PAR (Fig. 8c).

<<Figure 8>>

## 345 4 Discussion

### 4.1 Model accuracy analysis

Numerous studies have shown that most GPP models can reproduce the spatial changes in GPP but failed to reproduce the temporal variations (Keenan et al., 2012; Yuan et al., 2014). Therefore, the capacity to reproduce realistic interannual variations for a GPP model is significantly important. In our study, the revised EC-LUE model performed a higher accuracy in

350 reproducing the interannual variations in GPP than did the original EC-LUE model and other LUE models. Yuan et al. (2014) reported that the averaged slope of the regression relation between the mean annual GPP simulated by seven LUE models and the mean annual GPP estimated from EC tower ranged from 0.19 to 0.56 (Fig. 3c). While the revised EC-LUE model showed a higher slope of regression relation (0.60), which is much closer to 1 than that obtained from other LUE models (Fig. 3c). The VPM GPP showed less interannual variations across most biomes ( $R^2 < 0.5$ ), probably because of the insensitivity of the environmental stress factors at the interannual scale (Zhang et al., 2017). In contrast, 42% of the sites showed higher  $R^2$  values (>0.5) for the revised EC-LUE model. The improvements of the revised EC-LUE model in reproducing interannual variations are owing to the integration of several important environmental drivers for vegetation production (i.e., atmospheric  $\text{CO}_2$  concentration, radiation components, and VPD), which exhibited large variations and contributed significantly to vegetation production at interannual scale.

360 By integrating the atmospheric  $\text{CO}_2$  concentration, the revised EC-LUE model suggested a  $\text{CO}_2$  sensitivity ( $\beta_{\text{CO}_2}$ ) of  $12.31 \pm 0.61$  Pg C per 100 ppm (Fig. 8b), which indicates an increase of 11.6% in GPP with a rise of 100 ppm in atmospheric  $[\text{CO}_2]$ . Our estimate is comparable to the observed response of NPP to the increased  $\text{CO}_2$  in the FACE experiments (13% per 100 ppm) and estimates of other ecosystem models (5–20% per 100 ppm) (Piao et al., 2013). The elevated atmospheric  $\text{CO}_2$  concentration substantially contributes to vegetation productivity.

365 The evaporation fraction (EF), namely the ratio of evapotranspiration (ET) to net radiation (Rn), was used to indicate the water stress on vegetation growth in the original EC-LUE model (Yuan et al., 2007; 2010). While the atmospheric VPD was used to indicate water stress to avoid the aggregated errors from ET simulations in the revised EC-LUE model. Physiologically, vegetation production is sensitive to both atmospheric VPD and soil moisture availability to roots. [Several studies have reported highly consistent interannual variability of VPD and soil moisture \(Zhou et al., 2019a, b\). In addition,](#) recent studies highlighted that the increase in VPD had a larger limitation to the surface conductance and evapotranspiration than soil moisture over short time scales in many biomes (Novick et al., 2016; Sulman et al., 2016). Other studies have also suggested substantial impacts of VPD on vegetation growth (de Cárcer et al., 2018; Ding et al., 2018), forest mortality (Williams et al., 2013), and crop yields (Lobell et al., 2014). It is increasingly important to integrate the atmospheric water constraint to the carbon and water flux modeling.

## 375 **4.2 Comparison of global GPP products**

Global and regional GPP estimates remain highly uncertain despite the substantial advances in remote sensing technology, ground observations, and theory of carbon flux modeling (Zheng et al., 2018; Ryu et al., 2019). At regional scale, we compared the annual mean GPP between the revised EC-LUE model and other models across the bioclimatic zones in the Köppen-Geiger climate classification map (Beck et al., 2018) (Fig. 9). The GPP of the revised EC-LUE model was comparable to the mean value of other models for each bioclimatic zone (Fig. 9a). The GPP of different models exhibited large discrepancies in tropical regions (Af/Am/Aw) (Fig. 9a). The correlations ( $R^2$ ) of GPP across all the bioclimatic zones between the revised EC-LUE model and other models ranged from 0.73 (LPX-Bern) to 0.95 (FLUXCOM MARS, FLUXCOM RF) (Fig. 9b).

<<Figure 9>>

At global scale, our study showed large differences in the magnitude of global GPP estimated by various models varying from  
385 92.7 to 168.7 Pg C yr<sup>-1</sup> (Figs. 10–11). The LUE models simulated the global GPP ranging from 92.7 to 133.7 Pg C yr<sup>-1</sup> (Fig.  
11a1). Several machine learning approaches estimated the global GPP ranging from 111.0 to 144.2 Pg C yr<sup>-1</sup> (Fig. 11a2). A  
comparison of ten global terrestrial ecosystem models of TRENDY showed that the global GPP ranged from 107.8 to 154.9  
Pg C yr<sup>-1</sup> (Fig. 11a3). The revised EC-LUE model quantified the mean global GPP from 1982 to 2017 as 106.2 ± 2.9 Pg C  
390 yr<sup>-1</sup>. Other studies also support the conclusion that there are large uncertainties in the GPP estimates. By comparing diverse  
GPP models and products, Anav et al. (2015) reported that the global GPP ranged from 112 to 169 Pg C yr<sup>-1</sup>. Seven satellite-  
based LUE models estimated the global GPP ranged from 95.1 to 139.7 Pg C yr<sup>-1</sup> over the period of 2000–2010 (Cai et al.,  
2014).

<<Figure 10>>

The interannual variability and trend in GPP also vary substantially with different models. This study showed that the  
395 interannual variability (standard deviation) ranged from 0.32 to 5.89 Pg C yr<sup>-1</sup>, with the trends varying from -0.05 to 0.84 Pg  
C yr<sup>-1</sup> (Fig. 11). The biophysical models showed large interannual variability, with the standard deviation ranging from 1.38  
to 5.89 Pg C yr<sup>-1</sup>. The LUE models estimated the interannual variability varied from 1.30 to 3.13 Pg C yr<sup>-1</sup>. In contrast, the  
machine learning models exhibited less interannual variability with standard deviation under 1.0 Pg C yr<sup>-1</sup>. The interannual  
variability of the revised EC-LUE model was 2.9 Pg C yr<sup>-1</sup> (Figs. 11b1–b3). In general, the GPP interannual variability before  
400 the year 2000 year was greater than that after the year 2001 for most of the biophysical models and LUE models (Figs. 11b1–  
b3). Most GPP models showed an increased trend or insignificant trend during all valid years and before 2000. Similar to the  
standard deviation, the trends of machine learning models were less than other models. Compared with the other models,  
CLASS-CTEM and the revised EC-LUE model showed a significant decreasing trend after 2001 (Figs. 11c1–c3), probably  
because of the joint effect of increased VPD and decreased PAR (Fig. 8c).

405 <<Figure 11>>

### 4.3 Model uncertainty

The uncertainties of our GPP dataset were low in high and middle latitude areas but high in tropical areas (Fig. 7). This is  
consistent with the validations at site level that the revised EC-LUE model showed the lowest accuracy over the tropical  
evergreen broadleaf forest sites (Fig. 2). Similarly, other satellite-based models exhibited a large uncertainty in the GPP  
410 simulations over tropical forest areas (Ryu et al., 2011; Yuan et al., 2014). For example, MODIS GPP product (MOD17)  
underestimated GPP at high productivity sites over the tropical evergreen forests (de Almeida et al., 2018). Regarding the  
quality of satellite data, a high cloud cover exists over tropical regions, introducing large uncertainties to FAPAR/LAI and  
other vegetation indices (e.g., NDVI and EVI). As suggested by de Almeida et al. (2018), the lack of reliable MOD15 FAPAR  
data during January to April as a result of the cloudiness contamination could have substantially affected the seasonality of  
415 GPP estimates. Besides, the quality of satellite data can even affect the evaluation of the interannual variations in GPP. Using

MODIS EVI data, Saleska et al. (2007) reported a large-scale green up in the Amazon evergreen forests during the drought of 2005. However, an opposite conclusion was drawn when the cloud-contaminated data were excluded from the analysis (Samanta et al., 2010). ~~Subsequent studies noted an increase in LAI and EVI during the dry season in the Amazon evergreen forests; however, a recent study highlighted that the apparent seasonal changes in EVI were resulted from the variations in the sun sensor geometry rather than vegetation greenness (Morton et al., 2014).~~ In our study, a significant decrease of GPP was found in the Amazon evergreen forests, which may be resulted from the sharp increase in VPD after the late 1990s (Yuan et al., 2019). Studies using optical satellite data can be influenced by the cloudiness contamination. Recently studies using cloud free satellite-based microwave data also reported a carbon loss in tropic forest (Liu et al., 2015; Fan et al., 2019).

The latest study highlighted that the aggregate canopy phenology rather than the climate changes is the main cause of the seasonal changes in photosynthesis in evergreen broadleaf forests (Wu et al., 2016). In particular, the new leaf growing synchronously with dry season litterfall may shift the old canopy to be younger, which can explain the significant seasonal increase (~27%) in the ecosystem photosynthesis. Therefore, the vertical changes in leaf age and photosynthesis ability with canopy depth are important to simulate the seasonal variations in carbon flux in tropical forests (Wu et al., 2017). These leaf trait related parameters can be simulated from the narrow-band spectra of leaves (Serbin et al., 2012; Dechant et al., 2017). Nevertheless, because of the limitation in obtaining the large scale hyperspectral remote sensing data, regional or global estimation of these parameters are currently unavailable.

The revised EC-LUE model does not integrate the regulation of soil nitrogen content on vegetation production. Atmospheric nitrogen deposition has exhibited a large increasing trend in the past few decades because of the excessive fossil fuel combustion in the industrial and transportation sectors and the abuse of nitrogenous fertilizer in the agricultural practice (Galloway et al., 2004). And the global land atmospheric nitrogen deposition is expected to further increase dramatically from 25–40 Tg N yr<sup>-1</sup> in the 2000s to 60–100 Tg N yr<sup>-1</sup> in 2100 (Lamarque et al., 2005). A meta-analysis of worldwide nitrogen addition experiments found that nitrogen addition could have a significantly positive effect on vegetation productivity (Liu and Greaver, 2009). As most terrestrial ecosystems are nitrogen limited, quantifying the spatio-temporal distributions of vegetation nitrogen content at large scales is essential to improve the accuracy of carbon flux estimation. Several studies quantified the leaf nitrogen content by detecting the nitrogen absorption spectra from the narrow-band of hyperspectral data (Cho, 2007). However, leaf water, starch, lignin, and cellulose overlap with the absorption characters of nitrogen in the shortwave infrared bands, making it difficult to retrieve the nitrogen content (Kokaly and Clark, 1999). Additionally, canopy structures, background, and illumination/viewing geometry can further decrease the capacity to detect leaf nitrogen (Yoder and Pettigrew-Crosby, 1995; Knyazikhin et al., 2013). Advances in inversion and statistical models of leaf or canopy nitrogen have emerged (Asner et al., 2011; Dechant et al., 2017; Wang et al., 2018), but these methods require further evaluation over large regions and the global map of leaf or canopy nitrogen is not available yet.

Additionally, the uncertainty of the revised EC-LUE model may arise by scale mismatches between eddy covariance flux footprint and input dataset. The eddy covariance flux footprint is generally less than 3 km<sup>2</sup> and varies depending on the wind speed, wind direction and atmospheric stability (Tan et al., 2006). In our studies, the revised EC-LUE model was run at 0.05

450 degree (~5 km<sup>2</sup>) spatial resolution. The uncertainty of simulated GPP introduced by the scale effect is inevitable but smaller than that introduced by the model structures, parameters or input datasets (Sjostrom et al., 2013; Zheng et al., 2018).

## 5 Data availability

The 0.05° × 0.05° global GPP dataset for 1982-2017 is available at <https://doi.org/10.6084/m9.figshare.8942336> (Zheng et al., 2019). The dataset is provided in hdf format at 8-day interval. The valid value is ranged from 0 to 3000, and the background filled value is 65535. The scale factor of the data is 0.01. Each hdf file represents an 8-day GPP at daily value (unit: g C m<sup>-2</sup> day<sup>-1</sup>). To obtain the summation of each 8-day (or 5-day or 6-day) period, please multiply the GPP value by corresponding days (8 for the first 45 values, and 5 or 6 for the last value in a year).

## 6 Conclusion

In this study, we produced a long-term global GPP dataset by integrating several major long-term environmental variables into a light use efficiency model, including atmospheric CO<sub>2</sub> concentration, radiation components, and atmospheric water vapor pressure. These environmental variables showed substantial long-term changes and contributed significantly to vegetation production at interannual scale. The revised EC-LUE performed well in simulating the spatial, seasonal, and interannual variations in GPP across the globe. Particularly, it has a unique superiority in reproducing the interannual variations in GPP (R<sup>2</sup> = 0.44) compared with the original EC-LUE model (R<sup>2</sup> = 0.36) and other LUE models (R<sup>2</sup> ranged from 0.06 to 0.30 with an average value of 0.16). The GPP dataset derived from the revised EC-LUE model provides an alternative and reliable estimates of global GPP at the long-term scale by integrating the important environmental variables.

