



**Title:**

**Capillary rise affecting crop yields under different environmental conditions**

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*Abstract*

10 This paper describes analyses of different soil water flow regimes on growth and yields of  
grass, maize and potato crops in the Dutch delta, with a focus on the role of capillary rise.  
Different flow regimes are characterised by differences in soil composition and structure are  
derived from a national soil database. Capillary rise and its influence on crop growth and  
resulting yields is simulated using Swap-Wofost with different boundary conditions. Case  
15 studies and model experiments are used to illustrate the impact of capillary rise. This impact  
is clearly present in situations where a groundwater level is present (85% of NL) but also in  
other situations the impact of capillary rise on crop growth and production is considerable.  
When one compares situations with average groundwater levels with free drainage  
conditions without capillary rise yield-reductions of grassland, maize and potatoes are  
20 respectively 25, 4 and 15 % or respectively about 3.2, 0.5 and 1.6 ton dry Matter per ha.  
Neglecting capillary rise also has impact on the downward leaching water flux, the  
groundwater recharge. Impact can be considerable; for grassland and potatoes the  
reduction is 17 and 46% or 64 and 34 mm. Modelling of soil water flow should consider  
capillary rise of soil water which will results in improved yield and downward leaching  
25 simulations.

*1. Introduction*

30 Crop growth strongly depends on soil moisture conditions. Climate conditions determine  
these conditions through rain that penetrates directly into the root zone or comes available  
via lateral flow. The distribution of soil moisture strongly depends on soil physical properties  
that determine vertical flow. The upward capillary flow, capillary rise, becomes a vital supply  
term to a crop when it manages to bridge the distance between groundwater level and crop  
35 roots. The contribution of this capillary flow to the total water budget can be significant. For  
example Kowalik (2006) mentions that during the growing season, for grass the capillary  
rise was equal to 90–150 mm for Aquepts Inceptisols and 60–130 mm for Aquepts  
Histosols. In dry years the capillary supply can be 40–50% of the total supply for Histosols,



but close to zero for some Inceptisols (Kowalik, 2006). Babajimopoulos et al. (2007) found  
40 that under the specific field conditions about 3.6 mm/day of the water in the root zone  
originated from the shallow water table, which amounts to about 18% of the transpired water  
by a maize crop. Fan et al. (2013) analysed the groundwater depth globally and concluded  
that shallow groundwater influences 22 to 32% of global land area and 7 to 17% of this area  
has a water table or its capillary fringe within plant rooting depths, suggesting a widespread  
45 influence of groundwater on crops. This is especially the case in delta areas where high  
population densities occur and agriculture is the predominant landuse.

Wu et al. (2015) showed that capillary rise plays a main role in supplying the vegetation  
throughout the season, hence a strong dependence of vegetation upon groundwater.  
According to Geerts et al. (2008) the contribution of water from capillary rise to the quinoa  
50 production in the Irpani region (Peru), ranges from 8 to 25% of seasonal crop  
evapotranspiration (ET<sub>c</sub>) of quinoa, depending mostly on groundwater table depth and  
amount of rainfall during the rainy season.

In the Pampeana region of Argentina, more than 6 million hectares are subject to the  
influence of a water table that oscillates at depths reached by the roots of the plants (Martini  
55 and Baigorri, 2002). The water contribution from a water table located approximately 1.5 to 2  
m deep can represent up to 30% of the water requirements of soybeans in environments  
representative of the flooding sandy pampas, thus stabilizing the inter-annual variability of  
grain yield (Videla Mensegue et al., 2015).

60 In The Netherlands the average groundwater table is less than 2 meter below the soil  
surface in 85% of the area (De Vries, 2007) where capillary rise will reach the root zone. But  
also in the remaining 15% of the area with deep groundwater levels the effect of capillary  
rise is present, albeit to a more limited extent: in those areas the capillary rise is limited to  
moisture that has percolated to just below the root zone during a wet period, and which is  
65 then drawn back into the root zone as internal recirculation during an ensuing dry period as  
a result of capillary suction.

Wesseling and Feddes (2006) report that in summer when the evapotranspiration demand is  
high, crops partially depend on water supply from soil profile storage and capillary rise from  
70 the groundwater table in the Netherlands. Van der Gaast et al. (2009) applying the method  
of Wesseling (1991) found for the Netherlands a maximum capillary flow of 2 mm/d to the  
root zone in loamy soils where the groundwater level is at 2,5 meter below the soil surface.

Although the contribution of capillary rise to the total water budget can be significant, it is an  
often neglected part of the crop water demand in situations of shallow groundwater levels  
75 (Awan et al., 2014). The capillary properties of a soil strongly depend on soil type. Rijtema  
(1971) estimated that loamy soils have an almost 2 times higher capillary rise than sandy



soils. Integrated approaches are needed to relate water availability to crop yield prognosis (van der Ploeg & Teuling, 2013; Norman, 2013). The driving force for capillary rise is the difference in soil water potential, referred to as heads. There are several models available  
80 that solve these head differences in a numerical way, e.g. HYDRUS (Šimůnek et al., 2008) and SWAP (Feddes et al., 1988, Van Dam et al., 2008).

We applied the integrated model SWAP-WOFOST to solve head differences and crop yield simulations. Kroes and Supit (2011) applied the same model for grassland and recommended further analyses using different boundary conditions. We now apply this  
85 model with different boundary conditions using 45 years of observed weather climate weather and three different crops. For the lower boundary we using different hydrologic conditions that influence the vertical flow. For the soil system itself we will use a wide range of soil physical conditions. The importance of the soil system was already stated by several  
90 authors like Supit (2000). We build on their suggestions and apply the tools for different crops and boundary conditions. Before we applied the model to different boundary conditions we validated it at field scale.

This paper quantifies the effects of capillary rise on crop growth under different conditions of  
95 soil hydrology, soil type and weather. We hypothesize that neglecting capillary rise will result in neglecting a considerable amount of soil moisture that is available for crop growth. We quantify this amount and show the importance of including capillary rise modelling for crop growth modelling.

Our main research questions are: i) What is the impact of capillary rise on crop yield and  
100 downward leaching of water ? ii) What is the impact of soil types on capillary rise ?

## 2. Materials and methods

### 2.1 Modelling approach

105 We applied the coupled SWAP and WOFOST modeling system, using a one day time step. SWAP (Van Dam et al., 2008; Kroes et al., 2009) is a one-dimensional physically based transport model for water, heat and solute in the saturated and unsaturated zone, and includes modules for simulating irrigation practices. The first version of SWAP, called SWATRE, was developed by Feddes et al. (1978). SWAP simulates the unsaturated and  
110 saturated water flow in the upper part of the soil system, using the Richards' equation. Root water extraction and lateral exchange with surface water are accounted for. The soil hydraulics are described by the Mualem–Van Genuchten relations and the potential evapotranspiration is calculated with the Penman–Monteith equation (Allen et al., 1998). Hydraulic heads supplied by a separate regional hydrological model are used to simulate



