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Supplemental Online Material for

## Ocean and Coastal Acidification off New England and Nova Scotia

By D.K. Gledhill, M.M. White, J. Salisbury, H. Thomas, I. Mlsna, M. Liebman, B. Mook, J. Grear, A.C. Candelmo, R.C. Chambers, C.J. Gobler, C.W. Hunt, A.L. King, N.N. Price, S.R. Signorini, E. Stancioff, C. Stymiest, R.A. Wahle, J.D. Waller, N.D. Rebeck, Z.A. Wang, T.L. Capson, J.R. Morrison, S.R. Cooley, and S.C. Doney

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## Supplemental Methods

### Primary Controls on NECAN Seasonal Dynamics

Time-series analyses were evaluated at each of the locations denoted in Figure 4 using actual discrete time-series data where available (western Gulf of Maine, Long Island Sound) or by extracting estimates using the gridded data product produced by Signorini et al. (2013).

### Signorini-Based Estimates

Figure 4 denotes the estimated individual effects of CO<sub>2</sub> solubility (SOL), air-sea CO<sub>2</sub> flux (AS), mixing (MIX) and biologic activity (BIO) on  $\Omega_a$  at six New England/Nova Scotia locations. . . These estimates were derived from the calculations made using modeled 12-month climatologies of sea surface temperature (SST), sea surface salinity (SSS), total alkalinity (TA),  $p\text{CO}_{2,\text{air-sea}}$  CO<sub>2</sub> flux (FLUX), and surface mixed-layer depth (MLD). BIO was determined by combining December TA,  $p\text{CO}_{2,\text{sea}}$ , SST, and SSS to derive dissolved inorganic carbon ( $\text{DIC}_{\text{TA-}p\text{CO}_2}$ ) using the CO2SYS program (Lewis and Wallace, 1998;  $K_1$  and  $K_2$  of Millero, 2010;  $\text{KSO}_4$  of Dickson, 1990; TB [total boron] of Uppstrom, 1974), then deriving  $p\text{CO}_{2,\text{sea,TA-DIC}}$  and  $\Omega_{\text{arag,TA-DIC}}$  at monthly SST and SSS using December TA and  $\text{DIC}_{\text{TA-}p\text{CO}_2}$ . AS was determined by adding (for release of CO<sub>2</sub> to the atmosphere) or removing (for uptake of CO<sub>2</sub> from the atmosphere) the DIC (dissolved inorganic carbon) represented by each monthly FLUX out of or into the MLD ( $\text{FLUX}_{\text{DIC}}$ ) to the DIC calculated for the previous month, then deriving  $p\text{CO}_{2,\text{sea}}$  and  $\Omega_{\text{arag}}$  from monthly TA and  $\text{DIC}+\text{FLUX}_{\text{DIC}}$ . AS MLD depth was held to at least 5 m, which is was arbitrarily chosen as a logical minimum depth over which CO<sub>2</sub> is added or removed. We note that during the strongly stratified summer season, 5 m can be 1–3 m less than the pycnocline. MIX was determined using the Gulf of Maine TA-SSS and DIC-SSS using regressions taken

from Table 1 of Wang et al. (2013) using a TA:DIC ratio of 1:07  $TA = 37.3 \cdot SSS + 998$ ;  $DIC = 34.6 \cdot SSS + 933$ . Monthly TA and DIC were calculated according to these regressions ( $TA_{Wang}$  and  $DIC_{Wang}$ , respectively), with changes from month n-1 to month n calculated as  $\Delta TA_{Wang-n} = TA_{Wang-n} - TA_{Wang-n-1}$  and  $\Delta DIC_{Wang-n} = DIC_{Wang-n} - DIC_{Wang-n-1}$ . Then MIX  $pCO_2$  and  $\Omega_a$  were derived from monthly SSS together with TA and DIC at month n:  $TA_n = TA_{n=0} + \Delta TA_{Wang-n}$  and  $DIC_n = DIC_{n=0} + \Delta DIC_{Wang-n}$ . BIO was calculated as the residual between the climatologies of  $pCO_2$  and  $\Omega_a$  and the SOL, AS and MIX terms:  $BIO_{pCO_2} = pCO_2 - [(SOL - pCO_2) + (AS - pCO_2) + (MIX - pCO_2)]$ .

### Western Gulf of Maine Data

The calculated individual effects on  $pCO_{2,sw}$  and  $\Omega_a$  of changes in  $CO_2$  solubility (SOL), air-sea  $CO_2$  flux (AS), and mixing (MIX) at UNH Buoy D were calculated as above. SSS, SST,  $pCO_2$ , and FLUX source data were monthly climatologies derived from 2006–2014 buoy observations. MLD source data was a monthly climatology derived from 2004–2014 shipboard salinity and temperature profiles ( $n = 151$ ) in the region of UNH Buoy D, again limited to 5 m or deeper. TA was derived from salinity according to a locally derived regression:  $TA = (SSS \cdot 52.24) + 476.3$  (unpublished data from author Joe Salisbury).

### Long Island Sound Data

The calculated individual effects on  $pCO_2$  and  $\Omega_{arag}$  of changes in  $CO_2$  solubility (SOL), air-sea  $CO_2$  flux (AS), and mixing (MIX) in Long Island Sound (LIS) were calculated in the *Signorini-based estimate*. SSS, SST, and TA source data were monthly climatologies. A monthly climatology of pH (NBS scale), paired with TA, SSS, and SST was used with CO2SYS as above to generate a monthly  $pCO_2$  climatology. FLUX was calculated from  $pCO_2$ , atmospheric  $pCO_2$

at Mauna Loa (398  $\mu\text{atm}$ ), and monthly winds, according to the k660 parameterization of Ho et al. (2006). MLD for LIS was from the same model employed in the *Signorini-based estimate* at location 41.25°N –71.25°W, and again limited to 5 m or deeper.

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