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**Competing interests.** The authors declare that they have no conflict of interest.

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## 480 **References**

- Ainsworth, E.A., Long, S.P.: What have we learned from 15 years of free-air CO<sub>2</sub> enrichment (FACE)? A meta-analytic review of the responses of photosynthesis, canopy, *New Phytol.*, 165, 351-371, doi:10.1111/j.1469-8137.2004.01224.x, 2005.
- Alton, P.B., North, P.R., Los, S.O.: The impact of diffuse sunlight on canopy light-use efficiency, gross photosynthetic product and net ecosystem exchange in three forest biomes, *Global Change Biol.*, 13, 776-787, doi:10.1111/j.1365-2486.2007.01316.x, 485 2007.
- Anav, A., Friedlingstein, P., Beer, C., Ciais, P., Harper, A., Jones, C., Murray-Tortarolo, G., Papale, D., Parazoo, N.C., Peylin, P., Piao, S., Sitch, S., Viovy, N., Wiltshire, A., Zhao, M.: Spatiotemporal patterns of terrestrial gross primary production: A review, *Rev. Geophys.*, 53, 785-818, doi:10.1002/2015rg000483, 2015.
- Asner, G.P., Martin, R.E., Knapp, D.E., Tupayachi, R., Anderson, C., Carranza, L., Martinez, P., Houcheime, M., Sinca, F., 490 Weiss, P.: Spectroscopy of canopy chemicals in humid tropical forests, *Remote Sens. Environ.*, 115, 3587-3598, doi:10.1016/j.rse.2011.08.020, 2011.
- Beck, H.E., Zimmermann, N.E., McVicar, T.R., Vergopolan, N., Berg, A., Wood, E.F.: Present and future Koppen-Geiger climate classification maps at 1-km resolution, *Scientific Data*, doi:510.1038/sdata.2018.214, 2018.
- Cai, W., Yuan, W., Liang, S., Zhang, X., Dong, W., Xia, J., Fu, Y., Chen, Y., Liu, D., Zhang, Q.: Improved estimations of 495 gross primary production using satellite-derived photosynthetically active radiation, *J. Geophys. Res. G: Biogeosci.*, 119, 110-123, doi:10.1002/2013jg002456, 2014.
- Cai, W., Yuan, W., Liang, S., Liu, S., Dong, W., Chen, Y., Liu, D., Zhang, H.: Large Differences in Terrestrial Vegetation Production Derived from Satellite-Based Light Use Efficiency Models, *Remote Sens.*, 6, 8945-8965, doi:10.3390/rs6098945, 2014.
- 500 Canadell, J.G., Le Quere, C., Raupach, M.R., Field, C.B., Buitenhuis, E.T., Ciais, P., Conway, T.J., Gillett, N.P., Houghton, R.A., Marland, G.: Contributions to accelerating atmospheric CO<sub>2</sub> growth from economic activity, carbon intensity, and efficiency of natural sinks, *Proc. Natl. Acad. Sci. U.S.A.*, 104, 18866-18870, doi:10.1073/pnas.0702737104, 2007.
- Chen, J.M., Liu, J., Cihlar, J., Goulden, M.L.: Daily canopy photosynthesis model through temporal and spatial scaling for remote sensing applications, *Ecol. Modell.*, 124, 99-119, doi:10.1016/s0304-3800(99)00156-8, 1999.
- 505 Cho, M.A., Skidmore, A., Corsi, F., van Wieren, S.E., Sobhan, I.: Estimation of green grass/herb biomass from airborne hyperspectral imagery using spectral indices and partial least squares regression, *International Journal of Applied Earth Observation and Geoinformation*, 9, 414-424, doi:10.1016/j.jag.2007.02.001, 2007.

- Clark, D. B., Mercado, L. M., Sitch, S., Jones, C. D., Gedney, N., Best, M. J., Pryor, M., Rooney, G. G., Essery, R. L. H., Blyth, E., Boucher, O., Harding, R. J., Huntingford, C., and Cox, P. M.: The Joint UK Land Environment Simulator (JULES),  
510 model description – Part 2: Carbon fluxes and vegetation dynamics, *Geosci. Model Dev.*, 4, 701–722, doi:10.5194/gmd-4-701-2011, 2011.
- Collatz, G.J., Ball, J.T., Grivet, C., Berry, J.A.: Physiological and environmental regulation of stomatal conductance, photosynthesis and transpiration: a model that includes a laminar boundary layer, *Agric. For. Meteorol.*, 54, 107-136, 1991.
- de Almeida, C.T., Delgado, R.C., Galvao, L.S., de Oliveira Cruz e Aragao, L.E., Concepcion Ramos, M.: Improvements of  
515 the MODIS Gross Primary Productivity model based on a comprehensive uncertainty assessment over the Brazilian Amazonia, *ISPRS J. Photogramm. Remote Sens.*, 145, 268-283, doi:10.1016/j.isprsjprs.2018.07.016, 2018.
- de Cárcer, P.S., Vitasse, Y., Peñuelas, J., Jasey, V.E.J., Buttler, A., Signarbieux, C.: Vapor-pressure deficit and extreme climatic variables limit tree growth, *Global Change Biol.*, 24, 1108-1122, doi:10.1111/gcb.13973, 2018.
- Dechant, B., Cuntz, M., Vohland, M., Schulz, E., Doktor, D.: Estimation of photosynthesis traits from leaf reflectance spectra:  
520 Correlation to nitrogen content as the dominant mechanism, *Remote Sens. Environ.*, 196, 279-292, doi:10.1016/j.rse.2017.05.019, 2017.
- Ding, J., Yang, T., Zhao, Y., Liu, D., Wang, X., Yao, Y., Peng, S., Wang, T., Piao, S.: Increasingly Important Role of Atmospheric Aridity on Tibetan Alpine Grasslands, *Geophys. Res. Lett.*, 45, 2852-2859, doi:10.1002/2017gl076803, 2018.
- [Fan, L., Wigneron, J.-P., Ciais, P., Chave, J., Brandt, M., Fensholt, R., Saatchi, S. S., Bastos, A., Al-Yaari, A., Hufkens, K.,  
525 Qin, Y., Xiao, X., Chen, C., Myneni, R. B., Fernandez-Moran, R., Mialon, A., Rodriguez-Fernandez, N. J., Kerr, Y., Tian, F., and Penuelas, J.: Satellite-observed pantropical carbon dynamics, \*Nature Plants\*, 5, 944-951, 2019.](#)
- Farquhar, G.D., von Caemmerer, S., Berry, J.A.: A biochemical model of photosynthetic CO<sub>2</sub> assimilation in leaves of C<sub>3</sub> species, *Planta*, 149, 78-90, doi:10.1007/bf00386231, 1980.
- Fletcher, A.L., Sinclair, T.R., Allen, L.H.: Transpiration responses to vapor pressure deficit in well watered 'slow-wilting' and  
530 commercial soybean, *Environ. Exp. Bot.*, 61, 145-151, doi:10.1016/j.envexpbot.2007.05.004, 2007.
- Galloway, J.N., Dentener, F.J., Capone, D.G., Boyer, E.W., Howarth, R.W., Seitzinger, S.P., Asner, G.P., Cleveland, C.C., Green, P.A., Holland, E.A., Karl, D.M., Michaels, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J.: Nitrogen cycles: past, present, and future, *Biogeochemistry*, 70, 153-226, doi:10.1007/s10533-004-0370-0, 2004.
- Gilgen, H., Wild, M., Ohmura, A.: Means and trends of shortwave irradiance at the surface estimated from Global Energy  
535 Balance Archive data, *J. Clim.*, 11, 2042-2061, doi:10.1175/1520-0442-11.8.2042, 1998.
- Gu, L.H., Baldocchi, D., Verma, S.B., Black, T.A., Vesala, T., Falge, E.M., Dowty, P.R.: Advantages of diffuse radiation for terrestrial ecosystem productivity, *J. Geophys. Res. D: Atmos.*, 10710.1029/2001jd001242, 2002.
- Kanniah, K.D., Beringer, J., North, P., Hutley, L.: Control of atmospheric particles on diffuse radiation and terrestrial plant productivity: A review, *Progress in Physical Geography-Earth and Environment*, 36, 209-237,  
540 doi:10.1177/0309133311434244, 2012.

- Ju, W., Chen, J.M., Black, T.A., Barr, A.G., Liu, J., Chen, B.: Modelling multi-year coupled carbon and water fluxes in a boreal aspen forest, *Agric. For. Meteorol.*, 140, 136-151, doi:10.1016/j.agrformet.2006.08.008, 2006.
- Jain, A. K., Meiyappan, P., Song, Y., and House, J. I.: CO<sub>2</sub> Emissions from Land-Use Change Affected More by Nitrogen Cycle, than by the Choice of Land Cover Data, *Glob. Change Biol.*, 9, 2893–2906, 2013.
- 545 Jung, M., Reichstein, M., Schwalm, C.R., Huntingford, C., Sitch, S., Ahlstrom, A., Arneeth, A., Camps-Valls, G., Ciais, P., Friedlingstein, P., Gans, F., Ichii, K., Ain, A.K.J., Kato, E., Papale, D., Poulter, B., Raduly, B., Rodenbeck, C., Tramontana, G., Viovy, N., Wang, Y.-P., Weber, U., Zaehle, S., Zeng, N.: Compensatory water effects link yearly global land CO<sub>2</sub> sink changes to temperature, *Nature*, 541, 516-520, doi:10.1038/nature20780, 2017.
- Kanji, G.K., 1999. 100 Statistical Tests. SAGE Publications, London.
- 550 Kato, E., Kinoshita, T., Ito, A., Kawamiya, M., and Yamagata, Y.: Evaluation of spatially explicit emission scenario of land-use change and biomass burning using a process-based biogeochemical model, *Journal of Land Use Science*, 8, 104–122, 2013.
- Keenan, T.F., Baker, I., Barr, A., Ciais, P., Davis, K., Dietze, M., Dragon, D., Gough, C.M., Grant, R., Hollinger, D., Hufkens, K., Poulter, B., McCaughey, H., Raczka, B., Ryu, Y., Schaefer, K., Tian, H., Verbeeck, H., Zhao, M., Richardson, A.D.: Terrestrial biosphere model performance for inter-annual variability of land-atmosphere CO<sub>2</sub> exchange, *Global Change Biol.*,
- 555 18, 1971-1987, doi:10.1111/j.1365-2486.2012.02678.x, 2012.
- Keenan, T.F., Prentice, I.C., Canadell, J.G., Williams, C.A., Wang, H., Raupach, M., Collatz, G.J.: Recent pause in the growth rate of atmospheric CO<sub>2</sub> due to enhanced terrestrial carbon uptake, *Nat. Commun.*, 7, 1038/ncomms13428, 2016.
- Khair U., Fahmi H., Al Hakim S., Rahim R.: Forecasting Error Calculation with Mean Absolute Deviation and Mean Absolute Percentage Error. In *Journal of Physics: Conference Series*, vol. 930, No. 1, p. 012002. IOP Publishing, 2017.
- 560 King, D.A., Turner, D.P., Ritts, W.D.: Parameterization of a diagnostic carbon cycle model for continental scale application, *Remote Sens. Environ.*, 1157, 1653–1664, 2011.
- Knyazikhin, Y., Schull, M.A., Stenberg, P., Mottus, M., Rautiainen, M., Yang, Y., Marshak, A., Latorre Carmona, P., Kaufmann, R.K., Lewis, P., Disney, M.I., Vanderbilt, V., Davis, A.B., Baret, F., Jacquemoud, S., Lyapustin, A., Myneni, R.B.: Hyperspectral remote sensing of foliar nitrogen content, *Proc. Natl. Acad. Sci. U.S.A.*, 110, E185-E192,
- 565 doi:10.1073/pnas.1210196109, 2013.
- Kokaly, R.F., Clark, R.N.: Spectroscopic determination of leaf biochemistry using band-depth analysis of absorption features and stepwise multiple linear regression, *Remote Sens. Environ.*, 67, 267-287, doi:10.1016/s0034-4257(98)00084-4, 1999.
- Konings, A.G., Williams, A.P., Gentine, P.: Sensitivity of grassland productivity to aridity controlled by stomatal and xylem regulation, *Nat. Geosci.*, 10, 284-+, doi:10.1038/ngeo2903, 2017.
- 570 Korson, L., Drost-Hansen, W., Millero, F.J.: Viscosity of water at various temperatures, *J. Phys. Chem.*, 73, 34-39, doi:10.1021/j100721a006, 1969.
- Krinner, G., Viovy, N., de Noblet, N., Ogée, J., Friedlingstein, P., Ciais, P., Sitch, S., Polcher, J., and Prentice, I. C.: A dynamic global vegetation model for studies of the coupled atmosphere-biosphere system, *Global Biogeochem. Cy.*, 19, 1–33, 2005.