115 the interaction between the bottom boundary and the groundwater. Drainage and infiltration  
through the lateral boundary account for the flow to the surface water. The surface water  
system is simulated using a simplified, weir controlled, water balance. Note that the surface  
water system in its turn interact with the ground water system. In previous years, SWAP has  
120 been successfully used to study soil-water-atmosphere-plant relationships in many locations  
with various boundary conditions (e.g. Feddes et al., 1988; Bastiaanssen et al., 2007). See  
Van Dam et al. (2008) for an overview. A recent list is available at  
<http://www.swap.alterra.nl>. Eitzinger et al. (2004), Bonfante et al. (2010), Oster et al. (2012),  
Rallo et al. (2012) and Ahuja et al. (2014) amongst others tested the model performance.  
Van Keulen and Wolf (1986) explained the principles and Van Diepen et al. (1989)  
125 presented the first WOFOST version. WOFOST has been adapted and applied in many  
studies (e.g Rötter, 1993; Van Ittersum et al., 2003; de Wit and Van Diepen, 2008; Supit et  
al., 2012; De Wit et al., 2012). In WOFOST the crop assimilation is a function of the  
absorbed radiation and temperature. The assimilation is reduced in case water or nutrient  
stress occur. Subsequently, the maintenance respiration is subtracted and the remaining  
130 assimilates are partitioned over the plant organs (i.e. leaves, stems, roots and storage  
organs). For maize and potatoes the partitioning is development stage dependent. For  
perennial grass however, a constant partitioning factor is assumed. By integrating the  
difference between growth and senescence rates over time, dry weights of various plant  
organs are established.

135 In SWAP-WOFOST, crop assimilation depends on the ambient CO<sub>2</sub> concentration as well  
(see: Kroes and Supit, 2011; Supit et al., 2012). To account for unknown residual stress  
caused by diseases, pests and/or weeds an additional assimilation reduction factor is  
introduced. The rooting density decreases exponential with depth. To withdraw water for  
crop uptake from deeper soil layers if the upper part of the soil is very dry, a form of  
140 compensation root uptake (hydraulic lift) is used (Jarvis, 2011). The increasing atmospheric  
CO<sub>2</sub> concentrations during relatively long historical simulation periods (>20years) is  
accounted for.

## 2.2 Case studies

145 SWAP-WOFOST is validated using results of 7 case studies at 6 locations in The  
Netherlands (Figure 1) where grassland, maize and potatoes is grown and using  
hydrological, soil and crop observations. The main characteristics of the 7 cases are  
summarized in Table 1. The soil texture ranged from sand to clay. The observations  
included parameters, such as groundwater levels, yields and in some cases soil moisture  
150 contents, soil pressure head and evapotranspiration. The weather data were collected from  
nearby weather stations or from onsite measurements. Observations for case 1 and 3



(location De Marke) were available for a period of 22 years (1992-2013) from one field where grassland and maize was grown for respectively 7 and 15 years. Calibration resulted in parameter values for drought and management which are given in table 2. Planting and harvest dates were given. Oxygen stress was parameterised identic to 155 Hack et al. (2016). For all cases a so-called management factor was used to close the gap between observed and actual yield. Management factors were relatively high because the cases describe field studies that generally have a good management with limited yield losses. The input crop parameters for maize only differed with respect to the management 160 factor (MF) which ranges from 0.85-0.95. For potatoes the input crop parameters were kept the same for all 3 cases. Maximum rooting depth for grassland, maize and potatoes were respectively 40, 100 and 50 cm. Soil water conditions were different for all locations and boundary conditions varied, depending on local situation and available data (Table 3). In most cases a Cauchy bottom 165 boundary condition was applied using a hydraulic head based on piezometer observations from the Dutch Geological Survey (<https://www.dinoloket.nl/>). Observed groundwater levels were used as lower boundary condition for Borgerswold (crop: potato). In 2 cases a lateral boundary condition was applied with drainage to a surface water system (Table 3). The simulation results were analysed using an R-package (Bigiarini, 2013) and the statistics are 170 presented in Table 5.

### 2.3 Soil crop experiment

To analyse the impact of soil type on capillary rise we used 72 soils derived from a national soil data base (Wösten et al., 2013a). 175 These 72 soils were aggregated from 315 soil units of the 1:50000 Dutch Soil Map (Figure 2) using hydraulic clustering methods and considering following properties: maximum groundwater depth, saturation deficit between a certain depth and the soil surface, transmissivity for horizontal water flow, resistance for vertical water flow, the available water in the root zone, etc. These soil hydraulic properties were subsequently used as input in 180 SWAP-WOFOST. We calculated capillary rise as the upward flux at the bottom of the root zone. Root zone extension as a result of root growth is accounted for. For each soil we applied 3 hydrological boundary conditions: a) free drainage without capillary rise, b) free drainage with capillary rise and c) an average fluctuating groundwater level (Table 4). 185 Conditions a) and b) were simulated using free drainage at the bottom of the soil profile. Since the applied model for water flow solves the Richard equation in a numerical way this implies that shortages of water below the root zone will automatically be compensated by using stored water or water from deeper soil layers and thereby generating an upward capillary flux. Since we wanted to analyse the impact of capillary rise, we adjusted the model



190 SWAP in such a way that capillary rise could be switched on and off. Without capillary rise  
the hydraulic conductivity was minimized for upward flow in the soil layers just below the  
root zone. Condition c) approaches an average groundwater level by using conditions that  
were derived from a national study (Van Bakel et al., 2007). This national study used  
simulation units with are unique in groundwater dynamics and land use. We selected three  
195 large simulation units with long term average groundwater levels between 40 and 120 cm  
below the soil surface (Dutch groundwater class IV, units 2245, 3859 and 621 for grassland,  
maize and potato, covering respectively 1806, 794 and 58102 ha, data from Van Bakel et  
al., 2007). See also the supplementary material of Kroes and Supit (2011) for an additional  
explanation of the study from Van Bakel et al (2007).

200 The crop parameters were kept the same as for the case studies, with a few exceptions: i)  
for grassland an average management factor of 0.9 was used, ii) timing of grass-mowing  
was done when a dry matter threshold of 4200 kg/ha was exceeded, iii) for maize and  
potatoes the planting dates were respectively 25-Oct and 15-Oct.

205 The 3 crops and 3 lower boundary conditions resulted in 9 combinations. Each combination  
was simulated with 72 soils for a period of 45 years (1971-2015) with meteorological  
conditions from the station De Bilt (KNMI, 2016). In a subsequent analysis we grouped the  
results of these 72 soils to 5 main soil groups clay, loam, peat, peat-moor and sand to be  
able to analyse impact on grouped soil types.

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### 3 Results

#### 3.1 Case studies

215 The first 2 case studies are from one location (De Marke) where a grassland-maize rotation  
was practised. The results show that the hydrological conditions (Figure 3 and Table 5) were  
simulated accurately for those years for which observed data was available (1991-1995).  
From 1995-1997 the groundwater levels drop as a result of low precipitation (about 700  
mm/year). The fall of the year 1998 shows rising groundwater levels that correspond well  
220 with very wet conditions at that moment. The simulated grassland yields are overestimated  
and the simulated maize yields are underestimated however they are well within acceptable  
ranges (Figure 4 and Table 5).