- Krupkova, L., Markova, I., Havrankova, K., Pokorny, R., Urban, O., Sigut, L., Pavelka, M., Cienciala, E., Marek, M.V.:  
575 Comparison of different approaches of radiation use efficiency of biomass formation estimation in Mountain Norway spruce,  
Trees-Structure and Function, 31, 325-337, doi:10.1007/s00468-016-1486-2, 2017.
- Lamarque, J.F., Kiehl, J.T., Brasseur, G.P., Butler, T., Cameron-Smith, P., Collins, W.D., Collins, W.J., Granier, C.,  
Hauglustaine, D., Hess, P.G., Holland, E.A., Horowitz, L., Lawrence, M.G., McKenna, D., Merilees, P., Prather, M.J., Rasch,  
P.J., Rotman, D., Shindell, D., Thornton, P.: Assessing future nitrogen deposition and carbon cycle feedback using a  
580 multimodel approach: Analysis of nitrogen deposition, J. Geophys. Res. D: Atmos., 11010.1029/2005jd005825, 2005.
- Le Quéré, C., Andrew, R. M., Canadell, J. G., Sitch, S., Korsbakken, J. I., Peters, G. P., Manning, A. C., Boden, T. A., Tans,  
P. P., Houghton, R. A., Keeling, R. F., Alin, S., Andrews, O. D., Anthoni, P., Barbero, L., Bopp, L., Chevallier, F., Chini, L.  
P., Ciais, P., Currie, K., Delire, C., Doney, S. C., Friedlingstein, P., Gkritzalis, T., Harris, I., Hauck, J., Haverd, V., Hoppema,  
M., Klein Goldewijk, K., Jain, A. K., Kato, E., Körtzinger, A., Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S.,  
585 Lombardozzi, D., Melton, J. R., Metz, N., Millero, F., Monteiro, P. M. S., Munro, D. R., Nabel, J. E. M. S., Nakaoka, S.,  
O'Brien, K., Olsen, A., Omar, A. M., Ono, T., Pierrot, D., Poulter, B., Rödenbeck, C., Salisbury, J., Schuster, U., Schwinger,  
J., Séférian, R., Skjelvan, I., Stocker, B. D., Sutton, A. J., Takahashi, T., Tian, H., Tilbrook, B., van der Laan-Luijkx, I. T., van  
der Werf, G. R., Viovy, N., Walker, A. P., Wiltshire, A. J., and Zaehle, S.: Global Carbon Budget 2016, Earth Syst. Sci. Data,  
8, 605–649, <https://doi.org/10.5194/essd-8-605-2016>, 2016.
- 590 Li, X.L., Liang, S.L., Yu, G.R., Yuan, W.P., Cheng, X., Xia, J.Z., Zhao, T.B., Feng, J.M., Ma, Z.G., Ma, M.G., Liu, S.M.,  
Chen, J.Q., Shao, C.L., Li, S.G., Zhang, X.D., Zhang, Z.Q., Chen, S.P., Ohta, T., Varlagin, A., Miyata, A., Takagi, K., Saiqusa,  
N., Kato, T.: Estimation of gross primary production over the terrestrial ecosystems in China, Ecol. Modell., 261, 80-92,  
doi:10.1016/j.ecolmodel.2013.03.024, 2013.
- Liu, L., Greaver, T.L.: A review of nitrogen enrichment effects on three biogenic GHGs: the CO<sub>2</sub> sink may be largely offset  
595 by stimulated N<sub>2</sub>O and CH<sub>4</sub> emission, Ecol. Lett., 12, 1103-1117, doi:10.1111/j.1461-0248.2009.01351.x, 2009.
- Liu, S., Bond-Lamberty, B., Boysen, L.R., Ford, J.D., Fox, A., Gallo, K., Hatfield, J., Henebry, G.M., Huntington, T.G., Liu,  
Z., Loveland, T.R., Norby, R.J., Sohl, T., Steiner, A.L., Yuan, W., Zhang, Z., Zhao, S.: Grand Challenges in Understanding  
the Interplay of Climate and Land Changes, Earth Interactions, 21, 1-43, doi:10.1175/ei-d-16-0012.1, 2017.
- Liu, Y., Xiao, J., Ju, W., Zhu, G., Wu, X., Fan, W., Li, D., Zhou, Y.: Satellite-derived LAI products exhibit large discrepancies  
600 and can lead to substantial uncertainty in simulated carbon and water fluxes, Remote Sens. Environ., 206, 174-188,  
doi:10.1016/j.rse.2017.12.024, 2018.
- [Liu, Y. Y., van Dijk, A. I. J. M., de Jeu, R. A. M., Canadell, J. G., McCabe, M. F., Evans, J. P., and Wang, G.: Recent reversal  
in loss of global terrestrial biomass, Nat. Clim. Change, 5, 470-474, 2015.](#)
- Lobell, D.B., Roberts, M.J., Schlenker, W., Braun, N., Little, B.B., Rejesus, R.M., Hammer, G.L.: Greater Sensitivity to  
605 Drought Accompanies Maize Yield Increase in the US Midwest, Science, 344, 516-519, doi:10.1126/science.1251423, 2014.
- Monteith, J.: Solar radiation and productivity in tropical ecosystems, J. Appl. Ecol., 9, 747-766, 1972.

- Mahadevan, P., Wofsy, S.C., Matross, D.M., Xiao, X.M., Dunn, A.L., Lin, J.C., Gerbig, C., Munger, J.W., Chow, V.Y., Gottlieb, E.W.: A satellite-based biosphere parameterization for net ecosystem CO<sub>2</sub> exchange: Vegetation Photosynthesis and Respiration Model VPRM, *Global Biogeochem. Cycles*, 222, 1–17, 2008.
- 610 Melton, J. R. and Arora, V. K.: Competition between plant functional types in the Canadian Terrestrial Ecosystem Model (CTEM) v. 2.0, *Geosci. Model Dev.*, 9, 323–361, doi:10.5194/gmd-9-323-2016, 2016.
- ~~Morton, D.C., Nagol, J., Carabajal, C.C., Rosette, J., Palace, M., Cook, B.D., Vermote, E.F., Harding, D.J., North, P.R.J.: Amazon forests maintain consistent canopy structure and greenness during the dry season, *Nature*, 506, 221 +, doi:10.1038/nature13006, 2014.~~
- 615 Norby, R.J., DeLucia, E.H., Gielen, B., Calfapietra, C., Giardina, C.P., King, J.S., Ledford, J., McCarthy, H.R., Moore, D.J.P., Ceulemans, R., De Angelis, P., Finzi, A.C., Karnosky, D.F., Kubiske, M.E., Lukac, M., Pregitzer, K.S., Scarascia-Mugnozza, G.E., Schlesinger, W.H., Oren, R.: Forest response to elevated CO<sub>2</sub> is conserved across a broad range of productivity, *Proc. Natl. Acad. Sci. U.S.A.*, 102, 18052-18056, doi:10.1073/pnas.0509478102, 2005.
- Norby, R.J., Wullschleger, S.D., Gunderson, C.A., Johnson, D.W., Ceulemans, R.: Tree responses to rising CO<sub>2</sub> in field  
620 experiments: implications for the future forest, *Plant Cell Environ.*, 22, 683-714, doi:10.1046/j.1365-3040.1999.00391.x, 1999.
- Novick, K.A., Ficklin, D.L., Stoy, P.C., Williams, C.A., Bohrer, G., Oishi, A.C., Papuga, S.A., Blanken, P.D., Noormets, A., Sulman, B.N., Scott, R.L., Wang, L., Phillips, R.P.: The increasing importance of atmospheric demand for ecosystem water and carbon fluxes, *Nat. Clim. Change*, 6, 1023-1027, doi:10.1038/nclimate3114, 2016.
- Oleson, K. W., Lawrence, D. M., Bonan, G. B., Drewniak, B., Huang, M., Koven, C. D., Levis, S., Li, F., Riley, W. J., Subin,  
625 Z. M., Swenson, S. C., Thornton, P. E., Bozbiyik, A., Fisher, R., Heald, C. L., Kluzek, E., Lamarque, J., Lawrence, P. J., Leung, L. R., Lipscomb, W., Muszala, S., Ricciuto, D. M., Sacks, W., Tang, J., and Yang, Z.: Technical Description of version 4.5 of the Community Land Model (CLM), National Center for Atmospheric Research, Boulder, CO, USA, available at: [http://www.cesm.ucar.edu/models/cesm1.2/clm/CLM45\\_Tech\\_Note.pdf](http://www.cesm.ucar.edu/models/cesm1.2/clm/CLM45_Tech_Note.pdf), 2013.
- Piao, S., Sitch, S., Ciais, P., Friedlingstein, P., Peylin, P., Wang, X., Ahlstrom, A., Anav, A., Canadell, J.G., Cong, N.,  
630 Huntingford, C., Jung, M., Levis, S., Levy, P.E., Li, J., Lin, X., Lomas, M.R., Lu, M., Luo, Y., Ma, Y., Myneni, R.B., Poulter, B., Sun, Z., Wang, T., Viovy, N., Zaehle, S., Zeng, N.: Evaluation of terrestrial carbon cycle models for their response to climate variability and to CO<sub>2</sub> trends, *Global Change Biol.*, 19, 2117-2132, doi:10.1111/gcb.12187, 2013.
- Pierce, D.W., Westerling, A.L., Oyler, J.: Future humidity trends over the western United States in the CMIP5 global climate models and variable infiltration capacity hydrological modeling system, *Hydrol. Earth Syst. Sci.*, 17, 1833-1850,  
635 doi:10.5194/hess-17-1833-2013, 2013.
- Potter, C.S., Randerson, J.T., Field, C.B., Matson, P.A., Vitousek, P.M., Mooney, H.A., Klooster, S.A.: Terrestrial ecosystem production: A process model-based on global satellite and surface data, *Global Biogeochem. Cycles*, 7, 811-841, doi:10.1029/93gb02725, 1993.
- Prentice, I.C., Dong, N., Gleason, S.M., Maire, V., Wright, I.J.: Balancing the costs of carbon gain and water transport: testing  
640 a new theoretical framework for plant functional ecology, *Ecol. Lett.*, 17, 82-91, doi:10.1111/ele.12211, 2014.

- Rawson, H.M., Begg, J.E., Woodward, R.G.: The effect of atmospheric humidity on photosynthesis, transpiration and water use efficiency of leaves of several plant species, *Planta*, 134, 5-10, doi:10.1007/bf00390086, 1977.
- Reichstein, M., Falge, E., Baldocchi, D., Papale, D., Aubinet, M., Berbigier, P., Bernhofer, C., Buchmann, N., Gilmanov, T., Granier, A., Grunwald, T., Havrankova, K., Ilvesniemi, H., Janous, D., Knohl, A., Laurila, T., Lohila, A., Loustau, D.,  
645 Matteucci, G., Meyers, T., Miglietta, F., Ourcival, J.-M., Pumpanen, J., Rambal, S., Rotenberg, E., Sanz, M., Tenhunen, J., Seufert, G., Vaccari, F., Vesala, T., Yakir, D., Valentini, R.: On the separation of net ecosystem exchange into assimilation and ecosystem respiration: review and improved algorithm, *Global Change Biol.*, 11, 1424-1439, doi:10.1111/j.1365-2486.2005.001002.x, 2005.
- Reick, C. H., T. Raddatz, V. Brovkin, and Gayler, V.: The representation of natural and anthropogenic land cover change in  
650 MPIESM, *J. Adv. Model. Earth Syst.*, 5, 459–482, 2013.
- Rienecker, M.M., Suarez, M.J., Gelaro, R., Todling, R., Bacmeister, J., Liu, E., Bosilovich, M.G., Schubert, S.D., Takacs, L., Kim, G.-K., Bloom, S., Chen, J., Collins, D., Conaty, A., Da Silva, A., Gu, W., Joiner, J., Koster, R.D., Lucchesi, R., Molod, A., Owens, T., Pawson, S., Pegion, P., Redder, C.R., Reichle, R., Robertson, F.R., Ruddick, A.G., Sienkiewicz, M., Woollen, J.: MERRA: NASA's modern-era retrospective analysis for research and applications, *J. Clim.*, 24, 3624-3648,  
655 doi:10.1175/jcli-d-11-00015.1, 2011.
- Running, S.W., Nemani, R.R., Heinsch, F.A., Zhao, M.S., Reeves, M., Hashimoto, H.: A continuous satellite-derived measure of global terrestrial primary production, *Bioscience*, 54, 547-560, doi:10.1641/0006-3568(2004)054[0547:acsmog]2.0.co;2, 2004.
- Ryu, Y., Baldocchi, D.D., Kobayashi, H., van Ingen, C., Li, J., Black, T.A., Beringer, J., van Gorsel, E., Knohl, A., Law, B.E.,  
660 Roupsard, O.: Integration of MODIS land and atmosphere products with a coupled-process model to estimate gross primary productivity and evapotranspiration from 1 km to global scales, *Global Biogeochem. Cycles*, 2510.1029/2011gb004053, 2011.
- Ryu, Y., Berry, J.A., Baldocchi, D.D.: What is global photosynthesis? History, uncertainties and opportunities, *Remote Sens. Environ.*, 223, 95-114, doi:10.1016/j.rse.2019.01.016, 2019.
- Saleska, S.R., Didan, K., Huete, A.R., da Rocha, H.R.: Amazon forests green-up during 2005 drought, *Science*, 318, 612-612,  
665 doi:10.1126/science.1146663, 2007.
- Samanta, A., Ganguly, S., Hashimoto, H., Devadiga, S., Vermote, E., Knyazikhin, Y., Nemani, R.R., Myneni, R.B.: Amazon forests did not green-up during the 2005 drought, *Geophys. Res. Lett.*, 3710.1029/2009gl042154, 2010.
- Serbin, S.P., Dillaway, D.N., Kruger, E.L., Townsend, P.A.: Leaf optical properties reflect variation in photosynthetic metabolism and its sensitivity to temperature, *J. Exp. Bot.*, 63, 489-502, doi:10.1093/jxb/err294, 2012.
- 670 Simmons, A.J., Willett, K.M., Jones, P.D., Thorne, P.W., Dee, D.P.: Low-frequency variations in surface atmospheric humidity, temperature, and precipitation: Inferences from reanalyses and monthly gridded observational data sets, *J. Geophys. Res. D: Atmos.*, 11510.1029/2009jd012442, 2010.
- Sjostrom, M., Zhao, M., Archibald, S., Arneth, A., Cappelaere, B., Falk, U., de Grandcourt, A., Hanan, N., Kergoat, L., Kutsch, W., Merbold, L., Mougin, E., Nickless, A., Nouvellon, Y., Scholes, R.J., Veenendaal, E.M., Ardo, J.: Evaluation of MODIS

- 675 gross primary productivity for Africa using eddy covariance data, *Remote Sens. Environ.*, 131, 275-286, doi:10.1016/j.rse.2012.12.023, 2013.
- Smith, B., Wårlind, D., Arneth, A., Hickler, T., Leadley, P., Siltberg, J., and Zaehle, S.: Implications of incorporating N cycling and N limitations on primary production in an individualbased dynamic vegetation model, *Biogeosciences*, 11, 2027–2054, doi:10.5194/bg-11-2027-2014, 2014.
- 680 Smith, W.K., Reed, S.C., Cleveland, C.C., Ballantyne, A.P., Anderegg, W.R.L., Wieder, W.R., Liu, Y.Y., Running, S.W.: Large divergence of satellite and Earth system model estimates of global terrestrial CO<sub>2</sub> fertilization, *Nat. Clim. Change*, 6, 306-310, doi:10.1038/nclimate2879, 2016.
- Stocker, B. D., Feissli, F., Strassmann, K. M., Spahni, R., and Joos, F.: Past and future carbon fluxes from land use change, shifting cultivation and wood harvest, *Tellus B*, 66, 23188, doi:10.3402/tellusb.v66.23188, 2014.
- 685 Stocker, B.D., Zscheischler, J., Keenan, T.F., Prentice, I.C., Seneviratne, S.I., Peñuelas, J.: Drought impacts on terrestrial primary production underestimated by satellite monitoring, *Nat. Geosci.*, 12, 264-270, doi:10.1038/s41561-019-0318-6, 2019.
- Sulman, B.N., Roman, D.T., Yi, K., Wang, L., Phillips, R.P., Novick, K.A.: High atmospheric demand for water can limit forest carbon uptake and transpiration as severely as dry soil, *Geophys. Res. Lett.*, 43, 9686-9695, doi:10.1002/2016gl069416, 2016.
- 690 Tan, B., Woodcock, C.E., Hu, J., Zhang, P., Ozdogan, M., Huang, D., Yang, W., Knyazikhin, Y., Myneni, R.B.: The impact of gridding artifacts on the local spatial properties of MODIS data: Implications for validation, compositing, and band-to-band registration across resolutions, *Remote Sens. Environ.*, 105, 98-114, doi:10.1016/j.rse.2006.06.008, 2006.
- Tang, S., Chen, J.M., Zhu, Q., Li, X., Chen, M., Sun, R., Zhou, Y., Deng, F., Xie, D.: LAI inversion algorithm based on directional reflectance kernels, *J. Environ. Manage.*, 85, 638-648, doi:10.1016/j.jenvman.2006.08.018, 2007.
- 695 Turner, D.P., Ritts, W.D., Styles, J.M., Yang, Z., Cohen, W.B., Law, B.E., Thornton, P.E.: A diagnostic carbon flux model to monitor the effects of disturbance and interannual variation in climate on regional NEP, *Tellus B.*, 585, 476–490, 2006.
- Urban, O., Janous, D., Acosta, M., Czerny, R., Markova, I., Navratil, M., Pavelka, M., Pokorny, R., Sprtova, M., Zhang, R., Spunda, V., Grace, J., Marek, M.V.: Ecophysiological controls over the net ecosystem exchange of mountain spruce stand. Comparison of the response in direct vs. diffuse solar radiation, *Global Change Biol.*, 13, 157-168, doi:10.1111/j.1365-700 2486.2006.01265.x, 2007.
- Van Wijngaarden, W.A., Vincent, L.A.: Trends in relative humidity in Canada from 1953–2003. In 15th Symp. on Global Change and Climate Variations, 2004.
- Veroustraete, F., Sabbe, H., Eerens, H.: Estimation of carbon mass fluxes over Europe using the C-Fix model and Euroflux data, *Remote Sens. Environ.*, 833, 376–399, 2002.
- 705 Vuichard, N., Papale, D.: Filling the gaps in meteorological continuous data measured at FLUXNET sites with ERA-Interim reanalysis, *Earth Syst. Sci. Data*, 7, 157-171, doi:10.5194/essd-7-157-2015, 2015.

- Wang, Z., Skidmore, A.K., Darvishzadeh, R., Wang, T.: Mapping forest canopy nitrogen content by inversion of coupled leaf-canopy radiative transfer models from airborne hyperspectral imagery, *Agric. For. Meteorol.*, 253, 247-260, doi:10.1016/j.agrformet.2018.02.010, 2018.
- 710 Wild, M., Gilgen, H., Roesch, A., Ohmura, A., Long, C.N., Dutton, E.G., Forgan, B., Kallis, A., Russak, V., Tsvetkov, A.: From dimming to brightening: Decadal changes in solar radiation at Earth's surface, *Science*, 308, 847-850, doi:10.1126/science.1103215, 2005.
- Willett, K.M., Dunn, R.J.H., Thorne, P.W., Bell, S., de Podesta, M., Parker, D.E., Jones, P.D., Williams, C.N., Jr.: HadISDH land surface multi-variable humidity and temperature record for climate monitoring, *Climate of the Past*, 10, 1983-2006, 715 doi:10.5194/cp-10-1983-2014, 2014.
- Williams, A.P., Allen, C.D., Macalady, A.K., Griffin, D., Woodhouse, C.A., Meko, D.M., Swetnam, T.W., Rauscher, S.A., Seager, R., Grissino-Mayer, H.D., Dean, J.S., Cook, E.R., Gangodagamage, C., Cai, M., McDowell, N.G.: Temperature as a potent driver of regional forest drought stress and tree mortality, *Nat. Clim. Change*, 3, 292-297, doi:10.1038/nclimate1693, 2013.
- 720 Wu, J., Albert, L.P., Lopes, A.P., Restrepo-Coupe, N., Hayek, M., Wiedemann, K.T., Guan, K., Stark, S.C., Christoffersen, B., Prohaska, N., Tavares, J.V., Marostica, S., Kobayashi, H., Ferreira, M.L., Campos, K.S., da Silva, R., Brando, P.M., Dye, D.G., Huxman, T.E., Huete, A.R., Nelson, B.W., Saleska, S.R.: Leaf development and demography explain photosynthetic seasonality in Amazon evergreen forests, *Science*, 351, 972-976, doi:10.1126/science.aad5068, 2016.
- Wu, J., Guan, K., Hayek, M., Restrepo-Coupe, N., Wiedemann, K.T., Xu, X., Wehr, R., Christoffersen, B.O., Miao, G., da 725 Silva, R., de Araujo, A.C., Oliviera, R.C., Camargo, P.B., Monson, R.K., Huete, A.R., Saleska, S.R.: Partitioning controls on Amazon forest photosynthesis between environmental and biotic factors at hourly to interannual timescales, *Global Change Biol.*, 23, 1240-1257, doi:10.1111/gcb.13509, 2017.
- Xiao, X.M., Zhang, Q.Y., Hollinger, D., Aber, J., Moore, B.: Modeling gross primary production of an evergreen needleleaf forest using MODIS and climate data, *Ecol. Appl.*, 15, 954-969, doi:10.1890/04-0470, 2005.
- 730 Xiao, Z., Liang, S., Wang, J., Xiang, Y., Zhao, X., Song, J.: Long-Time-Series Global Land Surface Satellite Leaf Area Index Product Derived From MODIS and AVHRR Surface Reflectance, *IEEE Trans. Geosci. Remote Sens.*, 54, 5301-5318, doi:10.1109/tgrs.2016.2560522, 2016.
- Xu, B., Li, J., Park, T., Liu, Q., Zeng, Y., Yin, G., Zhao, J., Fan, W., Yang, L., Knyazikhin, Y., Myneni, R.B.: An integrated method for validating long-term leaf area index products using global networks of site-based measurements, *Remote Sens. Environ.*, 209, 134-151, doi:10.1016/j.rse.2018.02.049, 2018. 735
- Yoder, B.J., Pettigrew-Crosby, R.E.: Predicting nitrogen and chlorophyll content and concentrations from reflectance spectra (400-2500 nm) at leaf and canopy scales, *Remote Sens. Environ.*, 53, 199-211, doi:10.1016/0034-4257(95)00135-n, 1995.
- Yuan, W., Cai, W., Xia, J., Chen, J., Liu, S., Dong, W., Merbold, L., Law, B., Arain, A., Beringer, J., Bernhofer, C., Black, A., Blanken, P.D., Cescatti, A., Chen, Y., Francois, L., Gianelle, D., Janssens, I.A., Jung, M., Kato, T., Kiely, G., Liu, D., 740 Marcolla, B., Montagnani, L., Raschi, A., Rouspard, O., Varlagin, A., Wohlfahrt, G.: Global comparison of light use efficiency



- models for simulating terrestrial vegetation gross primary production based on the La Thuile database, *Agric. For. Meteorol.*, 192, 108-120, doi:10.1016/j.agrformet.2014.03.007, 2014.
- Yuan, W., Liu, S., Zhou, G., Zhou, G., Tieszen, L.L., Baldocchi, D., Bernhofer, C., Gholz, H., Goldstein, A.H., Goulden, M.L., Hollinger, D.Y., Hu, Y., Law, B.E., Stoy, P.C., Vesala, T., Wofsy, S.C., AmeriFlux, C.: Deriving a light use efficiency model from eddy covariance flux data for predicting daily gross primary production across biomes, *Agric. For. Meteorol.*, 143, 189-207, doi:10.1016/j.agrformet.2006.12.001, 2007.
- Yuan, W., Luo, Y., Li, X., Liu, S., Yu, G., Zhou, T., Bahn, M., Black, A., Desai, A.R., Cescatti, A., Marcolla, B., Jacobs, C., Chen, J., Aurela, M., Bernhofer, C., Gielen, B., Bohrer, G., Cook, D.R., Dragoni, D., Dunn, A.L., Gianelle, D., Gruenwald, T., Ibrom, A., Leclerc, M.Y., Lindroth, A., Liu, H., Marchesini, L.B., Montagnani, L., Pita, G., Rodeghiero, M., Rodrigues, A., Starr, G., Stoy, P.C.: Redefinition and global estimation of basal ecosystem respiration rate, *Global Biogeochem. Cycles*, 2510.1029/2011gb004150, 2011.
- Yuan, W., Liu, S., Yu, G., Bonnefond, J.-M., Chen, J., Davis, K., Desai, A.R., Goldstein, A.H., Gianelle, D., Rossi, F., Suyker, A.E., Verma, S.B.: Global estimates of evapotranspiration and gross primary production based on MODIS and global meteorology data, *Remote Sens. Environ.*, 114, 1416-1431, doi:10.1016/j.rse.2010.01.022, 2010.
- Yuan, W., Zheng, Y., Piao, S., Ciais, P., Lombardozzi, D., Wang, Y., Ryu, Y., Chen, G., Dong, W., Hu, Z., Jain, A.K., Jiang, C., Kato, E., Li, S., Lienert, S., Liu, S., Nabel, J.E.M.S., Qin, Z., Quine, T., Sitch, S., Smith, W.K., Wang, F., Wu, C., Xiao, Z., Yang, S.: Increased atmospheric vapor pressure deficit reduces global vegetation growth, *Science Advances*, [510.1126/sciadv.aax1396](https://doi.org/10.1126/sciadv.aax1396), [eaax1396](https://doi.org/10.1126/sciadv.aax1396), 2019.
- Zhang, H. Q., Pak, B., Wang, Y. P., Zhou, X. Y., Zhang, Y. Q., and Zhang, L.: Evaluating Surface Water Cycle Simulated by the Australian Community Land Surface Model (CABLE) across Different Spatial and Temporal Domains, *J. of Hydrometeorol.*, 14, 1119–1138, 2013.
- Zhang, Y., Xiao, X., Wu, X., Zhou, S., Zhang, G., Qin, Y., Dong, J.: Data Descriptor: A global moderate resolution dataset of gross primary production of vegetation for 2000-2016, *Scientific Data*, 410.1038/sdata.2017.165, 2017.
- Zhao, M., Running, S.W.: Drought-Induced Reduction in Global Terrestrial Net Primary Production from 2000 Through 2009, *Science*, 329, 940-943, doi:10.1126/science.1192666, 2010.
- Zheng, Y., Shen, R.; Wang, Y., Li, X., Liu, S., Liang, S., Chen, Jing M., Ju, W., Zhang, L., Yuan, W.: Improved estimate of global gross primary production for reproducing its long-term variation, 1982-2017. *figshare. Dataset*. doi:10.6084/m9.figshare.8942336, 2019.
- Zheng, Y., Zhang, L., Xiao, J., Yuan, W., Yan, M., Li, T., Zhang, Z.: Sources of uncertainty in gross primary productivity simulated by light use efficiency models: Model structure, parameters, input data, and spatial resolution, *Agric. For. Meteorol.*, 263, 242-257, doi:10.1016/j.agrformet.2018.08.003, 2018.
- [Zhou, S., Williams, A. P., Berg, A. M., Cook, B. I., Zhang, Y., Hagemann, S., Lorenz, R., Seneviratne, S. I., and Gentine, P.: Land-atmosphere feedbacks exacerbate concurrent soil drought and atmospheric aridity, \*Proc. Natl. Acad. Sci. U.S.A.\*, 116, 18848-18853, 2019a.](https://doi.org/10.1126/sciadv.aax1396)