For the maize case studies 3 and 4 groundwater level and soil moisture are well simulated  
(Figure 5 and Table 5). The simulated maize yields (Figure 6, Table 5) are less acceptable  
225 for case 3 as is indicated by a zero or negative Nash-Sutcliffe efficiency (NS) which  
suggests that the observed mean is a better predictor than the model. In 1976, a very dry



year, the soil hydrology dynamics and the resulting yield were well captured. The yield of case study 4 has a small bias of about  $300 \text{ kg}\cdot\text{ha}^{-1} \text{ DM}$  between observed and simulated. The simulated hydrological conditions for the 3 fields of the potato-cases 6 and 7 show a good fit with the observed (Figure 7, Tabel 5). The simulated yields (Figure 8) show the largest deviation from the observed for case 5 (location Borgerswold). The more recent experiments of cases studies 6 and 7 (Rusthoeve and Vredepeel) show differences between simulated and observed yields of respectively  $1400$  and  $300 \text{ kg}\cdot\text{ha}^{-1} \text{ DM}$ . These case studies unfortunately cover only one year. The Rusthoeve performs less due to the complex situation in the subsoil with drainage conditions that require more observations to improve the simulations.

Even though some yields are not accurate enough to satisfy statistical criteria for good model performance, we think that the global overall picture is acceptable. With more field information and calibration a better result may be achieved but we think that especially the good hydrological model performance allows an application at a larger scale with different hydrological boundary conditions.

Before carrying out an analyses at a larger scale we analysed the impact of capillary rise for the case studies. To analyse this impact we carried out new simulations without capillary rise towards the root zone for a grassland, maize and potato case. Results of these 3 cases are given in Table 6 for the situation with and without capillary rise. This shows that suppressing capillary rise lowers yields by 5, 2 and 22% respectively for a grassland, maize and potato case. The downward leaching of water of water from the root zone was reduced with respectively 3, 5 and 94% (Table 6). A larger scale experiment was then carried out to analyse this impact for different soil conditions.

### 3.2 Soil crop experiment

The 3 crops from the case studies were simulated with 72 soils from the national database using 3 different lower boundary conditions and 45 years with weather from 1970-2015. The capillary rise shows large variations among crops and bottom boundary conditions (Tables 7). The highest values for capillary rise were found for average groundwater conditions (Ave) with median for grassland, maize and potatoes of respectively 191, 79 and 115 mm/year. Differences among the lower boundary conditions are caused by differences in weather, growing season, dynamic position of the root zone and demand root water uptake.

Even in free drainage situations ( $\text{FD}_{\text{cr}}$ ) the capillary rise caused by the internal recirculation can be considerable, ranging from 20 – 78 mm long term average (Table 7). Results of simulated capillary rise of 45 years weather, 72 soils and 2 lower boundary conditions ( $\text{FB}_{\text{cr}}$  and Ave) are presented in Figures 9, 10, 11 (upper part) and for 5 grouped



soils (Figure 9, 10, 11, lower part). The differences in capillary rise between soils are  
265 relatively small compared to the differences among years and within one grouped soil type.  
For grassland the variation of capillary rise within years is larger than those of maize and  
potatoes because of the differences in rooting depth which is shortest for grassland and the  
continuous soil cover of grassland. The downward leaching of grassland is highest for  
grassland for the same reasons (Table 7).  
270 The highest median yields are simulated when average groundwater situations including  
capillary rise are considered (Table 7, Ave)

Results of the simulations with 3 different lower boundary conditions ( $FD_{nc}$ ,  $FD_{cr}$  and Ave)  
are also compared by subtraction. Results of the subtraction are given for capillary rise  
275 (Figures 12-14) and for yields (Figures 15-17) and summarized (Table 8). The elimination of  
capillary rise to the root zone in free drainage conditions reduces grassland, maize and  
potatoes yields with respectively 13, 1 and 8 % (Table 8). A comparison between situations  
with average groundwater levels and free drainage situations shows a similar yield  
reduction: respectively. 13, 3 and 8 %. When one compares situations with average  
280 groundwater levels with free drainage conditions without capillary rice yield-reductions of  
grassland, maize and potatoes are respectively 25, 4 and 15 % (Table 8) or respectively  
about 3.2, 0.5 and 1.6  $\text{ton}\cdot\text{ha}^{-1}$  Dry Matter (Table 7).

The impact of neglecting capillary rise on downward leaching flux, the groundwater  
recharge, is highest for potatoes and lowest for maize. For grassland, maize and potatoes  
285 differences were calculated of respectively 17, -6 and 46% (Table 8) or 64, -3 and 34 mm  
(Table 7). Low values for maize are caused by the deeper rooting systems which reduces  
the impact of relative differences because the groundwater level will be close to the bottom  
of the root zone.

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#### 4. Discussion

The case studies and soil-crop experiments in this paper clearly state the importance of  
capillary rise for simulations regarding impact of groundwater on crop yields. This impact is  
295 clearly present in situations where a groundwater level is present (85% of NL) but also in  
other situations the impact of capillary rise is considerable. According to our simulation  
experiments the impact of capillary rise on yield is about 50% due to internal recycling as  
occurs in free drainage conditions, another 50% of yield reduction is caused by direct  
upward capillary flow.

300 According to Supit & Van Der Goot (1999) water limited results are seldom chosen as  
predictors which suggests a minor contribution of the soil. Our analysis shows that the soil is





important. Several items can be addressed that have a high impact on capillary rise. The depth of the groundwater directly influences the size of the gap that the capillary flux has to bridge to reach to be able to reach the root zone. The vertical soil profile is important  
305 because differences in hydraulic properties influence vertical water flow.

Modelling concepts should consider upward capillary flow, variable rooting depth and a dynamic interactions between soil water and crop growth.

Precipitation, soil texture and water table depth jointly affected the amount of groundwater recharge and time-lag between water input and groundwater recharge (Ma et al., 2015).

310 We quantified some of these issues, but several items remain, such as the impact of rooting depth on crop yield and transpiration. Also management of soil and water, items like ploughing and irrigation, are not considered. Furthermore pattern of rooting needs a more detailed analyses; we applied an exponential decrease of rooting patterns and compensation of root uptake as hydraulic lift according to Jarvis (2011) but the actual  
315 rooting distributions is still simple and requires a more detailed analyses. Another item we neglected is the preferential flow of water by the occurrence of non-capillary sized macropores (Bouma, 1961, Feddes, 1988), which is relevant in especially clay soils. Hysteresis is also not considered. An additional analysis of these issues is recommended, preferably about the impact of different rooting patterns on capillary rise.  
320

### 5. Conclusions

We quantified the impact of capillary rise in layered soils on crop yields of grassland, maize and potatoes. We compared situations with average groundwater levels with free drainage  
325 conditions with and without capillary rise. The largest difference was found when one compares situations with average groundwater levels with free drainage conditions without capillary rise. From these differences one may conclude that neglecting capillary rise has a large impact on simulated yields and water balance calculations especially in regions where shallow groundwater occurs.

330 The comparison shows yield-reductions of grassland, maize and potatoes of respectively 25, 4 and 15 % or respectively 3.2, 0.5 and 1.6 ton Dry Matter per ha (Table 7). Reduction on the downward leaching water flux, the groundwater recharge can be considerable; for grassland and potatoes the reduction is 17 and 46% or 70 and 34 mm.

335 Improved modelling of soil water flow should consider capillary rise of soil water which will results in improved yield and downward leaching predictions.