775 [Zhou, S., Zhang, Y., Williams, A. P., and Gentine, P.: Projected increases in intensity, frequency, and terrestrial carbon costs of compound drought and aridity events. \*Science Advances\*, 5, eaau5740, 2019b.](#)

Zhu, Z., Piao, S., Myneni, R.B., Huang, M., Zeng, Z., Canadell, J.G., Ciais, P., Sitch, S., Friedlingstein, P., Arneeth, A., Cao, C., Cheng, L., Kato, E., Koven, C., Li, Y., Lian, X., Liu, Y., Liu, R., Mao, J., Pan, Y., Peng, S., Penuelas, J., Poulter, B., Pugh, T.A.M., Stocker, B.D., Viovy, N., Wang, X., Wang, Y., Xiao, Z., Yang, H., Zaehle, S., Zeng, N.: Greening of the Earth and its drivers, *Nat. Clim. Change*, 6, 791-796, doi:10.1038/nclimate3004, 2016.

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## Tables

**Table 1: Information on the eddy covariance (EC) sites used in this study.**

Site Name	Latitude	Longitude	Vegetation Type	Study Period
*DE-Kli	50.89°N	13.52°E	CRO	2004-2012
DE-RuS	50.87°N	6.45°E	CRO	2011-2012
FI-Jok	60.90°N	23.51°E	CRO	2001-2003
*FR-Gri	48.84°N	1.95°E	CRO	2005-2012
*US-ARM	36.61°N	97.49°W	CRO	2003-2012
*US-Ne1	41.16°N	96.47°W	CRO	2001-2012
*US-Ne2	41.16°N	96.47°W	CRO	2001-2012
*US-Ne3	41.17°N	96.43°W	CRO	2001-2012
CA-TPD	42.64°N	80.56°W	DBF	2012
*DE-Hai	51.08°N	10.45°E	DBF	2000-2012
*DK-Sor	55.49°N	11.64°E	DBF	2001-2012
*FR-Fon	48.48°N	2.78°E	DBF	2005-2012
IT-PT1	45.20°N	9.06°E	DBF	2002-2004
*IT-Ro2	42.39°N	11.92°E	DBF	2002-2008; 2010-2012
JP-MBF	44.39°N	142.32°E	DBF	2004-2005
*US-Ha1	42.54°N	72.17°W	DBF	1992-2012
*US-MMS	39.32°N	86.41°W	DBF	1999-2012
*US-Oho	41.55°N	83.84°W	DBF	2004-2012
*US-UMB	45.56°N	84.71°W	DBF	2000-2012
*US-UMd	45.56°N	84.70°W	DBF	2008-2012
*US-WCr	45.81°N	90.08°W	DBF	1999-2006; 2011-2012
*BR-Sa1	2.86°S	54.96°W	EBF	2002-2005; 2008-2011
BR-Sa3	3.02°S	54.97°W	EBF	2001-2003
CN-Din	23.17°N	112.54°E	EBF	2003-2005
*FR-Pue	43.74°N	3.60°E	EBF	2000-2012
*GF-Guy	5.28°N	52.92°W	EBF	2004-2012
*MY-PSO	2.97°N	102.31°E	EBF	2003-2009
CA-NS1	55.88°N	98.48°W	ENF	2002-2005
*CA-NS2	55.91°N	98.52°W	ENF	2001-2005
CA-NS3	55.91°N	98.38°W	ENF	2002-2005
CA-NS4	55.91°N	98.38°W	ENF	2003-2005
*CA-NS5	55.86°N	98.49°W	ENF	2001-2005

*CA-Qfo	49.69°N	74.34°W	ENF	2003-2010
CA-SF1	54.49°N	105.82°W	ENF	2003-2006
*CA-SF2	54.25°N	105.88°W	ENF	2001-2005
*CA-TP1	42.66°N	80.56°W	ENF	2003-2012
*CA-TP2	42.77°N	80.46°W	ENF	2003-2007
*CA-TP3	42.71°N	80.35°W	ENF	2003-2012
CN-Qia	26.74°N	115.06°E	ENF	2003-2005
*CZ-BK1	49.50°N	18.54°E	ENF	2004-2012
DE-Lkb	49.10°N	13.30°E	ENF	2009-2012
*DE-Obe	50.78°N	13.72°E	ENF	2008-2012
*DE-Tha	50.96°N	13.57°E	ENF	1996-2012
*FI-Hyy	61.85°N	24.30°E	ENF	1996-2012
IT-La2	45.95°N	11.29°E	ENF	2001
*IT-Lav	45.96°N	11.28°E	ENF	2003-2012
*IT-Ren	46.59°N	11.43°E	ENF	1999-2012
*IT-SRo	43.73°N	10.28°E	ENF	2001-2012
*NL-Loo	52.17°N	5.74°E	ENF	1996-2012
*RU-Fyo	56.46°N	32.92°E	ENF	1998-2012
*US-Blo	38.90°N	120.63°W	ENF	1997-2007
*US-Me2	44.45°N	121.56°W	ENF	2002-2012
US-Me6	44.32°N	121.61°W	ENF	2011-2012
*US-NR1	40.03°N	105.55°W	ENF	1999-2012
*CH-Cha	47.21°N	8.41°E	GRA	2006-2008; 2010-2012
*CH-Fru	47.12°N	8.54°E	GRA	2006-2008; 2010-2012
*CH-Oe1	47.29°N	7.73°E	GRA	2002-2008
CN-Cng	44.59°N	123.51°E	GRA	2007-2010
CN-Du2	42.05°N	116.28°E	GRA	2007-2008
CN-HaM	37.37°N	101.18°E	GRA	2002-2003
*CZ-BK2	49.49°N	18.54°E	GRA	2006-2011
*NL-Hor	52.24°N	5.07°E	GRA	2004-2011
RU-Ha1	54.73°N	90.00°E	GRA	2002-2004
US-AR1	36.43°N	99.42°W	GRA	2009-2012
US-AR2	36.64°N	99.60°W	GRA	2009-2012
*US-Goo	34.25°N	89.87°W	GRA	2002-2006
*US-IB2	41.84°N	88.24°W	GRA	2005; 2007-2011
*BE-Bra	51.31°N	4.52°E	MF	1999-2002; 2004-2012
*BE-Vie	50.31°N	6.00°E	MF	1997-2012

*CA-Gro	48.22°N	82.16°W	MF	2004-2012
CN-Cha	42.40°N	128.10°E	MF	2003-2005
JP-SMF	35.26°N	137.08°E	MF	2003-2006
*US-PFa	45.95°N	90.27°W	MF	1996-2012
*US-Syv	46.24°N	89.35°W	MF	2001-2006; 2012
AU-Ade	13.08°S	131.12°E	SAV	2007-2009
AU-Cpr	34.00°S	140.59°E	SAV	2011-2012
*AU-DaS	14.16°S	131.39°E	SAV	2008-2012
AU-Dry	15.26°S	132.37°E	SAV	2009-2012
AU-RDF	14.56°S	132.48°E	SAV	2011-2012
SD-Dem	13.28°N	30.48°E	SAV	2007-2009
*US-Ton	38.43°N	120.97°W	SAV	2001-2012
ZA-Kru	25.02°S	31.50°E	SAV	2009-2012
CA-NS6	55.92°N	98.96°W	SRH	2002-2005
CA-NS7	56.64°N	99.95°W	SRH	2003-2005
*CA-SF3	54.09°N	106.01°W	SRH	2002-2006
ES-LgS	37.10°N	2.97°W	SRH	2007-2009
US-KS2	28.61°N	80.67°W	SRH	2003-2006
CN-Ha2	37.61°N	101.33°E	WET	2003-2005
DE-Akm	53.87°N	13.68°E	WET	2010-2012
DE-SfN	47.81°N	11.33°E	WET	2012
DE-Spw	51.89°N	14.03°E	WET	2010-2012
RU-Che	68.61°N	161.34°E	WET	2002-2004
US-Ivo	68.49°N	155.75°W	WET	2004-2007
*US-Los	46.08°N	89.98°W	WET	2001-2008; 2010
US-WPT	41.46°N	83.00°W	WET	2011-2012

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785 \* indicates the site used to investigate the interannual variations in GPP with observations greater than 5-years.

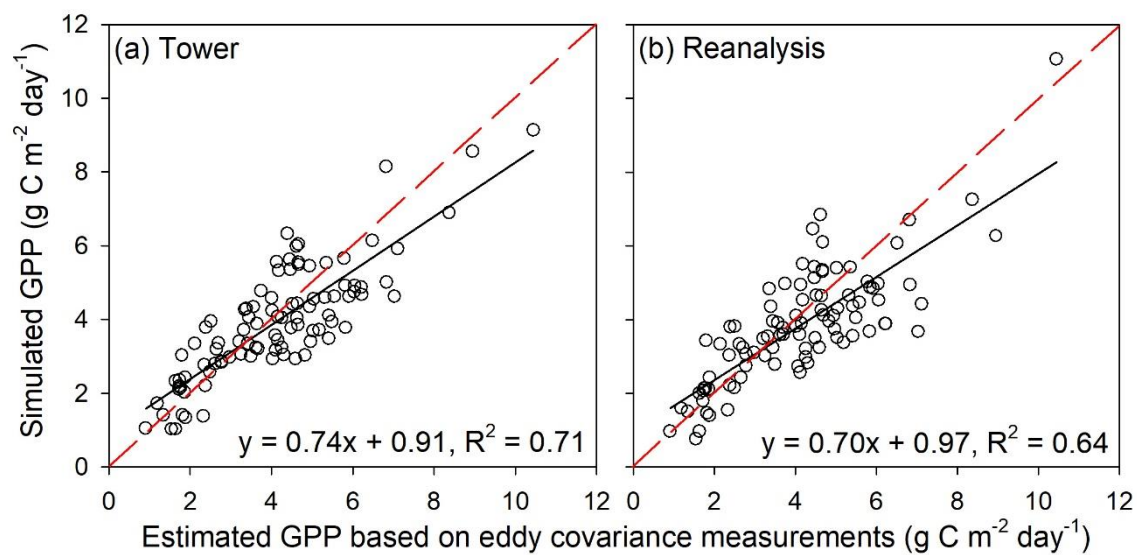
**Table 2: Input datasets used to drive the revised EC-LUE model.**