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## 515 Tables

*Table 1 Main characteristics of case studies used to verify setup of model combination SWAP-WOFOST*

Case study	Crop	Location	Period	Soil type	Observations*	Reference
1	Grassland	De Marke	1995-1996, 2005-2008, 2013	dry sandy soil	Gwl, Yield, Theta20cm	Hack et al. (1996)
2	Silage maize	De Marke	1992-1994, 1997-2003, 2009-2012	dry sandy soil	Gwl, Yield, Theta20cm	Hack et al. (1996)
3	Silage maize	Cranendonck	1974-1982	Cumulic Anthrosol	Gwl, Yield	Schröder (1985)
4	Silage maize	Dijkgraaf	2007	Umbric Gleysol	Gwl, Yield, ET,Theta20cm	Elbers et al. (2010)
5	Potato	Borgerswold	1992, 1994	Sandy loam	Gwl, Yield	Dijkstra et al., 1995
6	Potato	Rusthoeve	2013	lichte kleibodem	Gwl, Yield, Qdrain	Van Den Brande (2013)
7	Potato	Vredepeel	2002	Sandy loam	Gwl, Yield	De Vos et al., 2006

\* Gwl = Ground water level, Yield = Actual Yield as Dry Matter of Harvested product, Theta20cm= Soil moisture content at a depth of 20cm below surface, Qdrain = drainage from field to surface water via tube drains, ET = Evapotranspiration measured via Eddy Correlation method.

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*Table 2 Important crop parameter for the 3 crops in the 7 case studies*

Case study	Crop	Drought stress*	MF**	Maximum depth of root zone (cm)
1	Grassland	h3h = -200.0 cm h3l = -800.0 cm h4 = -8000.0 cm	0.8	40
2	Silage	h3h = -400.0 cm	0.90	100
3	maize	h3l = -500.0 cm	0.85	
4		h4 = -10000.0 cm	0.95	
5	Potato	h3h = -300.0 cm	0.8	50
6		h3l = -500.0 cm	0.8	
7		h4 = -10000.0 cm	0.8	

\* h3h = h below which water uptake reduction starts at high Tpot; h3l = h below which water uptake red. starts at low Tpot; h4 = No water extraction at lower pressure heads; \*\* MF=Management Factor to account for imperfect management

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*Table 3 Boundary conditions for vertical water flow at 6 locations*

Case study	Location	Bottom boundary condition*	Lateral Boundary
1,2	De Marke	Cauchy	No Drainage
3	Cranendonck	Cauchy	No Drainage
4	Dijkgraaf	Cauchy	No Drainage
5	Borgerswold	Observed groundwater	No Drainage
6	Rusthoeve	Cauchy	Drainage, tube at -90 cm
7	Vredepeel	Closed	Drainage, ditch at -100 cm

\* The Cauchy bottom boundary condition uses a hydraulic head based on piezometer observations from an open data portal (see text)





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**Table 4 Characteristics of 3 different hydrological bottom boundary conditions**

Condition	Bottom Boundary	Description
a	FD <sub>nc</sub>	Free drainage without capillary rise towards root zone
b	FD <sub>cr</sub>	Free drainage with capillary rise towards root zone
c	Ave	An average fluctuating groundwater level

**Table 5 Results of Case studies: simulated and observed values**

Case study	Name	unit	Simulated mean	Observed mean	ME <sup>1</sup>	RMSE <sup>2</sup>	NS <sup>3</sup>	d <sup>4</sup>	n <sup>5</sup>
1.									
Grassland	Yield	kg/ha/yr DM	11728	11049	679	1343	0.6	0.9	7
	Gwl	m-soil	-1.34	-1.30	-0.04	0.46	0.3	0.9	77
	Theta20cm	m3/m3	0.27	0.27	0.01	0.06	0.5	0.9	43
2. Maize	Yield	kg/ha/yr DM	11564	11850	-286	2825	-3.2	0.4	14
3. Maize	Yield	kg/ha/yr DM	14054	13788	266	2587	-1.1	0.7	9
	Gwl	m-soil	-1.42	-1.36	-0.05	0.25	0.4	0.9	61
4. Maize	Yield	kg/ha/yr DM	15974	16306	-332				1
	LAI	m2/m2	2.1	2.5	-0.3	0.6	0.7	0.9	10
	ETact	mm/yr	1.4	1.9	-0.6	0.9	0.5	0.9	232
	Gwl	m-soil	-1.03	-1.07	0.03	0.06	0.9	1.0	112
5. Potato	Yield	kg/ha/yr DM	10532	9246	1286	1350	-31.9	0.3	2
	Gwl	m-soil	-1.10	-1.10	0.00	0.03	1.0	1.0	123
6. Potato	Yield	kg/ha/yr DM	10019	8610	-1409				1
	Gwl	m-soil	-1.07	-1.10	0.02	0.19	0.6	0.9	887
	qDrain	mm	1.1	0.6	0.4	1.4	0.4	0.8	1084
7. Potato	Yield	kg/ha/yr DM	11071	11359	-288				1
	Gwl	m-soil	-1.04	-1.07	0.03	0.11	0.8	0.9	353

535 <sup>1</sup>ME: Mean Error between simulated (sim) and observed (obs), in the same units of sim and obs, with treatment of missing values. A smaller value indicates better model performance

<sup>2</sup>RMSE: Root Mean Square Error between sim and obs, in the same units of sim and obs, with treatment of missing values. RMSE gives the standard deviation of the model prediction error. A smaller value indicates better model performance.

540 <sup>3</sup>NS: Nash-Sutcliffe efficiencies range from -Inf to 1. Essentially, the closer to 1, the more accurate the model is. NS = 1, corresponds to a perfect match of modelled to the observed data. NS = 0, indicates that the model predictions are as accurate as the mean of the observed data. -Inf < NS < 0, indicates that the observed mean is better predictor than the model.

545 <sup>4</sup>d: The Index of Agreement (d) developed by as a standardized measure of the degree of model prediction error and varies between 0 and 1. A value of 1 indicates a perfect match, and 0 indicates no agreement at all. The index of agreement can detect additive and proportional differences in the observed and simulated means and variances; however, it is overly sensitive to extreme values due to the squared differences.;

<sup>5</sup>n: the number of values used with the previous 4 statistical criteria to compare simulated and observed results.



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*Table 6 Results of case studies: values and differences of yield, capillary rise and leaching fluxes, resulting from simulations with and without capillary rise*

Case study***	Model Result	Condition		Differences		Differences (%) 100*(A-B)/A
		A*	B**	A-B	Unit	
1. Grassland	Y <sub>act</sub>	11904	11353	551	kg/ha/yr DM	5
	q <sub>caprise</sub>	28	0	28	mm/yr	100
	q <sub>leaching</sub>	314	306	9	mm/yr	3
2. Maize	Y <sub>act</sub>	12504	12257	247	kg/ha/yr DM	2
	q <sub>caprise</sub>	8	0	8	mm/yr	100
	q <sub>leaching</sub>	81	77	4	mm/yr	5
7. Potato	Y <sub>act</sub>	11071	8665	2406	kg/ha/yr DM	22
	q <sub>caprise</sub>	105	0	105	mm/yr	100
	q <sub>leaching</sub>	15	1	14	mm/yr	94

\* Condition A has actual bottom boundary conditions (according to table 3); \*\* Condition B has actual bottom boundary conditions (table 3) but without capillary rise to root zone;

\*\*\* Cases studies 1 and 2 were simulated for limited periods of respectively 2005-2008 and 1991-1994 to have a continuous sequence of years, Case study 7 was simulated for one year

560 *Table 7 Results of soil crop experiments: values of 5 parameters resulting from 3 different hydrological bottom boundary conditions (FD<sub>nc</sub>, FD<sub>cr</sub> and Average)*