Variable	Dataset/provider	Source
Air temperature	MERRA2	<a href="https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/">https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/</a>
Dew point temperature	MERRA2	<a href="https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/">https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/</a>
Direct PAR	MERRA2	<a href="https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/">https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/</a>
Diffuse PAR	MERRA2	<a href="https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/">https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/</a>
LAI	GLASS	<a href="http://www.glass.umd.edu/Download.html">http://www.glass.umd.edu/Download.html</a>
Landcover map	MCD12Q1	<a href="https://lpdaac.usgs.gov/products/mcd12q1v006/">https://lpdaac.usgs.gov/products/mcd12q1v006/</a>
C4 crop percentage	ISLSCP II C4 Vegetation Percentage	<a href="https://doi.org/10.3334/ORNLDAAAC/932">https://doi.org/10.3334/ORNLDAAAC/932</a>
CO <sub>2</sub> concentration	NOAA's Earth System Research Laboratory	<a href="http://www.esrl.noaa.gov/gmd/ccgg/trends/">www.esrl.noaa.gov/gmd/ccgg/trends/</a>

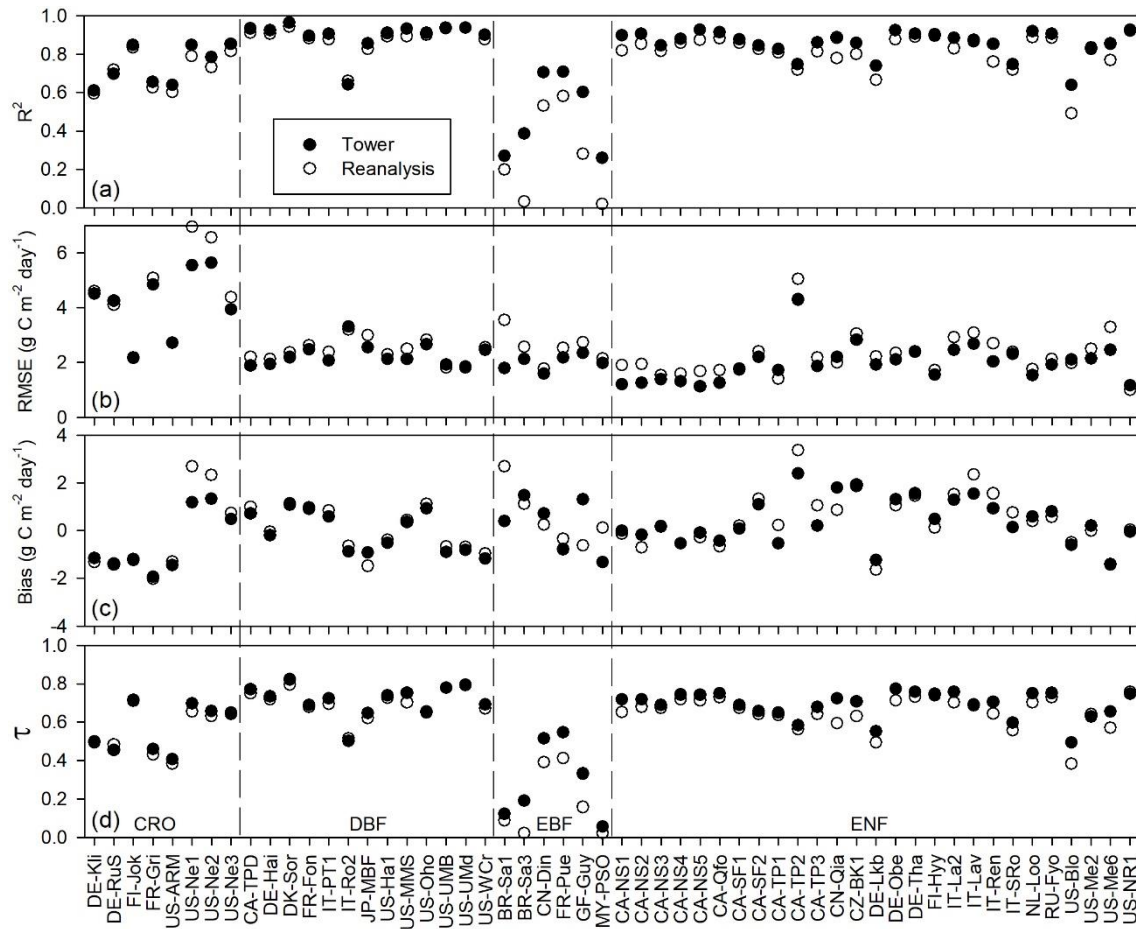
790 **Table 3: Optimized parameters ( $\epsilon_{msu}$ ,  $\epsilon_{msh}$ ,  $\phi$ , and  $VPD_0$ ) of the revised EC-LUE model for different vegetation types.**

Vegetation Types	DBF	ENF	EBF	MF	GRA	CRO-C3	CRO-C4	SAV	SHR	WET
$\epsilon_{msu}$ (g C MJ <sup>-1</sup> )	1.28 ± 0.36	1.72 ± 0.42	1.67 ± 0.85	1.38 ± 0.21	1.16 ± 0.15	1.25 ± 0.42	2.46 ± 0.78	2.24 ± 0.68	1.21 ± 0.25	1.34 ± 0.26
$\epsilon_{msh}$ (g C MJ <sup>-1</sup> )	3.59 ± 0.66	3.87 ± 0.58	4.35 ± 0.72	3.29 ± 0.63	1.91 ± 0.46	2.46 ± 0.52	5.64 ± 1.02	4.26 ± 0.95	2.71 ± 0.52	2.62 ± 0.49
$\phi$ (ppm)	32 ± 8.25	25 ± 7.59	20 ± 6.36	49 ± 11.25	57 ± 11.85	43 ± 9.56	54 ± 15.36	54 ± 12.23	34 ± 7.59	36 ± 10.32
$VPD_0$ (k Pa)	1.15 ± 0.25	1.34 ± 0.26	0.57 ± 0.15	0.62 ± 0.14	1.69 ± 0.35	1.02 ± 0.19	1.53 ± 0.31	1.65 ± 0.26	1.34 ± 0.21	0.62 ± 0.12

## Figures



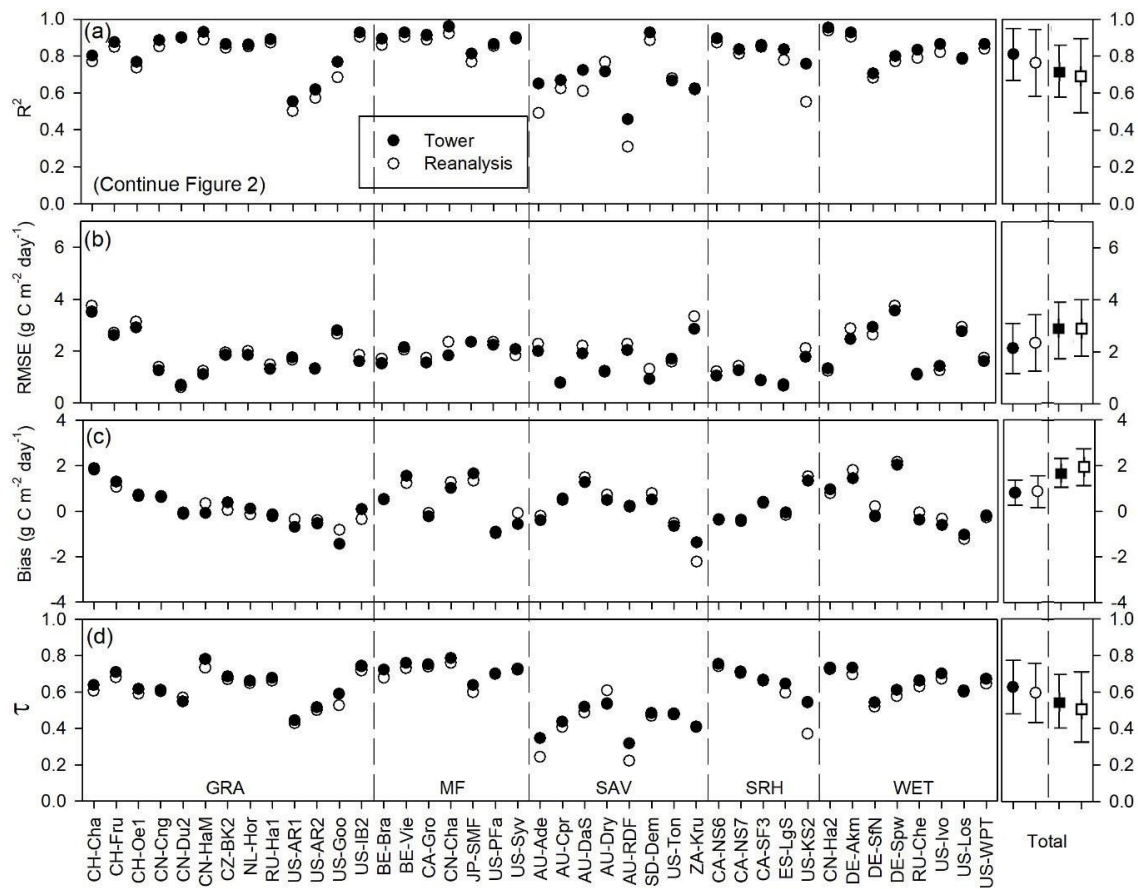
795 **Figure 1: Comparisons between annual mean GPP estimated from EC towers and annual mean GPP simulated by the revised EC-LUE model. The modeled GPP were simulated using (a) tower-derived meteorology and (b) global reanalysis meteorology. The black lines are the regression lines, and the red dash lines are the 1:1 lines. The insert equations are the regression equations derived from all the sites.**

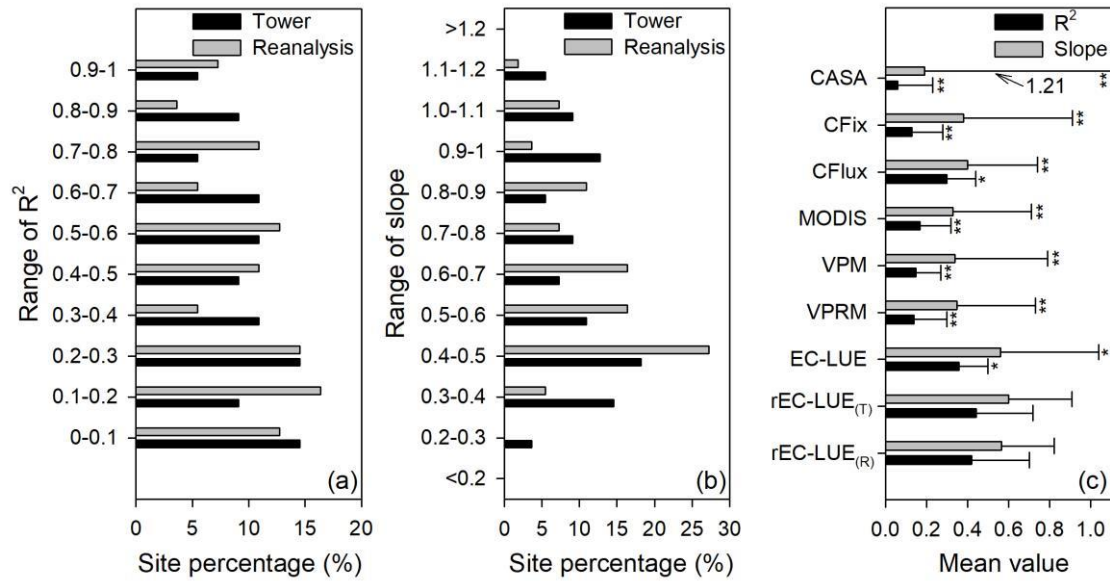


**Figure 2: Comparisons of 8-day mean GPP between the model simulated GPP and tower estimated GPP. Solid and open dots indicate the GPP simulations of the revised EC-LUE model derived from tower-derived meteorology data and meteorological reanalysis dataset, respectively; solid and open squares indicate the GPP simulations of the original EC-LUE model derived from tower-derived meteorology data and meteorological reanalysis dataset, respectively.**



Figure 2 (continue)





815 **Figure 3: Site percentage of (a) correlation coefficients ( $R^2$ ), and (b) regression slopes between the model simulated and tower-based interannual variabilities in GPP. (c) Averaged values (error bars represent the standard deviations) of  $R^2$  and slope for various LUE models. rEC-LUE<sub>(T)</sub> and rEC-LUE<sub>(R)</sub> indicate the revised EC-LUE models derived from tower-derived meteorology data and meteorological reanalysis dataset. The  $R^2$  and slopes of the other seven LUE models (i.e., EC-LUE, VPRM, VPM, MODIS, CFlux, CFix, and CASA) in the figure were obtained from the study by Yuan et al. (2014). \*\* and \* indicate a significant difference in statistic variables ( $R^2$  and slope) between the rEC-LUE<sub>(T)</sub> and other LUE models (i.e., rEC-LUE<sub>(T)</sub> and other seven LUE models) at  $p$ -value < 0.01 and  $p$ -value < 0.05, respectively.**