Crop	Model Result	FD <sub>nc</sub>	FD <sub>cr</sub>	Average	Unit
	q <sub>caprise</sub>	0	78	191	mm
	q <sub>leaching</sub>	318	340	382	mm
Maize	Y <sub>act</sub>	12127	12239	12626	kg/ha/yr DM
	q <sub>caprise</sub>	0	20	79	mm
	q <sub>leaching</sub>	48	53	45	mm
Potato	Y <sub>act</sub>	8764	9482	10342	kg/ha/yr DM
	q <sub>caprise</sub>	0	44	115	mm
	q <sub>leaching</sub>	39	50	73	mm



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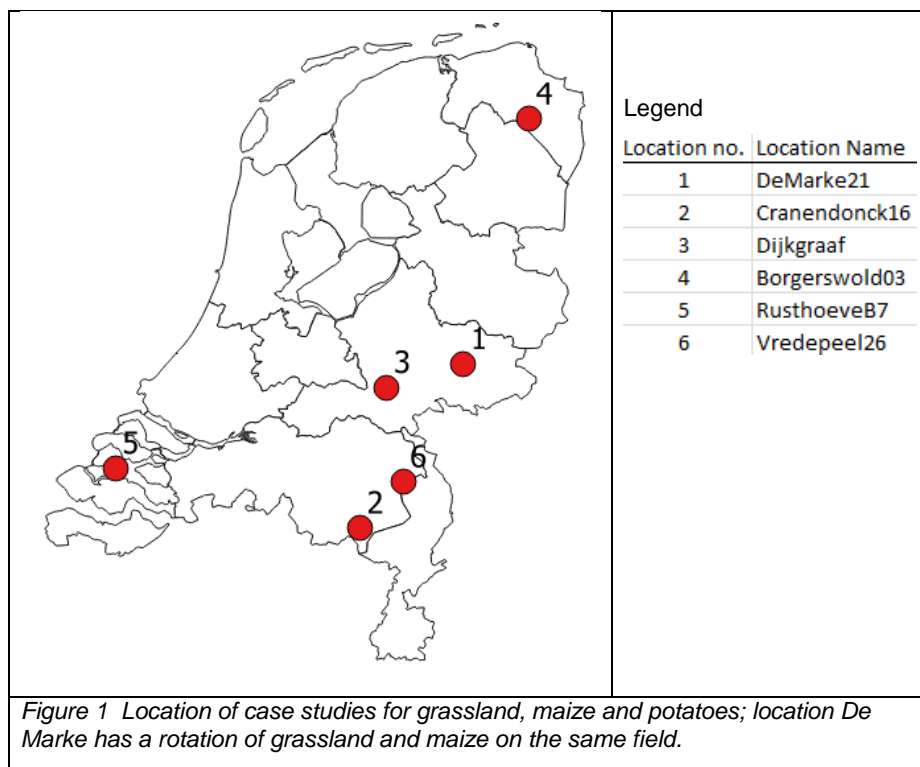
*Table 8 Results of soil crop experiments: Differences (%) between results from hydrological bottom boundary conditions ( $FD_{nc}$ ,  $FD_{cr}$  and Ave)*

Crop	Model Result	Differences (%)		
		$FD_{cr} - FD_{nc}$	Ave- $FD_{cr}$	Ave- $FD_{nc}$
Grassland	$Y_{act}$	13	13	25
	$q_{caprise}$			
	$q_{leaching}$	6	11	17
Maize	$Y_{act}$	1	3	4
	$q_{caprise}$			
	$q_{leaching}$	10	-18	-6
Potato	$Y_{act}$	8	8	15
	$q_{caprise}$			
	$q_{leaching}$	22	32	46

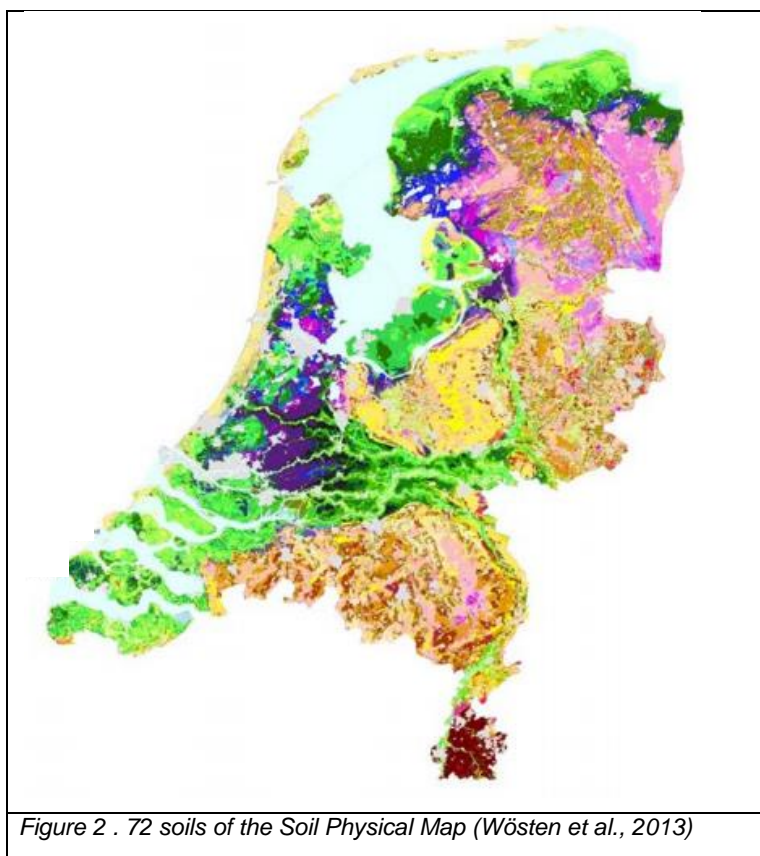
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## Figures

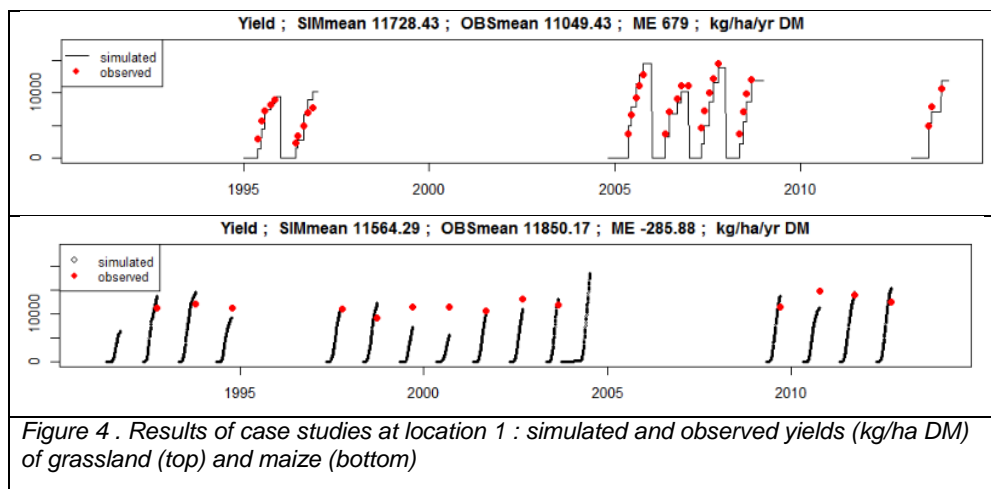
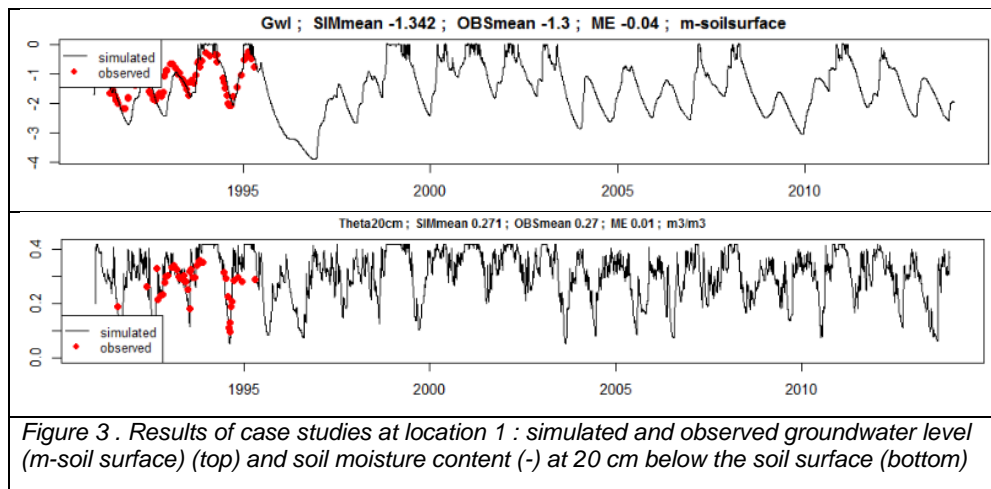


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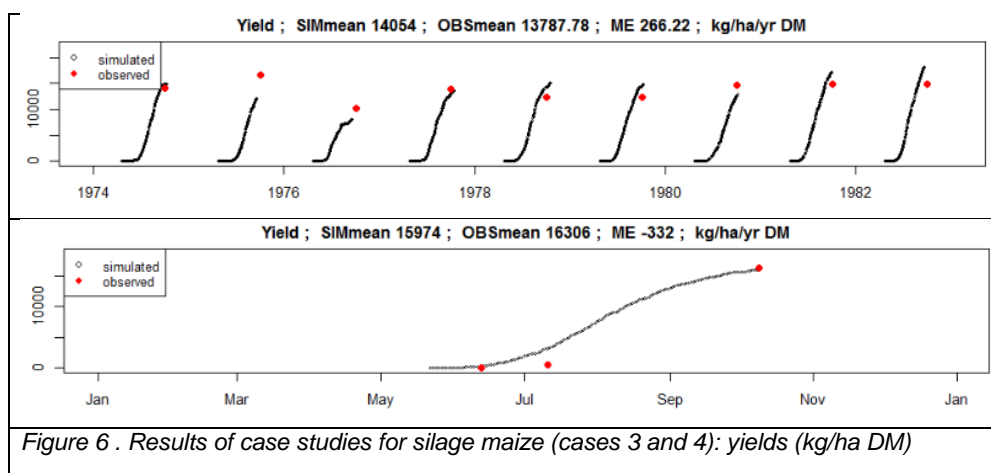
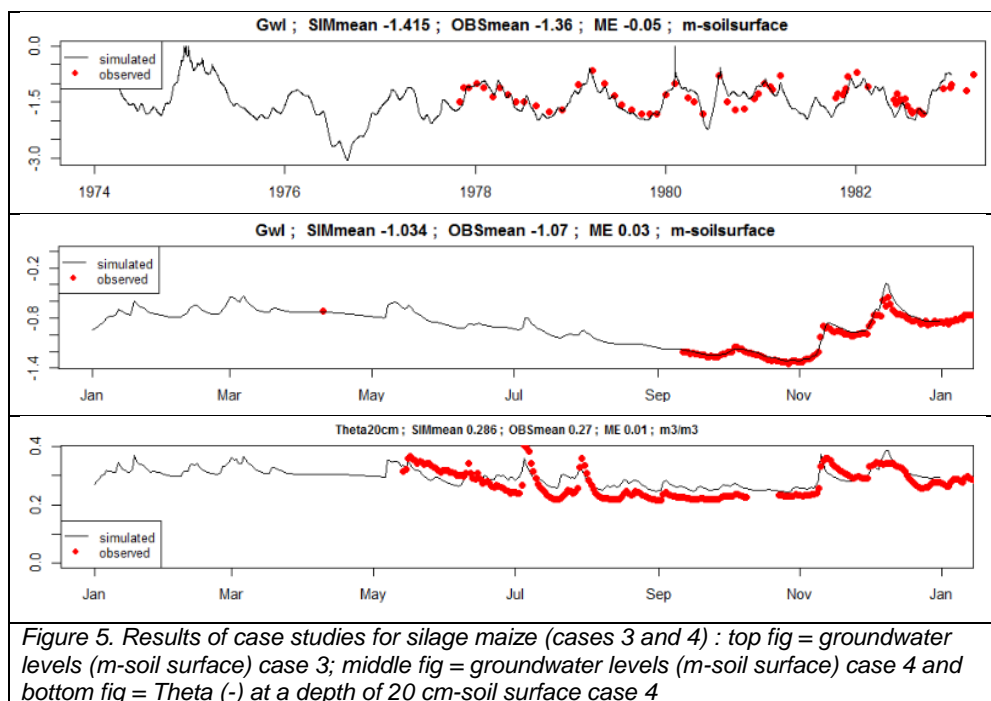