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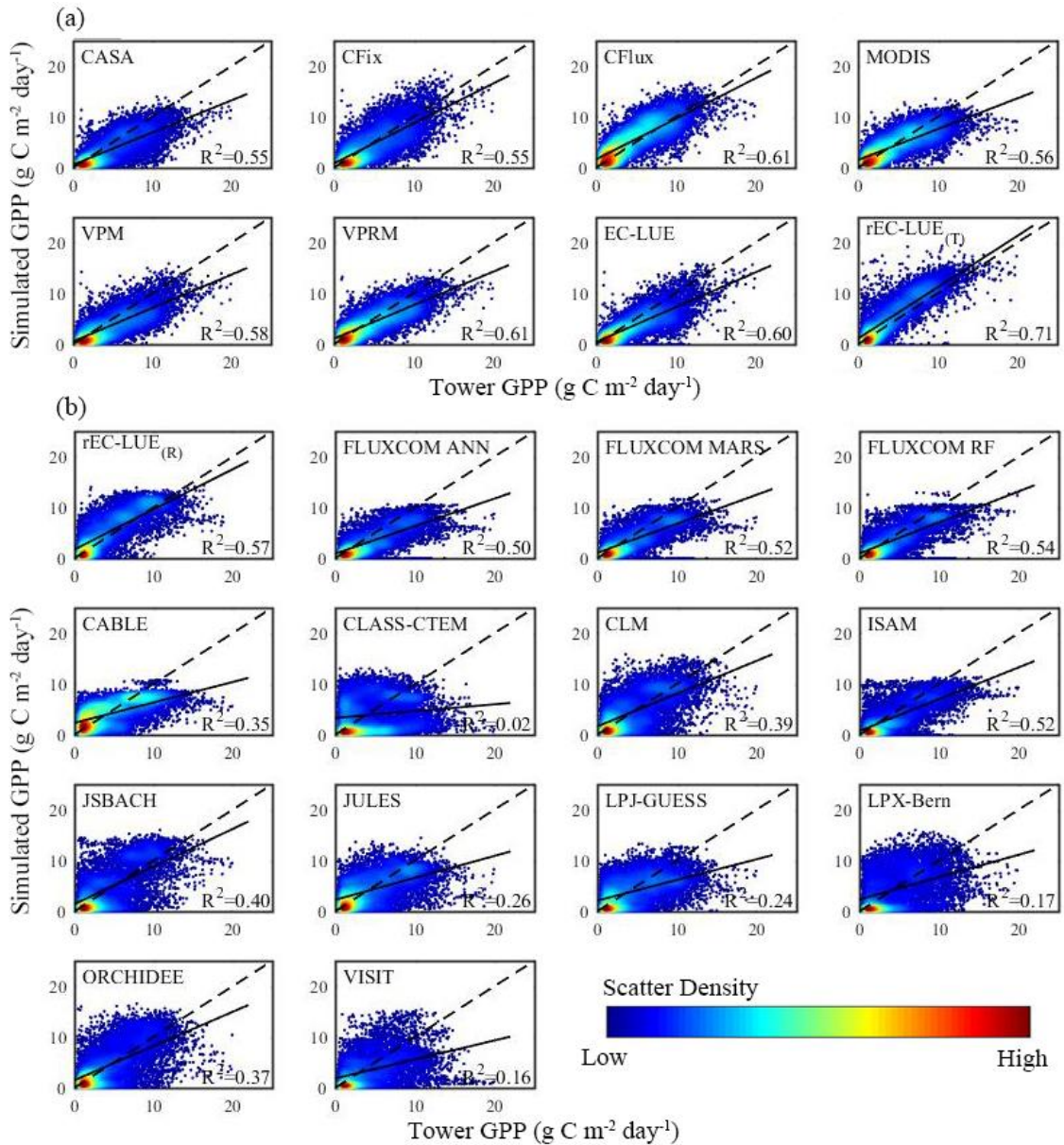
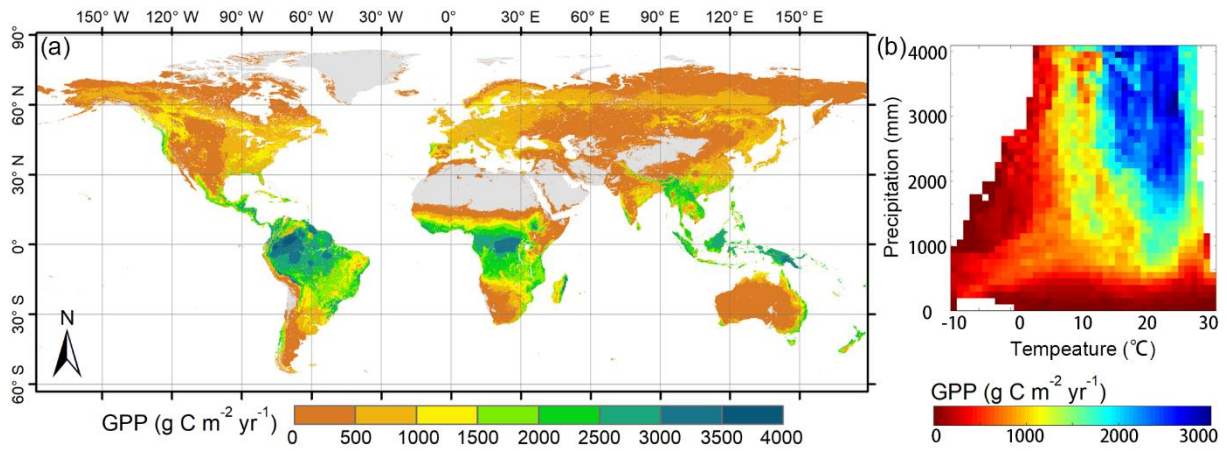
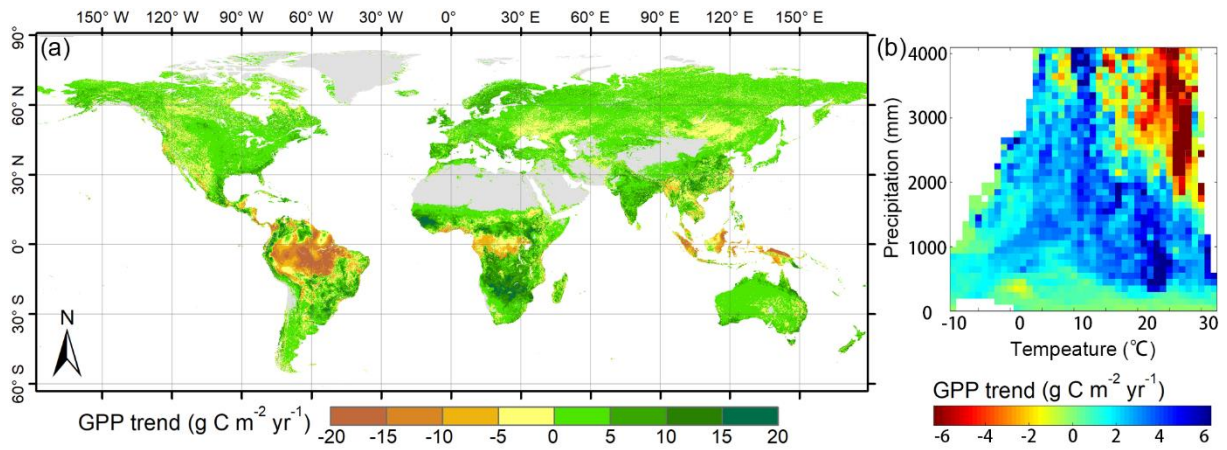


Figure 4: Comparisons between estimated GPP based on EC measurements and GPP simulations in growing season (defined as temperature larger than 0°C) by the various models (including LUE models, machine learning models, and process-based models in TRENDY) at monthly scale. The comparison of GPP simulations were simulated using (a) tower-derived meteorology data and (b) global reanalysis meteorology data, respectively (see method 2.4). rEC-LUE<sub>(T)</sub> and rEC-LUE<sub>(R)</sub> indicate the simulations of the revised EC-LUE model derived from tower-derived meteorology data and global reanalysis meteorology data, respectively.



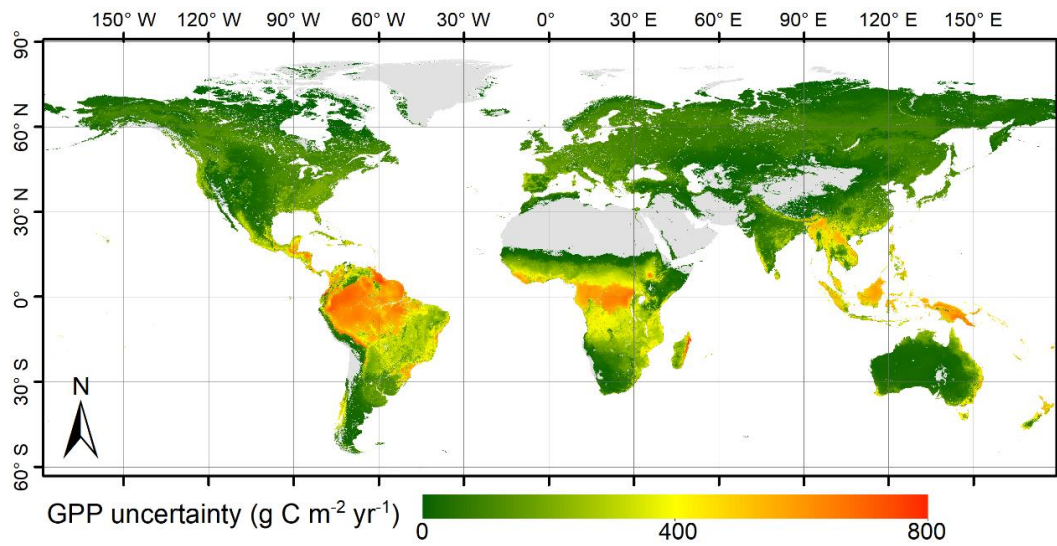
**Figure 5: Spatial pattern of global GPP simulated by the revised EC-LUE model during 1982–2017: (a) averaged annual GPP, (b) averaged annual GPP at different temperature and precipitation gradients.**

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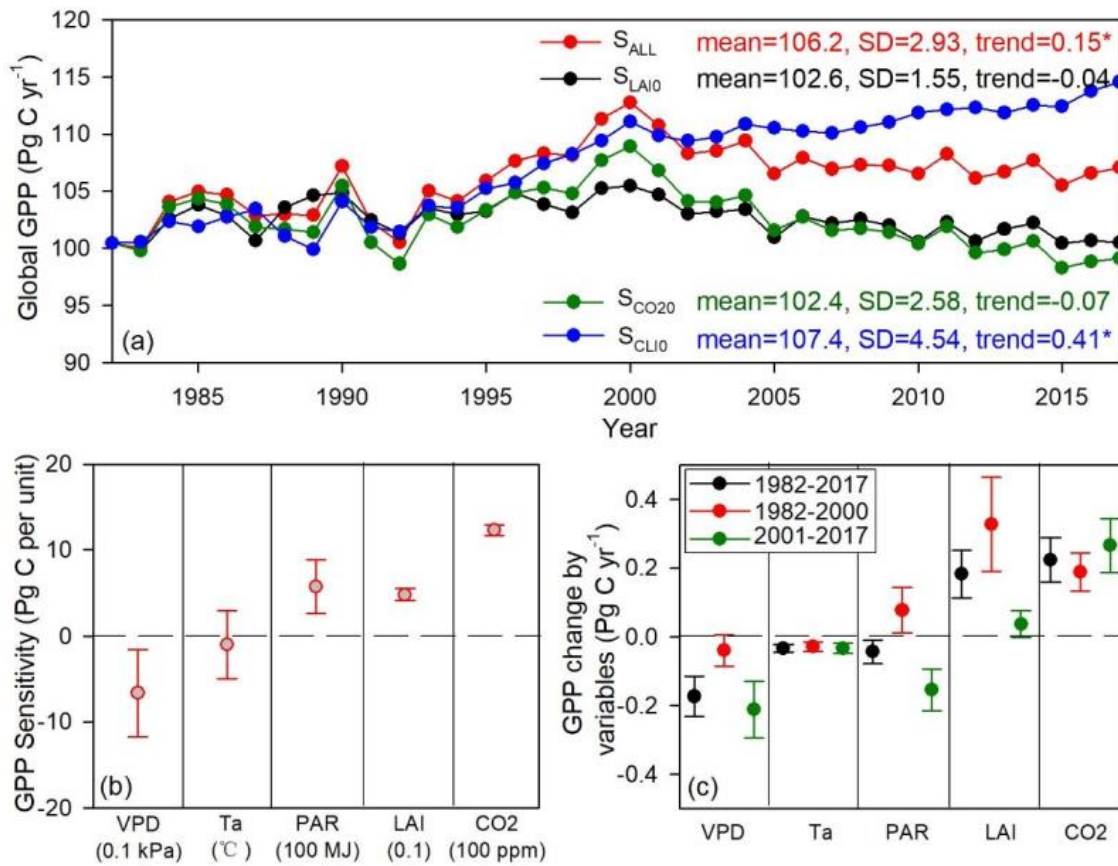
**Figure 6: Spatial pattern of global GPP trend simulated by the revised EC-LUE models during 1982–2017: (a) trend of annual GPP, (b) trend of annual GPP at different temperature and precipitation gradients.**