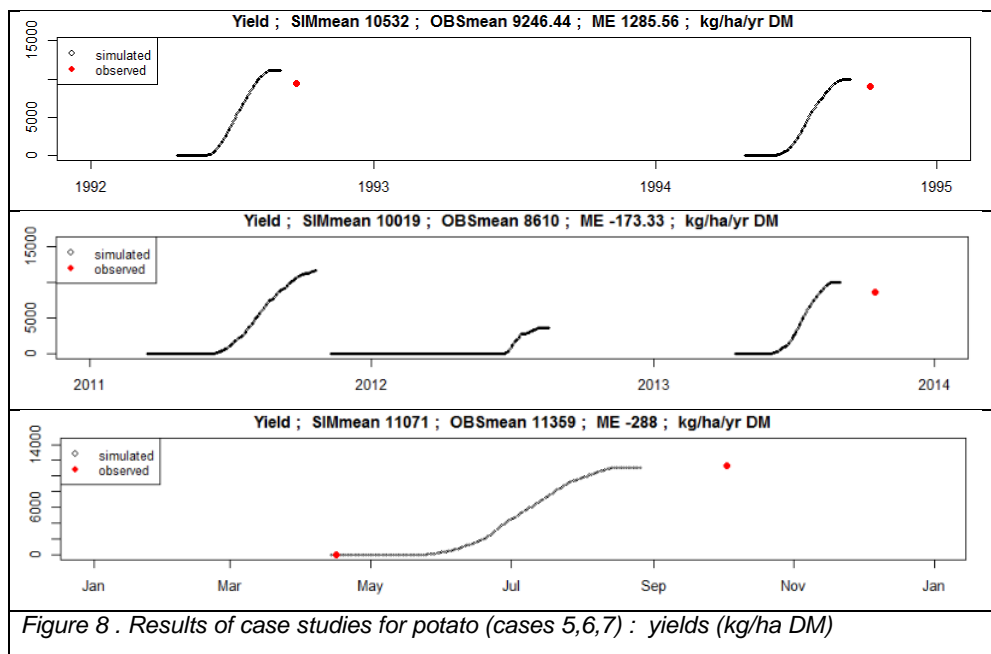
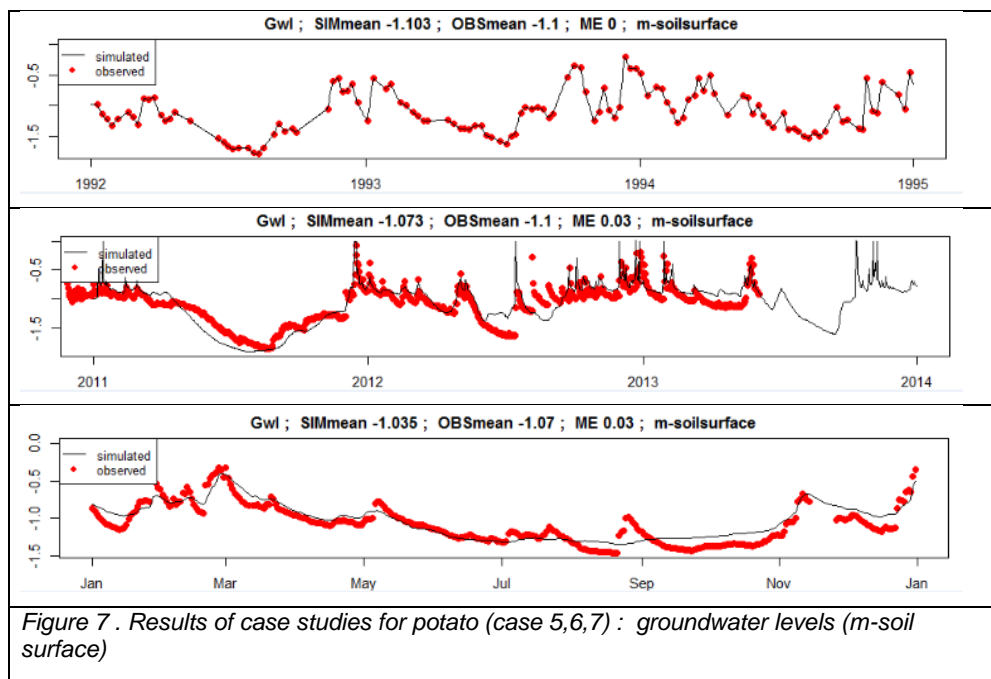


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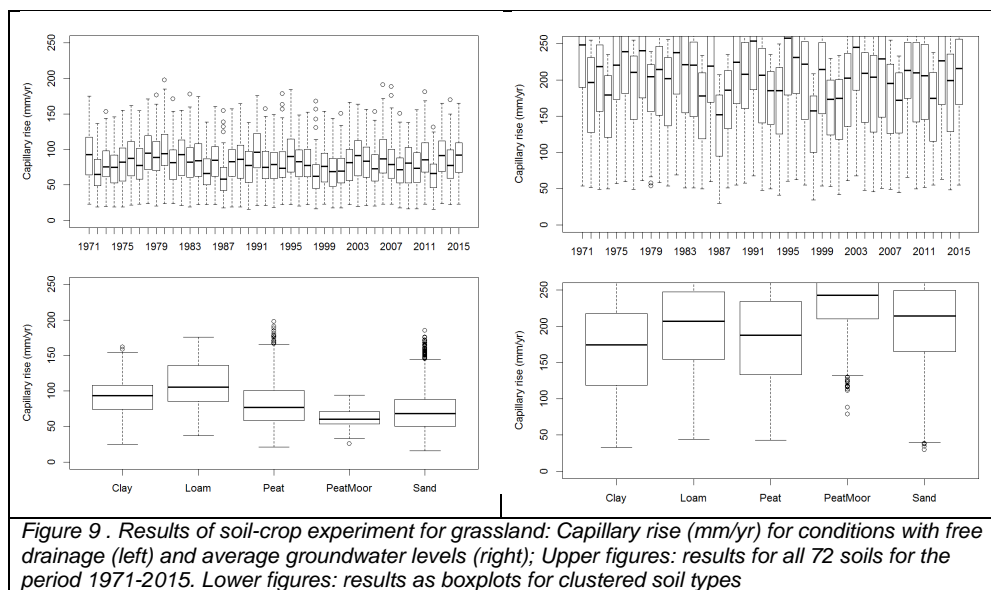


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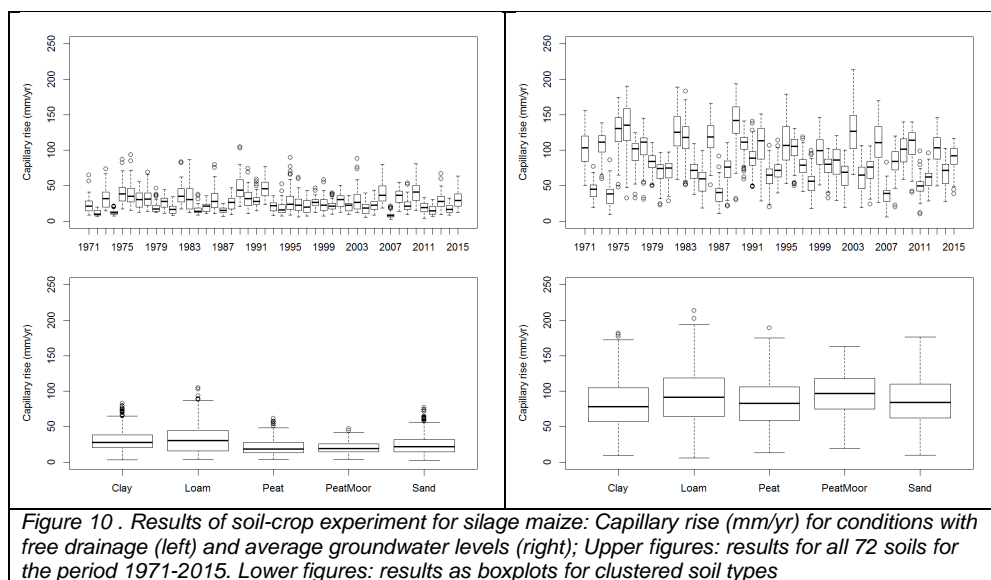




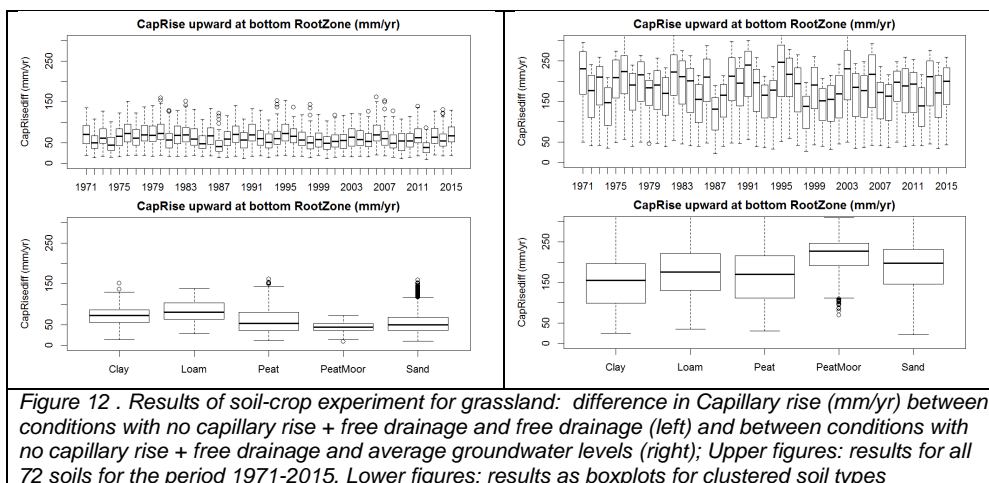
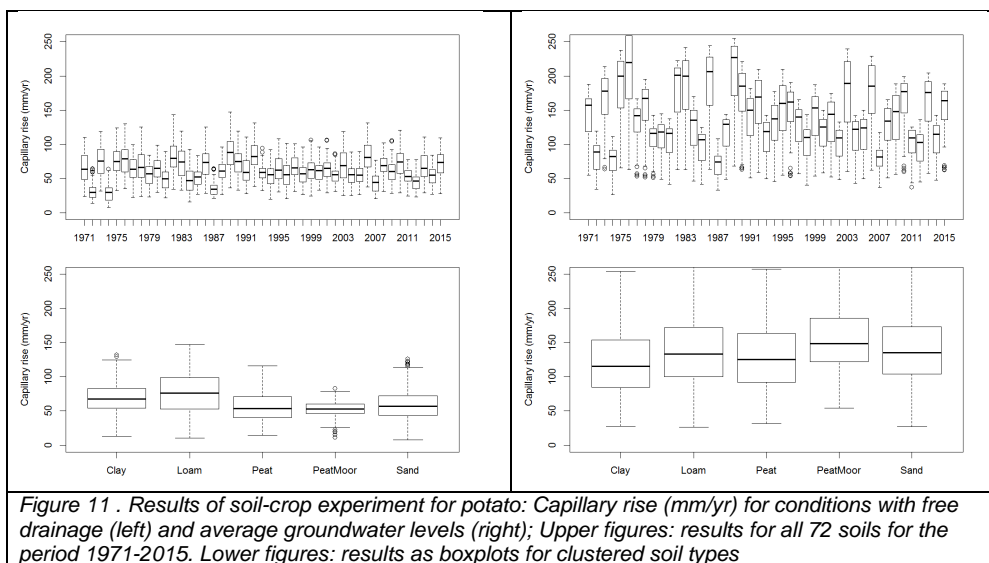




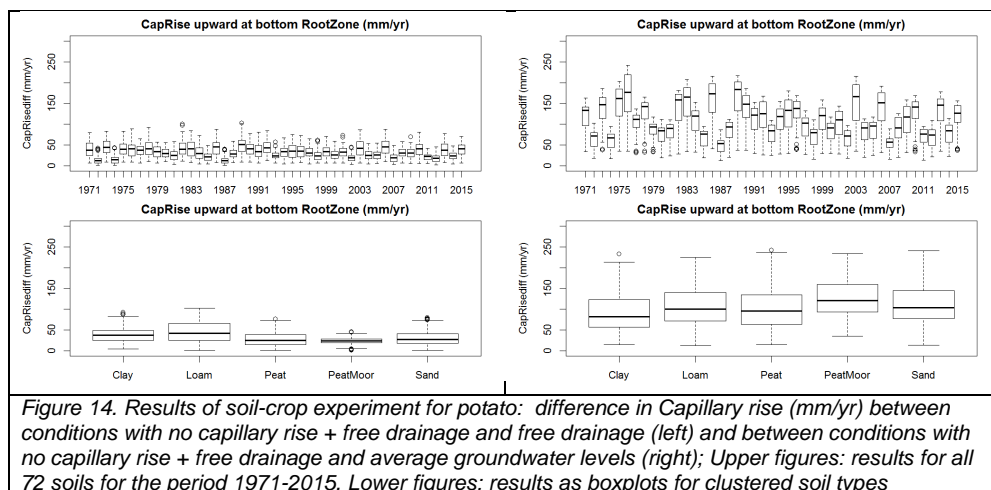
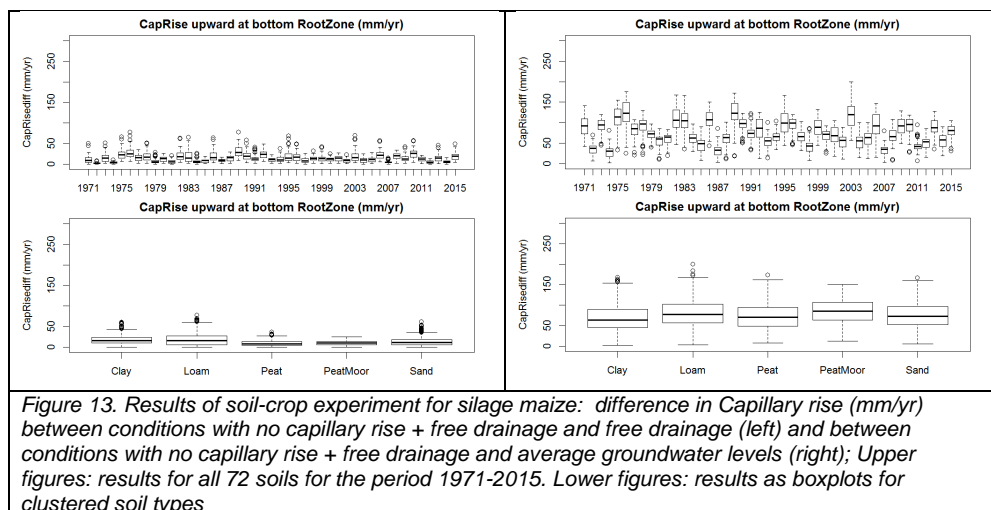
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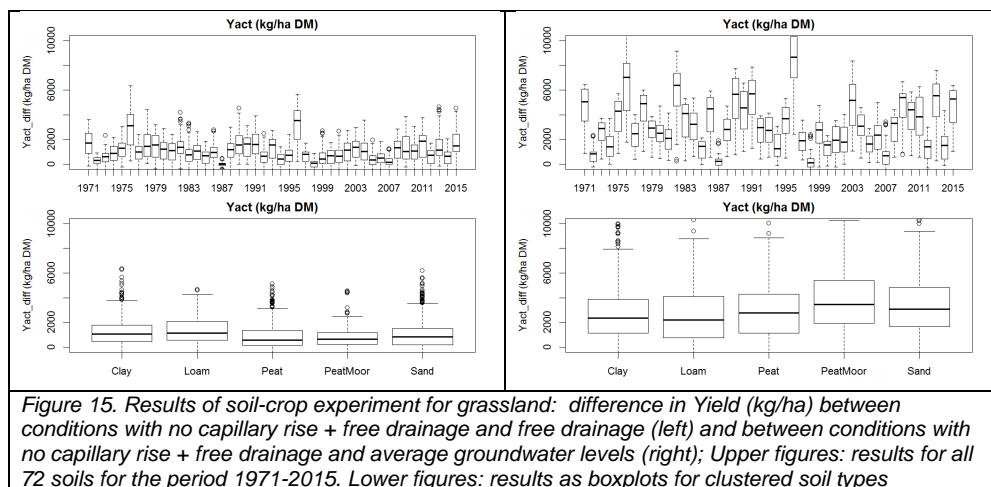
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