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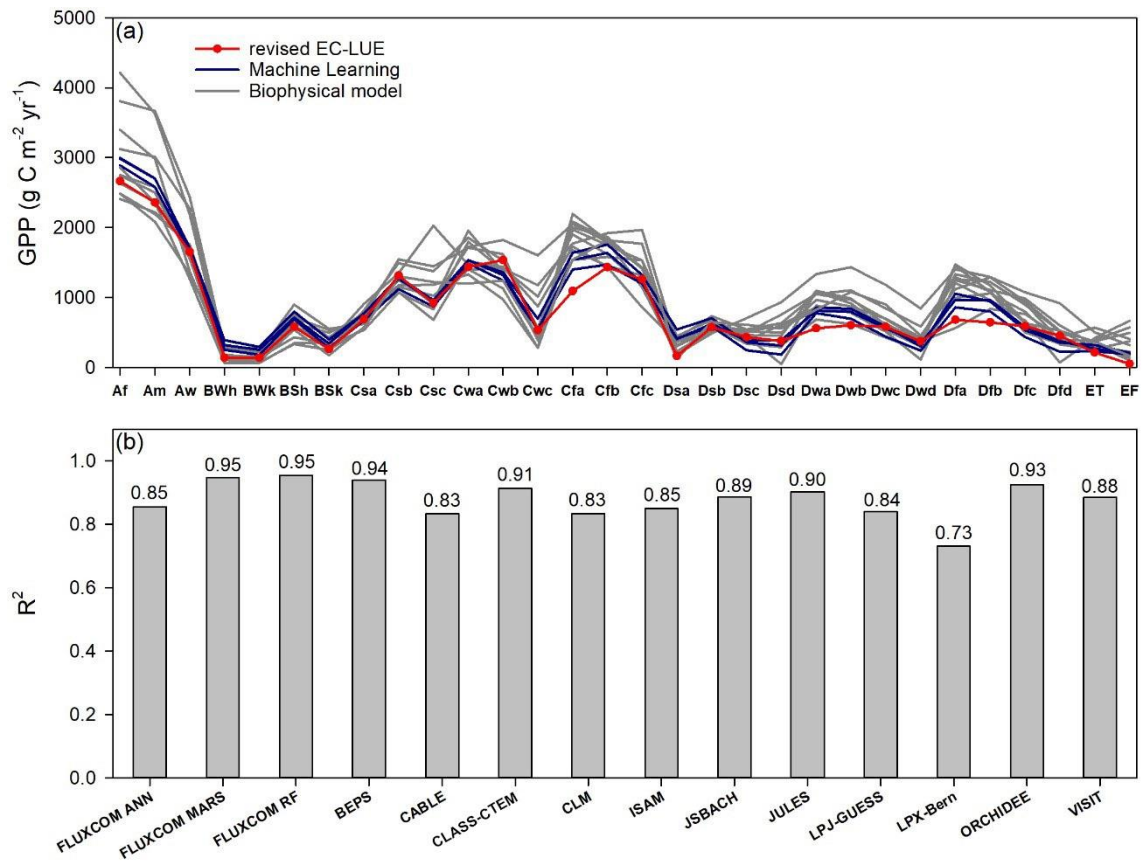


**Figure 7: Spatial pattern of the uncertainty in global GPP simulated by the revised EC-LUE model.**





840 **Figure 8: Long-term changes in global GPP and the environmental regulations: (a) Global summed GPP derived from the four experimental simulations in section 2.5, (b) GPP sensitivity to climate variables (i.e., VPD, Ta, and PAR), LAI, and atmospheric CO<sub>2</sub>, (c) contributions of climate variables (i.e., VPD, Ta, and PAR), LAI, and atmospheric CO<sub>2</sub> to GPP changes over 1982–2017, 1982–2000, and 2001–2017. \* indicates the significant level at  $p$ -value<0.05.**

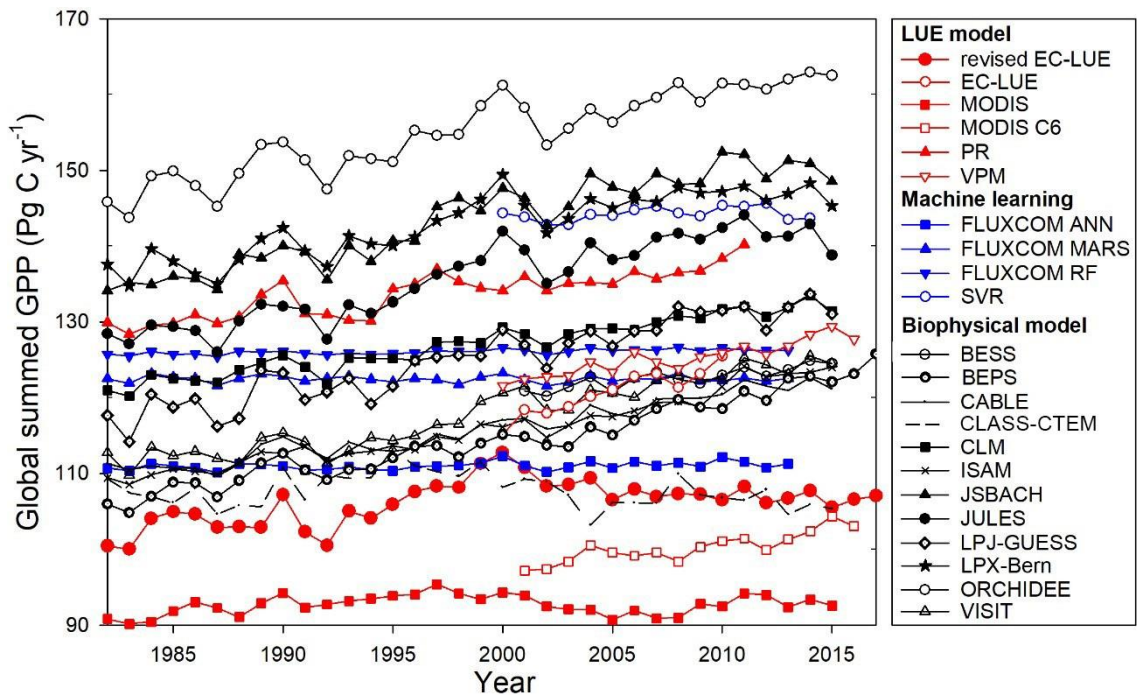


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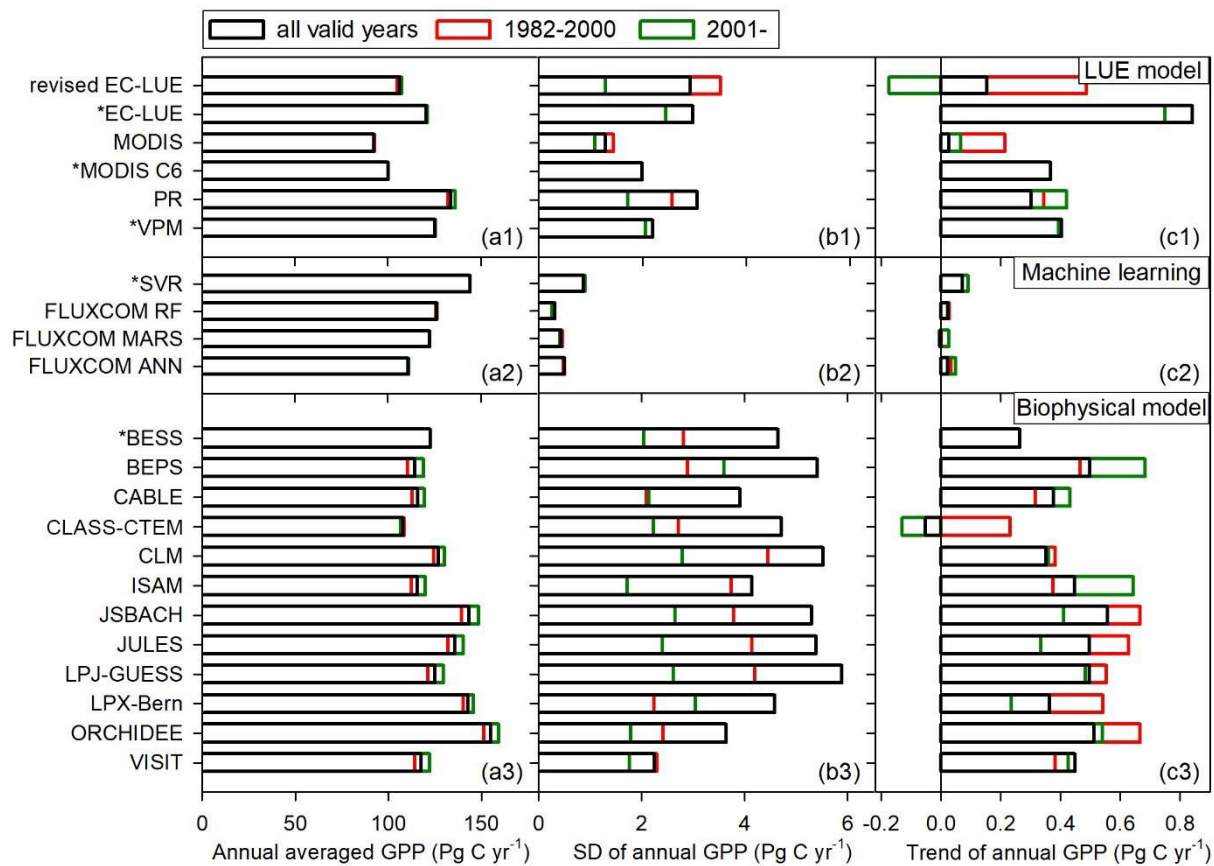
Figure 9: Comparisons of long-term (1982 to 2010s) averaged GPP between the revised EC-LUE model and other models across bioclimatic zones in the Köppen-Geiger climate classification map (Beck et al., 2018). (a) the regional averaged value (b) correlation coefficients ( $R^2$ ) of GPP across all the bioclimatic zones between the revised EC-LUE model and other models. These models including machine learning models (FLUXCOM ANN, FLUXCOM MARS, FLUXCOM RF; Jung et al., 2017), biophysical models BEPS (Ju et al., 2006; Liu et al., 2018), and ten biophysical models in TRENDY (CABLE, CLASS-CTEM, CLM, ISAM, JSBACH, JULES, LPJ-GUESS, LPX-Bern, ORCHIDEE, and VISIT). The abbreviations for the bioclimatic zones are as follows: Af, tropical, rainforest; Am, tropical, monsoon; Aw, tropical, savannah; BWh, arid, desert, hot; BWk, arid, desert, cold; BSh, arid, steppe, hot; BSk, arid, steppe, cold; Csa, temperate, dry summer, hot summer; Csb, temperate, dry summer, warm summer; Csc, temperate, dry summer, cold summer; Cwa, temperate, dry winter, hot summer; Cwb, temperate, dry winter, warm summer; Cwc, temperate, dry winter, cold summer; Cfa, temperate, no dry season, hot summer; Cfb, temperate, no dry season, warm summer; Cfc, temperate, no dry season, cold summer; Dsa, cold, dry summer, hot summer; Dsb, cold, dry summer, warm summer; Dsc, cold, dry summer, cold summer; Dsd, cold, dry summer, very cold winter; Dwa, cold, dry winter, hot summer; Dwb, cold, dry winter, warm summer; Dwc, cold, dry winter, cold summer; Dwd, cold, dry winter, very cold winter; Dfa, cold, no dry season, hot summer; Dfb, cold, no dry season, warm summer; Dfc, cold, no dry season, cold summer; Dfd, cold, no dry season, very cold winter; ET, polar, tundra; EF, polar, frost.

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865 **Figure 10: Comparisons of annual global summed GPP estimates from various models. The datasets or model algorithms were obtained from: EC-LUE (Cai et al., 2014), MODIS (Smith et al., 2016), MOD17 C6 (Running et al., 2004), PR (Keenan et al., 2016), VPM (Zhang et al., 2017), FLUXCOM (Jung et al., 2017), SVR (Kondo et al., 2015), BESS (Jiang and Ryu, 2016), BEPS (Ju et al., 2006; Liu et al., 2018), and models in TRENDY (CABLE, CLASS-CTEM, CLM, ISAM, JSBACH, JULES, LPJ-GUESS, LPX-Bern, ORCHIDEE, and VISIT).**



**Figure 11: Comparison of (a1)–(a3) averaged annual GPP, (b1)–(b3) interannual variability in annual GPP represented by standard deviation (SD), and (c1)–(c3) annual GPP trend among different GPP datasets or models. The references of these models are the same as in Figure 9. \* indicates that the valid period of the dataset begins from 2000 or 2001.**